

3. SUMMARY OF WESTERN NORTH PACIFIC AND NORTH INDIAN OCEAN TROPICAL CYCLONES

3.1 GENERAL

For the western North Pacific 1990 became the busiest in JTWC's history - 794 warnings were issued on 32 tropical cyclones (Table 3-1). This was slightly more than the climatological mean of 31 tropical cyclones noted in Table 3-2. The North Indian Ocean was moderately active with 4 tropical cyclones which is just below the average of five per year. During the year a record 841 warnings were issued on 36 tropical cyclones in the Northern Hemisphere. A chronology of the activity is provided in Figure 3-1.

In the western North Pacific, JTWC was in warning status 165 days compared to 154 in 1989 and 114 in 1988. Again, considering only

the western North Pacific, there were 54 days when the Center issued warnings on two cyclones and 3 days when it warned on three cyclones (Table 3-3). There were no days with warnings were issued on four or more tropical cyclones at once. When the North Indian Ocean is included in the total, there were 180 days with warnings on one cyclone and 10 days with warnings on two. Thirty-three initial Tropical Cyclone Formation Alerts were issued on western North Pacific tropical disturbances (Table 3-4) and 8 on disturbances in the North Indian Ocean. Alerts preceded warnings on all significant tropical cyclones in the western North Pacific and North Indian Oceans with the exception of Tropical Depression 04W.

TABLE 3-1

**NORTHWEST PACIFIC OCEAN
SIGNIFICANT TROPICAL CYCLONES
FOR 1990**

<u>TROPICAL CYCLONE</u>	<u>PERIOD OF WARNING</u>	<u>NUMBER OF WARNINGS ISSUED</u>	<u>MAXIMUM SURFACE WINDS KT (M/SEC)</u>	<u>ESTIMATED MSLP (MB)</u>
(01W) TY Koryn	12 Jan - 17 Jan	19	75 (39)	967
(02W) TS Lewis	29 Apr - 03 May	15	35 (18)	997
(03W) TY Marian	15 May - 19 May	17	90 (46)	954
(04W) TD 04W	14 Jun - 15 Jun	4	30 (15)	1000
(05W) TS Nathan	15 Jun - 19 Jun	14	55 (28)	984
(06W) TY Ofelia	17 Jun - 25 Jun	31	90 (46)	954
(07W) TY Percy	21 Jun - 30 Jun	36	115 (59)	927
(08W) TS Robyn	07 Jul - 11 Jul	18	45 (23)	991
(09W) TY Steve	25 Jul - 02 Aug	31	115 (59)	927
(10W) TS Tasha	28 Jul - 31 Jul	12	55 (28)	984
(11W) TY Vernon	29 Jul - 07 Aug	39	95 (49)	948
(12W) TY Winona	06 Aug - 11 Aug	20	65 (33)	976
(01C) TS Aka*	07 Aug - 15 Aug	32	45 (23)	991
(13W) TY Yancy	13 Aug - 21 Aug	31	90 (46)	954
(14W) TY Zola	17 Aug - 23 Aug	23	100 (51)	944
(15W) TY Abe	24 Aug - 01 Sep	36	90 (46)	954
(16W) TY Becky	24 Aug - 30 Aug	25	70 (36)	972
(17W) TY Dot	03 Sep - 09 Sep	25	80 (41)	963
(18W) TS Cecil	04 Sep - 05 Sep	5	45 (23)	991
(19W) TY Ed	10 Sep - 20 Sep	40	90 (46)	954
(20W) STY Flo	12 Sep - 20 Sep	31	145 (75)	891 **
(21W) TY Gene	23 Sep - 30 Sep	30	80 (41)	963
(22W) TY Hattie	30 Sep - 08 Oct	31	90 (46)	954
(23W) TS Ira	02 Oct - 03 Oct	7	35 (18)	997
(24W) TS Jeana	13 Oct - 15 Oct	6	35 (18)	997
(25W) TY Kyle	16 Oct - 22 Oct	28	90 (46)	954
(26W) TS Lola	17 Oct - 18 Oct	7	40 (21)	994
(27W) STY Mike	07 Nov - 18 Nov	43	150 (77)	885
(28W) TS Nell	10 Nov - 12 Nov	7	50 (26)	987
(29W) STY Page	19 Nov - 30 Nov	45	140 (72)	898
(30W) STY Owen	21 Nov - 03 Dec	48	140 (72)	898
(31W) TY Russ***	14 Dec - 24 Dec	38	125 (64)	916

TOTAL: 794

* 24 WARNINGS ISSUED BY NWOC

** BASED ON AIRCRAFT DATA.

*** TWO WARNINGS ISSUED BY AJTWC.

The criteria used in Table 3-2 are as follows:

1. If a tropical cyclone was first warned on during the last two days of a particular month and continued into the next month for longer than two days, then that system was attributed to the second month.
2. If a tropical cyclone was warned on prior to the last two days of a month, it was attributed to the first month, regardless of how long the system lasted.
3. If a tropical cyclone began on the last day of the month and ended on the first day of the next month, that system was attributed to the first month. However, if a tropical cyclone began on the last day of the month and continued into the next month for only two days, then it was attributed to the second month.

TABLE 3-2 LEGEND

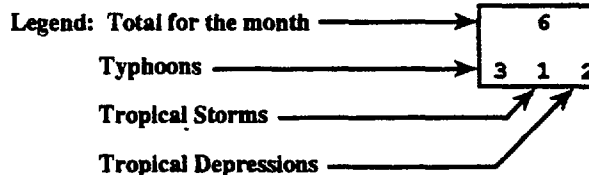


TABLE 3-2 WESTERN NORTH PACIFIC TROPICAL CYCLONE DISTRIBUTION

YEAR	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTALS
1960	1	0	1	1	1	3	3	9	5	4	1	1	30
	001	000	001	100	010	210	210	810	041	400	100	100	19 8 3
1961	1	1	1	1	4	6	5	7	6	7	2	1	42
	010	010	100	010	211	114	320	313	510	322	101	100	20 11 11
1962	0	1	0	1	3	0	8	8	7	5	4	2	39
	000	010	000	100	201	000	512	701	313	311	301	020	24 6 9
1963	0	0	1	1	0	4	5	4	4	6	0	3	28
	000	000	001	100	000	310	311	301	220	510	000	210	19 6 3
1964	0	0	0	0	3	2	8	8	8	7	6	2	44
	000	000	000	000	201	200	611	350	521	331	420	101	26 13 5
1965	2	2	1	1	2	4	6	7	9	3	2	1	40
	110	020	010	100	101	310	411	322	531	201	110	010	21 13 6
1966	0	0	0	1	2	1	4	9	10	4	5	2	38
	000	000	000	100	200	100	310	531	532	112	122	101	20 10 8
1967	1	0	2	1	1	1	8	10	8	4	4	1	41
	010	000	110	100	010	100	332	343	530	211	400	010	20 15 6
1968	0	1	0	1	0	4	3	8	4	6	4	0	31
	000	001	000	100	000	202	120	341	400	510	400	000	20 7 4
1969	1	0	1	1	0	0	3	3	6	5	2	1	23
	100	000	010	100	000	000	210	210	204	410	110	010	13 6 4
1970	0	1	0	0	0	2	3	7	4	6	4	0	27
	000	100	000	000	000	110	021	421	220	321	130	000	12 12 3
1971	1	0	1	2	5	2	8	5	7	4	2	0	37
	010	000	010	200	230	200	620	311	511	310	110	000	24 11 2
1972	1	0	1	0	0	4	5	5	6	5	2	3	32
	100	000	001	000	000	220	410	320	411	410	200	210	22 8 2
1973	0	0	0	0	0	0	0	6	3	4	3	0	23
	000	000	000	000	000	000	430	231	201	400	030	000	12 9 2
1974	1	0	1	1	1	4	5	7	5	4	4	2	35
	010	000	010	010	100	121	230	232	320	400	220	020	15 17 3
1975	1	0	0	1	0	0	1	6	5	6	3	2	25
	100	000	000	001	000	000	010	411	410	321	210	002	14 6 5
1976	1	1	0	2	2	2	4	4	5	0	2	2	25
	100	010	000	110	200	200	220	130	410	000	110	020	14 11 0
1977	0	0	1	0	1	1	4	2	5	4	2	1	21
	000	000	010	000	001	010	301	020	230	310	200	100	11 8 2
1978	1	0	0	1	0	3	4	8	4	7	4	0	32
	010	000	000	100	000	030	310	341	310	412	121	000	15 13 4
1979	1	0	1	1	2	0	5	4	6	3	2	3	28
	100	000	100	100	011	000	221	202	330	210	110	111	14 9 5
1980	0	0	1	1	4	1	5	3	7	4	1	1	28
	000	000	001	010	220	010	311	201	511	220	100	010	15 9 4
1981	0	0	1	1	1	2	5	8	4	2	3	2	29
	000	000	100	010	010	200	230	251	400	110	210	200	16 12 1
1982	0	0	3	0	1	3	4	5	6	4	1	1	28
	000	000	210	000	100	120	220	500	321	301	100	100	19 7 2
1983	0	0	0	0	0	1	3	6	3	5	5	2	25
	000	000	000	000	000	010	300	231	111	320	320	020	12 11 2
1984	0	0	0	0	0	2	5	7	4	8	3	1	30
	000	000	000	000	000	020	410	232	130	521	300	100	16 11 3
1985	2	0	0	0	1	3	1	7	5	5	1	2	27
	020	000	000	000	100	201	100	520	320	410	010	110	17 9 1
1986	0	1	0	1	2	2	2	5	2	5	4	3	27
	000	100	000	100	110	110	200	410	200	320	220	210	19 8 0
1987	1	0	0	1	0	2	4	4	7	2	3	1	25
	100	000	000	010	000	110	400	310	511	200	120	100	18 6 1
1988	1	0	0	0	1	3	2	5	8	4	2	1	27
	100	000	000	000	100	111	110	230	260	400	200	010	14 12 1
1989	1	0	0	1	2	2	6	8	4	6	3	2	35
	010	000	000	100	200	110	231	332	220	600	300	101	21 10 4
1990	1	0	0	0	2	4	4	5	5	5	4	1	32
	100	000	000	000	110	211	220	500	410	230	310	100	21 10 1
(1960-1990)													
MEAN:	0.6	0.2	0.5	0.7	1.3	2.2	4.5	6.1	5.5	4.6	2.8	1.4	30.8
CASES:	19	8	17	22	41	68	141	190	172	144	88	44	954

Figure 3-1. Chronology of western North Pacific and North Indian Ocean tropical cyclones for 1990.

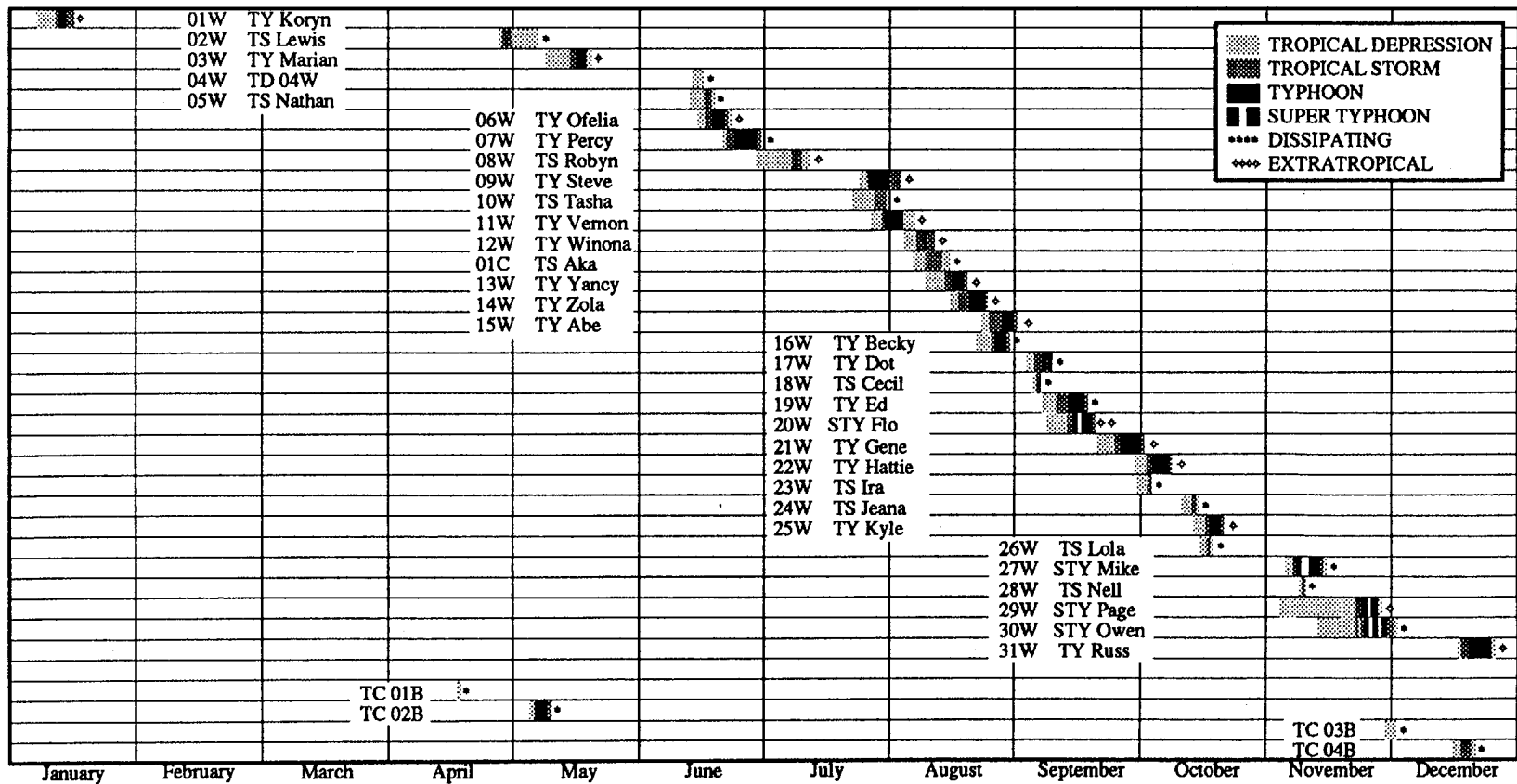


TABLE 3-3

WESTERN NORTH PACIFIC TROPICAL CYCLONES

TYPHOONS
(1945 - 1959)

	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTALS
MEAN:	0.3	0.1	0.3	0.4	0.7	1.0	2.9	3.1	3.3	2.4	2.0	0.9	16.4
CASES:	5	1	4	6	10	15	29	46	49	36	30	14	245

(1960 - 1990)

	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTALS
MEAN:	0.3	0.1	0.2	0.5	0.7	1.1	2.7	3.2	3.2	3.1	1.7	0.6	17.5
CASES:	9	2	6	15	23	34	84	99	100	97	54	20	543

TROPICAL STORMS AND TYPHOONS
(1945 - 1959)

	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTALS
MEAN:	0.4	0.1	0.5	0.5	0.8	1.6	2.9	4.0	4.2	3.3	2.7	1.2	22.2
CASES:	6	2	7	8	11	22	44	60	64	49	41	18	332

(1960 - 1990)

	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTALS
MEAN:	0.6	0.3	0.4	0.7	1.1	1.9	4.2	5.3	4.9	4.2	2.6	1.2	27.3
CASES:	18	8	13	21	35	58	129	165	153	130	82	38	850

1990 FORMATION ALERTS: 30 OF 33 INITIAL FORMATION ALERTS DEVELOPED INTO SIGNIFICANT TROPICAL CYCLONES. NO TROPICAL CYCLONE FORMATION ALERT WAS ISSUED FOR TROPICAL CYCLONE 04W.

WARNINGS DAYS:

NUMBER OF CALENDAR WARNING DAYS: 165

NUMBER OF CALENDAR WARNING DAYS WITH TWO TROPICAL CYCLONES: 54

NUMBER OF CALENDAR WARNING DAYS WITH THREE TROPICAL CYCLONES: 3

TABLE 3-4

TROPICAL CYCLONE FORMATION ALERTS
WESTERN NORTH PACIFIC OCEAN

YEAR	INITIAL TCFAS	TROPICAL CYCLONES WITH TCFAS	TOTAL TROPICAL CYCLONES	FALSE ALARM RATE	PROBABILITY OF DETECTION
1976	34	25	25	26%	100%
1977	26	20	21	23%	95%
1978	32	27	32	16%	84%
1979	27	23	28	15%	82%
1980	37	28	28	24%	100%
1981	29	28	29	3%	96%
1982	36	26	28	28%	93%
1983	31	25	25	19%	100%
1984	37	30	30	19%	100%
1985	39	26	27	33%	96%
1986	38	27	27	29%	100%
1987	31	24	25	23%	96%
1988	33	26	27	21%	96%
1989	51	32	35	32%	91%
1990	33	30	31	9%	97%
(1976-1990) MEAN:	34.3	26.5	27.9	21%	95%

3.2 WESTERN NORTH PACIFIC TROPICAL CYCLONES

1990 was an average year with 31 tropical cyclones - 4 super typhoons, 17 typhoons, 10 tropical storms and one tropical depression. This was above average for the number of typhoons and super typhoons, similar to 1989, but below in tropical depressions. All the tropical cyclones formed in the monsoon, or near-equatorial, trough even though the TUTT was much in evidence during the summer.

The year started off with a bang with Typhoon Koryn (01W) forming below 5° north latitude in a near-equatorial trough the second week of January. After a three month break in activity, Lewis (02W) flared up in low latitudes the last week of April, and Marian (03W) followed, finishing up by mid-May. The rest of May and first half of June were quiet, then Tropical Depression 04W formed in the South China Sea. As the monsoon trough extended eastward, so did the area for development of the next four tropical cyclones. First Nathan (05W) started just east of Mindanao, then Ofelia (06W) a little farther east, and Percy (07W), and finally, the last two days of June, Robyn (08W) in the eastern Caroline Islands.

During the first two weeks of July, a change took place in the synoptic pattern. A large TUTT low became dominant just west of the date line and drifted westward. Deep convection extended in a hook-like pattern south and east of the TUTT low and the low level monsoon trough became oriented northeast to southwest. A three-storm multiple outbreak followed during the third week of July. Three small tropical cyclones - Steve (09W), Tasha (10W) and Vernon (11W) - formed as the trough continued moving northwestward towards Asia. By the first week of August, the trough axis was near 25° north latitude and supported Winona's (12W) development near Okinawa.

After Winona (12W) the monsoon trough reestablished a normal orientation, extending southeastward from Asia in the southern Philippine Sea and large tropical cyclones generated one at a time, starting with Yancy

(13W) in the second week of August. The process continued through Hattie (22W) which started the last week of September. Cecil (18W), a midget tropical storm, was the only exception. During the second week of October northeasterly low-level flow surged into the northern South China Sea, as Hattie (22W) recurved. Three South China Sea cyclones - Ira (23W), Jeana (24W) and Lola (26W) - followed. After Kyle (25W), which began the middle of October, recurved just east of Iwo Jima, the summer monsoon weakened as winter set in and the axis of the monsoon trough shifted equatorward.

Following a two week break in activity, Mike (27W) formed in the eastern Caroline Islands at the end of the first week of November and became the first of three super typhoons to occur during the month. Nell (28W) developed in the South China Sea in association with the enhanced monsoonal flow into Mike (27W). The winter monsoon became established across Southeast Asia, however activity continued in the near-equatorial trough to the east. Initially tropical cyclone development was slow with both Page (29W) and Owen (30W) remaining as tropical disturbances for over a week. The pair intensified as Sina (03P) generated in the Southern Hemisphere near the date line. Almost two weeks of relative quiet followed before Russ (31W) formed in the near-equatorial trough below 5° north latitude with a twin, Joy (06P) forming in the Southern Hemisphere.

JANUARY THROUGH MAY

The first tropical cyclone of 1990 in the western North Pacific, **Koryn (01W)** also became the third typhoon to occur in January in the past eleven years. Unlike Typhoon Jack (1989), which two weeks earlier came to an abrupt halt and rapidly dissipated just east of Guam, this typhoon turned northward and tracked through the Mariana Islands. Koryn brought the strongest sustained winds to the Marianas since Roy (1988), another January typhoon. After a three month lull, **Lewis (02W)** developed 200 nm south of Chuuk and moved

north, passing directly over the island. After continuing its northward trek for four more days, it was sheared apart by a digging midlatitude trough, and the low-level remnants of the tropical cyclone drifted west-northwestward for several more days before completely dissipating. **Marian (03W)** followed and persisted in low latitudes for almost a week before intensifying to become the second typhoon of 1990 and the only significant tropical cyclone to form in May. It tracked from the Philippine Sea across the Philippine Islands and into the South China Sea, where recurved and merged with a frontal system to form an extratropical low.

JUNE

Following a one month break in tropical cyclone activity, **Tropical Depression 04W**, became the first significant tropical cyclone to form in the South China Sea this year. Because satellite and synoptic fix positions disagreed throughout the depression's life, the depression proved to be very difficult to locate and forecast. As **Tropical Storm Nathan (05W)** crossing into the South China Sea, Tropical Depression 04W was drawn into the larger circulation and absorbed. Nathan, then executed an abrupt track change and stalled before tracking off to the north. Both the track and intensity of TD04W and Nathan were dominated by a larger monsoon circulation in the South China Sea. **Ofelia (06W)** became the third typhoon of 1990 and the first for the month of June. It moved toward the Philippine Islands, then slowed and turned to the northwest. It was the second tropical cyclone of the year to strike Taiwan and the first to affect the east coast of China. After recurvature, the extratropical remnants of Ofelia crossed Korea; an unusual characteristic for a June system. **Percy (07W)** followed as the fourth and last tropical cyclone in June. After forming southeast of Guam, it turned on an unusual track to the southwest for 36 hours before paralleling Ofelia's (06W) track to the west-northwest around the western periphery of the subtropical ridge. Percy damaged the western Caroline

Islands and became the second typhoon within a week to batter northern Luzon before recurving over eastern China.

JULY THROUGH OCTOBER

The first significant tropical cyclone of July, **Robyn (08W)** followed what at first glance might appear to be a typical recurvature track. However, Robyn's motion was actually a classic example of the response of a tropical cyclone to the establishment of an omega block in the westerlies to the north, and thus was significant as a case study of an infrequent, but complex, synoptic influence on tropical cyclone motion. The monsoon trough activity substantially increased and **Steve (09W)** along with Tropical Storm Tasha (10W) and Typhoon Vernon (11W) combined into the only three storm tropical cyclone outbreak to occur in the northwest Pacific this year. Steve persisted on an atypical northeastward track throughout its existence. **Tasha (10W)**, the third of four western Pacific tropical cyclones to occur in July, developed in the monsoon trough, but instead of following Steve (09W) and Vernon (11W) to the northeast, it made only a brief start in that direction before curving to the west and entering the South China Sea. After erratic motion and slow intensification, Tasha finally reached tropical storm intensity before slamming into the southern coast of China. **Vernon (11W)**, the last of four tropical cyclones to develop during July, followed Steve's northward-oriented track, as the monsoon trough underwent a major displacement to the north. The first typhoon of 1990 to hit Japan, **Winona (12W)** was the only tropical cyclone to form poleward of 25° north latitude this year. It formed in August from the remnants of Tropical Storm Tasha (10W) in a monsoon trough that was displaced northward of its normal location. Winona had an unusual track to the southeast before it turned northward to cross the southern portion of the Kanto Plain. In the central Pacific, **Aka (01C)** developed and remained embedded in the trade wind trough. It tracked steadily west-northwestward and never developed beyond tropical storm intensity. Aka

was the only tropical cyclone of 1990 to be in warning status when it crossed the date line from the Central into the Western Pacific Ocean. As Aka was dissipating, Yancy (13W) generated in the monsoon trough. It became JTWC's best forecast tropical cyclone of the year, and although the track was generally toward the northwest, it contained several interesting features, including interaction with a strengthening subtropical ridge, the effects of a passing mid-latitude shortwave trough and land interaction with the mountainous terrain of Taiwan. In the wake of Typhoon Yancy (13W), a surge in the southwesterly monsoon flow developed and Zola (14W) formed west of Guam in the monsoon trough. The depression initially tracked northeastward in response to a monsoon surge and slowly intensified. Zola then broke away from the monsoon trough and intensified to a typhoon. The typhoon recurved over western Honshu, moved into the Sea of Japan and accelerated east-northeastward. Typhoon Abe (15W), the fourth of five tropical cyclones in August, caused extensive damage from the Republic of the Philippines to northern China during its nine day life. Abe was also noteworthy as a classic example of the erratic motion and rapid reorganization that can occur in association with an intense monsoon surge. Becky (16W), a midget typhoon and the eleventh typhoon of 1990, generated in the monsoon trough and tracked south of the subtropical ridge throughout its existence. After initially moving west-northwestward, the storm took a southwestward track across the northwestern tip of Luzon before heading westward across the South China Sea. Becky hit northern Luzon with typhoon-force winds and later slammed into northern Vietnam as a severe tropical storm. Dot (17W) developed in the monsoon trough at the same time as Tropical Storm Cecil (18W) and brought enhanced southwesterly wind flow and heavy rains across Guam. Later, as Dot crossed central Taiwan, torrential monsoon rains from the associated monsoon surge caused extensive flooding in northern Luzon. During its passage across Taiwan and the Fujian Province of China, surface winds in

the Formosa Strait exceeded 50 kt (26 m/sec) for 30 hours. Tropical Storm Cecil (18W) was a short-lived, midget tropical cyclone that formed in the wake of Typhoon Abe (15W). As Abe raced poleward, the monsoon trough reestablished itself over northern Luzon, and Cecil formed at the northeast end of the trough. Cecil tracked northward and skirted the northern coast of Taiwan before making landfall in southeastern China. Ed (19W), which had the second longest track (3150 nm (5830 km)) of any "straight runner" in 1990, formed in the Marshall Islands and continued westward for nearly two weeks before finally making landfall in northern Vietnam. It was the third of six tropical cyclones to form in September. Flo (20W) was the fourth of six tropical cyclones to develop in September, the first of four super typhoons this year, and the object of over three consecutive days of upper-tropospheric aircraft reconnaissance missions during the TCM-90 field experiment. Flo formed in the wake of Typhoon Ed (19W), passed close by Guam, then rapidly intensified into a super typhoon as it approached Okinawa. Recurvature was slow before the tropical cyclone accelerated northeastward towards the Japanese mainland where it was called the most powerful typhoon to hit Honshu in 19 years. At least 38 people were reported dead or missing, and damage was estimated in the millions of dollars. Transportation, communications and power were also disrupted. Gene (21W) was the fifth significant tropical cyclone to form in September and the fifteenth of the year to reach typhoon intensity. The initial disturbance formed 250 nm (465 km) west-southwest of Guam and tracked westward for three days before turning northwestward. Gene followed a classic recurvature pattern, passing west of Okinawa and skirting southern Japan. The orientation of Gene's recurvature track resulted in sustained radar contact from 251400Z to 300400Z and an excellent, high quality set of 250 position reports from land radar sites in the islands nearby. Hattie (22W), the last of six tropical cyclones to form in September, was the fourth tropical cyclone in a six-week period to affect Okinawa and southern Japan. It also

followed a classic recurvature track. **Ira (23W)** was the eighth tropical cyclone to hit Vietnam in 1990 and the last in a series of weak, highly sheared tropical systems in the South China Sea. It formed in a broad area of convection near Palawan Island. The convective cloud mass tracked steadily westward in the deep easterly flow and made landfall at Qui Nhon, Vietnam on the third of October. **Jeana (24W)**, the second of four tropical cyclones to form in October, was the fifth to churn across the South China Sea in 1990. This minimal tropical storm proved to be as difficult to estimate intensity for, as it was to position. **Kyle (25W)** generated from a disturbance in the monsoon trough 600 nm (1110 km) east of Guam. Separating from the trough, the cloud system gained organization and began to track along the southern edge of the subtropical ridge to its northeast. The subtropical ridge and a series of fast moving mid-latitude short-wave troughs strongly influenced Kyle's track. The tropical cyclone passed through the northern Mariana Islands, causing minimal damage, intensified into a typhoon, and recurved. **Lola (26W)**, the last of four tropical cyclones to develop in October, formed in the South China Sea. It tracked westward along the same path taken by Tropical Storm Jeana (24W) four days earlier.

NOVEMBER THROUGH DECEMBER

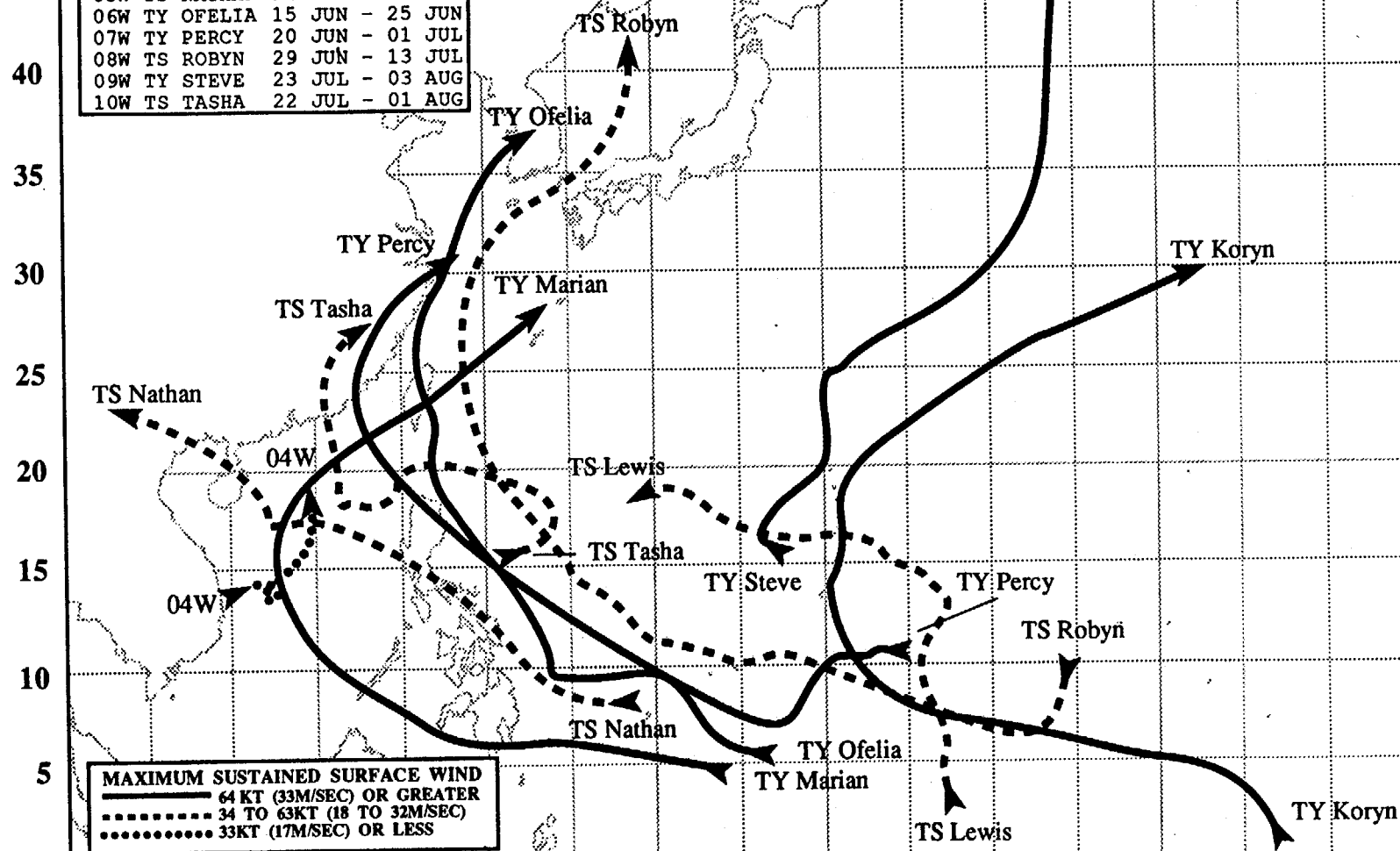
Mike (27W), one of the most intense and destructive tropical cyclones of 1990, caused havoc in western Carolines and in the central Philippine islands. Although basically a west-northwestward "straight runner," it posed numerous forecast challenges due to frequent direction, speed and intensity changes. As a result of the devastation and death in the Republic of the Philippines, Super Typhoon Mike's name was retired from the JTWC list of tropical cyclone names. **Nell (28W)**, the second of four November tropical cyclones, intensified in the South China Sea and tracked westward, making landfall in Vietnam. **Page (29W)** was the third of four tropical cyclones to form in November, the second super typhoon of the month, and part of the three-storm outbreak

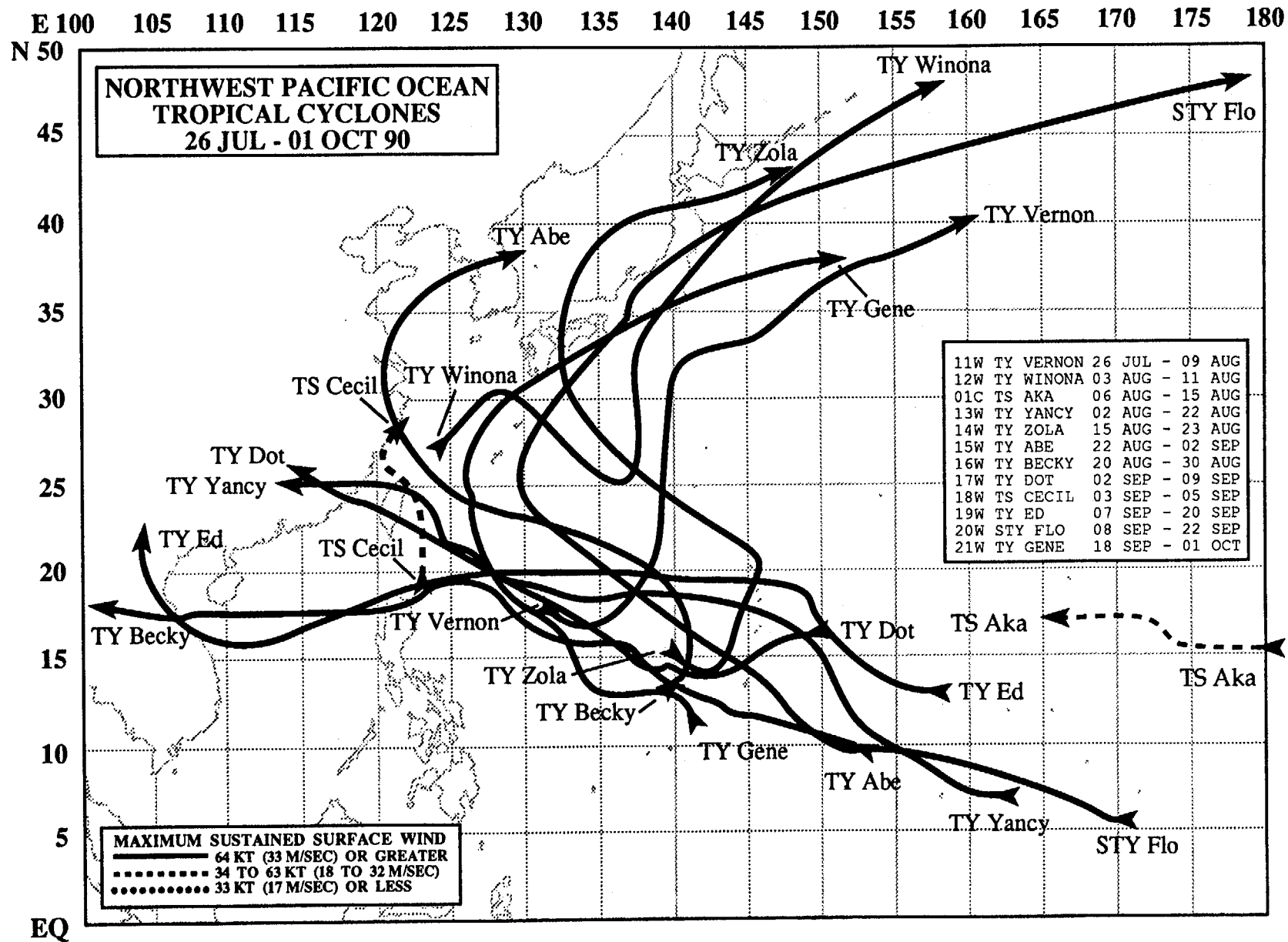
which included a pair of tropical cyclones near the dateline: **Owen (30W)** in the northern hemisphere and **Sina (TC 03P)** in the southern hemisphere. Persisting as a discrete disturbance for nearly two weeks before the first warning was issued, Page took only three days to intensify to 140 kt (70 m/sec) once development commenced. **Owen (30W)** was both the longest lasting and one of the most interesting tropical cyclones of 1990. It started to rapidly intensify while still a tropical depression, explosively deepened to super typhoon intensity, weakened and then reintensified to a super typhoon. Owen started as a discrete cloud mass southwest of Hawaii, maintained its integrity as it tracked westward in the trade wind trough, but did not intensify until it crossed the dateline and passed north of Kwajalein in the Marshall Islands. It then reached typhoon intensity in less than 18 hours and continued westward over the central Caroline Islands until its deep convection was sheared away southeast of Ulithi Island in the western Carolines. The exposed low-level remained organized for six more days as it moved north, then west, and finally southwestward before dissipating over the Celebes Sea after crossing Mindanao. **Russ (31W)**, the last western North Pacific tropical cyclone of 1990, was the most severe to strike Guam in 14 years. Damage was estimated as high as 120 million dollars. Russ formed in the Marshall Islands, tracked west-northwestward and intensified to near super typhoon intensity as it approached Guam. The typhoon passed within 30 nm (55 km) of the southern tip of Guam and brought typhoon force winds which caused extensive damage, especially to the southern portion of the island. After leaving Guam, Russ slowly weakened, recurved and became an extratropical cyclone.

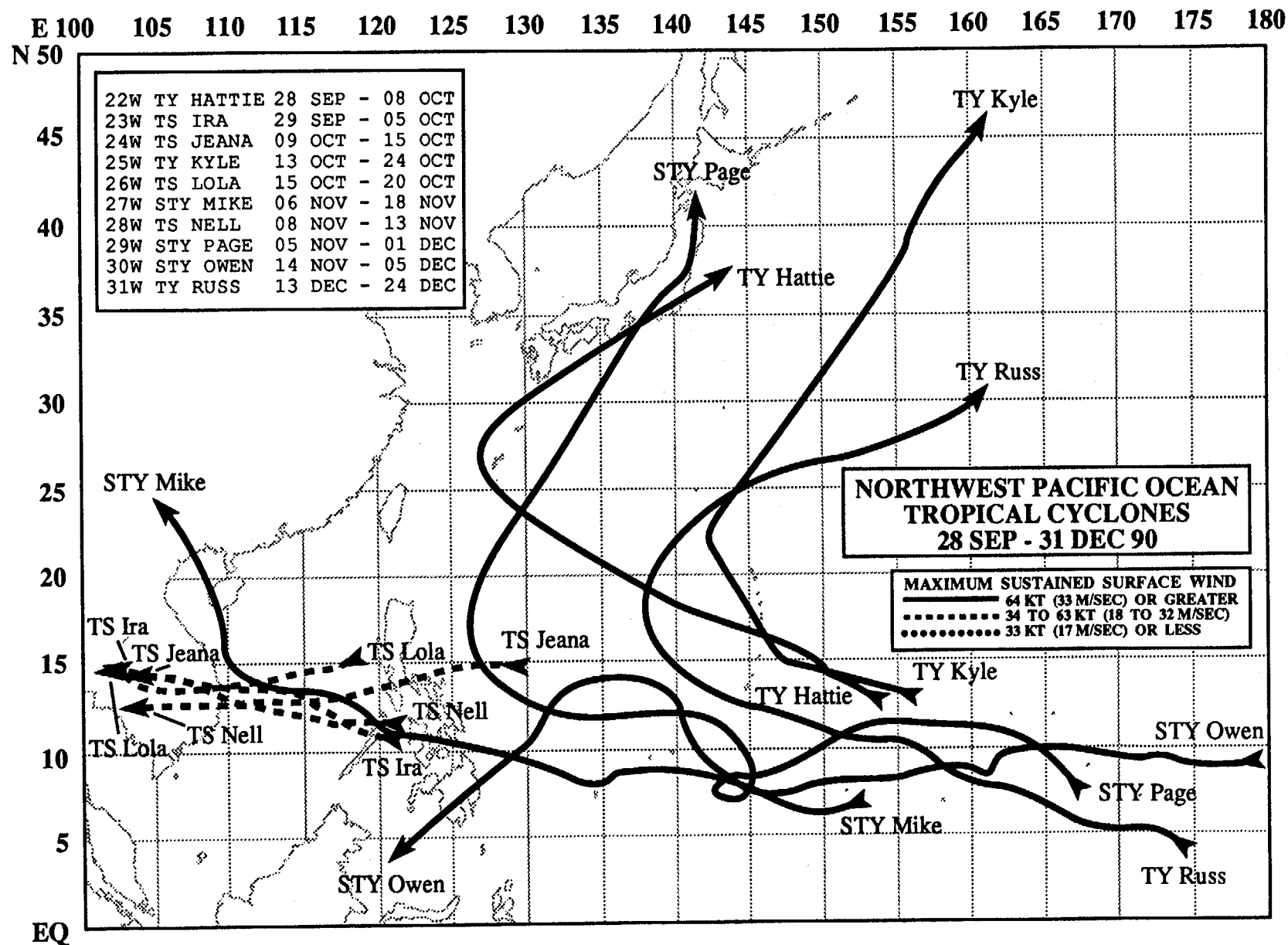
E 100 105 110 115 120 125 130 135 140 145 150 155 160 165 170 175 180
N 50

01W	TY KORYN	08 JAN - 17 JAN
02W	TS LEWIS	27 APR - 07 MAY
03W	TY MARIAN	09 MAY - 20 MAY
04W	TD 04W	13 JUN - 16 JUN
05W	TS NATHAN	12 JUN - 19 JUN
06W	TY OFELIA	15 JUN - 25 JUN
07W	TY PERCY	20 JUN - 01 JUL
08W	TS ROBYN	29 JUN - 13 JUL
09W	TY STEVE	23 JUL - 03 AUG
10W	TS TASHA	22 JUL - 01 AUG

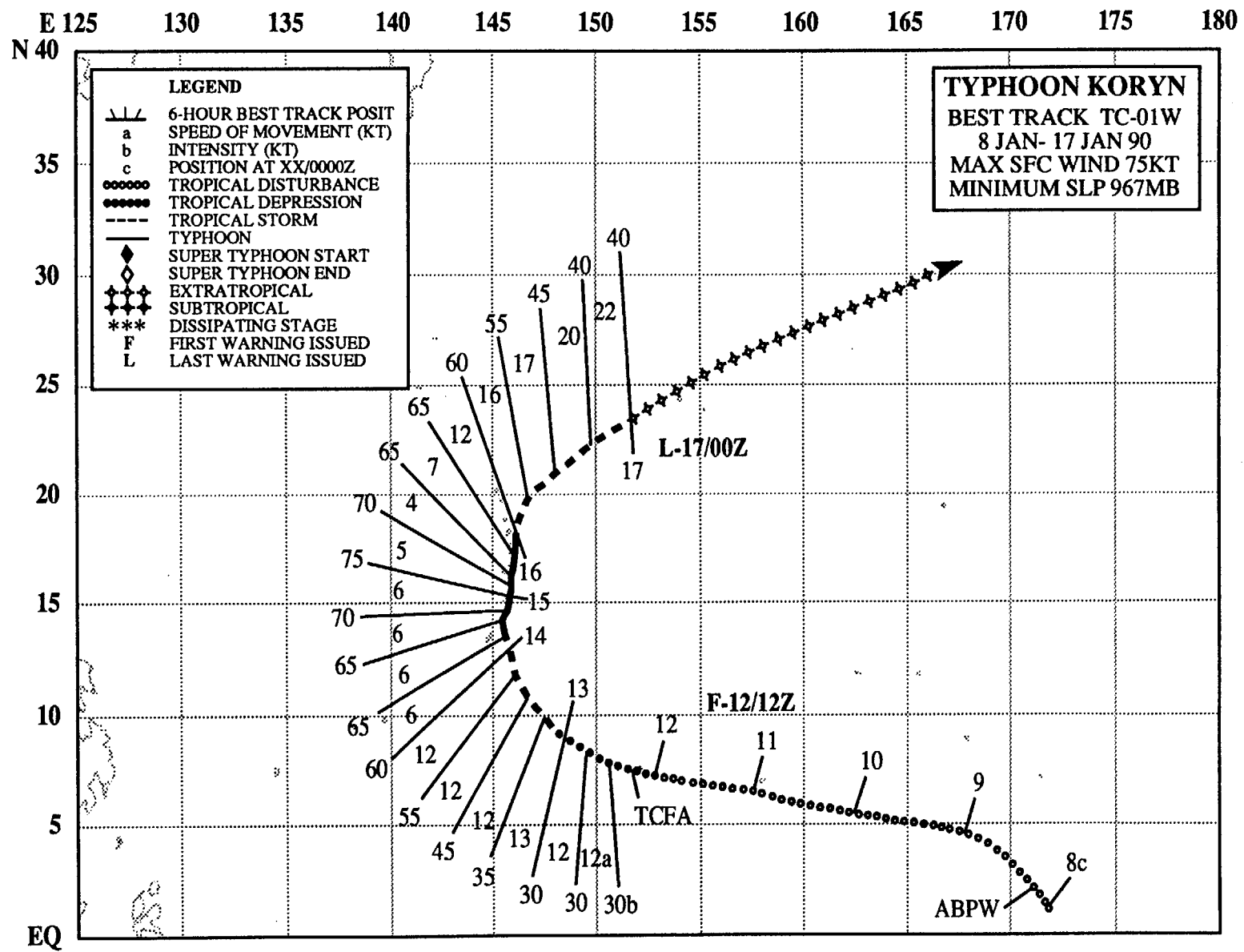
**NORTHWEST PACIFIC OCEAN
TROPICAL CYCLONES
01 JAN - 03 AUG 90**







Intentionally left blank.



TYPHOON KORYN (01W)

I. HIGHLIGHTS

Koryn, the first tropical cyclone of 1990 in the western North Pacific, became the third typhoon to occur in January in the past eleven years. It developed at an unusually low latitude. Unlike Typhoon Jack (1989), which two weeks earlier came to an abrupt halt and rapidly dissipated just east of Guam, this typhoon turned northward and tracked through the Mariana Islands. Koryn brought the strongest sustained winds to the Marianas since Roy (1988), another January typhoon.

II. CHRONOLOGY OF EVENTS

- 081500Z - First mentioned on the Significant Tropical Weather Advisory due to persistence of convection.
- 120430Z - Tropical Cyclone Formation Alert followed 4 mb pressure falls with strong easterly flow to the north, weak westerlies to the south and a CI 1.5.
- 121200Z - First warning based on increased convective curvature and outflow aloft.
- 130600Z - Upgraded to tropical storm intensity following improved organization of convection and good outflow aloft in all quadrants which resulted in a CI 2.5.
- 140600Z - Upgraded to typhoon based on the appearance of an eye and a CI 4.0.
- 150000Z - Peak intensity 75 kt (39 m/sec) with a ragged eye and a CI 4.5.
- 160000Z - Downgraded to tropical storm with signs of extratropical transition, shearing-type cloud pattern and restricted outflow.
- 170000Z - Final warning. Koryn extratropical with exposed low-level circulation center displaced to southwest of central cloud mass.

III. TRACK AND MOTION

Koryn originated as a disturbance (Figure 3-01-1) near the Gilbert Islands. The cyclonic circulation formed in sympathetic response to enhanced westerly monsoonal flow extending from the Solomon Islands eastward along 5° south latitude to a low pressure system near the Fiji Islands. While Koryn was embedded in the flow south of the subtropical ridge, it moved west-northwestward to Chuuk (Truk) in the eastern Caroline Islands. The subtropical ridge was north of the tropical cyclone along 20°

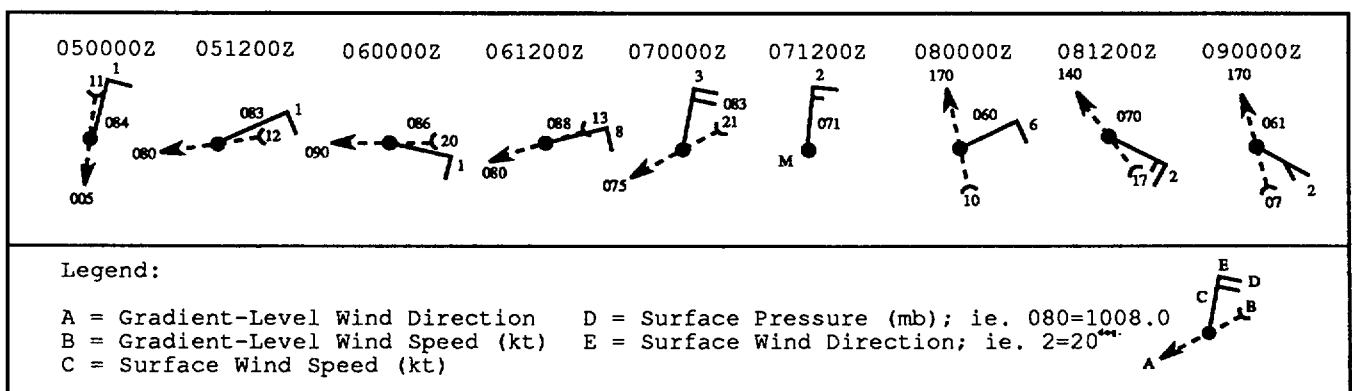


Figure 3-01-1. Surface pressure, gradient and surface wind reports for Tarawa (WMO 91610) in the Gilbert Islands reflect the formation of Koryn just to the west of the station. From 050000Z until 071200Z, the gradient-level wind is the normal cross-equatorial flow from the Northern Hemisphere, around a buffer system on the equator to the monsoon westerlies farther south. Note that on 8 January, the pressure in the past 24 hours fell over 2 mb and the gradient-level flow abruptly shifted to south-southeasterly. This supports the formation of a cyclonic circulation just to the west of the station.

north latitude; but lower pressures aloft in the northern Marianas indicated a break in the ridge. Koryn turned to a more northward track toward this break in the ridge and Guam. The typhoon slowed, passed just east of Guam and directly over Saipan. The slow forward motion and prolonged northward track appear related to the weaker steering flow associated with the break in the ridge and with the relative broad character of the ridge itself (Figure 3-01-2). Strong zonal westerlies aloft resulted in recurvature and a northeastward acceleration. Koryn's residual circulation and associated cloudiness continued northeastward along the edge of the maritime polar air and linked up to a passing short wave.

IV. INTENSITY

Koryn's weak low-level circulation first appeared just to the north of a broad area of cloudiness that stretched along and south of the equator. As this circulation moved west-northwestward, convection flared-up to its north and east. This enhanced cloudiness (Figure 3-01-3) became more organized and developed into a tropical cyclone as the low-level circulation center moved beneath an area of upper-level divergence. The synoptic scale upper-level anticyclone remained displaced to the east. Although upper tropospheric southeasterlies restricted Koryn's outflow to the southeast, the upper-level anticyclone of the typhoon continued to provide good outflow until the system reached its peak intensity (Figure 3-01-4). As Koryn moved northward, increasing vertical wind shear in the mid-

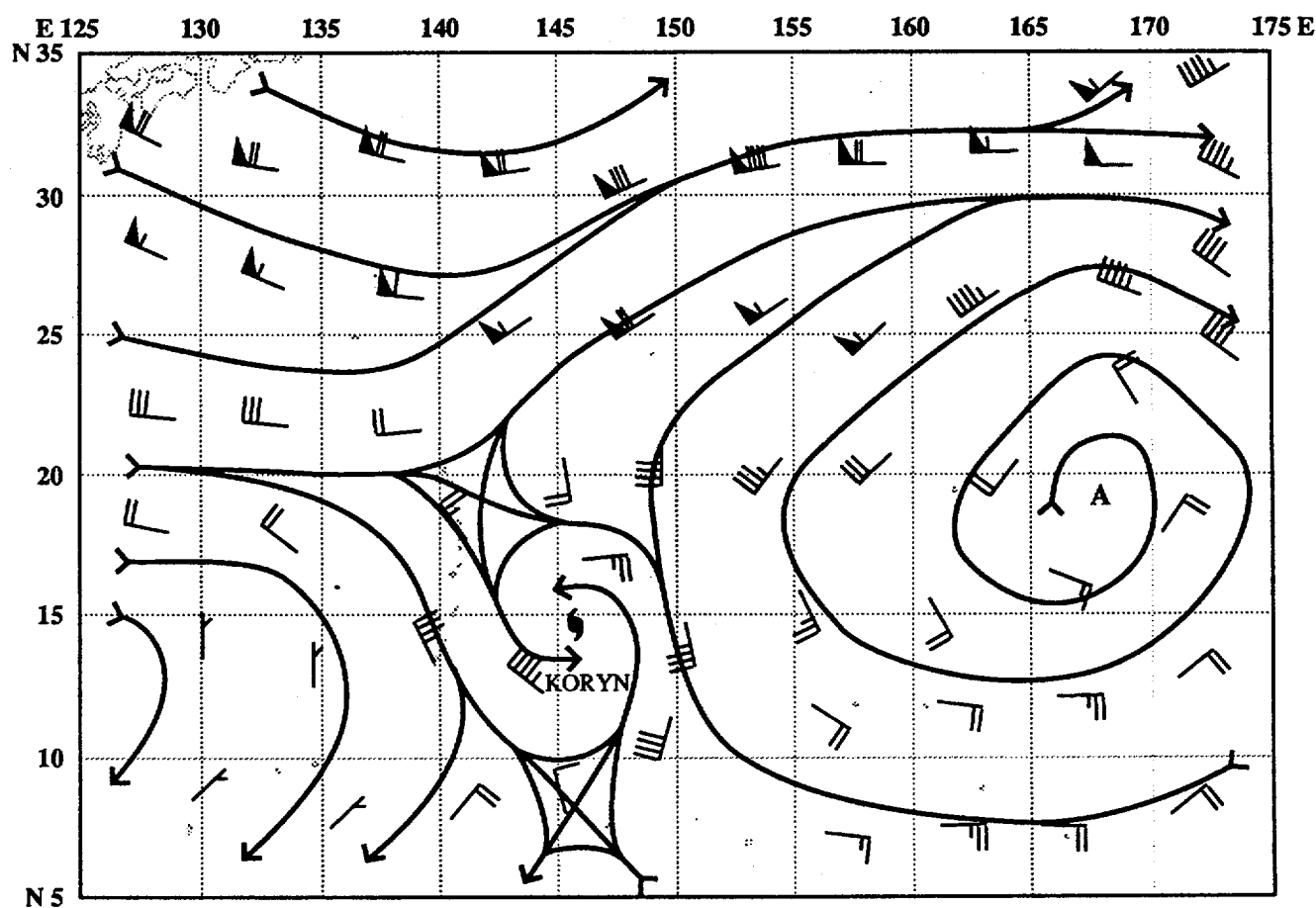
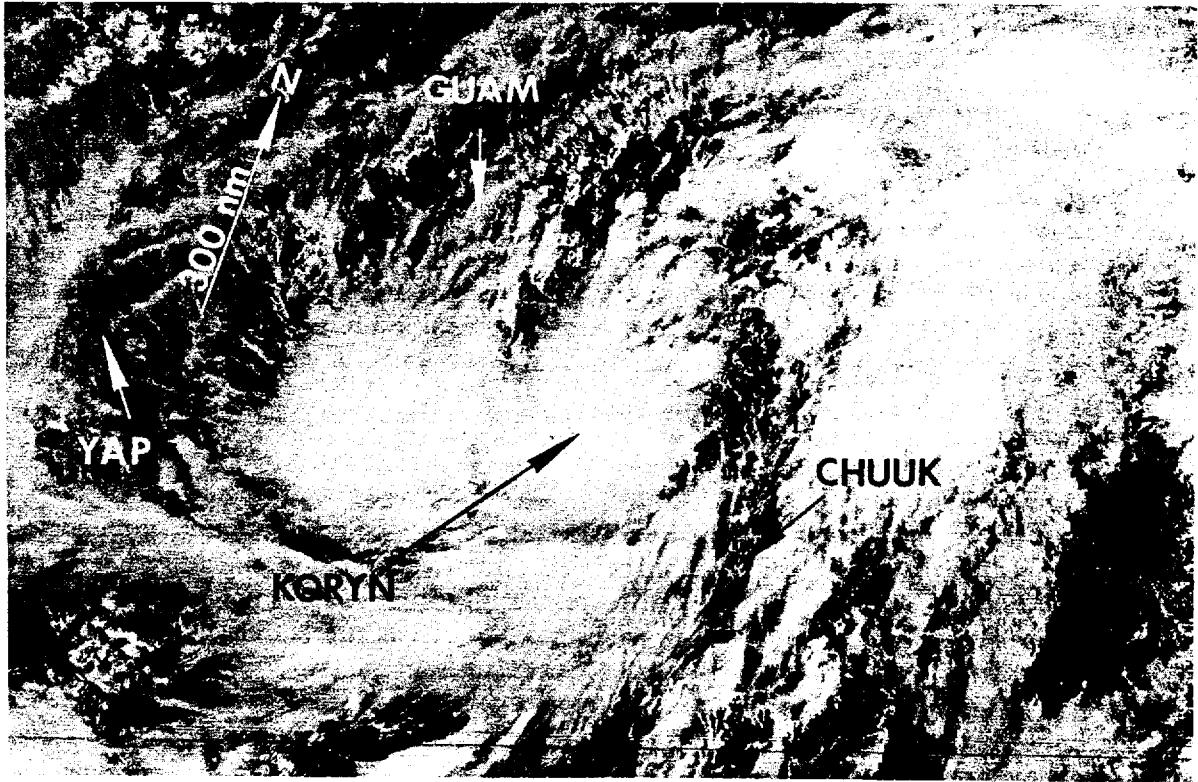
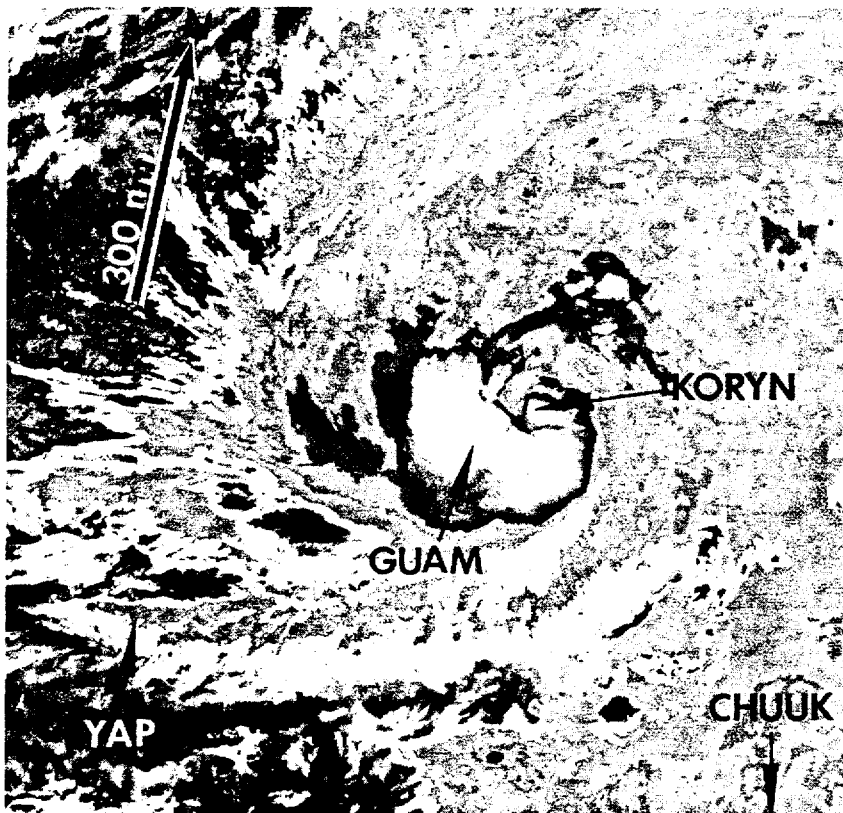


Figure 3-01-2. NOGAPS 500 mb analysis for 141200Z January shows the typhoon in the relative broad subtropical ridge. The ridge axis is at approximately 20° north latitude.



Above: Figure 3-01-3. Tropical Depression 01W's poorly defined cloudiness southeast of Guam (122330Z January DMSP visual imagery).



Left: Figure 3-01-4. Typhoon Koryn near peak intensity and just before maximum surface wind gusts to 70 kt (36 m/sec) were recorded on Guam (140958Z January NOAA enhanced infrared imagery).

latitude westerlies weakened the system. After recurvature, the cyclone's acceleration retarded the penetration of cooler low-level air into the center maintaining the intensity. Extratropical transition was completed a day after recurvature started.

V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-01-5. Initially, the weakness in the subtropical ridge was not expected to influence the track. As a result, JTWC forecast a westward track instead of recurvature near Guam. OTCM guidance (Figure 3-01-6) at first indicated a west-northwest track. However, on 13 January OTCM began to hint at recurvature. At 131800Z, JTWC included recurvature (Figure 3-01-7) as an alternate scenario, and it became the primary on the next warning. The forecast track might have been adjusted sooner, but disagreement among radar and satellite fixes resulted in the initial working best track being more westward and slower than the actual track as the system approached Guam.

VI. IMPACT

The forecasting difficulties mentioned above reduced on-island preparation time for Koryn's closest approach to Guam. Andersen AFB suspended aircraft evacuation and only one Navy ship sortied from Apra Harbor. The aircraft and ships remaining in Guam did not sustain any damage. Although Koryn passed within 50 nm (93 km) east of Guam, the island suffered only slight damage. Maximum winds reported at Andersen AFB were 40 kt (21 m/sec) gusting to 55 kt (28 m/sec). NAS Agana reported 54 kt (28 m/sec) gusting to 70 kt (36 m/sec). Koryn passed directly over Saipan, which also sustained only minor damage. Maximum winds at the Saipan Airport were 32 kt (16 m/sec), and the minimum sea-level pressure was 981 mb.

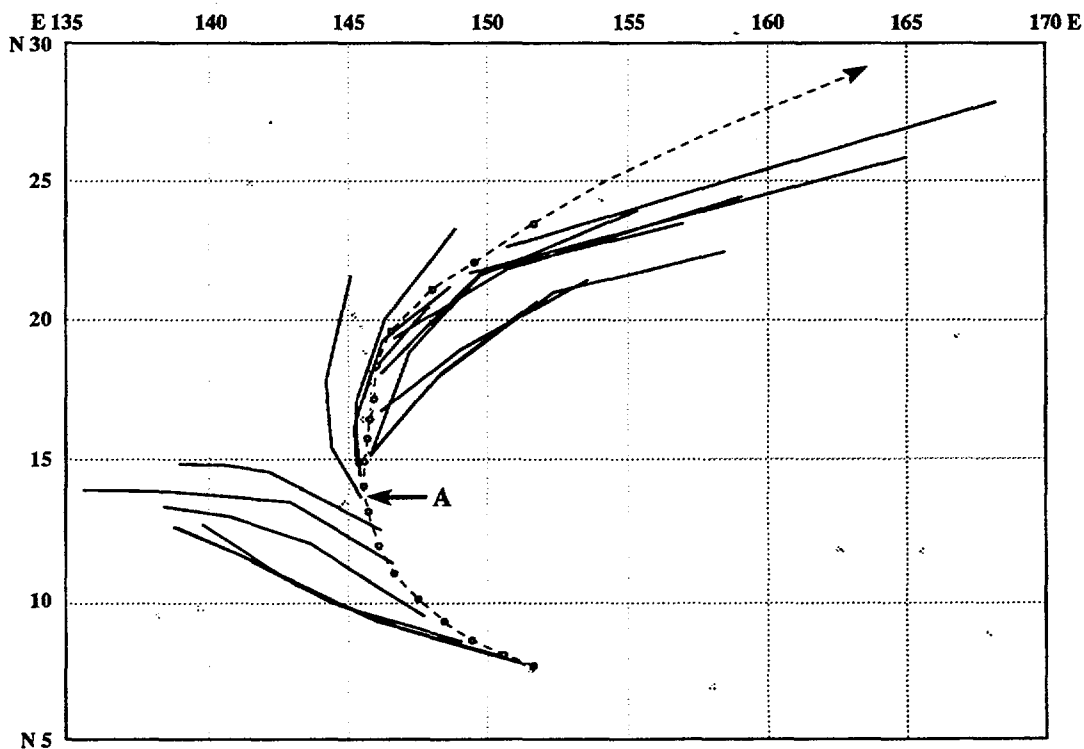


Figure 3-01-5. Summary of forecasts (solid lines) for Koryn superimposed on the final best track (dashed line). Point A identifies the first recurvature forecast at 140000Z.

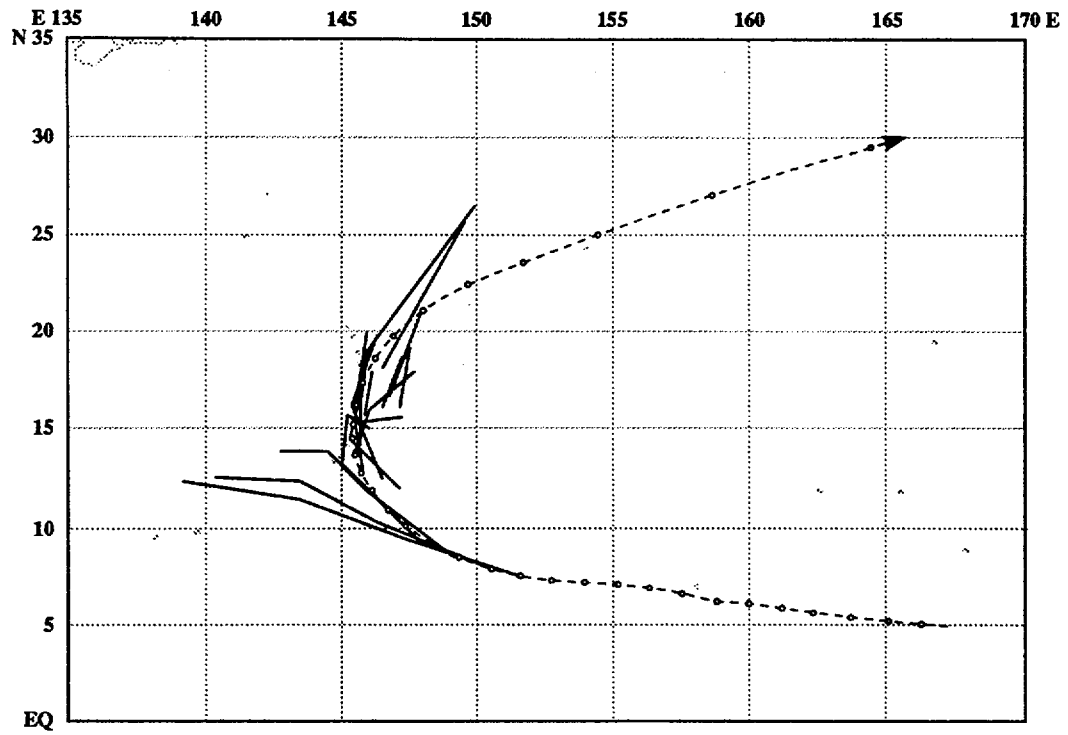


Figure 3-01-6. OTCM guidance (solid lines) superimposed on the final best track for Koryn (dashed line). OTCM started to hint at recurvature early on 13 January.

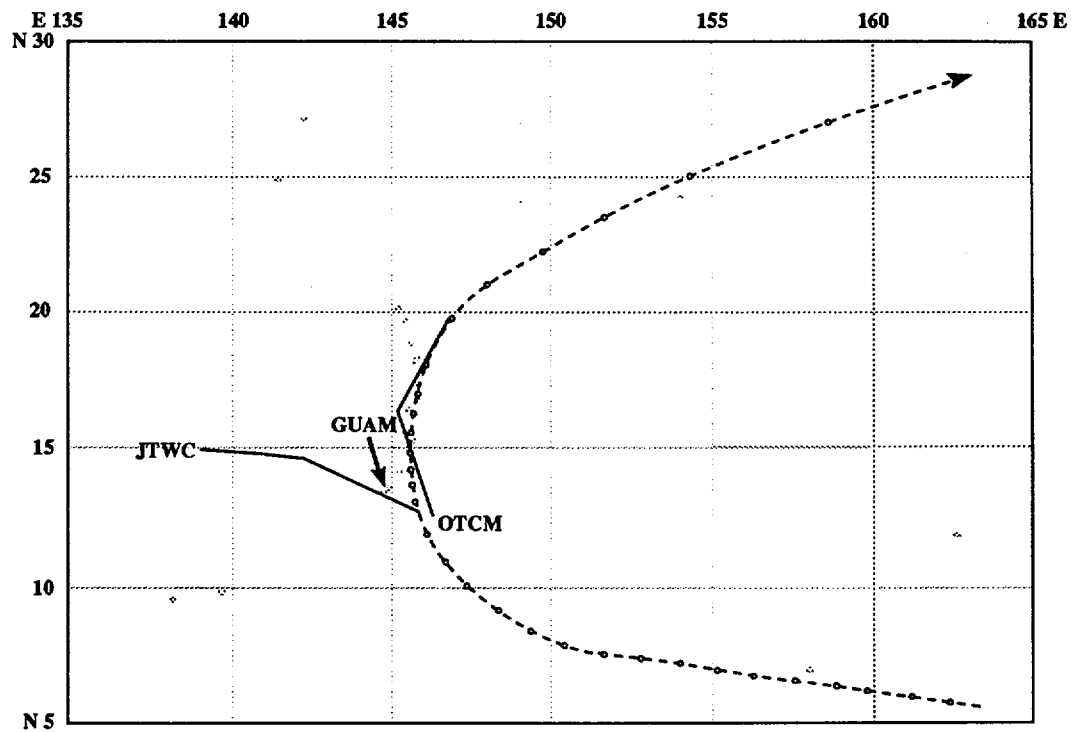
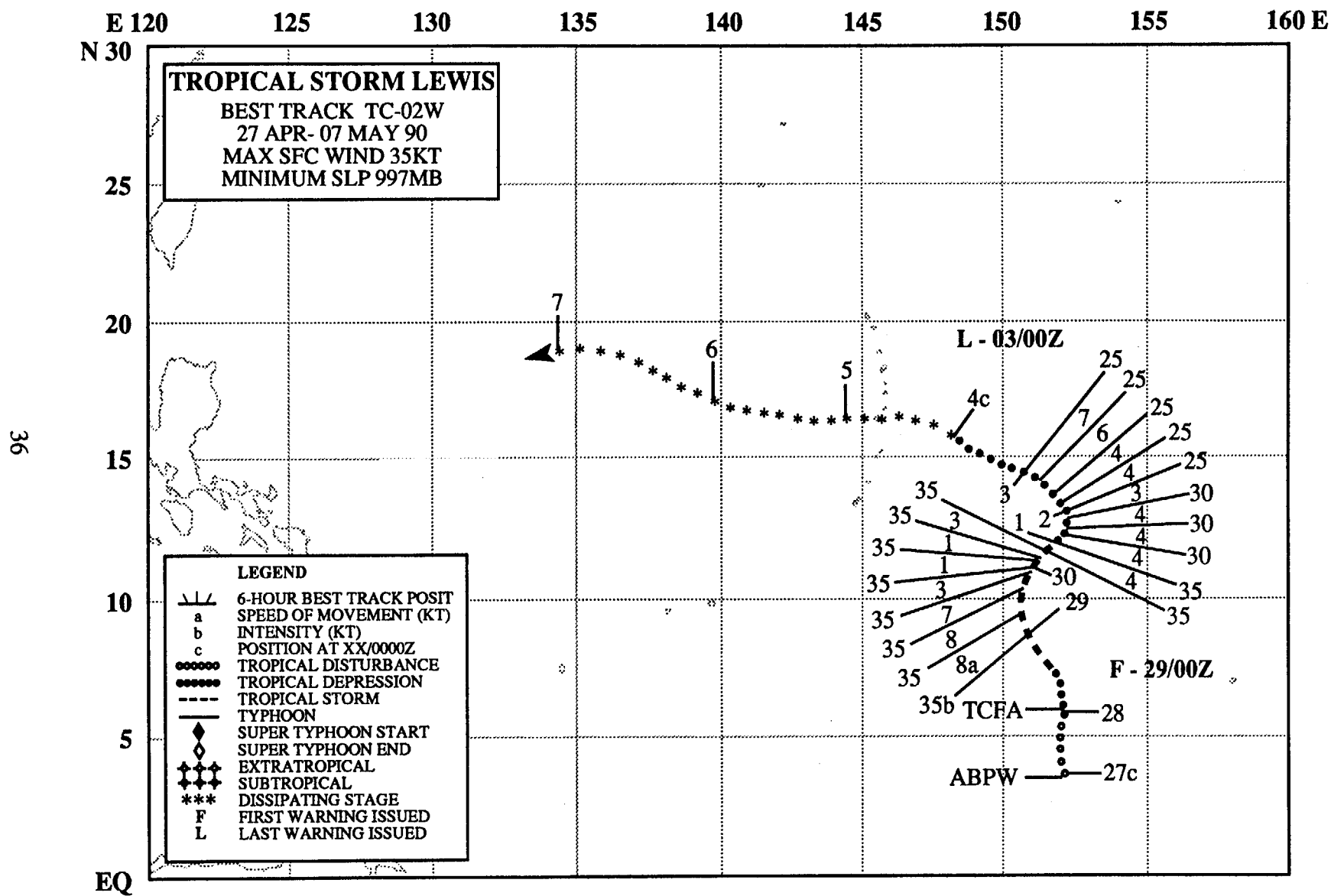


Figure 3-01-7. Comparison of the JTWC forecast (solid line) and OTCM guidance (solid line) at 131800Z.



TROPICAL STORM LEWIS (02W)

I. HIGHLIGHTS

Lewis ended the two and a half month lull in northern hemisphere tropical cyclone activity that followed Typhoon Koryn (01W) in January. Developing from a tropical disturbance 200 nm south of Chuuk in the central Caroline Islands, Lewis passed directly over Chuuk while still a tropical depression and continued a northward trek for four more days. After being sheared apart by a digging midlatitude trough, the low-level remnants of the tropical cyclone drifted west-northwestward for several more days before completely dissipating.

II. CHRONOLOGY OF EVENTS

- 262330Z - The Significant Tropical Weather Advisory was reissued to address the redevelopment of an area of persistent convection with an estimated minimum sea-level pressure of 1009 mb.
- 280300Z - Tropical Cyclone Formation Alert based on increased convection, organization, and outflow aloft.
- 290000Z - First warning due to continued improvement in organization of the convection. Initial intensity based on synoptic data vice Dvorak intensity which had been CI 2.5 for approximately six hours.
- 290600Z - Upgrade to tropical storm prompted by improved upper-level organization. Peak intensity never exceeded 35 knots.
- 011800Z - Downgrade to tropical depression based on visual satellite imagery which showed partially exposed low-level.
- 030000Z - Final warning - dissipating over water - due to fully exposed low-level circulation.

III. TRACK AND MOTION

During initial development, Lewis tracked northward due to southerly flow associated with a mid-level anticyclone over the Marshall Islands. The anticyclone was separate from the subtropical ridge that was located near 20° north latitude. The initial northward motion changed to northwestward at 281200Z (Figure 3-02-1). This synoptic adjustment resulted in Lewis passing directly over Chuuk. As a midlatitude trough began to dig to the northwest of Lewis, the steering flow veered from southeasterly to southwesterly (Figure 3-02-2) and caused the tropical cyclone to begin recurving at

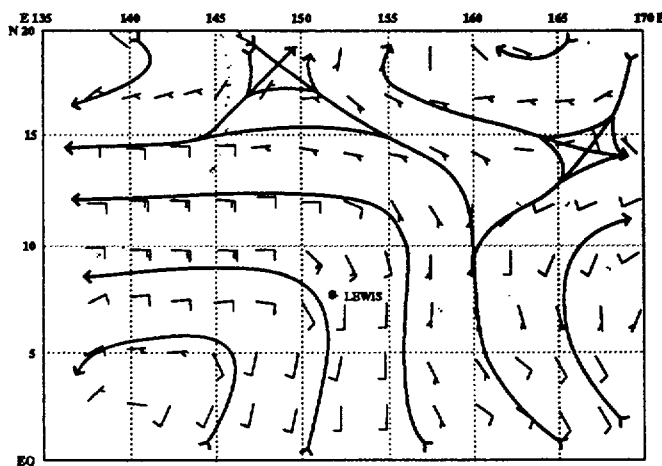


Figure 3-02-1. Lewis' turn to the northwest appears related to the subtle change of the steering flow from south to southeast on the 281200Z deep layer mean analysis.

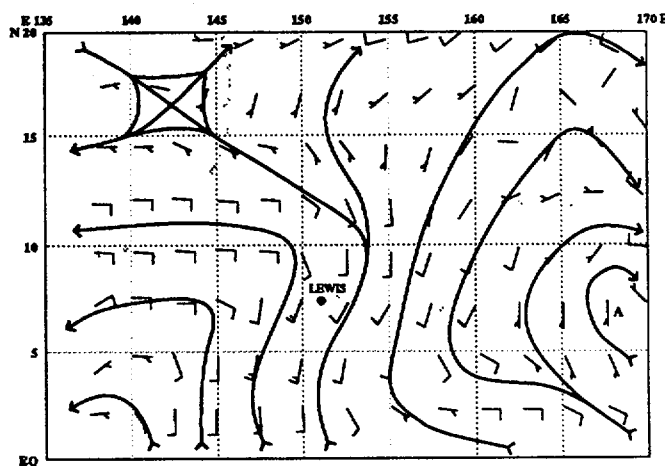


Figure 3-02-2. The 290000Z deep layer mean analysis shows the weakening of the ridge north-northeast of Lewis and maintenance of the anticyclonic circulation east of the tropical cyclone. This synoptic change, plus Lewis' continued movement to the north, brought Lewis into an area of light southwesterly steering flow.

291200Z. However, by 300000Z, the upper-level trough dug so far equatorward (to 10° north latitude) that the top of Lewis was sheared off by stronger westerlies aloft. Although Lewis' central convective activity intermittently flared up, the low-level circulation became exposed at 020000Z, and the low-level remnants of the cyclonic circulation drifted west-northwestward in response to the steering flow under the 850-mb ridge.

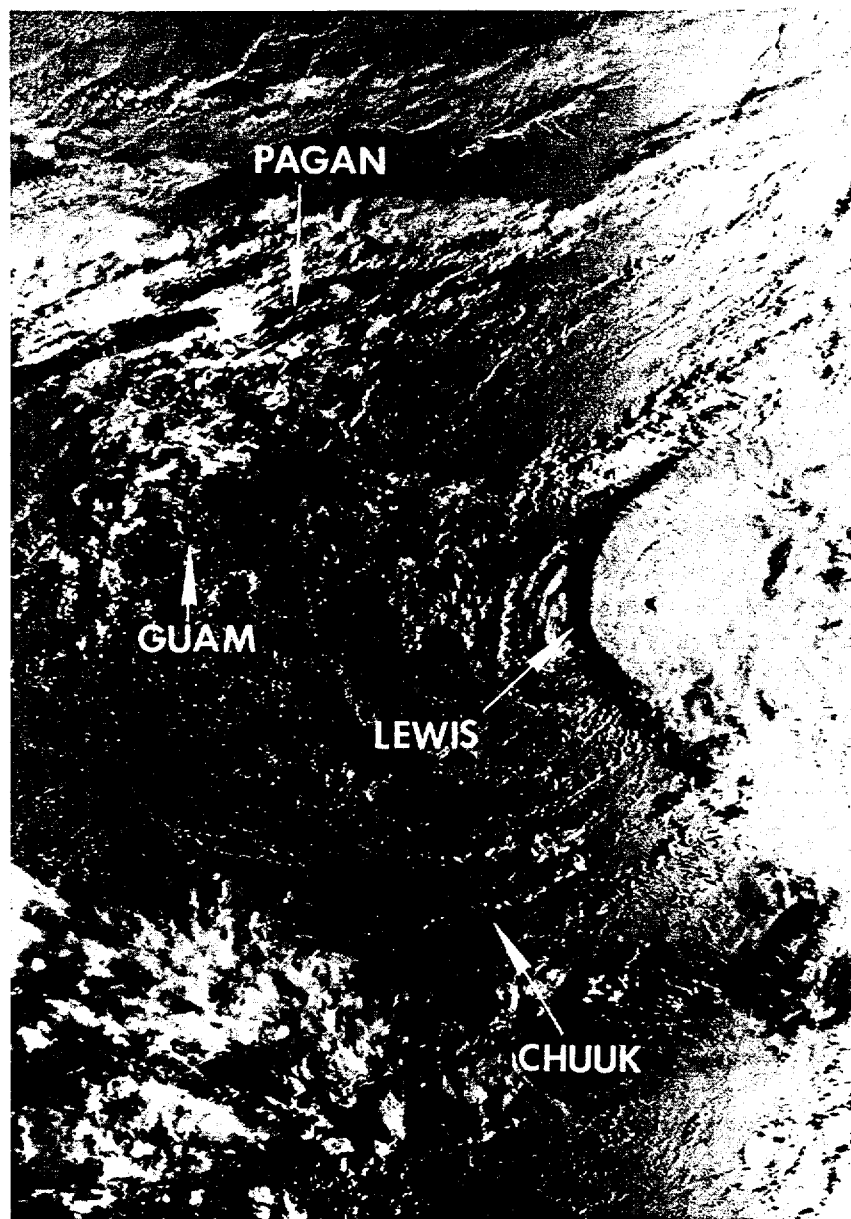


Figure 3-02-3. The sheared condition of Lewis (02W) is strikingly emphasized by the low sun-angle (012022Z May DMSP visual imagery).

IV. INTENSITY

In the early stages of its development, Lewis exhibited sufficient outflow to support moderate development. However, after reaching minimal tropical storm intensity, Lewis' further development was arrested by the encroaching 200-mb westerlies associated with the digging midlatitude trough (Figure 3-02-3). Two days later, the system began to slowly dissipate.

V. FORECASTING PERFORMANCE

Figure 3-02-4 shows the JTWC forecast performance for Lewis. Although the early forecasts anticipated the track change to the northwest followed by a change to the northeast, the forecasts were slow to anticipate the recurving effect of the digging midlatitude trough. Since neither subjective guidance nor the objective forecast aids available to JTWC were able to precisely address a shear-induced decoupling of the low-level circulation from its upper-level, the official forecasts incorrectly presumed continued recurvature. However, as early as 300600Z forecasters included an alternate scenario of shear-induced decoupling followed by west-northwestward movement of the low-level circulation in the prognostic reasoning.

VI. IMPACT

No information received.

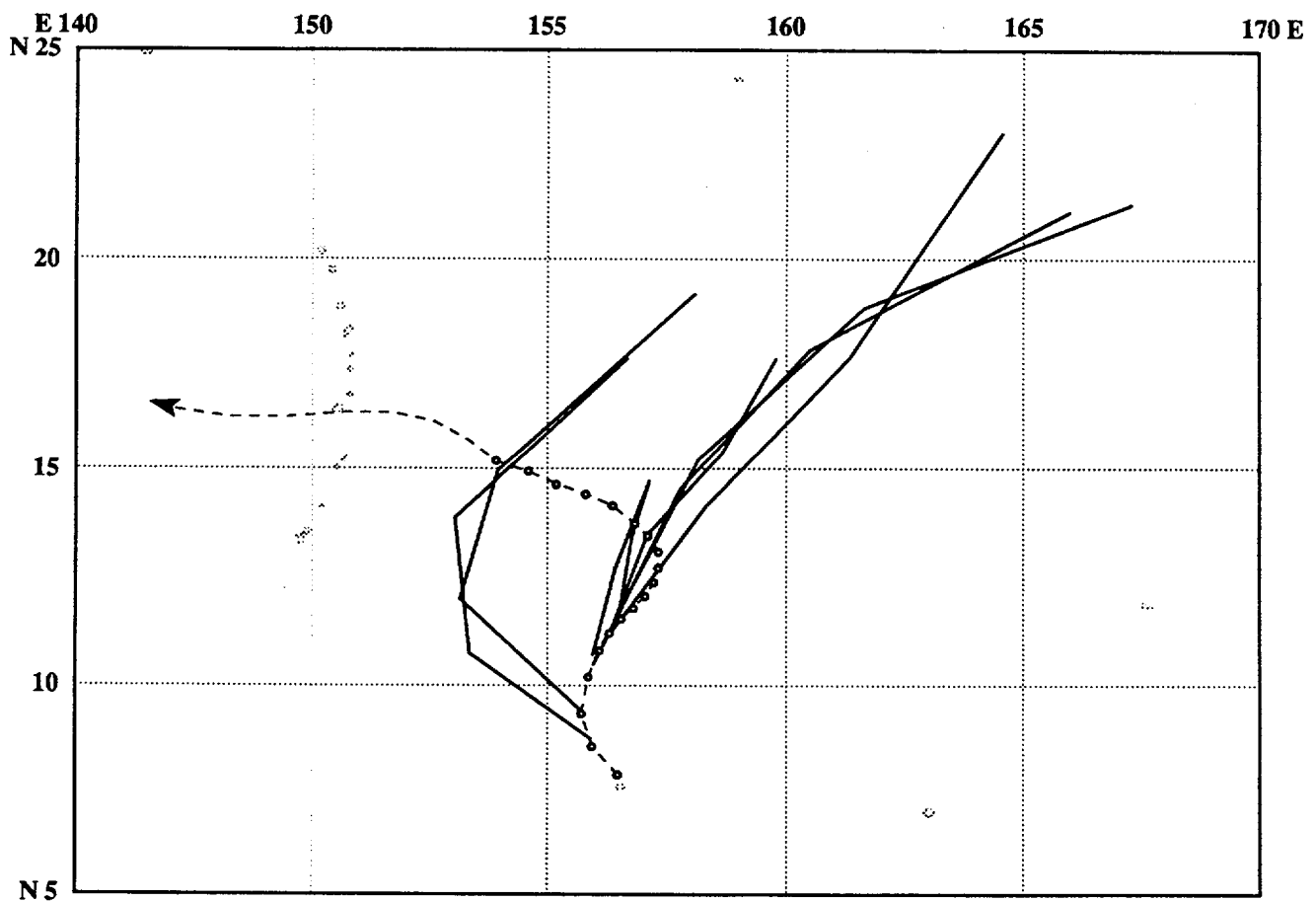
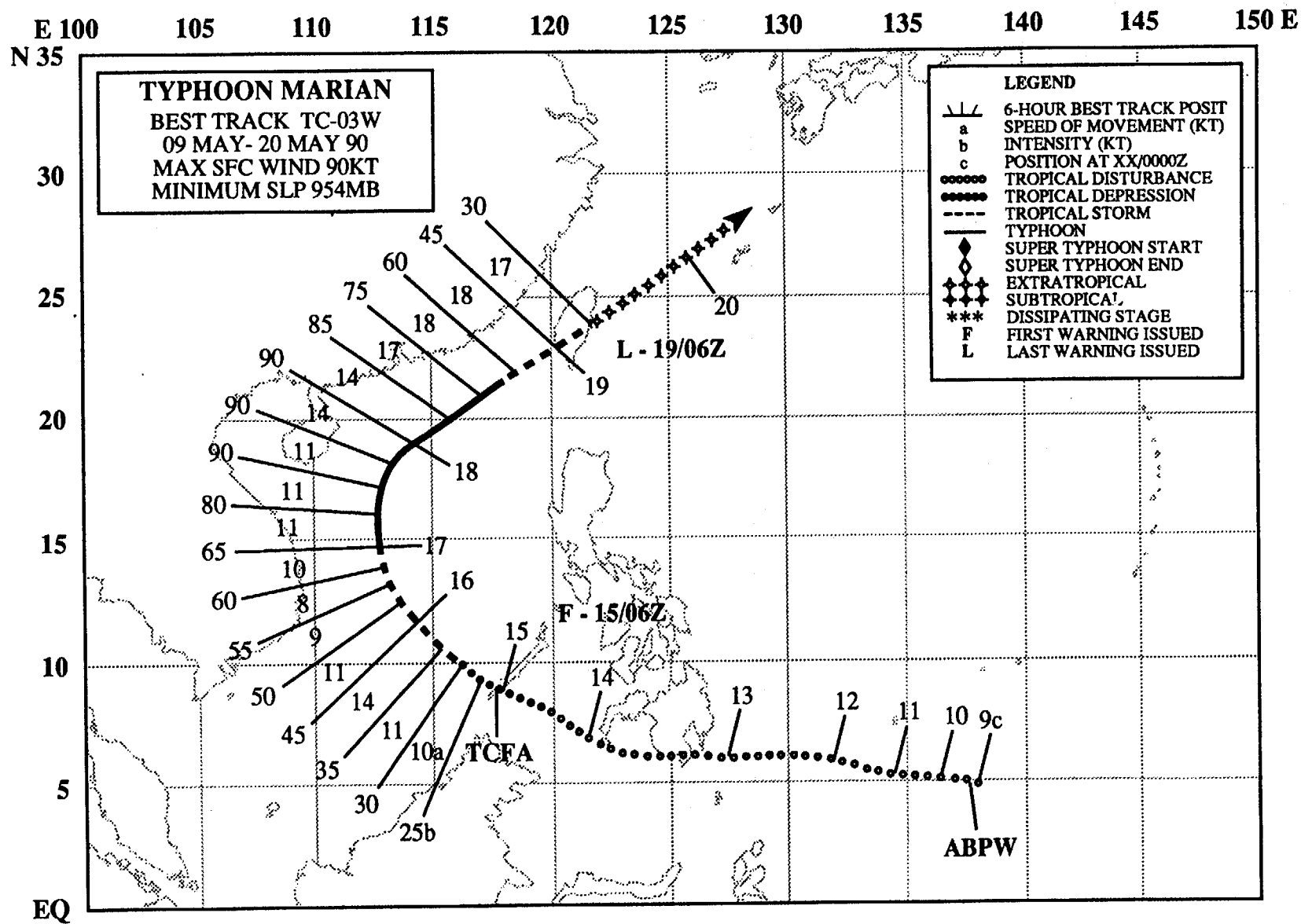


Figure 3-02-4. Summary of JTWC forecasts (solid lines) for Lewis (02W) superimposed on the final best track (dashed line).



TYPHOON MARIAN (03W)

I. HIGHLIGHTS

Marian, the second typhoon of 1990 in the western North Pacific and the only significant tropical cyclone to form in May, persisted in low latitudes for almost a week before intensifying. Its convective cloud mass tracked westward initially, passing south of Yap and Palau in the western Caroline Islands. After entering the South China Sea, the system finally developed into a typhoon. Marian then recurved and merged with a frontal system to form an extratropical low.

II. CHRONOLOGY OF EVENTS

- 090600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1006 mb.
- 150230Z - Tropical Cyclone Formation Alert based on better convective organization with increased low-level inflow and outflow aloft.
- 150600Z - First warning due to increased amount of central convection and cloud organization.
- 151800Z - Upgraded to a tropical storm prompted by steady intensification, favorable outflow aloft in all quadrants and the first intensity estimate of CI 2.5.
- 170000Z - Upgraded to typhoon following improved outflow, expected formation of an eye and the first CI 4.0.
- 171800Z - Peak intensity - 90 kt (46 m/sec) - coincident with visible eye with intensity estimate of CI 5.0.
- 181800Z - Downgraded to tropical storm because of increased vertical wind shear and start of extratropical transition. Convection decreased in amount and organization.
- 190600Z - Final warning - (extratropical) - followed interaction with rugged mountains of Taiwan. Principle low-level circulation center passed east of the island.

III. TRACK AND MOTION

The system developed in low latitudes in the central Caroline Islands and tracked slowly westward on the south side of the subtropical ridge. After passing over Mindanao in the southern Philippine Islands, Marian tracked around the western end of the subtropical ridge (Figure 3-03-1). As

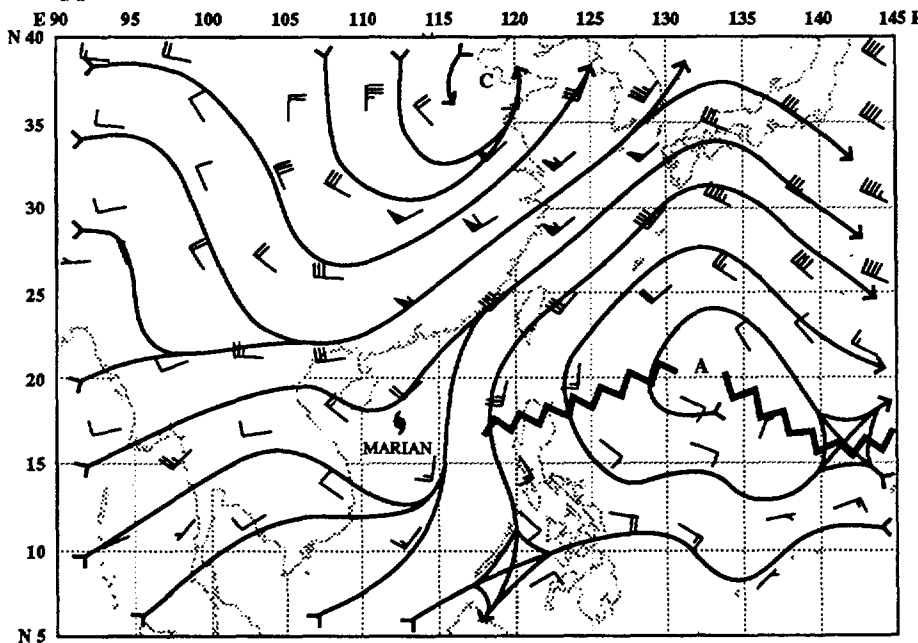


Figure 3-03-1. 500 mb NOGAPS analysis from 171200Z May, showing the cutoff low over eastern China, subtropical ridge to the east of Marian's surface position. The tropical cyclone, is tracking around the western periphery of the subtropical ridge and beginning to accelerate.

the tropical cyclone approached the south coast of China, increased southwesterlies aloft accelerated Marian northeastward along the edge of the modifying polar air.

IV. INTENSITY

The convective cloud mass that eventually developed into Typhoon Marian remained intact, but relatively unorganized, for almost a week. Brisk easterly trade winds (Figure 3-03-2) to the north and light cross-equatorial flow to the south supported the circulation, but outflow aloft was restricted by zonal westerly winds to the north. As the disturbance passed over the southern islands of the Philippine archipelago, interaction with land further inhibited low-level development. Upper-level conditions became favorable for intensification as a new outflow channel to the north combined with the preexisting weak one to the south and west. As the cyclone entered the South China Sea, it developed into Tropical Storm Marian. Steady intensification continued until an eye formed (Figure 3-03-3). After reaching peak intensity on 17 May, increased southwesterly flow aloft ahead of a shortwave

trough began to strip away the convection. As the system recurved, it was caught up in the approaching cold front and commenced extratropical transition (Figure 3-03-4).



V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-03-5. The initial forecasts did not call for recurvature. The NOGAPS prognostic series retained a weak mid-level ridge over the South China Sea, suggesting continued west-northwestward motion and eventual landfall in Vietnam. Because of the proximity of the shortwave trough over China, an alternate scenario was developed to weaken the subtropical ridge, allowing Marian to recurve. This alternate soon became the primary forecast, as the ridge did weaken and Marian recurved.

VI. IMPACT

No information was received.

Figure 3-03-2. Marian approaches the southern Philippine Islands. To the north of the central cloud mass at point A, low-level cloud arcs can be seen in the brisk easterly trade flow. Towering cumulus and cumulonimbus forming on these arcs were sheared away by westerly winds aloft (110051Z May DMSP visual imagery).

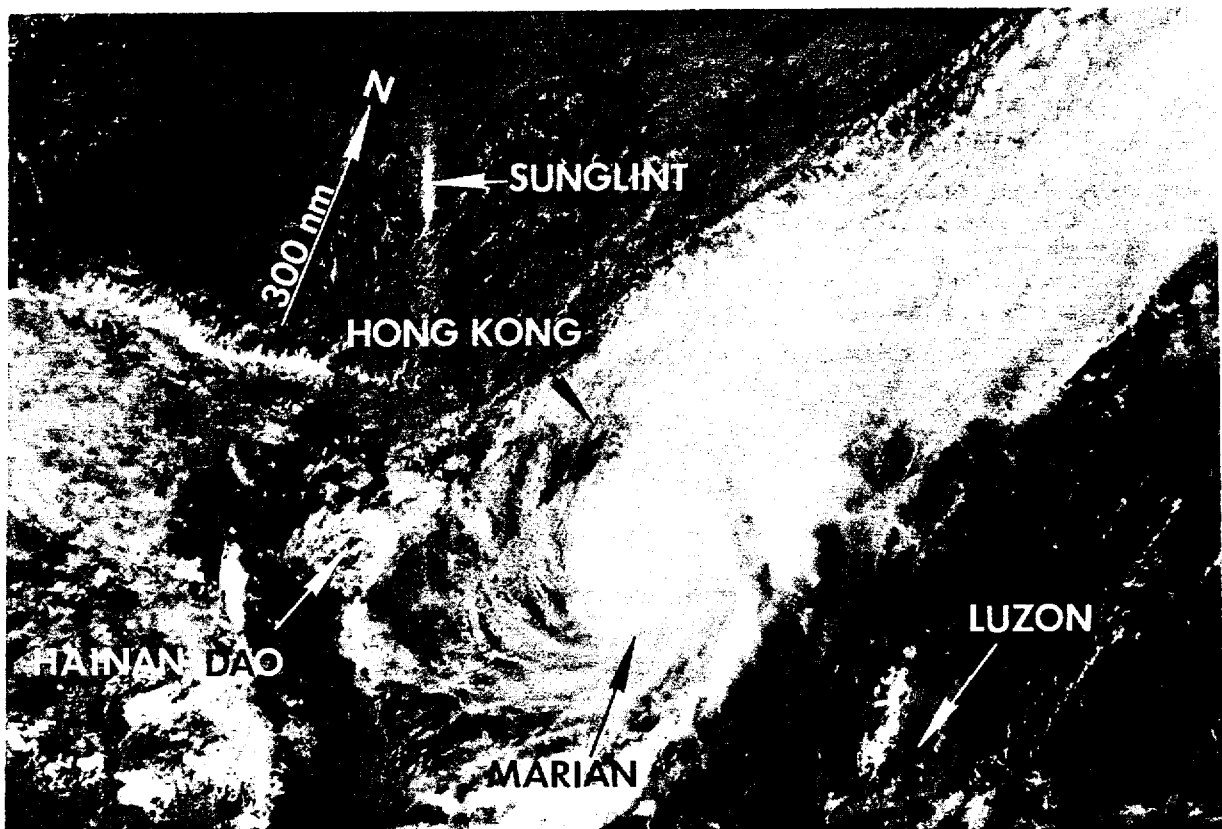


Figure 3-03-3. Typhoon Marian with a small eye interacts with a frontal system that is moving seaward from eastern Asia (180608Z May NOAA visual imagery).

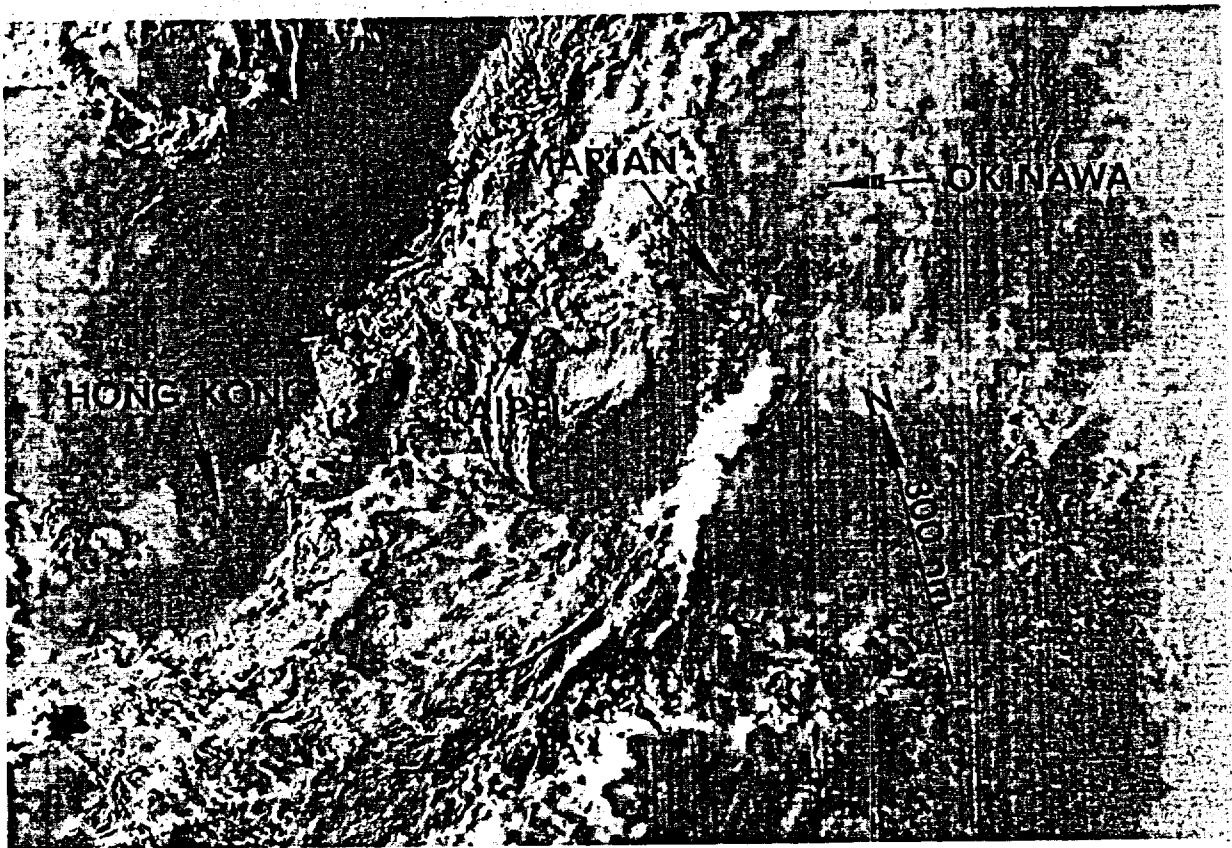


Figure 3-03-4. The remnants of Marian are embedded in the frontal zone just east of Taiwan. There appears to be no middle or high cloud in the subsiding air over the center of the vortex (191022Z May DMSP visual imagery).

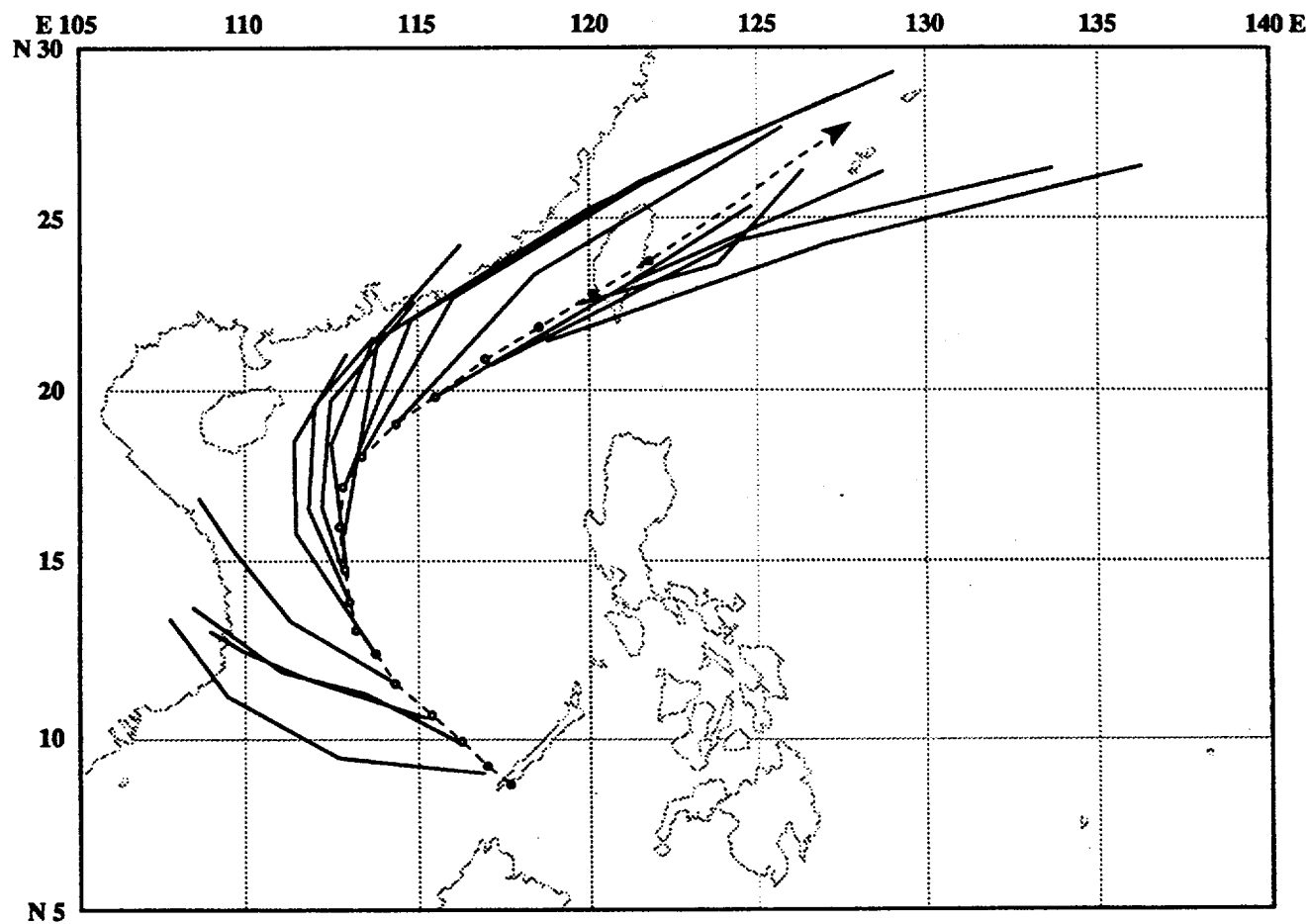
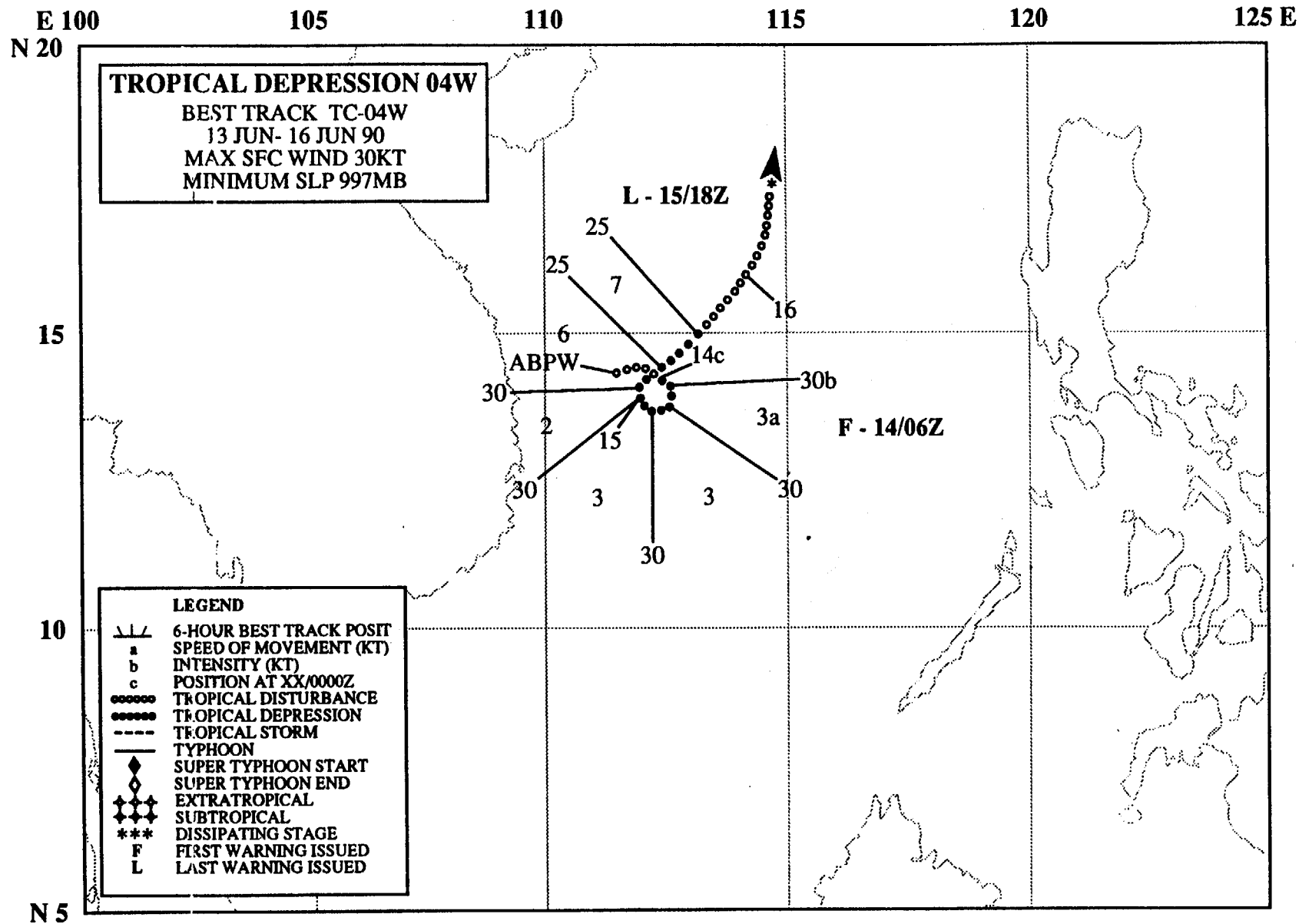


Figure 3-03-5. Summary of JTWC forecasts (solid lines) for Marian is superimposed on the final best track (dashed line).



TROPICAL DEPRESSION 04W

I. HIGHLIGHTS

Tropical Depression 04W, the first significant tropical cyclone to form in the South China Sea this year, proved to be very difficult to locate and forecast. Satellite and synoptic fix positions disagreed throughout the depression's life. As the convection flared near the center of the system, the mid-level and upper-level prevailing east-northeasterly flow moved the convection toward the coast of Vietnam. The satellite analysts tracked the convection onto the coast of Vietnam. However, as the area of convection over Vietnam dissipated a new area of convection developed near the circulation center indicated in the synoptic data. As Tropical Storm Nathan (05W) continued to develop, Tropical Depression 04W was drawn into the larger circulation and absorbed.

II. CHRONOLOGY OF EVENTS

- 130600Z - First mentioned on Significant Tropical Weather Advisory due to weak low-level circulation center in the synoptic data and 1004 mb pressure.
- 140600Z - First Warning based on the low-level circulation center exposed to the east of the poorly organized central cloud mass. Synoptic data indicated the presence of 20-30 kt (10-15 m/sec) winds.
- 140600Z - Peak Intensity of 30 kt (15 m/sec) established in synoptic data.
- 151200Z - Final warning followed the loss of convective signature as the low level circulation was absorbed by Nathan (05W).

III. MOTION

Tropical Depression 04W proved to be a significant motion forecast problem. From the beginning, the 850 mb wind patterns in the area indicated that the vortex was located along the western side of the low-level mean wind flow of approximately 30 kt (15 m/sec) from the west-southwest associated with the summer monsoon. The depression remained quasi-stationary for the first two days. As Tropical Storm Nathan (05W) moved into the South China Sea, strong southwesterly monsoon flow began to feed into it. Tropical Depression 04W (Figure 3-04-1) became involved in the associated broad scale flow and was absorbed by the larger cyclone.

IV. INTENSITY

The strong vertical wind shear always restricted Tropical Depression 04W development. The 200-mb winds over the area were 30 to 35 kt (15 to 18 m/sec) and the low-level monsoonal flow was of equal intensity and opposing direction. As a result of the strong shear, JTWC did not expect intensification above 30 kt (15 m/sec) and issued only 36-hour tropical depression warnings.

V. FORECASTING PERFORMANCE

Superimposed on the final best track are the JTWC forecasts (Figure 3-04-2). Due to the lack of synoptic data in the early portions of the forecast scenario, JTWC depended primarily on satellite fixes to determine Tropical Depression 04W's location. In this high vertical wind shear environment the satellite fixes indicated an apparent westward motion of the system. Thus, JTWC forecast aids and the official forecast track indicated westward motion for most of the life of the depression.

VI. IMPACT

No impact was reported in association with Tropical Depression 04W.

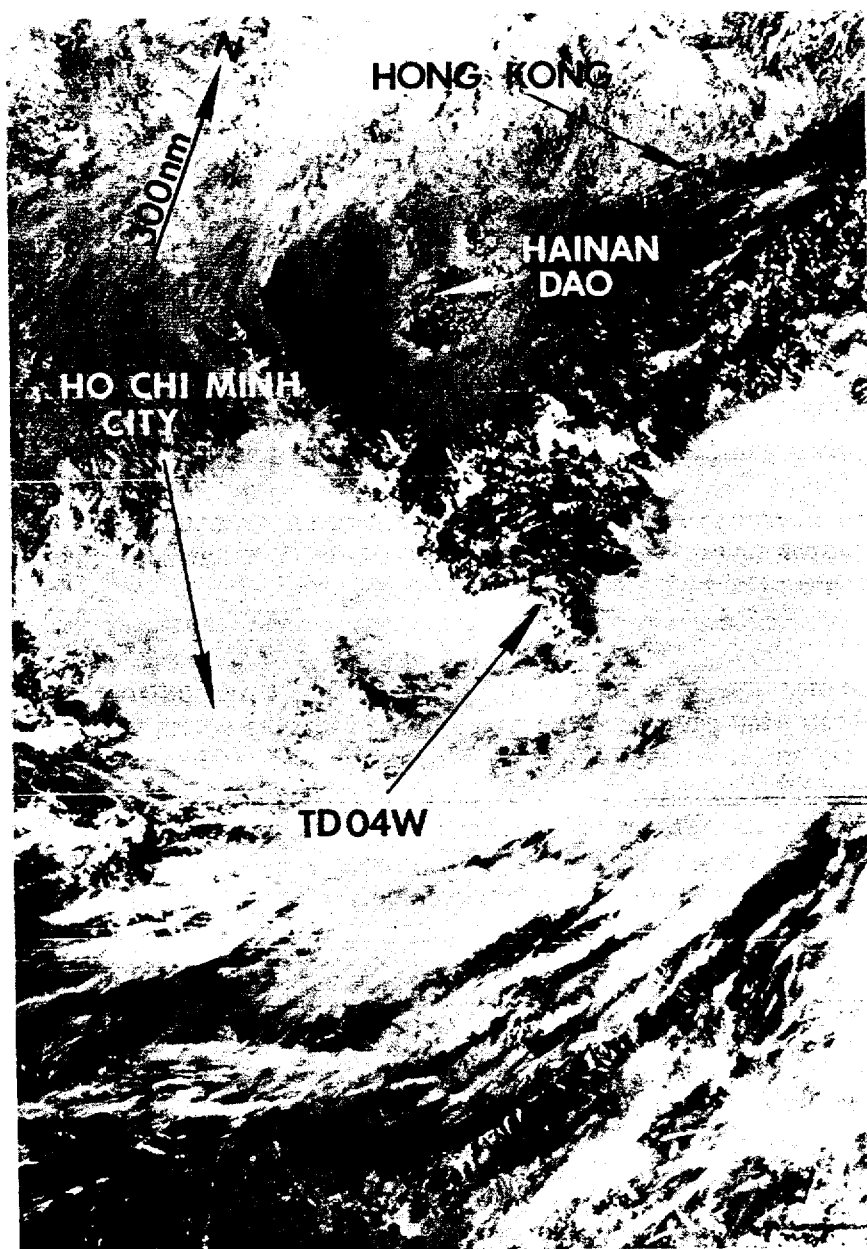


Figure 3-04-1. Tropical Depression 04W, which is south-southeast of Hainan Dao, becomes involved with Tropical Storm Nathan (05W) (150210Z June DMSP visual imagery).

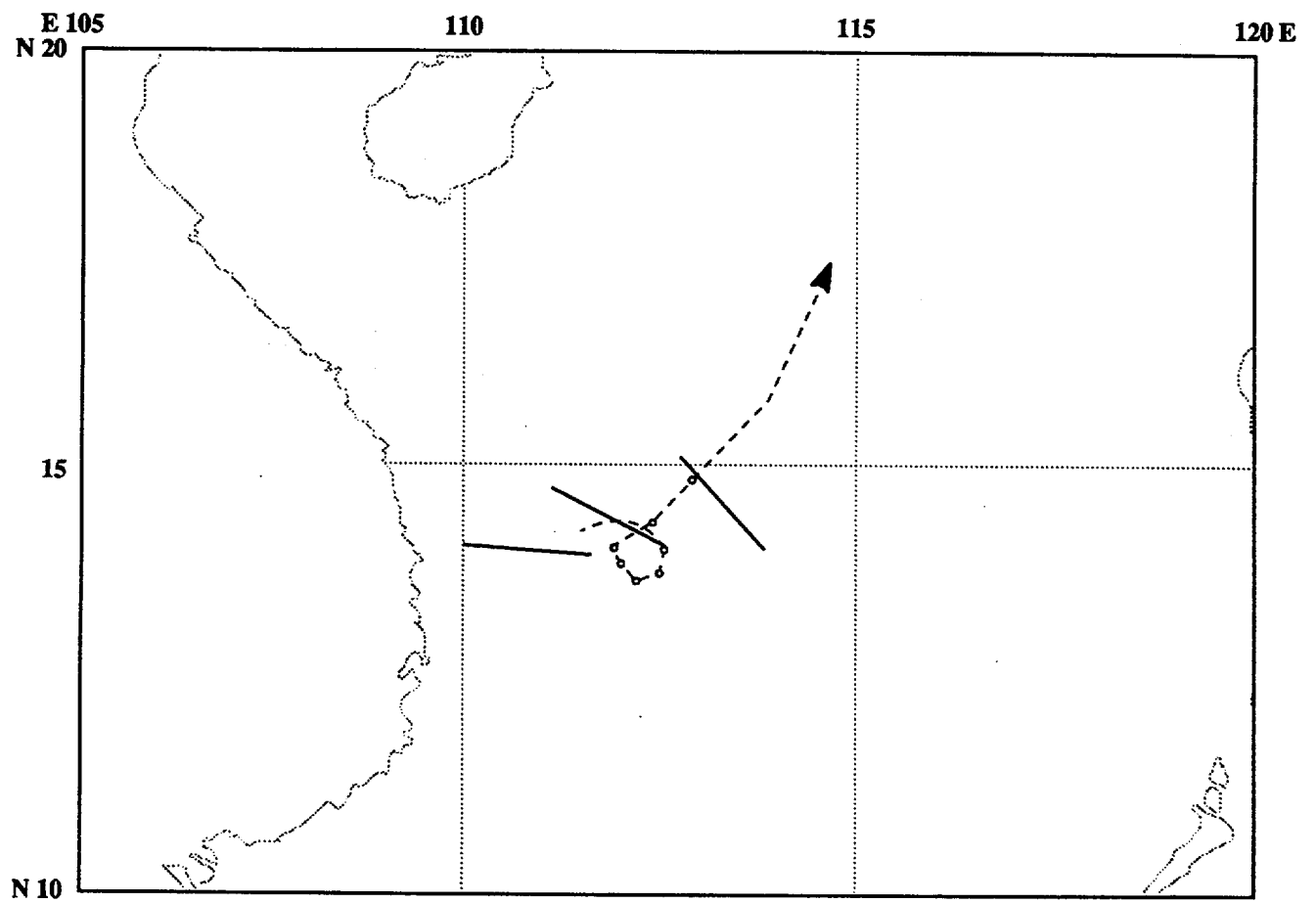
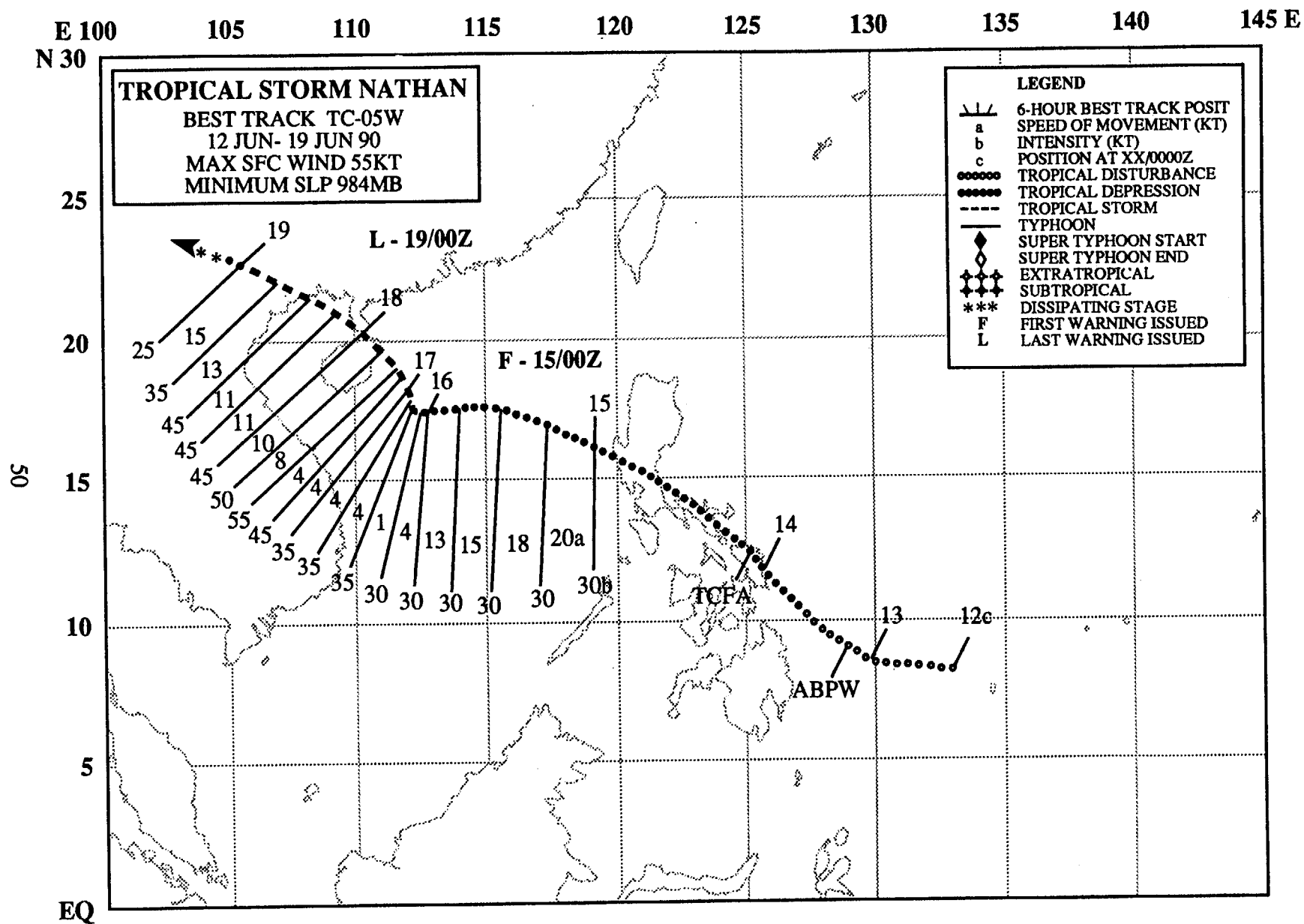


Figure 3-04-2. Summary of JTWC forecasts (solid lines) superimposed on the final best track (dashed line).



TROPICAL STORM NATHAN (05W)

I. HIGHLIGHTS

Nathan, the second tropical cyclone to form in June, crossed the Philippine island of Luzon as a disturbance, executed an abrupt track change and stalled in the South China Sea. Both the track and intensity of TD04W and Nathan were dominated by a larger monsoon circulation in the South China Sea.

II. CHRONOLOGY OF EVENTS

- 130600Z - First mentioned on Significant Tropical Weather Advisory as an area of weak circulation with an estimated minimum sea-level pressure of 1004 mb embedded in the monsoon trough.
- 140300Z - Tropical Cyclone Formation Alert based on improved organization with increased low-level inflow and increased outflow aloft.
- 150000Z - First warning due to increased winds as the system came off Luzon and entered the warm waters of the South China Sea.
- 161200Z - Upgraded to tropical storm after system became quasi-stationary and the exposed low-level became more aligned with the deep convection; the first intensity estimate of CI 2.5 received.
- 171200Z - Peak intensity - 55 kt (28 m/sec) -based on a ship report of 50 kt (26 m/sec) winds within 55 nm (100 km) of cloud system center.
- 181200Z - Landfall along Chinese/Vietnamese border, 100 nm (185 km) east-northeast of Hanoi.
- 190000Z - Final warning - (dissipated over land)- followed rapid weakening as Nathan encountered the mountains of northern Vietnam.

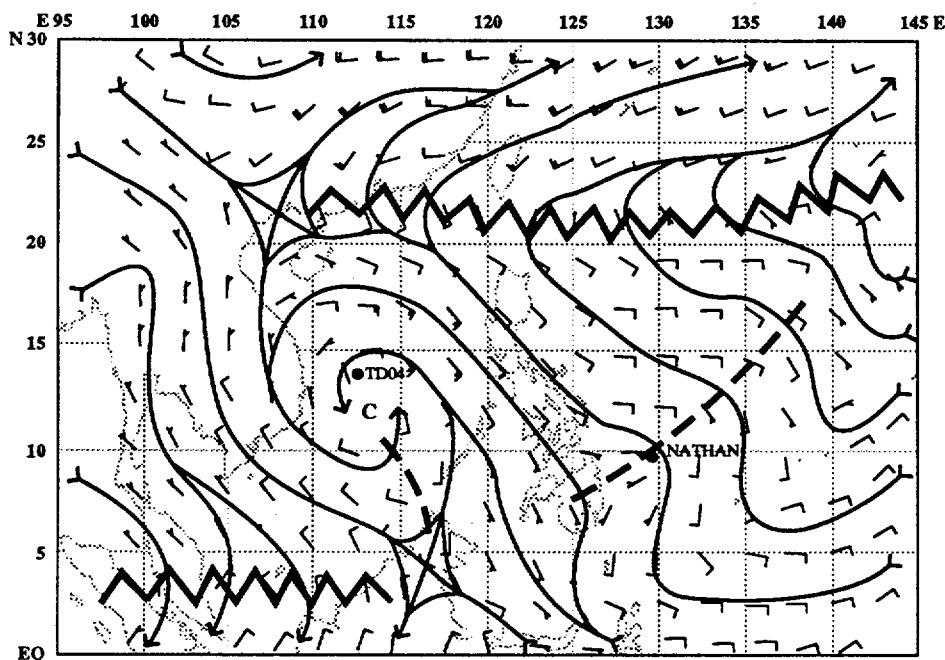


Figure 3-05-1. The 131200Z June deep layer mean analysis shows the large monsoon circulation (LMC) near 11° north latitude in the South China Sea and the subtropical ridge near 20° north latitude. Tropical Depression 04W is northwest of the center of the LMC and Nathan appears as an inverted trough east of Mindanao.

III. TRACK AND MOTION

A large monsoon circulation (hereafter called LMC) in the South China Sea and the subtropical ridge along 20° north latitude set the stage for Nathan's unusual track. Initially Nathan was reflected in the deep layer mean analysis (Figure 3-05-1) as a wave in the easterlies. Farther to the west Tropical Depression 04W was a smaller shallow circulation embedded within the synoptic scale LMC. As Nathan moved northwestward and crossed southern Luzon, both the subtropical ridge and the LMC began shifting

northward (Figure 3-05-2). The curved best track reflects both Nathan's westward movement into the LMC in the South China Sea and the displacement to the north of the entire synoptic pattern. For a time, Tropical Depression 04W was expected to be drawn into Nathan; however, as Nathan sped by, Tropical Depression 04W dissipated. Nathan's abrupt track change and stall on 16 June was the result of a binary interaction with the LMC. The tropical cyclone separated from the LMC core and continued northwestward (Figure 3-05-3).

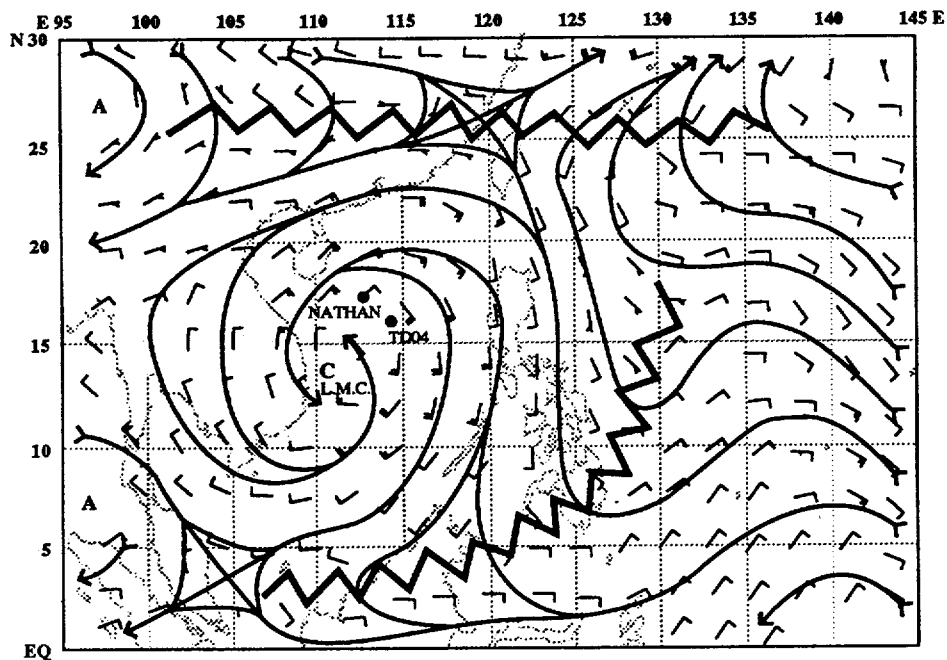


Figure 3-05-2. Both TD04W and Nathan are embedded in the flow near the center of the LMC on the 160000Z June deep layer mean analysis.

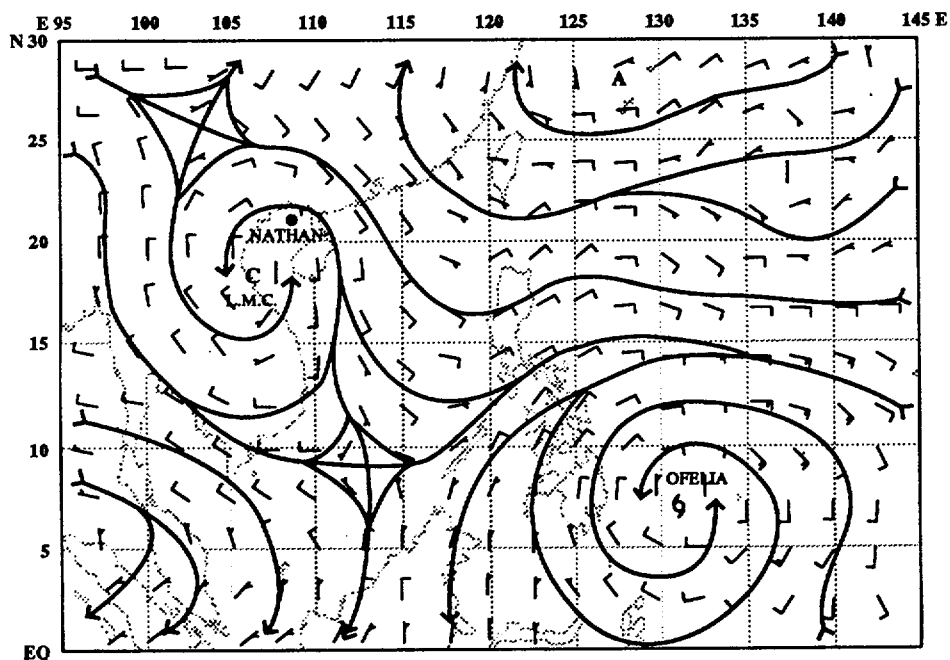


Figure 3-05-3. The 181200Z June deep layer mean analysis shows Nathan north-northeast of the center of the LMC.

IV. INTENSITY

Nathan slowly consolidated from multiple low-level circulations in an area of poorly organized convection. Convection continued to increase in amount and organization as the system approached the Philippine Islands (Figure 3-05-4). Nevertheless, passage across Luzon, rapid motion toward the LMC in the South China Sea and strong vertical wind shear all kept Nathan below tropical storm intensity. Intensification finally occurred when Nathan entered the core of the LMC on 16 June. The shear-type cloud pattern with its exposed low-level circulation center gave way to a central dense overcast, and

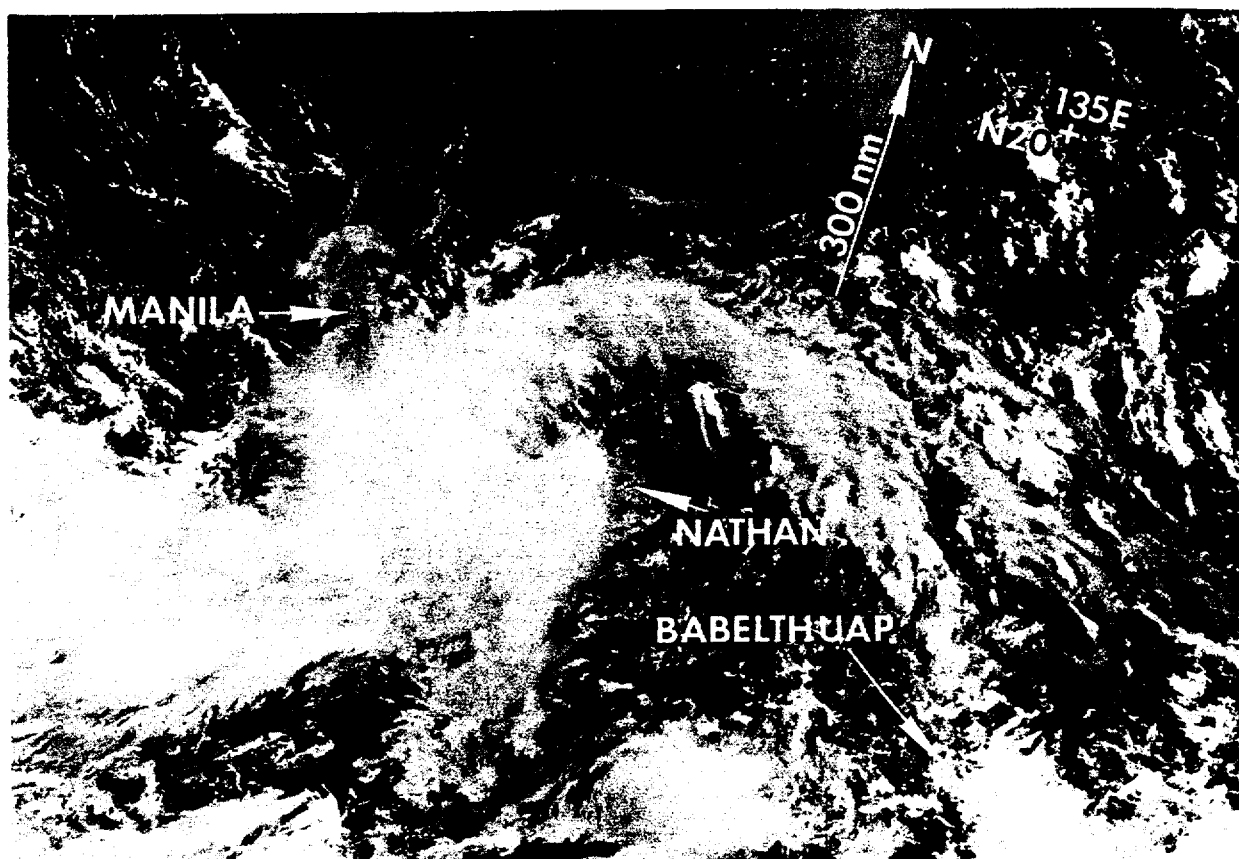


Figure 3-05-4. Nathan consolidates as it approaches the Philippine Islands (140049Z June DMSP visual imagery).

Nathan intensified into a tropical storm (Figure 3-05-5). Slow intensification continued until the tropical cyclone began interacting with land. Nathan weakened and dissipated rapidly after crossing Hainan Dao and making landfall on the coast of Vietnam on 18 June.

V. FORECASTING PERFORMANCE

Plots of JTWC's forecasts on the best track are presented in Figure 3-05-6. JTWC had a difficult time with this tropical cyclone in the South China Sea. Nathan's interaction with the center of the LMC and the northward shift of the entire synoptic pattern became apparent only after the fact. OTCM had a better handle on the overall northwestward track (Figure 3-05-7), but it did not reflect the interaction with the core of the LMC, as can be seen by the OTCM guidance to the south on 15 June.

VI. IMPACT

In Hong Kong, according to the "Monthly Weather Summary June 1990" published by the Royal Observatory, 13 people were killed, 5 were missing and 15 injured as a result of Nathan. Minor mudslides were reported throughout the area and scaffoldings collapsed in Kowloon. The cargo ship "Tien Fu" sank in the South China Sea on the night of 16 June with the loss of the captain and three of its crew. Along China's southern coast, torrential rain associated with Nathan caused 10 deaths and flooded 5,000 hectares of farmland in eastern Guangdong. In Zhanjiang, 100,000 hectares of paddy fields were destroyed. Two men were reported missing in Macao after being swept overboard from a dredger on 17 June.

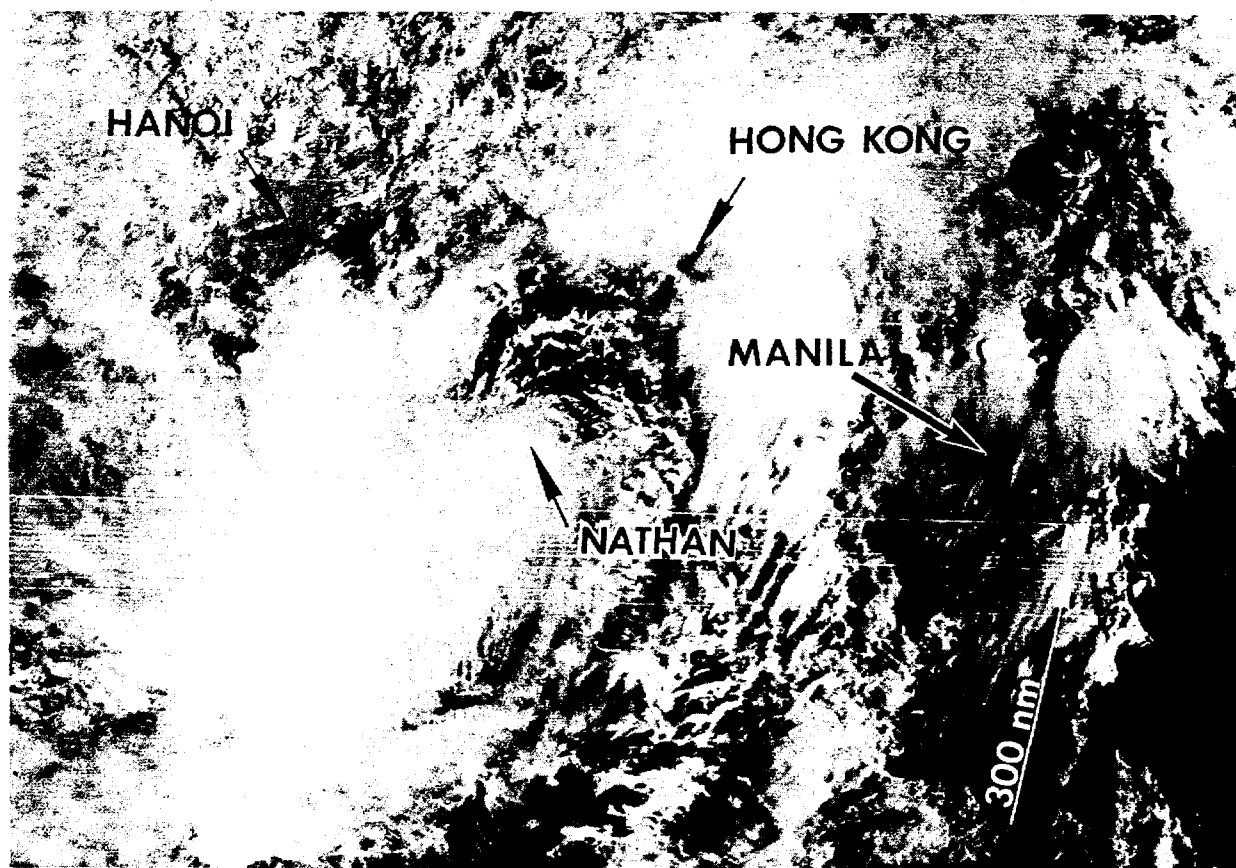


Figure 3-05-5. Tropical Storm Nathan with a ragged central dense overcast churns towards Hainan Dao (170128Z June DMSP visual imagery).

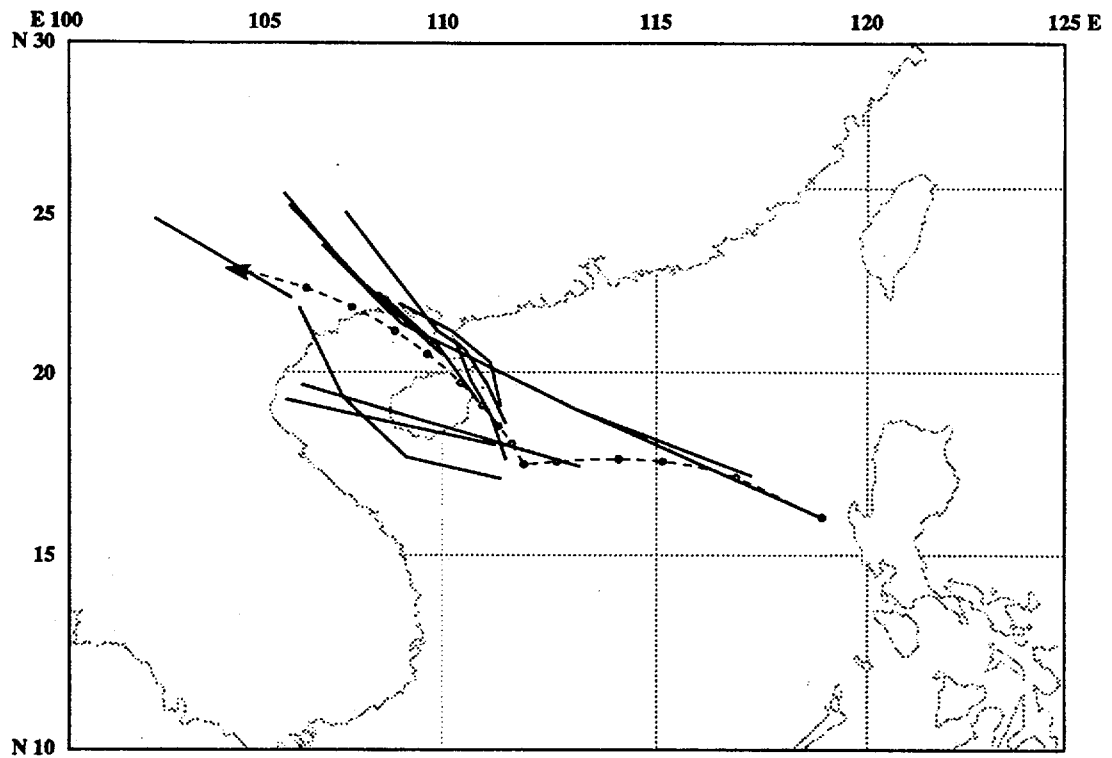


Figure 3-05-6. JTWC forecasts (solid lines) for Nathan are superimposed on the final best track (dashed line). The abrupt track change and stall in the South China Sea were difficult to forecast.

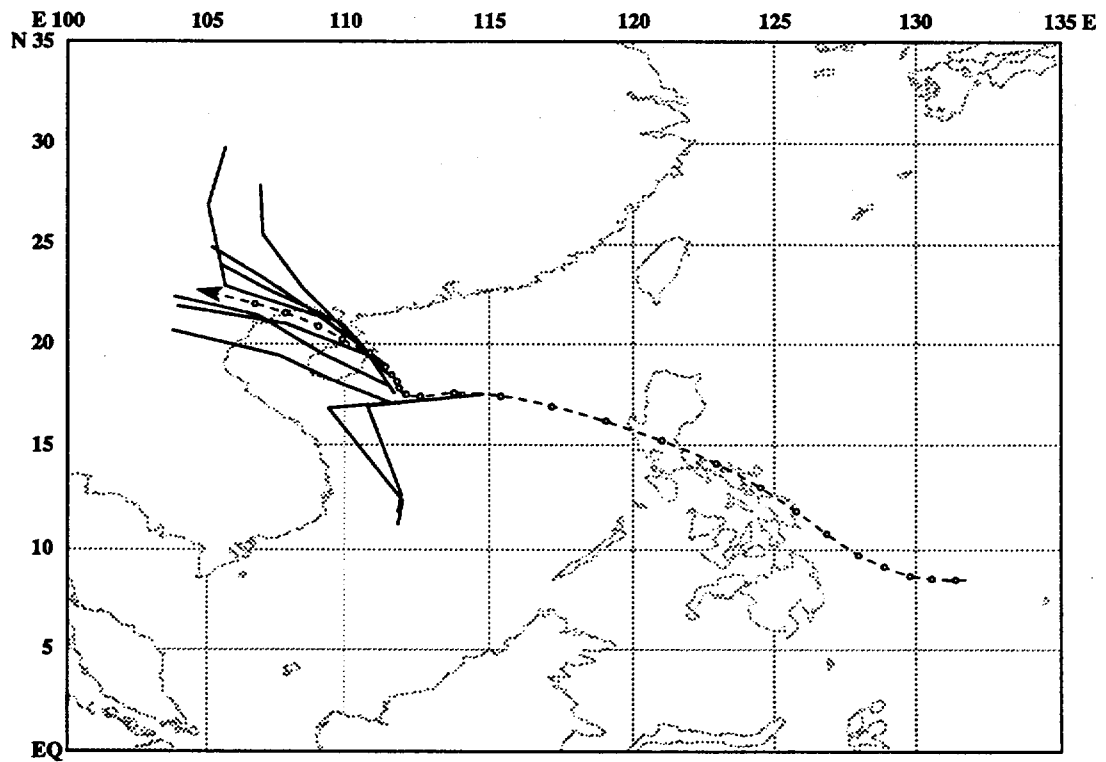
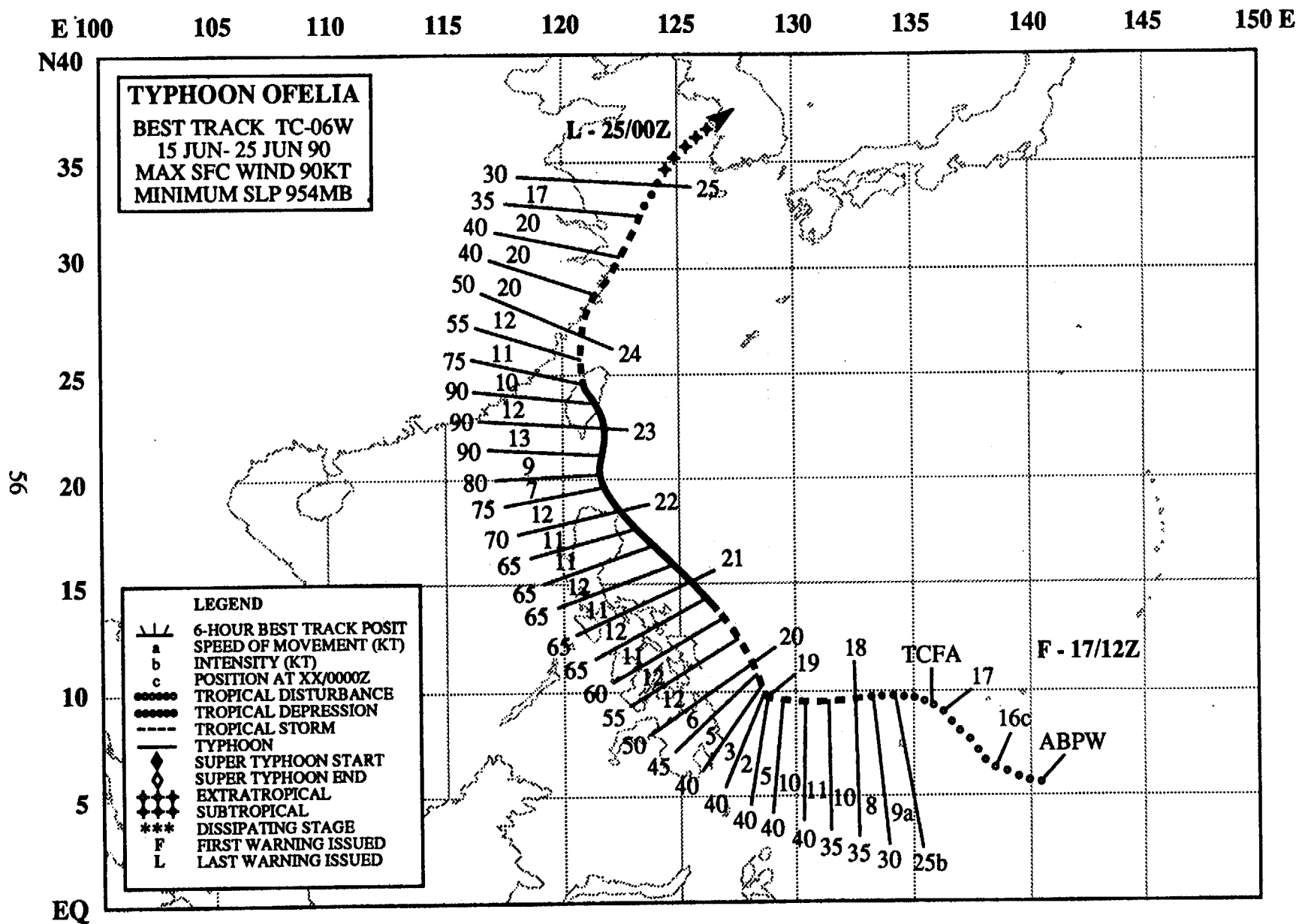


Figure 3-05-7. When Nathan interacted with the core of the LMC, OTCM guidance had difficulties, as indicated by the two solid lines that abruptly turn southward on 15 June.



TYPHOON OFELIA (06W)

I. HIGHLIGHTS

Ofelia was the third western North Pacific typhoon of 1990 and the first for the month of June. It moved toward the Philippine Islands, then slowed and turned to the northwest. Ofelia became the second tropical cyclone of the year to hit Taiwan and the first to affect the east coast of China. After recurvature, the extratropical remnants of Ofelia crossed Korea, unusual for a June system.

II. CHRONOLOGY OF EVENTS

- 150600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with estimated maximum winds of 15 kt.
- 170430Z - Tropical Cyclone Formation Alert based on increased convection during diurnal minimum, more curvature to the cloud bands, and better outflow aloft.
- 171200Z - First warning due to improved cloud signature.
- 180000Z - Upgraded to tropical storm prompted by an intensity estimate of CI 2.5.
- 201800Z - Upgraded to typhoon based on well-defined central dense overcast and overshooting cloud tops.
- 230000Z - Peak intensity - 90 kt (46 m/sec) - based on appearance of an eye and a CI 5.0 estimate.
- 231800Z - Downgraded to tropical storm after crossing Taiwan and weakening due to land effects.
- 250000Z - Final warning - (extratropical) - as cyclone merged with a frontal boundary while approaching the Korean Peninsula.

III. TRACK AND MOTION

Ofelia developed in the monsoon trough in the central Caroline Islands and tracked westward along the periphery of the subtropical ridge. On 19 June the tropical cyclone slowed and executed an abrupt track change to the northwest. Although the NOGAPS 500-mb analysis (Figure 3-06-1) at 190000Z June failed to show any significant reason for the track anomaly, the 850-mb analysis (Figure 3-06-2) revealed the presence of 30 to 40 kt (15 to 21 m/sec) southwesterly flow. Since the heights and

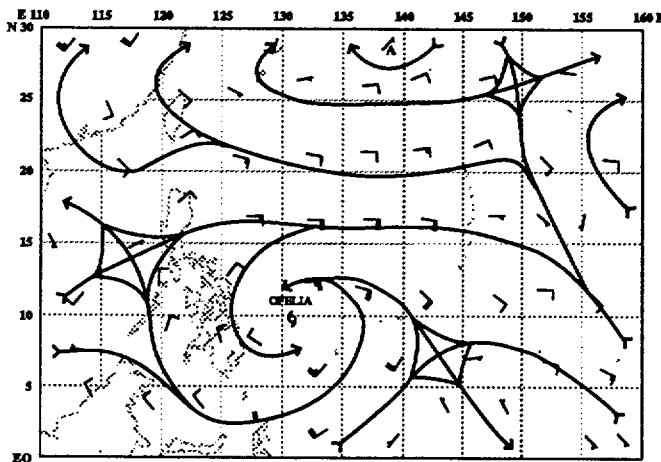


Figure 3-06-1. The 190000Z June NOGAPS 500-mb analysis shows a roughly balanced flow around Ofelia.

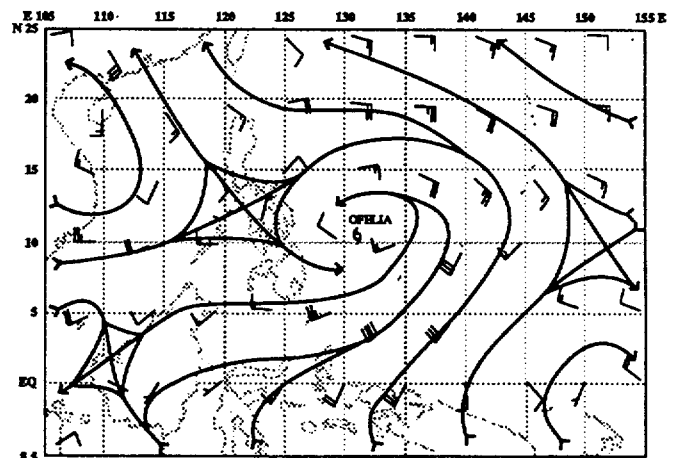


Figure 3-06-2. The 190000Z June NOGAPS 850-mb analysis reveals a stronger southwesterly inflow into the tropical cyclone.

patterns of the subtropical ridge to the north were relatively unchanged, it appears that the start of a shallow monsoon surge from the southwest into Ofelia disrupted the normal steering current. By 20 June a balance between the monsoon steering and the ridge steering had returned, and the tropical cyclone continued tracking around the ridge. On 22 June, when Ofelia was in the Bashi Channel between Luzon and Taiwan, the southwesterly monsoon flow at 850mb (Figure 3-06-3) broadened and reached 50 kt (26 m/sec) over the central Philippine Islands. This flow also deepened through the middle troposphere, where 40 kt (21 m/sec) winds appeared on the 500-mb analysis (Figure 3-06-4). It appears that as Ofelia approached 20° north latitude, the strength of the surge temporarily resulted in a more northward track. Soon after, the typhoon took a northwestward slide across Taiwan, then reached the axis of the subtropical ridge and began recurving toward Korea.

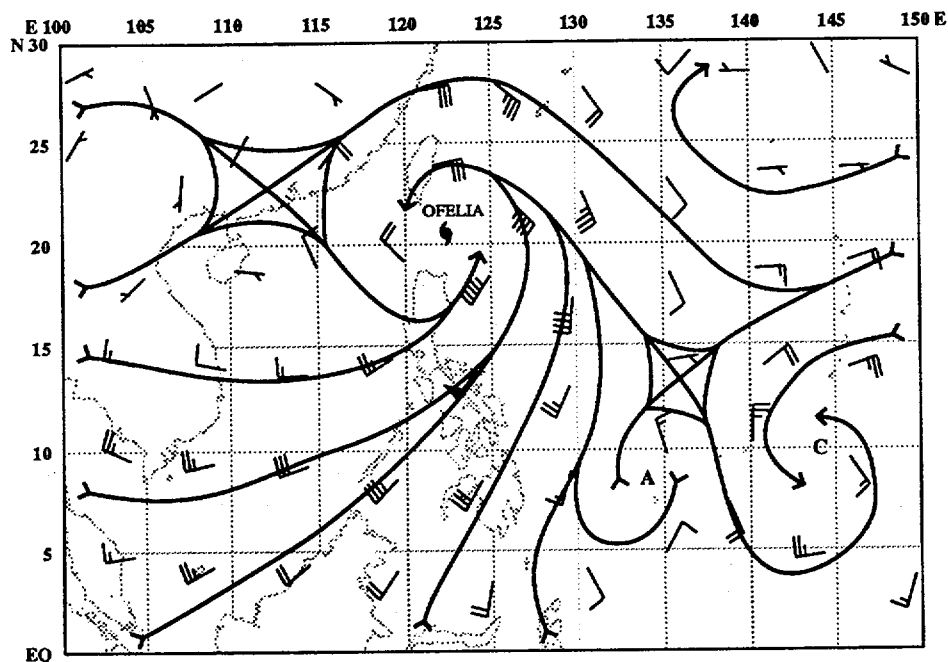


Figure 3-06-3. The 221200Z June NOGAPS 850-mb analysis shows the broad southwest monsoon flow with 50 kt (26 m/sec) across the central Philippine Islands.

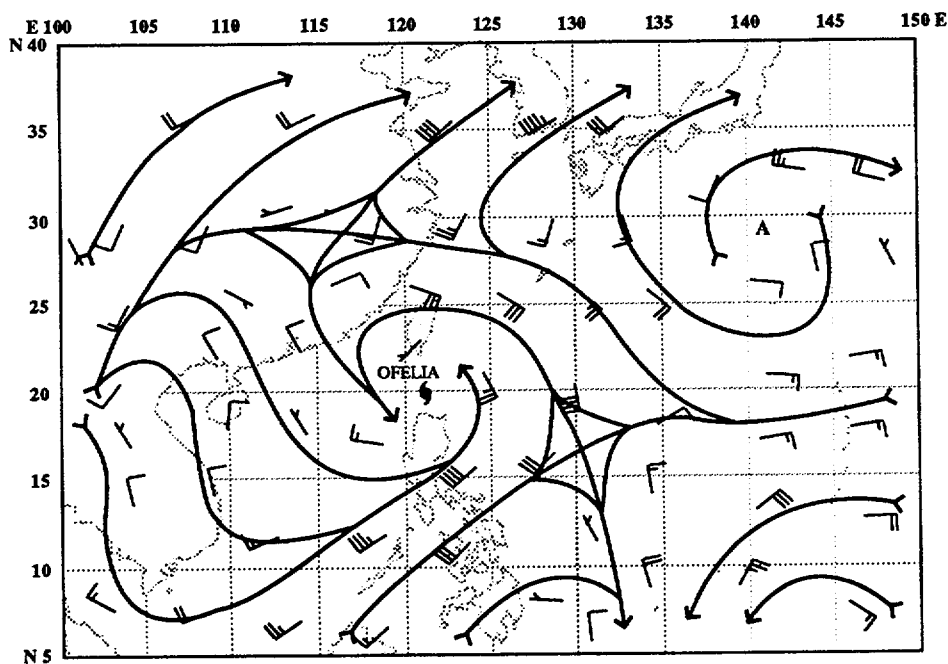


Figure 3-06-4. The 221200Z June NOGAPS 500-mb analysis indicates that the southwesterly flow extends well up into the middle troposphere.

IV. INTENSITY

The tropical depression which was to become Ofelia was initially slow to develop due to vertical wind shear from the northeast. As the southwesterly inflow into the tropical cyclone increased and deepened, an anticyclone formed aloft and the vertical wind shear decreased. Ofelia (Figure 3-06-5) intensified at a slower than average rate and peaked at 90 kt (46 m/sec) (Figure 3-06-6), five days after reaching tropical storm intensity. Part of this slower than average rate was caused by land influences from the Philippine Islands to the west of track. Rapid weakening after 230000Z was caused by land interaction, as the cloud system crossed the mountainous island of Taiwan and moved northward over the China coast.

V. FORECASTING PERFORMANCE

The NOGAPS series kept the subtropical ridge across the Philippine Sea north of the cloud system and linked it to the ridge over central China. JTWC initially expected a more westward track for the system, and continued to forecast the track too far to the west until the system approached Taiwan (Figure 3-06-7). The bias to the west of track appeared in the NOGAPS guidance and suggests that the influence of the strength and depth of the southwest monsoonal flow on Ofelia may not have been correctly addressed by the NOGAPS model.

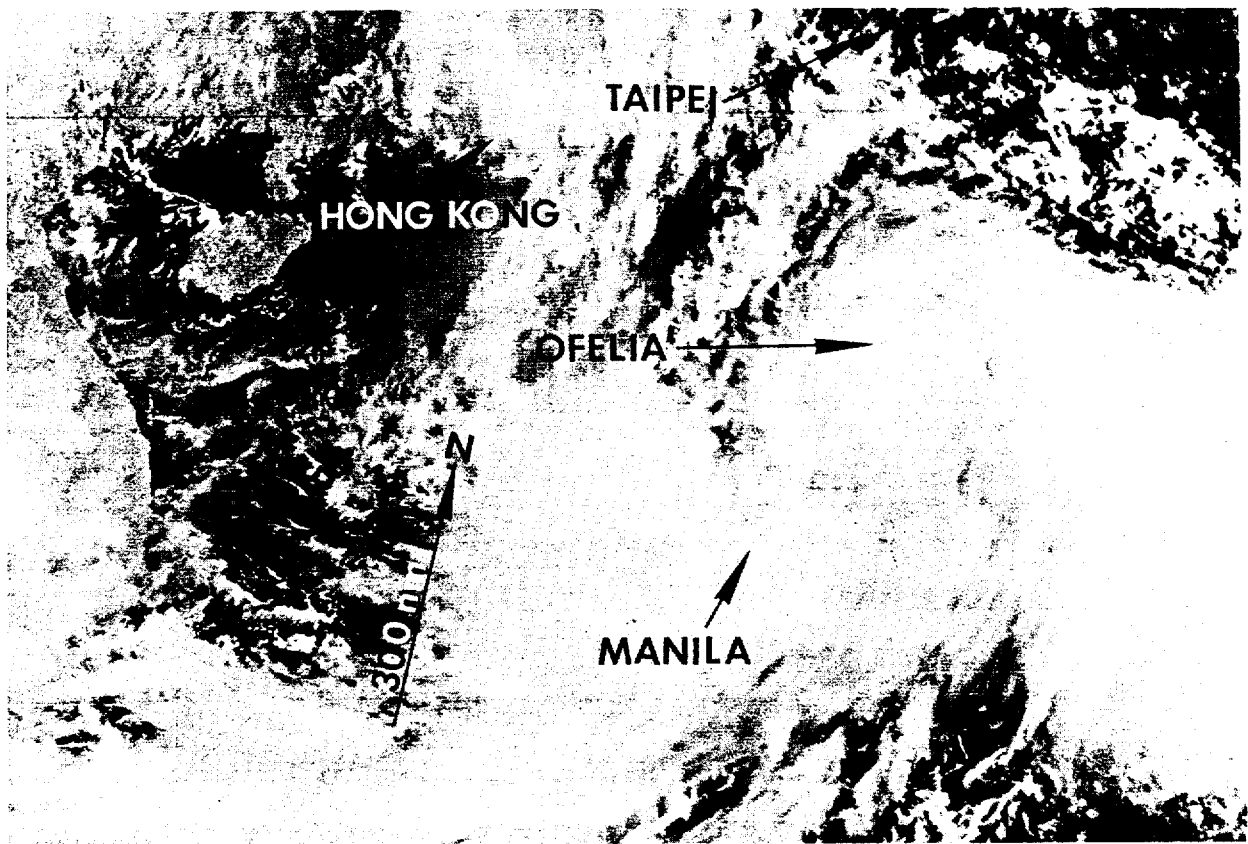


Figure 3-06-5. Typhoon Ofelia is located north of Luzon. The overcast conditions over the Philippine Islands are associated with the deep southwesterly inflow into the typhoon (220124Z June DMSP visual imagery).

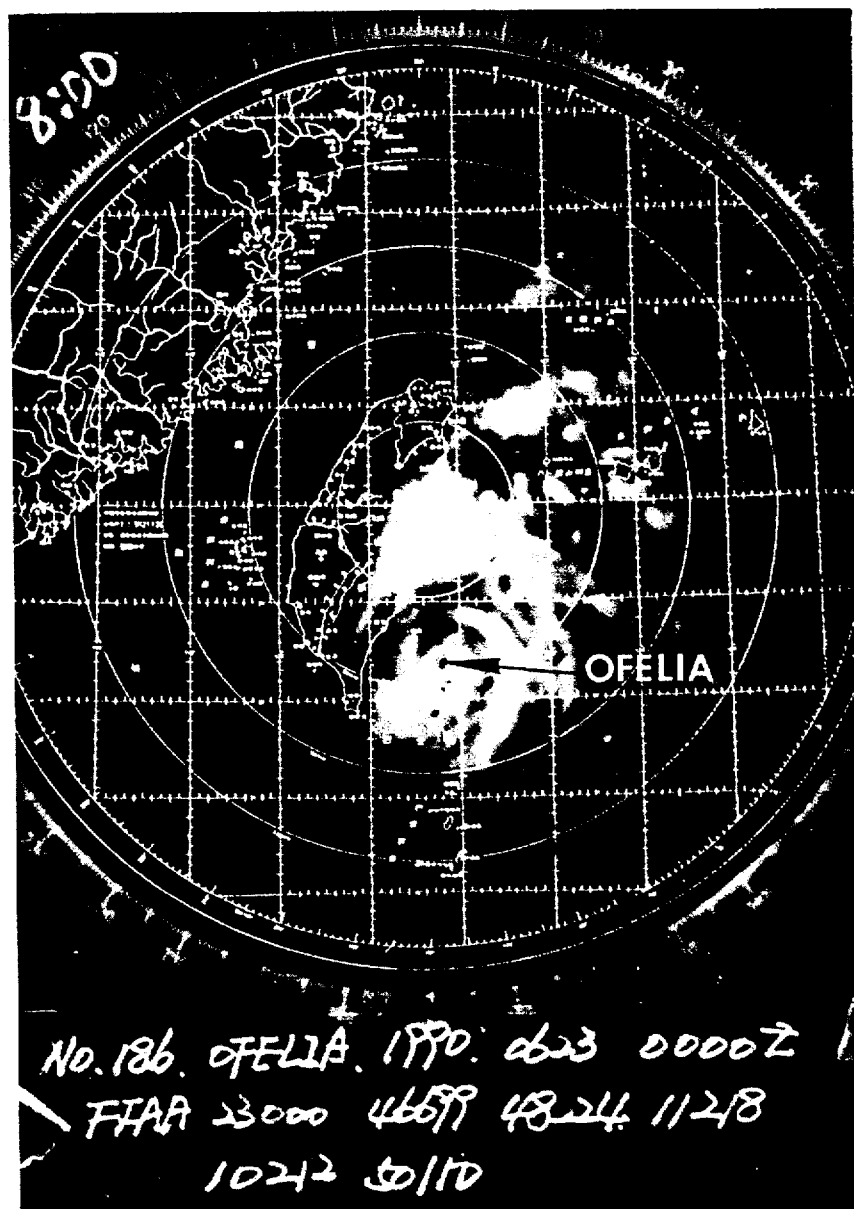


Figure 3-06-6. The 230000Z June radar image from Haulien, Taiwan (WMO 46699) of Ofelia at peak intensity. A small eye is present (radar photo courtesy of the Central Weather Bureau, Taipei, Taiwan).

VI. IMPACT

Ofelia was a destructive system. Although it didn't cross directly over northern Luzon, the system caused a surge in the southwest monsoon which resulted in torrential rains and widespread flooding in the northern Philippine Islands. Newspaper reports indicated that more than 25 people died and over 84,000 were forced to flee their homes. Taiwan took a direct hit from Ofelia. Media releases said the storm was the worst to hit eastern Taiwan in 30 years. Seventeen people died and 23 were missing due to floods and mud slides. In central China, at least 22 were killed as Ofelia, which caused flooding to low-lying provinces, moved up the coast.

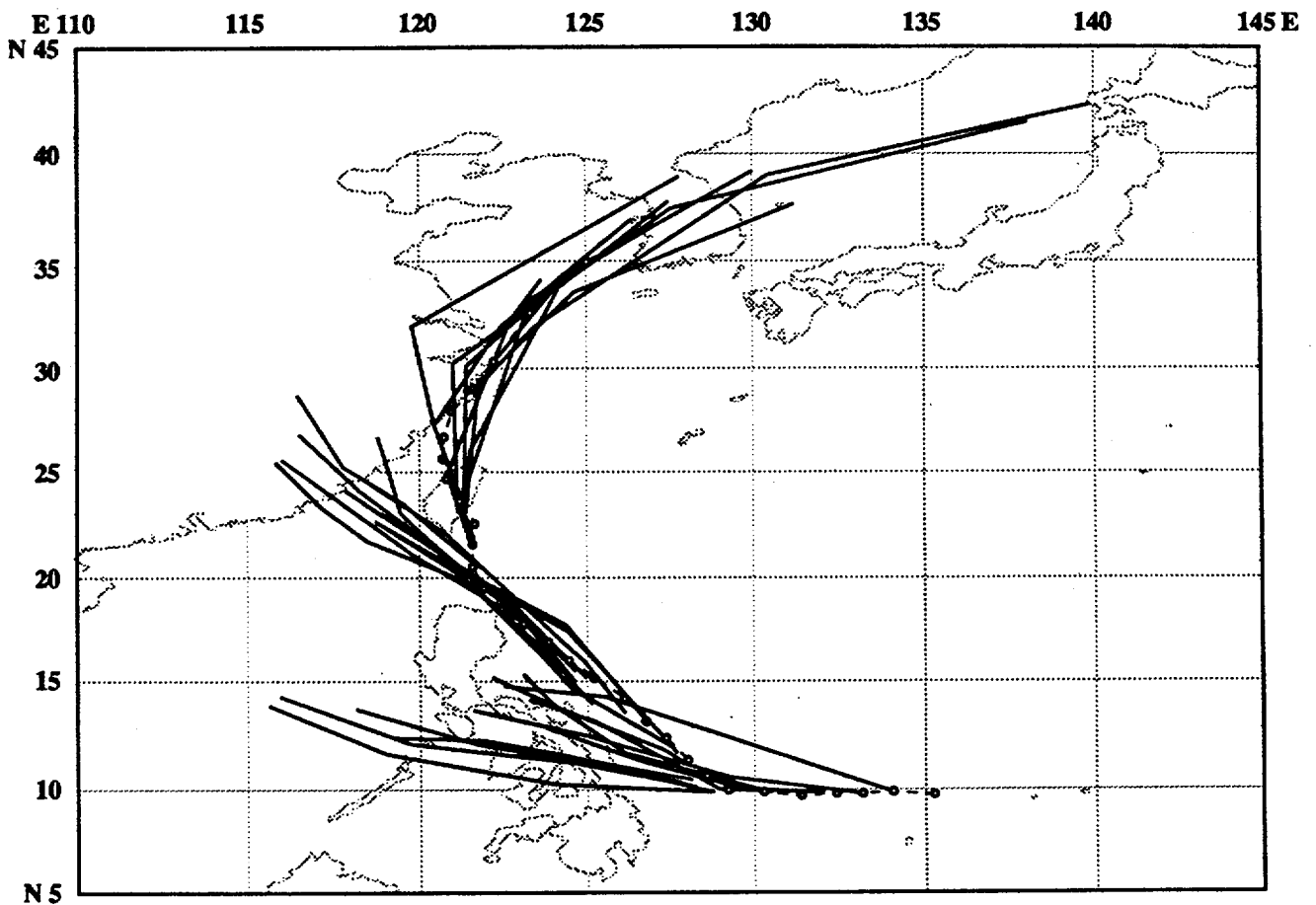
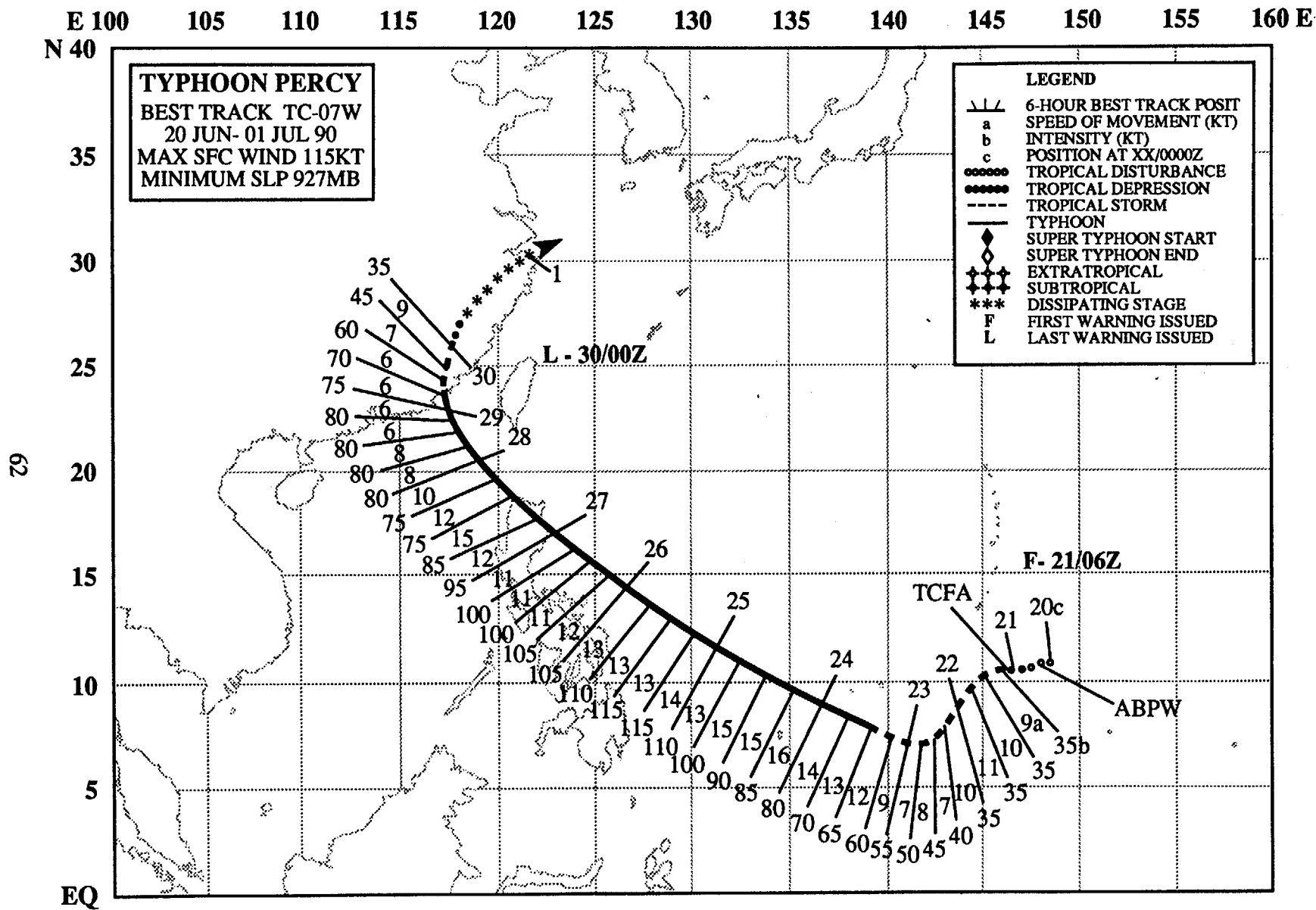


Figure 3-06-7. Summary of JTWC forecasts (solid lines) for Ofelia are superimposed on the final best track (dashed line).



TYPHOON PERCY (07W)

I. HIGHLIGHTS

Percy was the fourth and last tropical cyclone in June. After forming southeast of Guam, it executed an unusual track to the southwest for 36 hours before paralleling Ofelia's (06W) track to the west-northwest around the western periphery of the subtropical ridge. Percy damaged the western Caroline Islands and became the second typhoon within a week to batter northern Luzon before recurving over eastern China.

II. CHRONOLOGY OF EVENTS

- 200600Z - First mentioned on the Significant Tropical Weather Advisory as an area of convection that had persisted for 12 hours. A cyclonic circulation was present in the low-level wind field under weakly divergent flow aloft.
- 202230Z - Advisory reissued to upgrade system's potential for development from poor to fair as outflow and cloud signature improved.
- 210300Z - Tropical Cyclone Formation Alert due to significant increase in organized convection and improved outflow aloft during the past 24 hours.
- 210600Z - First warning and upgrade to tropical storm prompted by receipt of 35 kt (18m/sec) ship report.
- 231800Z - Upgraded to typhoon followed initial signs of eye formation within the central dense overcast and first intensity estimate of T4.0.
- 250600Z - Peak intensity - 115 kt (59 m/sec) - with 25 nm (46 km) diameter eye and T6.0.
- 291200Z - Downgraded to tropical storm resulted from weakened convective signature following cyclone's interaction with the coast of southeastern China.
- 300000Z - Final warning - (dissipating over land) - followed further loss of convective organization as system underwent increased vertical wind shear and loss of latent and sensible heat.

III. TRACK AND MOTION

After initially tracking westward, Percy turned and tracked southwestward for approximately 36 hours. Since the extent of the subtropical ridge and its axis along 28°N remained relatively unchanged during this period, the track change must have resulted from activity near the monsoon trough. An anticyclone had formed southeast of Typhoon Ofelia (06W) and was tracking west-northwestward in tandem with it. As Percy formed, subsidence associated with the converging outflow aloft from both Percy and Ofelia strengthened the anticyclone which resulted in northerly steering flow across Percy (Figure 3-07-1). Percy tracked around the east side of this anticyclone until approximately 221200Z. As Ofelia moved northwestward away from Percy, the anticyclone between them tracked northwestward, weakened and merged with the subtropical ridge to its north. By 231200Z it was only evident as a southwestward extension of the subtropical ridge between Ofelia and Percy (Figure 3-07-2), and by 251200Z it was no longer discernible. Percy then tracked west-northwestward around the subtropical ridge (Figure 3-07-3). After making landfall on the southeast coast of China, Percy was picked up by a mid-latitude short wave trough and finally dissipated as it recurved over eastern China.

IV. INTENSITY

Starting as a low-level circulation at the eastern end of the monsoon trough, Percy quickly developed into a tropical storm as it moved into an area of upper-level divergence. An upper-level anticyclone soon developed over the low-level circulation center. The vertically aligned system intensified into a typhoon as it obtained an outflow channel to the south. As Percy cleared the western Caroline Islands, it developed an additional outflow channel to the north and further intensified, reaching its maximum intensity of 115 kt (59 m/sec) at 250600Z (Figure 3-07-4). The typhoon

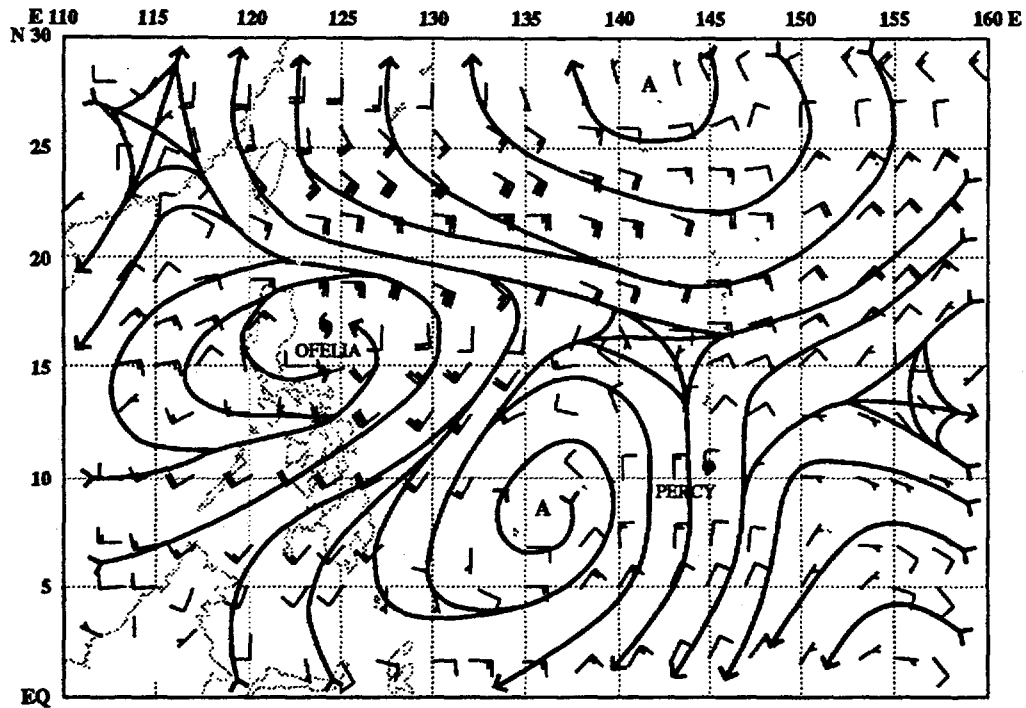


Figure 3-07-1. The 211200Z June deep layer mean analysis shows Percy embedded in northerly flow with a anticyclone to its west. Ofelia's (06W) circulation is located to the northwest of the anticyclone.

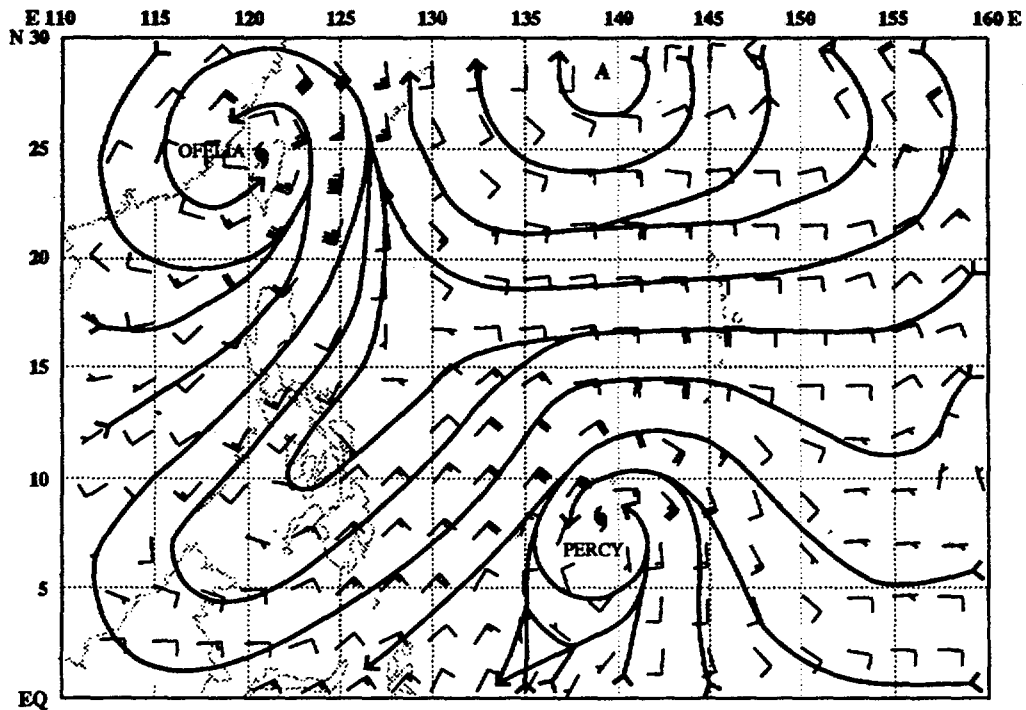


Figure 3-07-2. The 231200Z June deep layer mean analysis indicates the ridge between Ofelia (06W) and Percy has weakened and become a southwestward extension of the subtropical ridge.

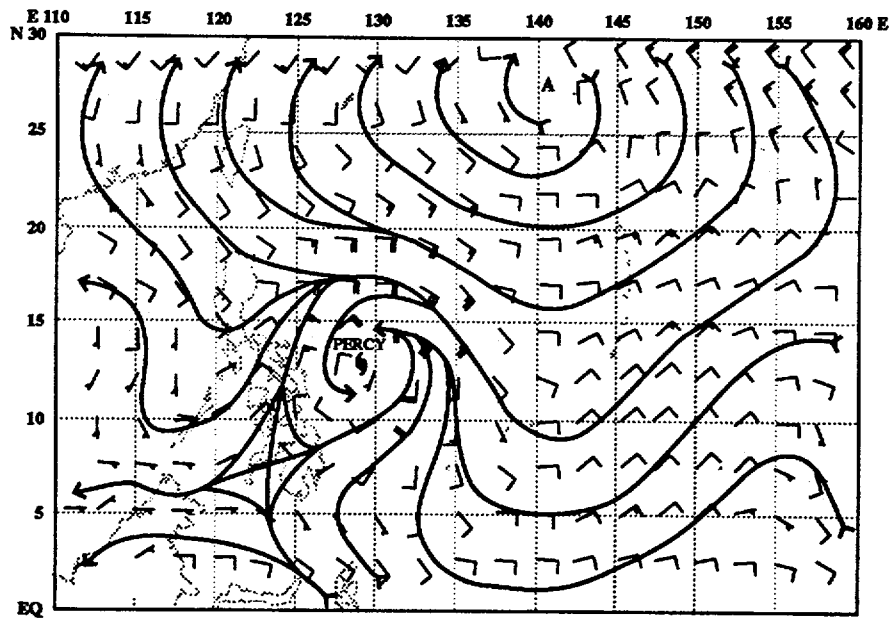


Figure 3-07-3. The 251200Z June deep layer mean analysis shows Percy embedded in the flow around the western end of the subtropical ridge.

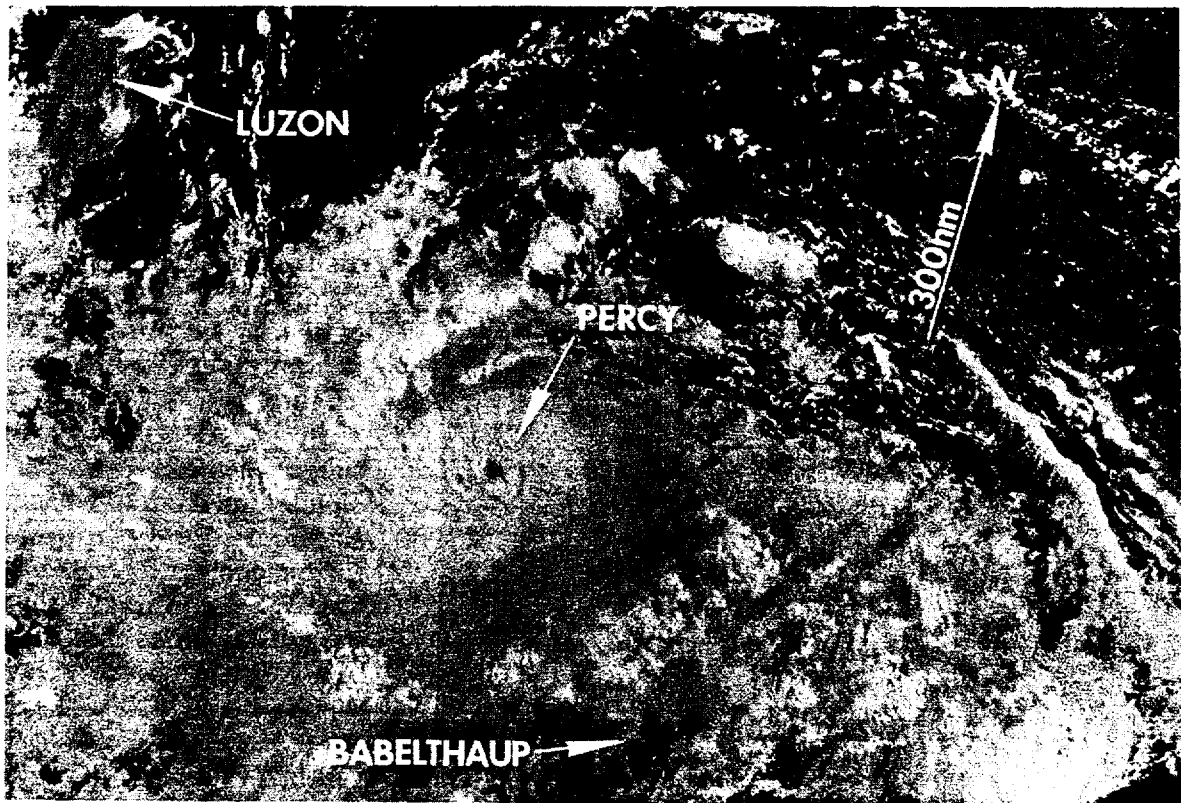


Figure 3-07-4. Typhoon Percy just prior to reaching maximum intensity. Northern Luzon is visible at the top left of the image (250021Z June DMSP visual imagery).

weakened initially due to increasing vertical wind shear from the northeast, and later, from land interactions with northern Luzon (Figure 3-07-5). After moving into the South China Sea and reintensifying slightly, Percy's eye wall (Figure 3-07-6) assumed a polygonal structure (Lewis and Hawkins, 1982). Further weakening resulted from additional vertical wind shear and passage over China.

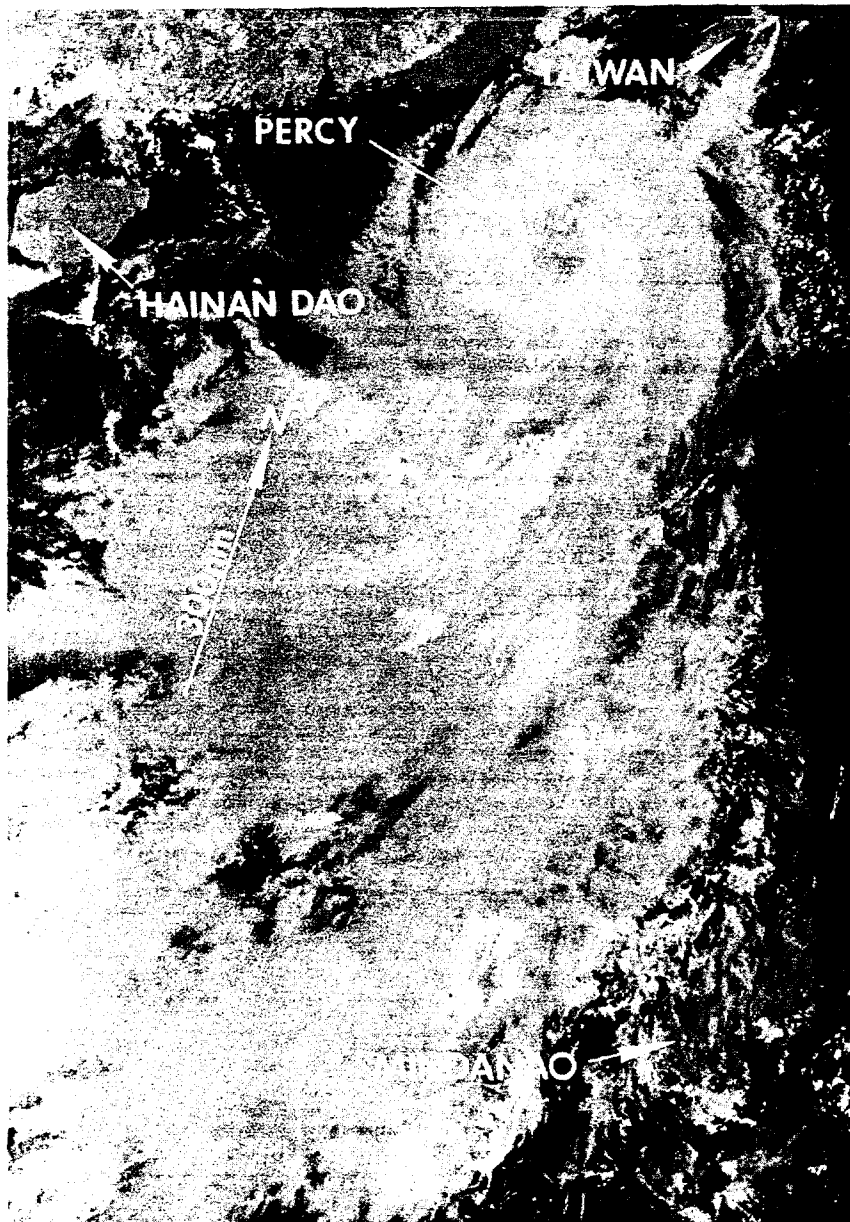


Figure 3-07-5. A ragged, cloud-filled eye reformed after the typhoon collided with northern Luzon. Percy is one day before making landfall in southeastern China. Taiwan is at top right and Hainan Dao at top left (280100Z June DMSP enhanced infrared imagery).

V. FORECASTING PERFORMANCE

Of particular interest was the southwestward portion of Percy's track. Initially, JTWC thought the dynamic high pressure system between Ofelia (06W) and Percy was too weak to influence Percy's track. Forecasters favored persistence and climatology for a west-northwestward track. Forecasters assumed that any departure from this track would be short lived as a result of interactions with a vorticity center associated with a mass of convection to the southeast of Percy. A binary interaction (Figure 3-07-7), when added to the translation of the overall system, would cause a net displacement of Percy to the southwest. This would only last until the two vortices merged. In contrast, OTCM guidance (Figure 3-07-8), which agreed with the deep layer mean, suggested a track south of west which turned out to be accurate. Later, as Percy approached Luzon, another forecast problem arose. The NOGAPS prognostic series indicated that the subtropical ridge would weaken and allow Percy to recurve east of Taiwan. As a result, JTWC and a number of the objective aids forecast recurvature at that longitude. However, the subtropical ridge

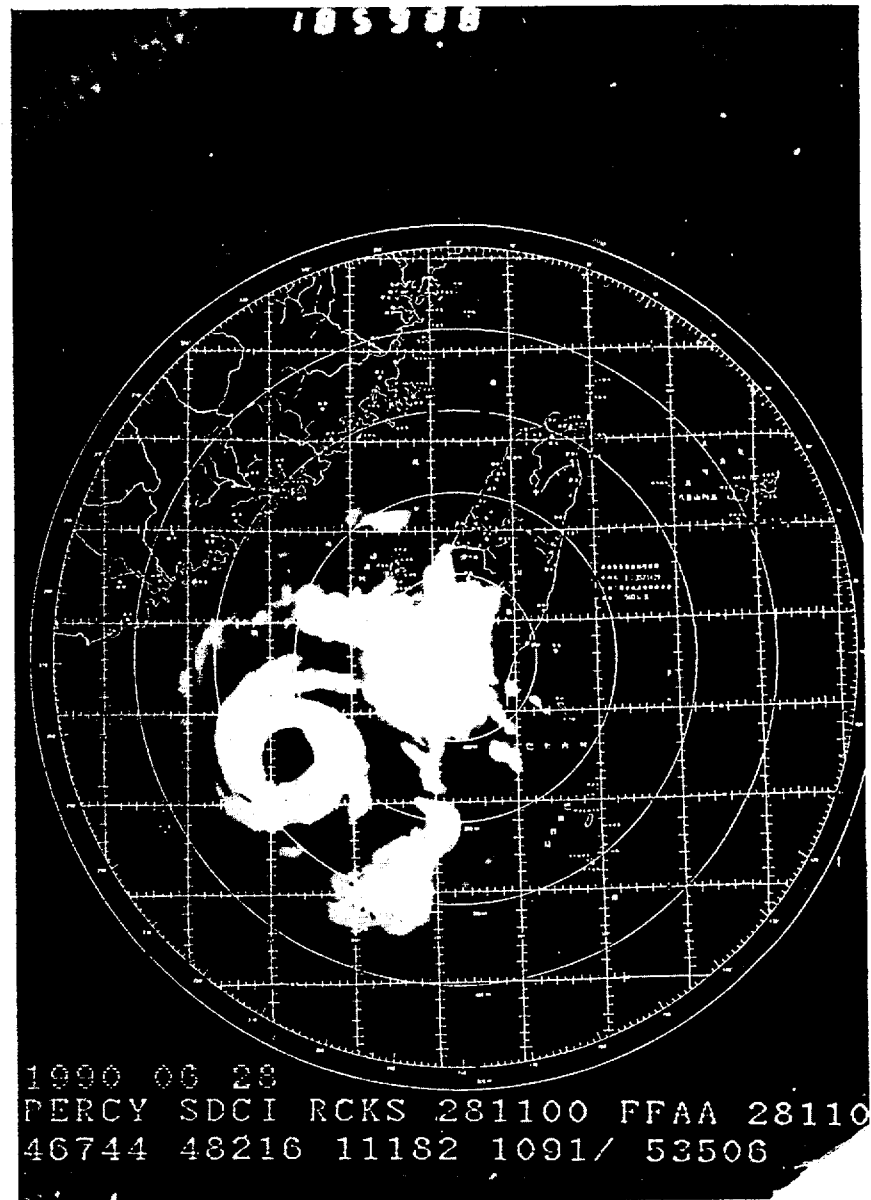


Figure 3-07-6. The polygonal structure of Percy's eye wall at 281100Z June as viewed by the zohsiung (WMO 46744) radar (photograph courtesy of the Central Weather Bureau, Taipei, Taiwan).

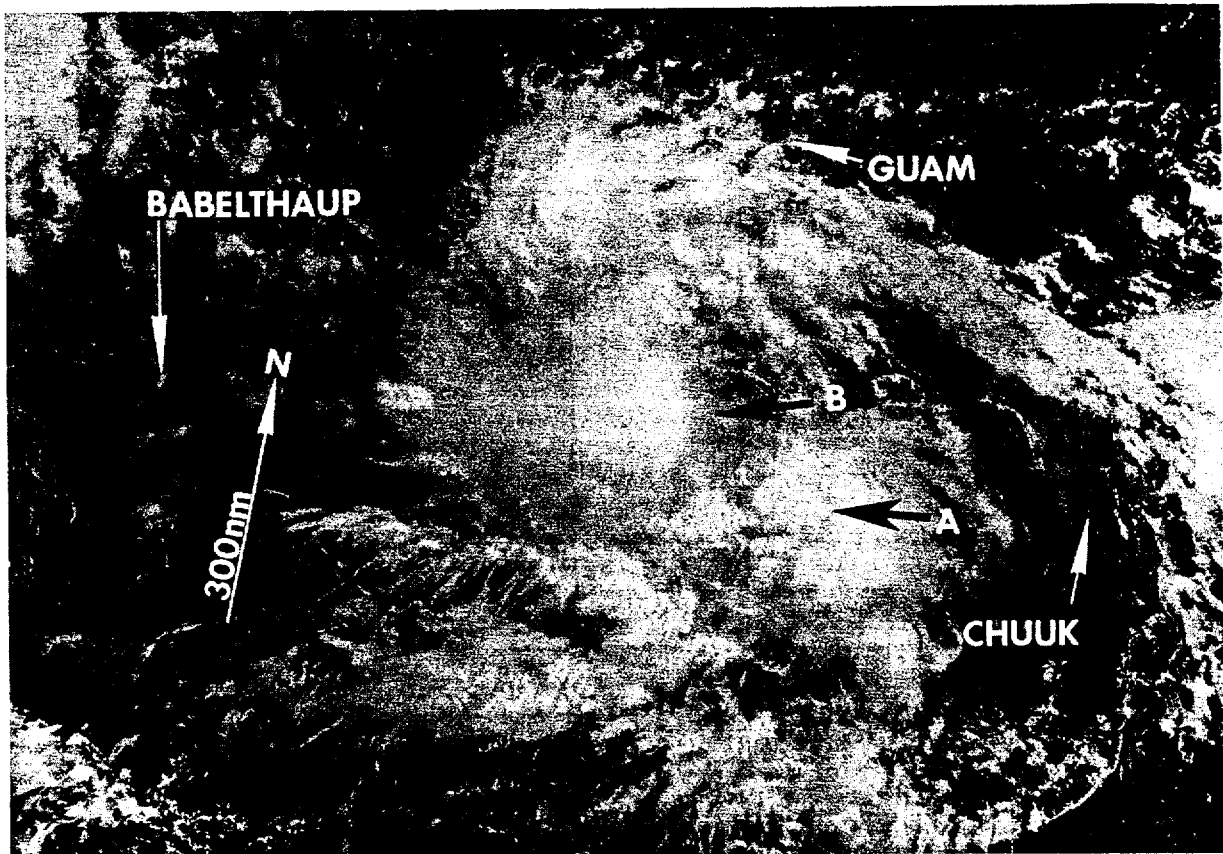


Figure 3-07-7. The vorticity associated with the convective mass (at point A) may have joined with the vorticity associated with the low-level circulation center (at point B), to interact as a binary pair. Babelthaup in the Palau Islands can be seen to the west of Percy's cloudiness (212343Z June DMSP visual imagery).

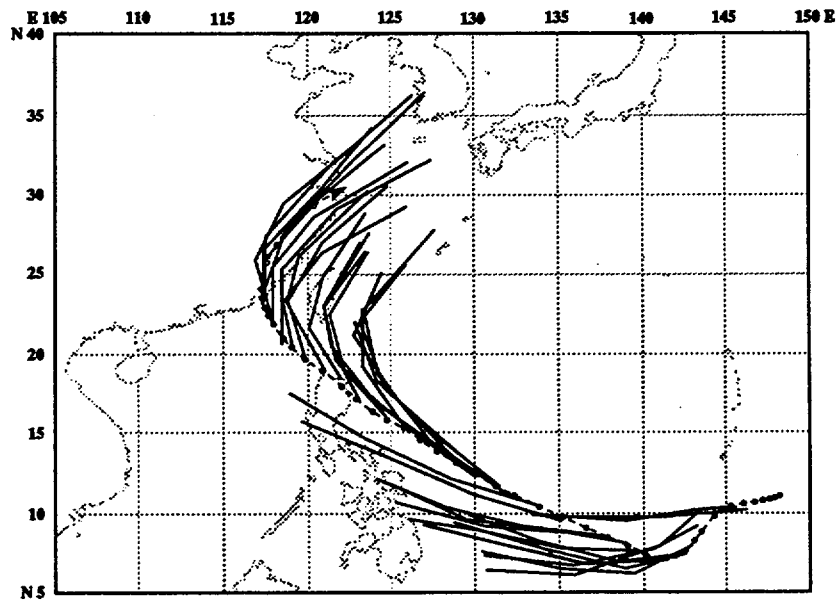


Figure 3-07-8. OTCM guidance and the JTWC forecasts compared to the final best track near the start of the unusual southwesterly motion.

did not weaken and Percy tracked further west before recurving. All the JTWC forecasts are plotted on the best track in Figure 3-07-9.

VI. IMPACT

Percy seriously affected several islands in the western Carolines. The first of these was Sorol, an atoll located 150 nm (280 km) southeast of Yap. As the tropical storm passed 40 nm (75 km) to the south of Sorol, the second largest island in the lagoon, Pegelmol, was almost cut in half and another island lost one third of its area due to wave action. Taro patches, coconut trees and other vital crops were essentially wiped out and will take years to replace. After reaching typhoon intensity, Percy passed 55 nm (100 km) south-southwest of Yap, which received sustained winds of 35-45 kt (18-23 m/sec) with gusts to 55 kt (28 m/sec). In addition, Yap suffered extensive flooding along its eastern shore. Most roads were blocked by water and later by debris and flooding from the unusually high tide. Nugulu, 60 nm (110 km) to the south-southwest of Yap, took the brunt of the typhoon. Maximum gusts estimated at 70 kt (35 m/sec) totally destroyed all crops. Seven homes were completely demolished; others lost their roofs. Fortunately there were no fatalities. Palau was not as lucky; one child's death was attributed to the typhoon as Percy passed 125 nm (230 km) to the north-northeast of Koror. Power, radio and TV were knocked out as winds ripped off tin roofs and snapped power poles. Broken limbs took out power lines. Once past the Caroline islands, Percy became the second storm in less than a week to devastate northern Luzon. The resulting landslides and floods left at least 8 people dead and 31,206 homeless, adding to the misery left behind by Ofelia (06W).

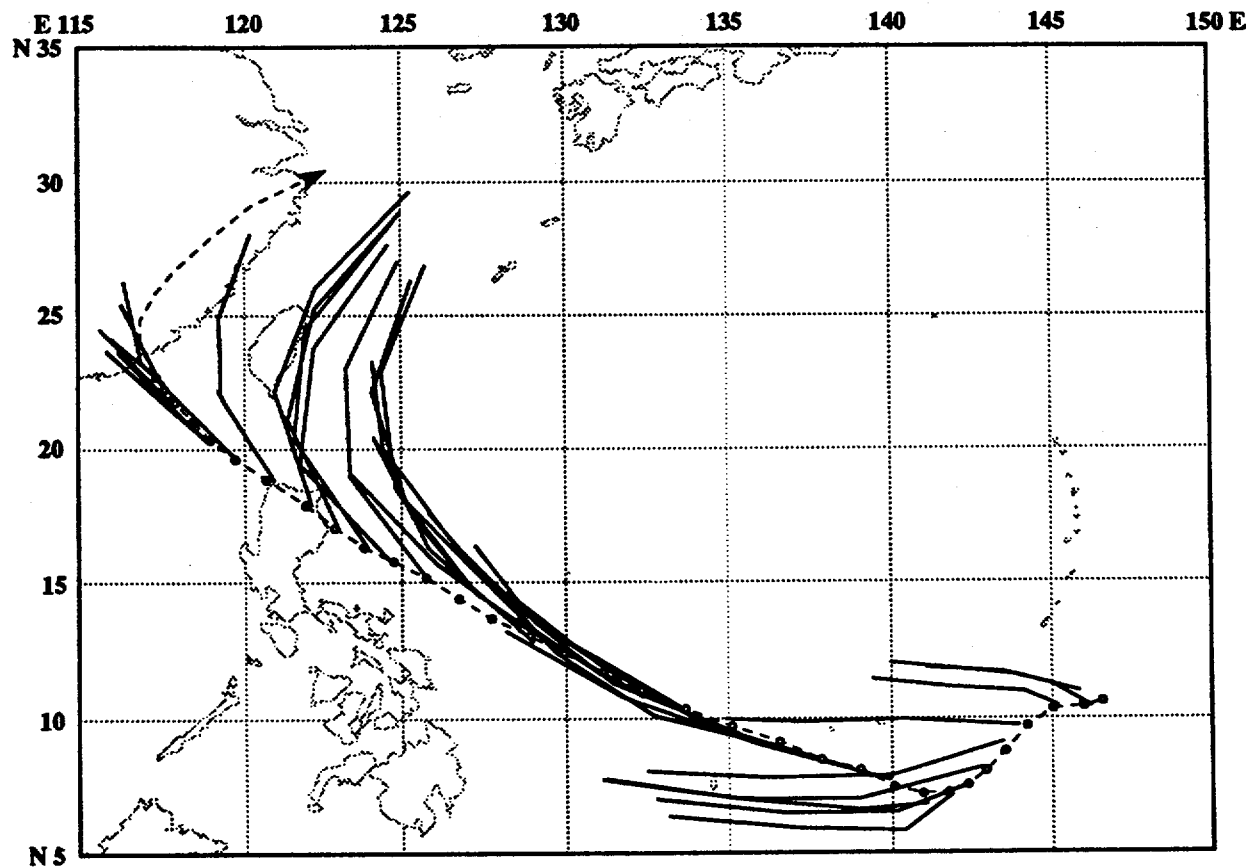


Figure 3-07-9. A plot of all the JTWC forecasts (solid lines) with the best track (dashed lines).

E 110 115 120 125 130 135 140 145 150 155 160 165 170 175 E

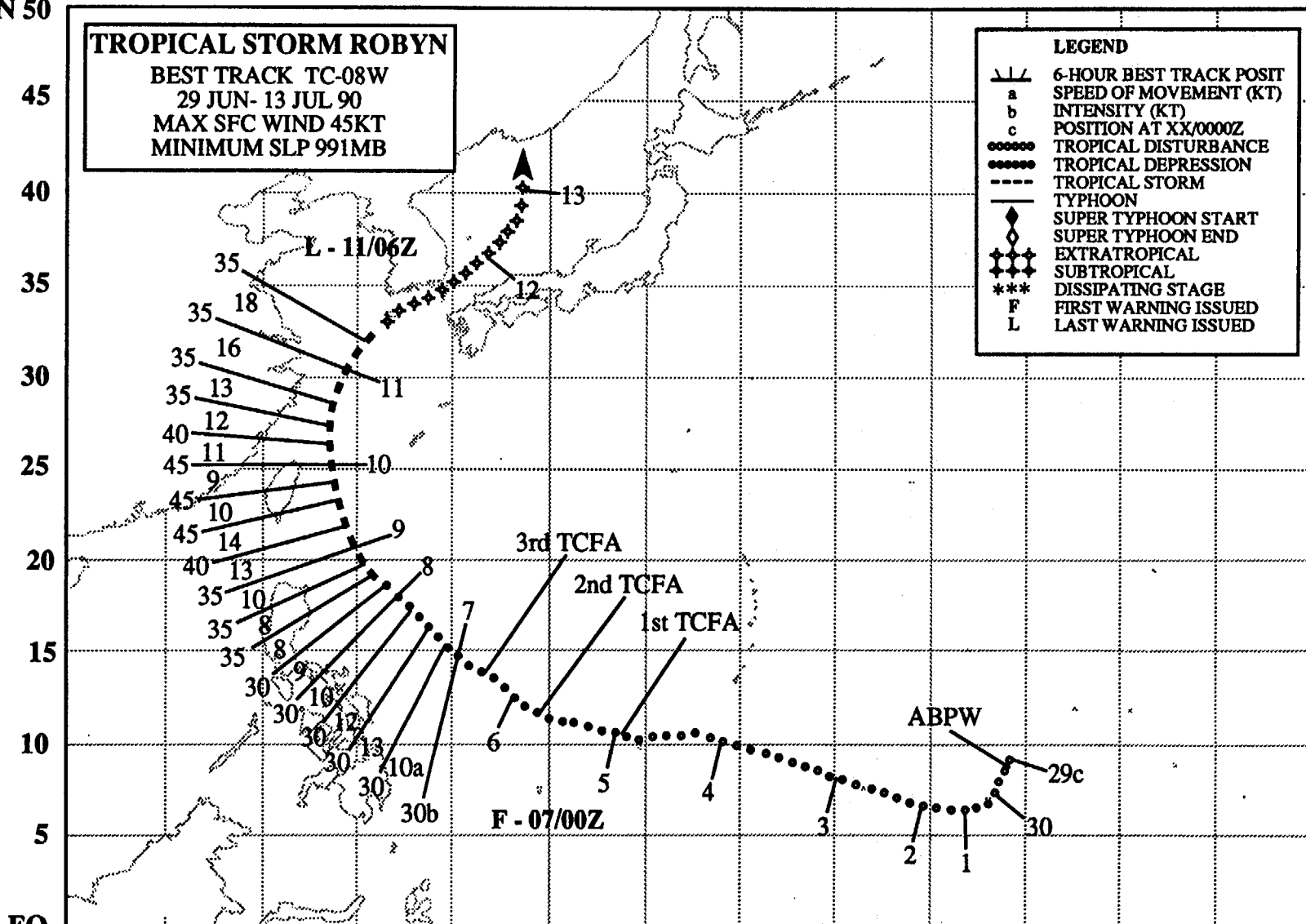
N 50

TROPICAL STORM ROBYN

BEST TRACK TC-08W
29 JUN- 13 JUL 90
MAX SFC WIND 45KT
MINIMUM SLP 991MB

LEGEND

- 6-HOUR BEST TRACK POSIT
- a SPEED OF MOVEMENT (KT)
- b INTENSITY (KT)
- c POSITION AT XX/0000Z
- TROPICAL DISTURBANCE
- TROPICAL DEPRESSION
- TROPICAL STORM
- TYPHOON
- SUPER TYPHOON START
- SUPER TYPHOON END
- EXTRATROPICAL
- SUBTROPICAL
- DISSIPATING STAGE
- F FIRST WARNING ISSUED
- L LAST WARNING ISSUED



70

EQ

TROPICAL STORM ROBYN (08W)

I. HIGHLIGHTS

Robyn, the first significant tropical cyclone of July, followed what at first glance might appear to be a typical recurvature track. However, Robyn's motion was actually a classic example of the response of a tropical cyclone to the establishment of an omega block, and thus is significant as a case study of an infrequent, but complex, synoptic influence on tropical cyclone motion.

II. CHRONOLOGY OF EVENTS

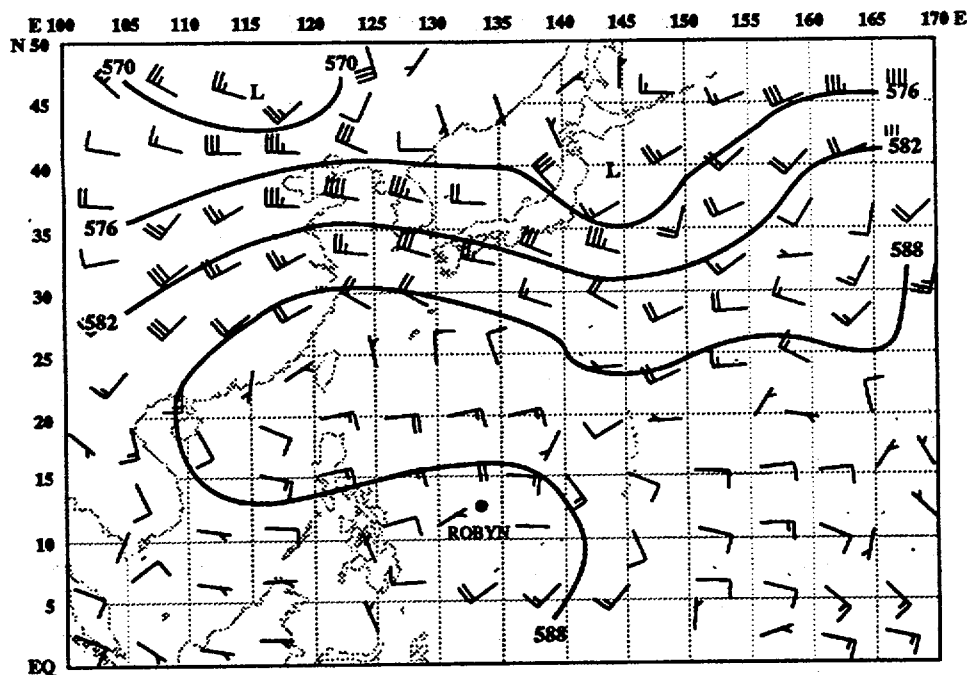
- 290600Z - (June) First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1006 mb.
- 042300Z - (July) First Tropical Cyclone Formation Alert based on increased convection, organization, and outflow aloft.
- 051530Z - Second Tropical Cyclone Formation Alert issued. Organization temporarily delayed due to upper-level wind shear.
- 061530Z - Third Tropical Cyclone Formation Alert issued. Convection still consolidating during diurnal fluctuations.
- 070000Z - First warning based on improved outflow to the southeast and anticipated reduction of vertical wind shear.
- 081800Z - Upgraded to tropical storm based on enhanced convection and improved organization.
- 091800Z - Peak intensity of 45 kt (23 m/sec) based on synoptic data.
- 110000Z - Downgraded to tropical depression.
- 110600Z - Final warning - (extratropical) - due to the loss of persistent central convection.

III. TRACK AND MOTION

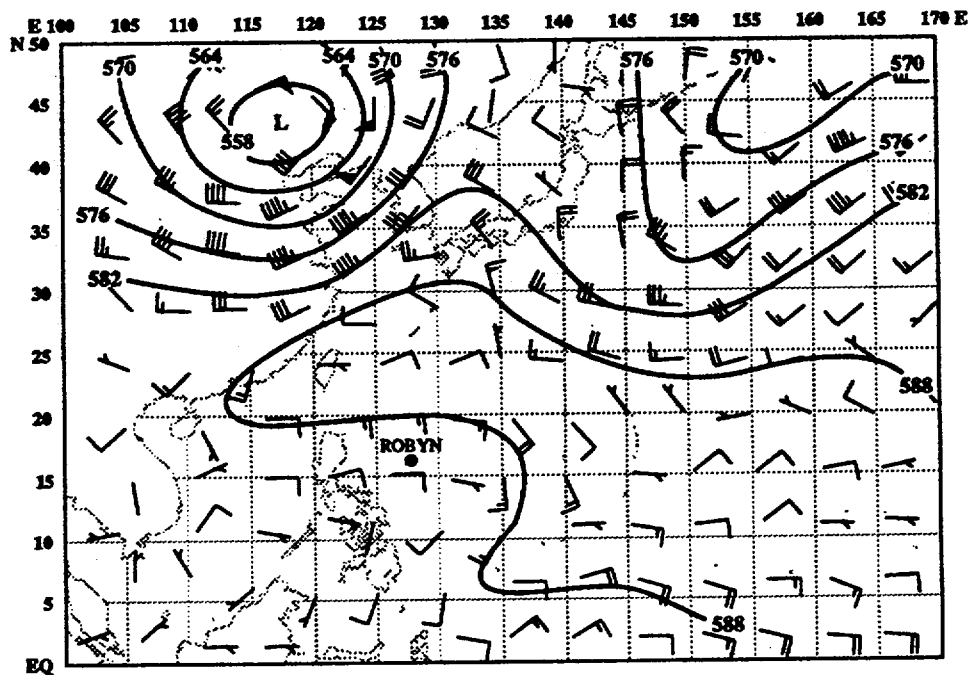
From the initial mention on the Significant Tropical Weather Advisory until the first warning at 070000Z, Robyn tracked essentially west-northwestward under the subtropical ridge. However, on 6 July an omega block began to form with the digging lows located about 49°N 117°E and 51°N 153°E as shown in Figure 3-08-1a. The 5880 meter height contour identified significant ridging poleward of Robyn, which under normal circumstances would imply continued westward movement. As shown in Figure 3-08-1b, the omega block was firmly established at 080000Z, and the digging lows had dramatically eroded the ridge north of Robyn causing the increase in its northward motion component. At 100000Z, the ridge was fully eroded permitting Robyn to pass Taiwan to the east (Figure 3-08-1c). The downwind digging low had penetrated more equatorward than its upwind counterpart causing the omega block to tilt eastward. This shift signaled the beginning of the breakdown of the block. Still, the ridging directly east of Robyn, associated with the central axis of the block, was sufficient to keep Robyn on a northward track, delaying recurvature. At 120000Z, the central ridging of the omega block (Figure 3-08-1d) had broken down sufficiently for Robyn to recurve and significantly accelerate as it moved into the Sea of Japan as an extratropical low.

IV. INTENSITY

The delayed development of Robyn and its subsequent intensification to only a nominal tropical storm (Figure 3-08-2) was due to moderate but persistent vertical wind shear associated with the eastern periphery of the summertime 200-mb easterly jet over southern Asia. In addition, the ridging to the north of Robyn for much of its life-cycle restricted outflow. When the ridge broke down, Robyn briefly intensified to 45 kt (23 m/sec) in response to outflow into the midlatitude westerlies.

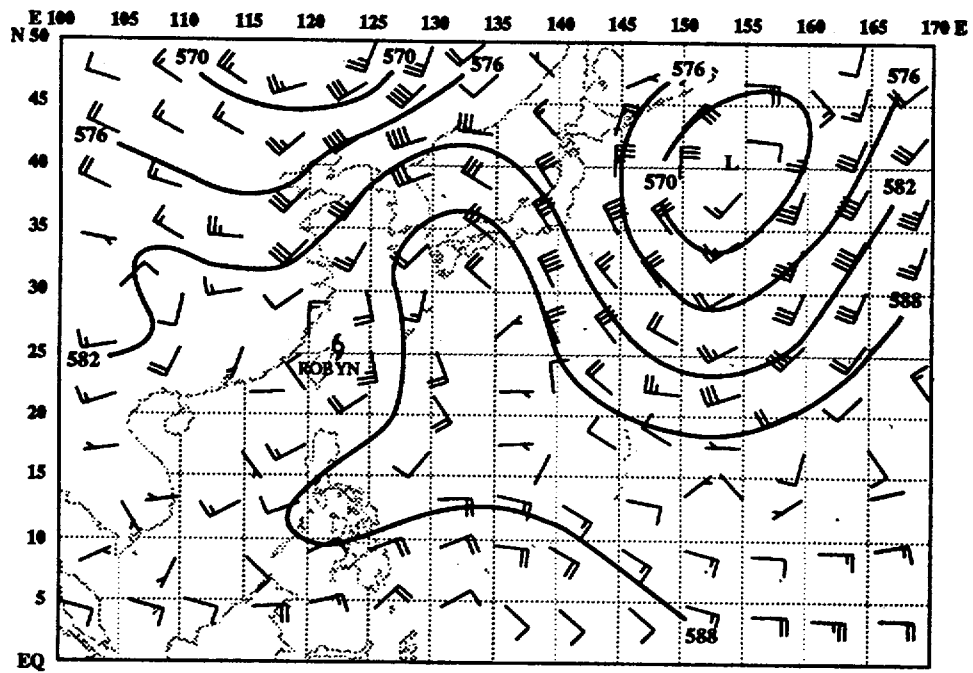


a.

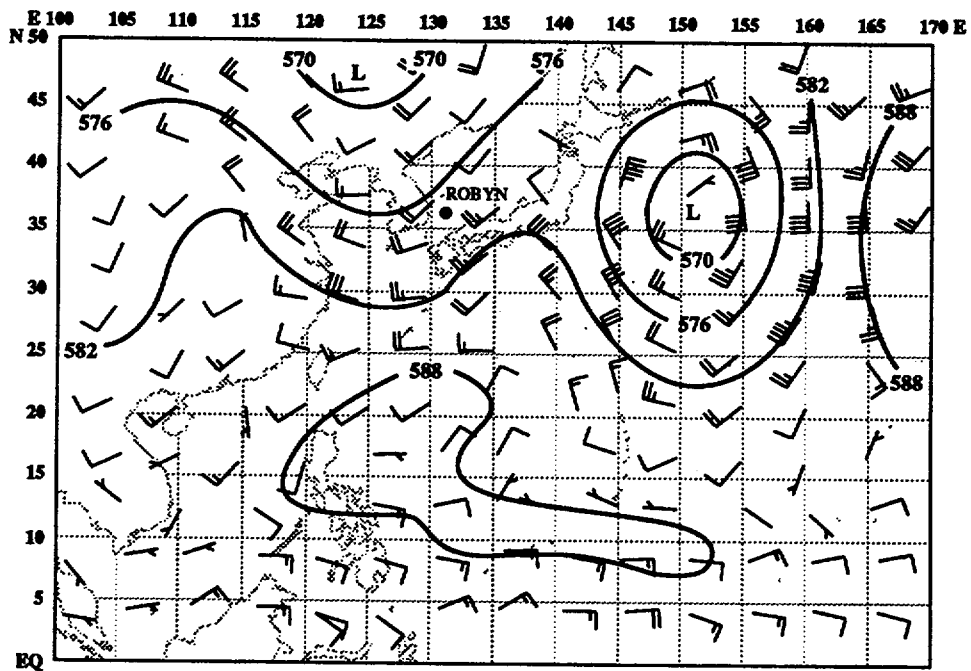


b.

Figure 3-08-1 a, b. NOGAPS 500-mb height analyses (in decameters) with the corresponding positions of Robyn for valid times a.) 060000Z and b.) 080000Z



c.



d.

Figure 3-08-1 c, d. NOGAPS 500-mb height analyses (in decameters) with the corresponding positions of Robyn for valid times c.) 100000Z and d.) 120000Z.

V. FORECASTING PERFORMANCE

The key variable in forecasting the motion of Robyn was the rapidity with which the low upwind of the omega block would break down the mid-level ridge to the north of Robyn. As Figure 3-08-3 illustrates, guidance available to JTWC between 070000Z and 080000Z did not indicate that the breakdown would proceed in time to permit Robyn to recurve east of Taiwan. JTWC relies heavily on the dynamic models OTCM and FBAM, the accuracies of which in turn depend heavily on the accuracy of the NOGAPS prognoses. By comparing Figure 3-08-3 with Figure 3-08-1c, it is evident that the NOGAPS 500-mb 72-hour prognosis for 100000Z had prematurely weakened the upwind low of the omega block. As a result, the NOGAPS 500-mb 72-hour prognoses that verified between 090000Z and 100000Z retained ridging north of Robyn that did not verify. This, in turn, caused NOGAPS-dependent objective techniques such as OTCM and FBAM to forecast continued west-northwestward movement for Robyn, which contributed to JTWC's left-of-track bias during that same period.

VI. IMPACT

No information received.

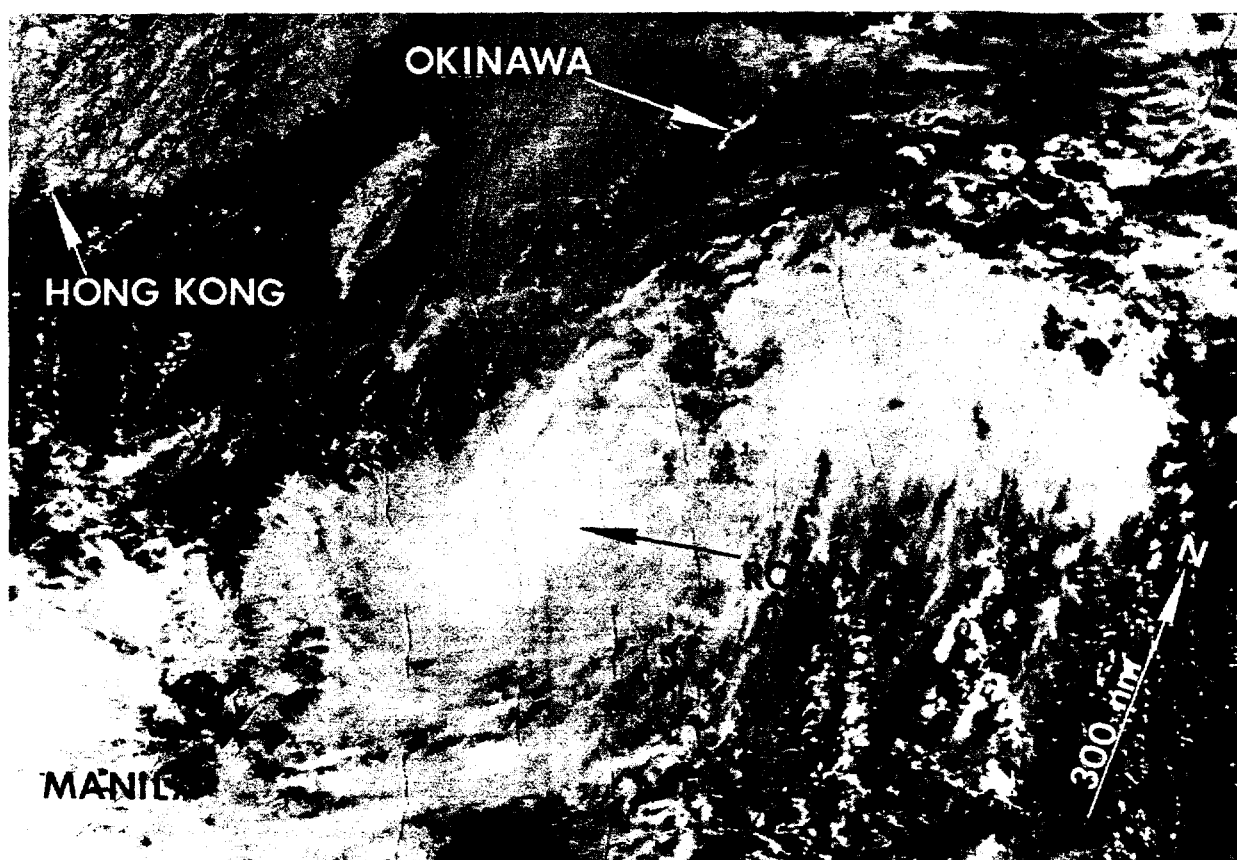


Figure 3-08-2. Robyn just before reaching tropical storm intensity (080514Z July NOAA visual imagery).

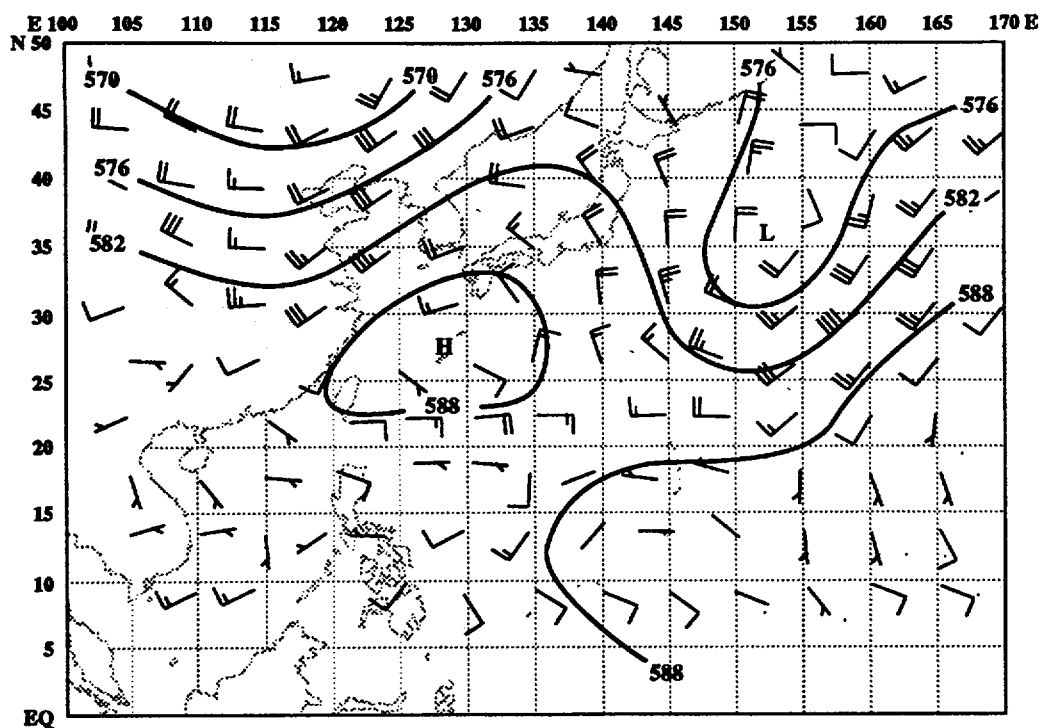


Figure 3-08-3. NOGAPS 500-mb 72-hr prognosis in decameters valid at 100000Z July.

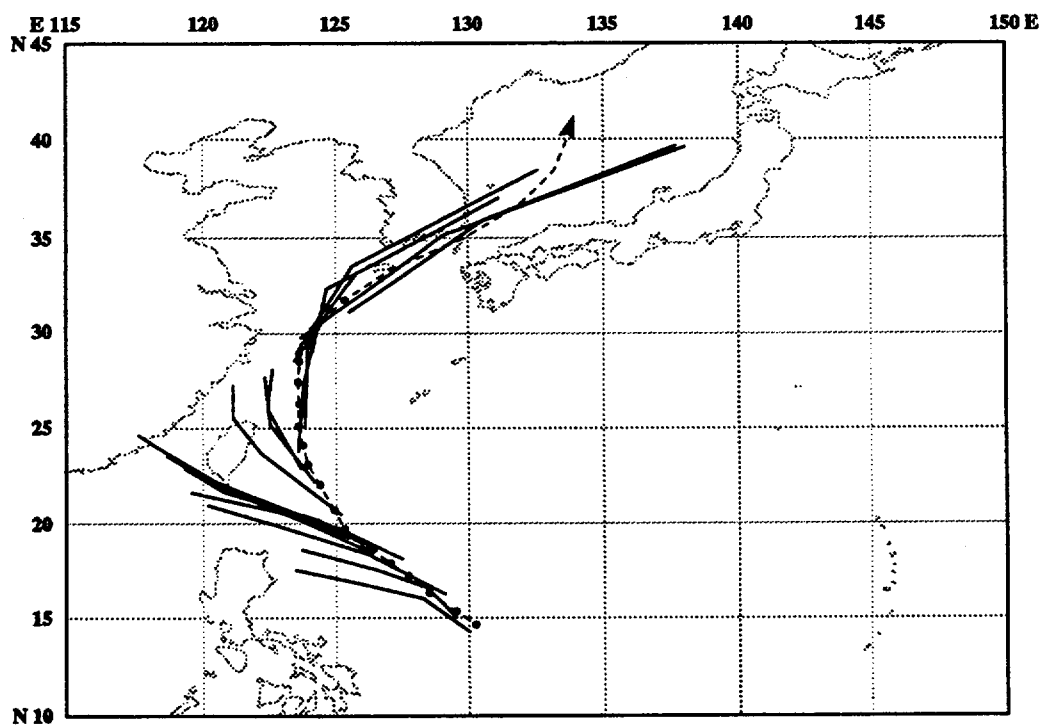
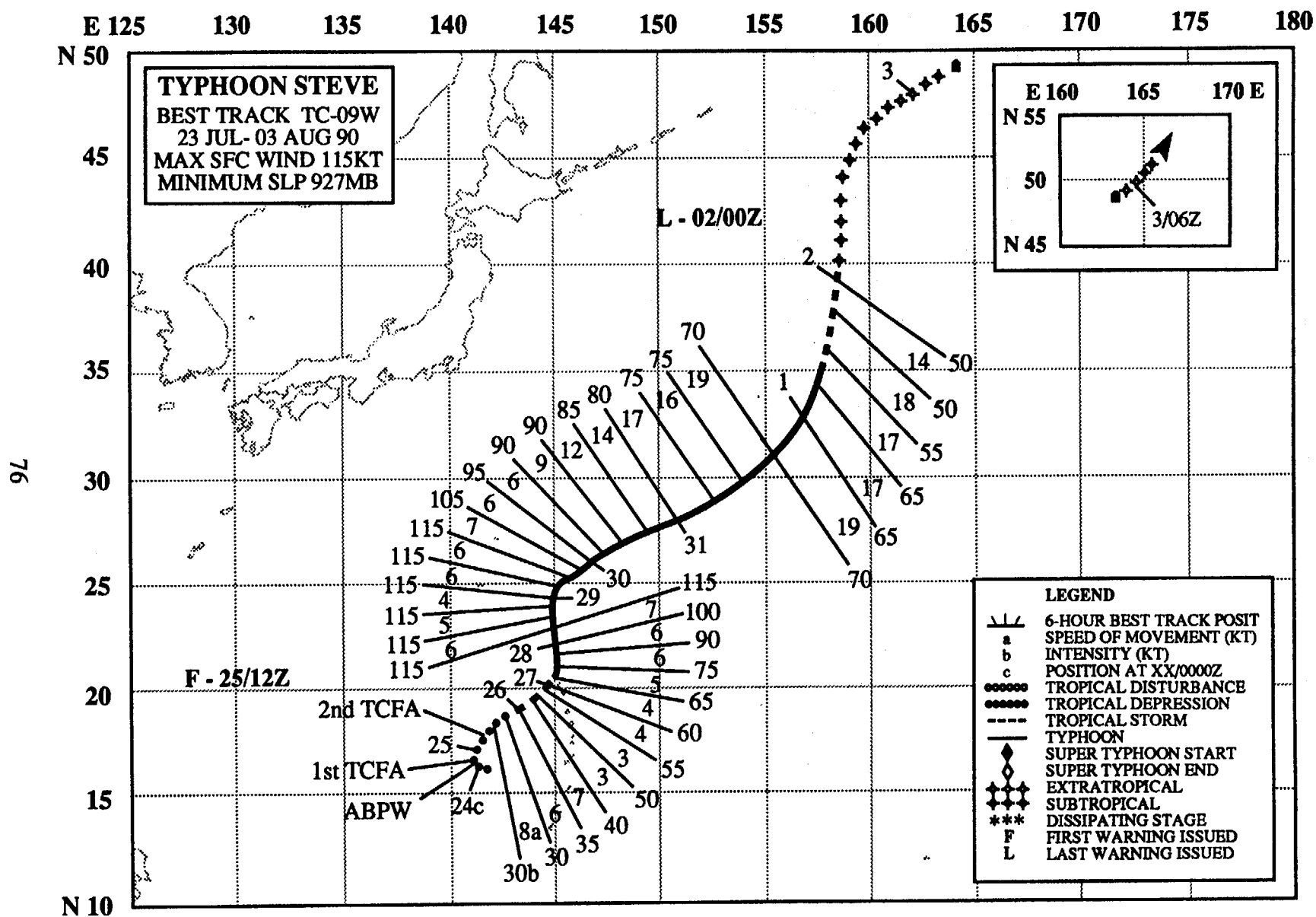


Figure 3-08-4. Summary of JTWC forecasts (solid lines) for Robyn superimposed on the best track (dashed line).



TYPHOON STEVE (09W)

I. HIGHLIGHTS

Steve, with Tropical Storm Tasha (10W) and Typhoon Vernon (11W), made up the only three storm tropical cyclone outbreak to occur in the northwest Pacific this year. Steve persisted on an atypical northeastward track throughout its existence. The orientations of the monsoon trough and the subtropical ridge influenced the track of this system.

II. CHRONOLOGY OF EVENTS

- 240600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1004 mb.
- 240800Z - First Tropical Cyclone Formation Alert based on continued development of the convection.
- 250800Z - Tropical Cyclone Formation Alert reissued due to slow development.
- 251200Z - First warning based on development of persistent central convection.
- 260000Z - Upgraded to tropical storm after restriction to outflow eased.
- 270600Z - Upgraded to typhoon based on eye formation.
- 281800Z - Peak intensity - 115 kt (59 m/sec) - restricted outflow to the west, preventing further intensification.
- 011200Z - Downgraded to tropical storm because of decreased convection and increased vertical wind shear.
- 020000Z - Final warning - (extratropical) - based on the loss of central convection.

III. TRACK AND MOTION

On 19 July, several days before Steve formed, a large TUTT low appeared near the dateline and was reflected in the deep-layer mean analysis (Figure 3-09-1) as an inverted trough. By 23 July, the TUTT low became associated with the eastern extension of the Asian monsoon trough in the deep layer mean analysis (Figure 3-09-2). This synoptic-scale trough segmented the subtropical ridge into an

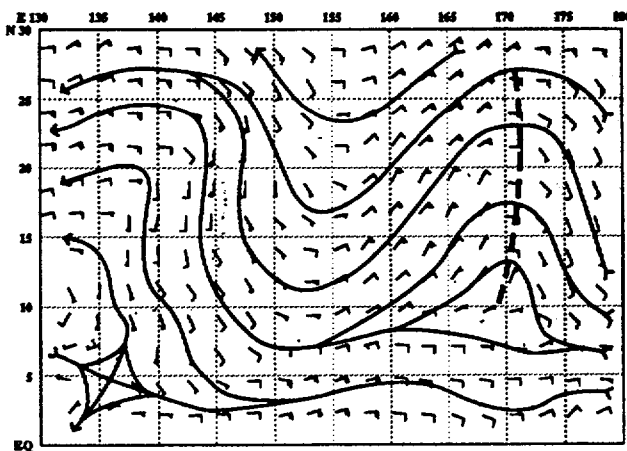


Figure 3-09-1. Deep-layer mean analysis at 190000Z July, showing the reflection of the TUTT low as an inverted trough oriented north-south along 170° East longitude.

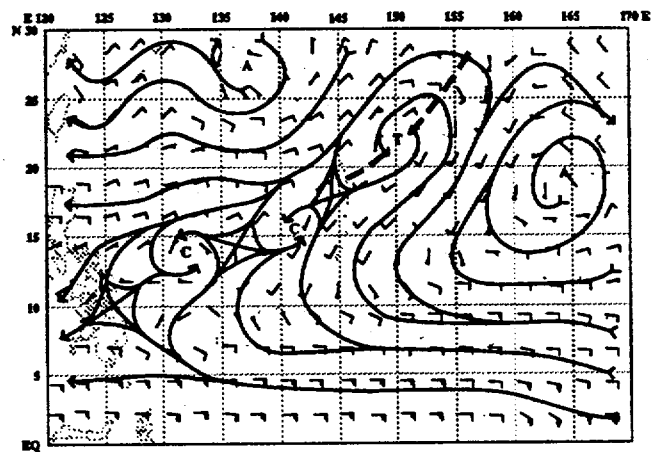


Figure 3-09-2. Deep layer mean analysis at 231200Z July indicates that the TUTT low at point T has elongated northeast-southwest and appears as an eastward extension of the Asian monsoon.

Asian cell, extending eastward from Asia, and a maritime cell, southeast of and parallel to the trough axis. Once Steve formed in the low-level monsoon trough, its basic track was to the northeast, roughly parallel to the axis of the monsoon trough. Short term speed and direction changes appeared to be related to the interaction between Steve, a midget typhoon, and the larger cyclonic circulation in the trough. Note in Figure 3-09-3 that Steve was east of a larger cyclonic circulation as shown on the deep layer mean analysis for 261200Z. It was also under southwesterly mid-tropospheric flow. The track change to the north at 271200Z was related to the change in steering from southwesterly to southerly on the analysis (Figure 3-09-4). After Steve reached higher latitudes and began to weaken, it became the dominant cyclonic circulation. As the system took on extratropical characteristics and increased in size, it filled and accelerated northeastward.

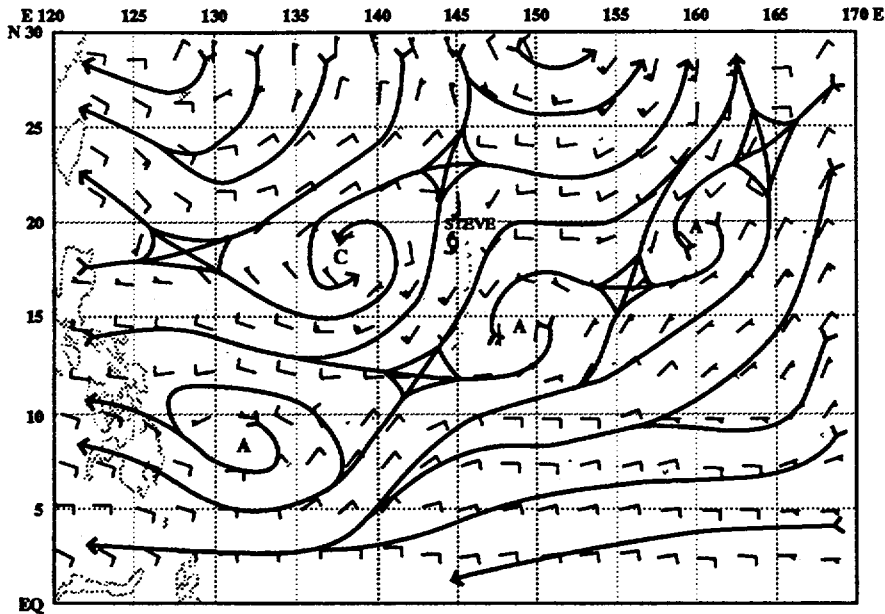


Figure 3-09-3. Deep layer mean analysis at 261200Z July depicts Steve east of a larger cyclonic circulation in the monsoon trough, embedded in southwesterly flow.

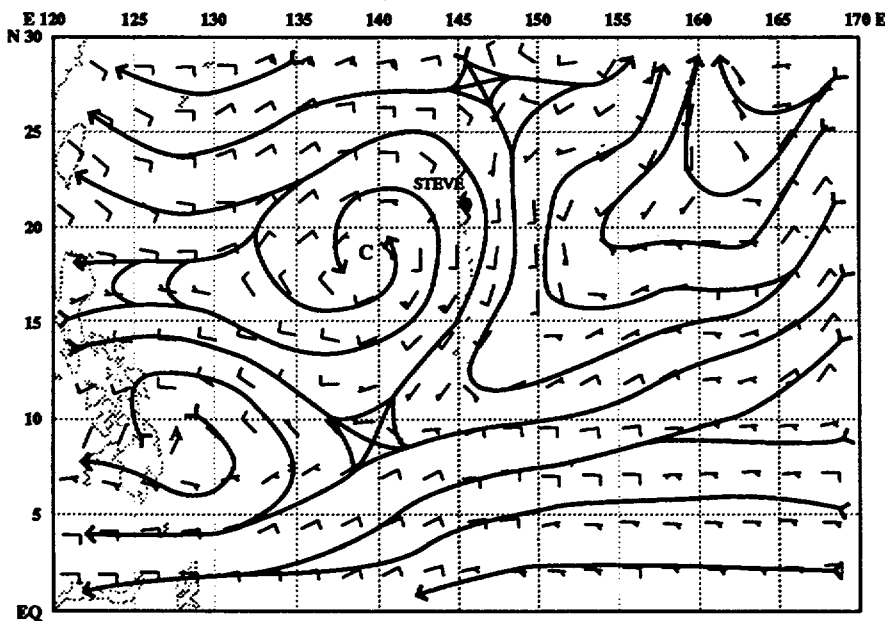


Figure 3-09-4. Deep layer mean analysis at 271200Z July shows Steve embedded in southerly flow.

IV. INTENSITY

The area of convection that eventually became Typhoon Steve formed in the monsoon trough and moved under strong upper-level divergence. Once the convection consolidated, the system (Figure 3-09-5) developed rapidly but remained relatively small -- its deep convection was confined to within 90 nm (165 km) of the center. With no restriction to its outflow, Steve quickly developed an eye and reached typhoon intensity. The typhoon intensified to 115 kt (59 m/sec) and remained at peak intensity



Figure 3-09-5. Steve intensifies as it tracks northward (280811Z July DMSP visual imagery).

for 24 hours until upper-level outflow from Tropical Storm Vernon (11W) to the southwest began to restrict Steve's outflow to the west (Figure 3-09-6). Steve weakened slowly as its deep convection gradually decreased. Its circulation expanded in size and retained storm-force winds as an extratropical system.

V. FORECAST PERFORMANCE

Steve's atypical track produced a difficult forecasting situation. The synoptic features that influenced the track, the monsoon trough and subtropical ridge, were themselves difficult to forecast. To further complicate matters, Steve was one of three tropical cyclones active in the Northwest Pacific at the same time. JTWC's track forecasts were based on Steve's location relative to the subtropical ridge to the northwest. Steve's northeastward movement put the JTWC forecasts significantly left of the actual track (Figure 3-09-7). Dynamical, statistical and climatological objective aids all predicted initial northwestward motion for Steve. The 72-hour forecast errors for Steve averaged 556 nm (1030 km).

VI. IMPACT

No information was received.

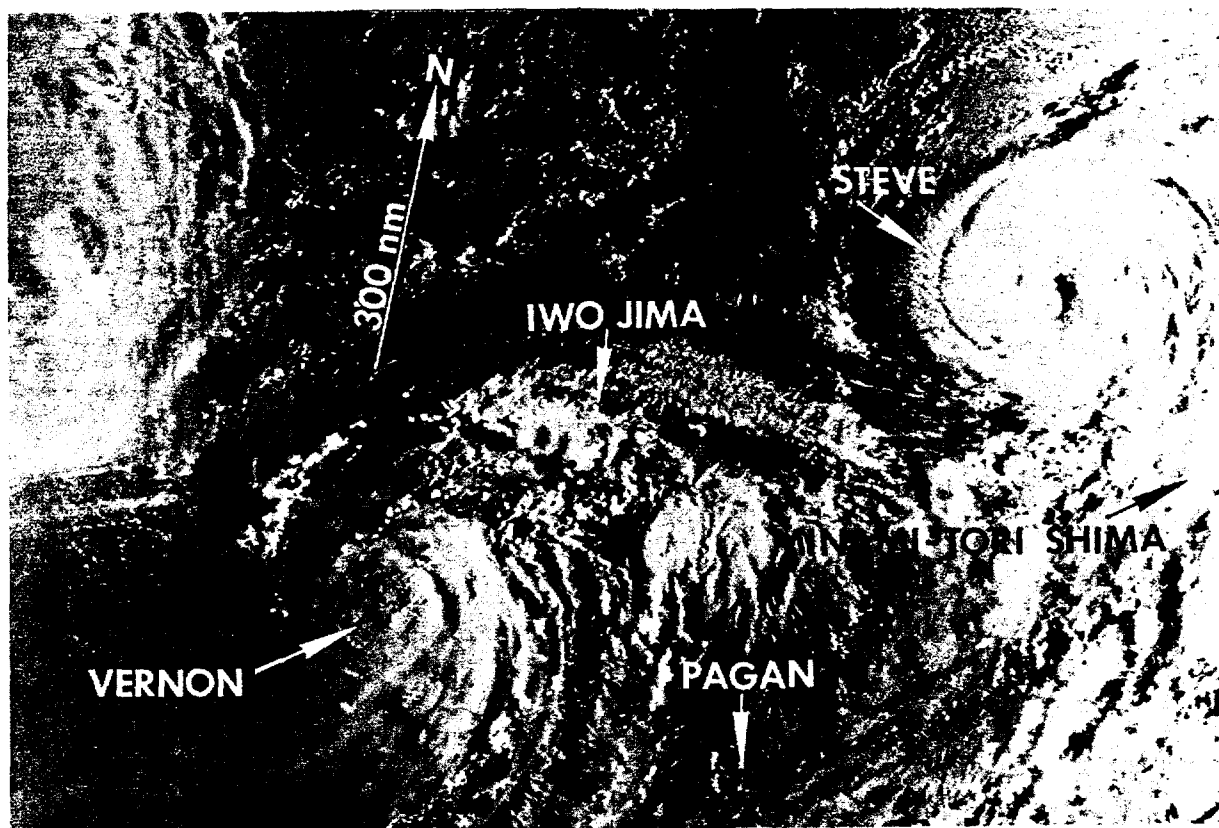


Figure 3-09-6. As Steve weakens, Tropical Storm Vernon (11W) intensifies to the southwest (302028Z July DMSP visual imagery).

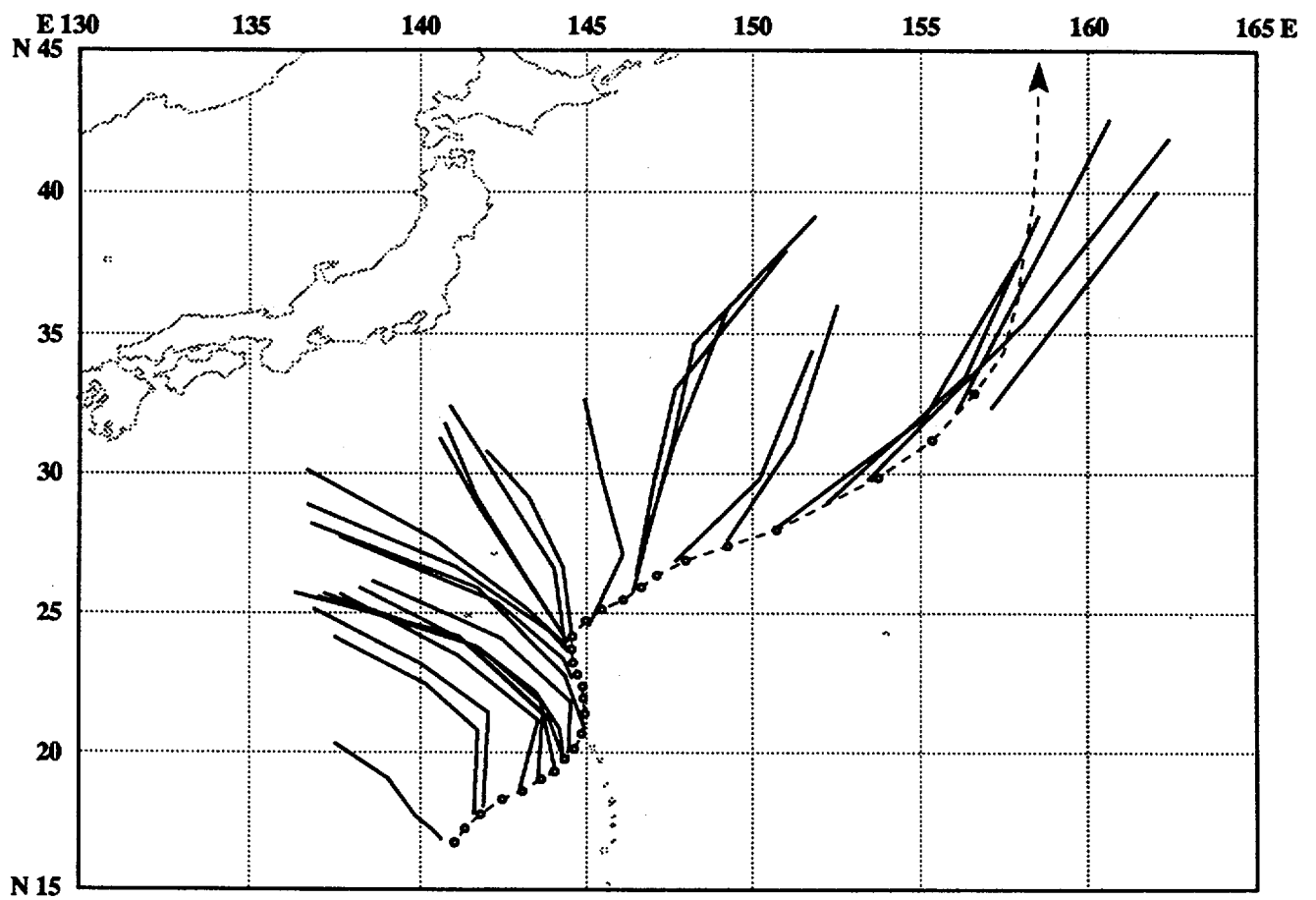
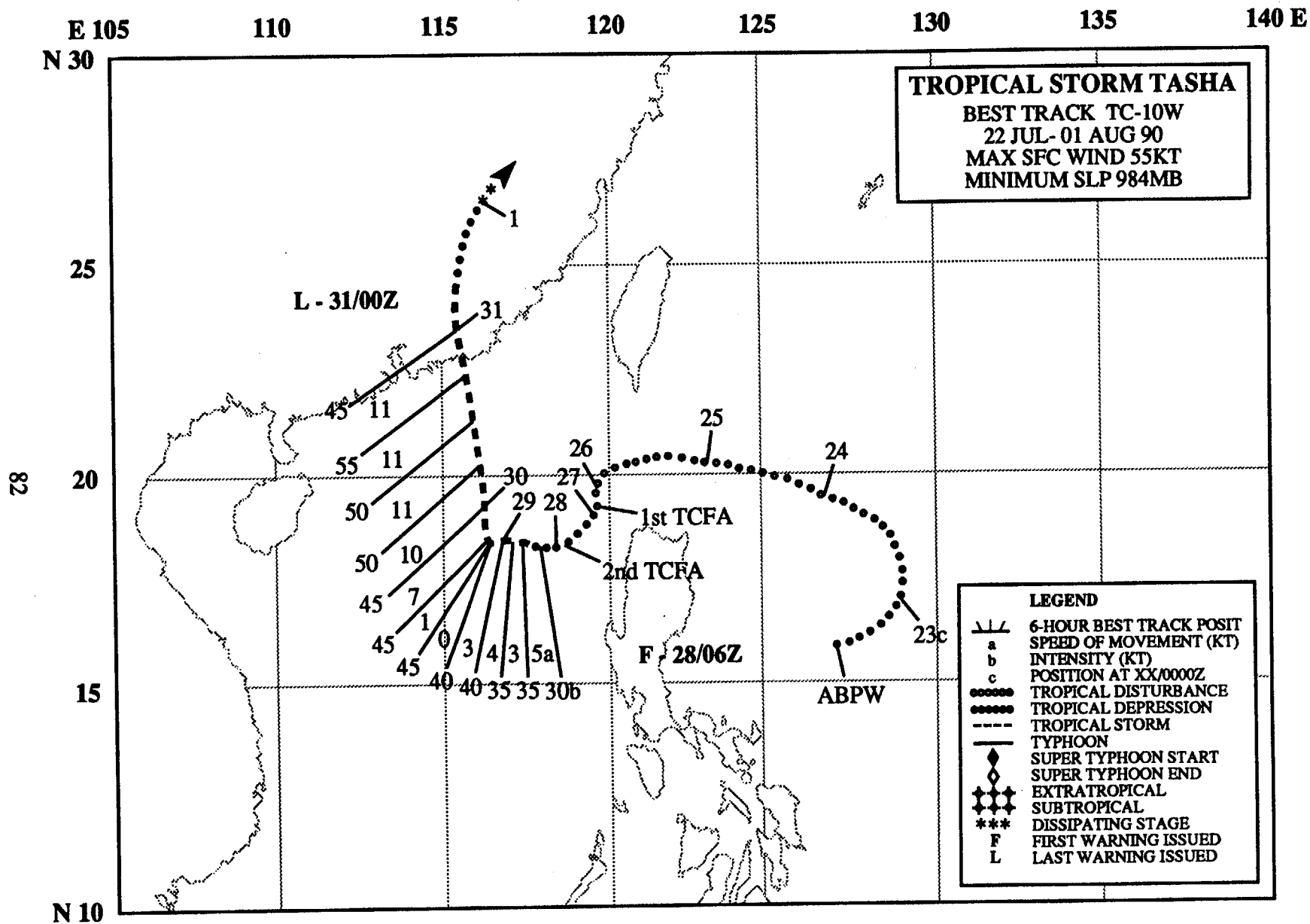


Figure 3-09-7. Summary of JTWC forecasts (solid lines) for Steve is superimposed on the best track (dashed line).



TROPICAL STORM TASHA (10W)

I. HIGHLIGHTS

Tasha, the third of four western Pacific tropical cyclones to occur in July, developed in the monsoon trough. Instead of following Steve (09W) and Vernon (11W) to the northeast, it made only a brief start in that direction before curving to the west and entering the South China Sea. After erratic motion and slow intensification, Tasha finally reached tropical storm intensity before slamming into the southern coast of China.

II. CHRONOLOGY OF EVENTS

- 220600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1006 mb.
- 262000Z - Tropical Cyclone Formation Alert issued based on indications in the synoptic data of increased organization of the low-level circulation and upper-level outflow.
- 272000Z - Tropical Cyclone Formation Alert reissued based on increased central convection and falling surface pressures.
- 280600Z - First warning issued due to preliminary appearance of a central dense overcast.
- 281200Z - Upgraded to tropical storm based on a ship report of 35 kt (20 m/sec) and a minimum sea-level pressure of 995 mb.
- 301800Z - Peak intensity of 55 kt (28 m/sec) coincident with increased size of the central dense overcast and an intensity estimate of CI 3.5.
- 310000Z - Final warning - dissipating over land - followed landfall 75 nm (140 km) east of Hong Kong.

III. TRACK AND MOTION

Tasha, developed in the monsoon trough over the warm 84°F (29°C) waters of the Philippine Sea. The low-level cyclonic circulation initially tracked northeastward in response to shallow southwesterly wind flow that extended up to 700 mb (Figure 3-10-1). As the pre-Tasha disturbance

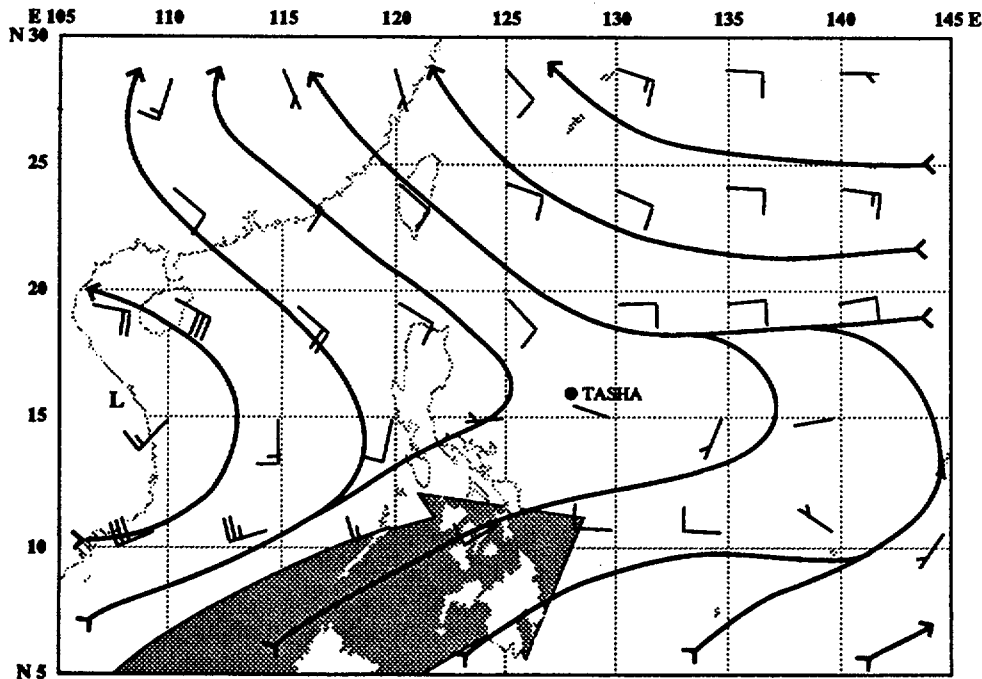


Figure 3-10-1. 700-mb NOGAPS streamline analysis at 221200Z July, showing the southwesterly steering flow over the southern Philippine Sea.

continued to develop, it turned westward in response to easterly flow associated with the an extension of the subtropical ridge centered over the East China Sea (Figure 3-10-2). For the next several days, the disturbance drifted slowly westward and passed through the Luzon Strait. At this point, Tasha moved slowly southward and westward, interacting with a larger, synoptic-scale cyclonic circulation to the southwest in the monsoon trough (Figure 3-10-3). By 29 July, Tasha had intensified and become the dominant vortex in the South China Sea. After a 12-hour period of quasi-stationary motion, Tasha then commenced a northward track at 291800Z in response to a moderately strong (up to 35 kt (18 m/sec) surface winds) and deep (1000 to 700 mb) surge in the monsoonal flow to the south (Figure 3-10-4) which was accompanied by a northward shift of the synoptic-scale monsoon trough axis (Figure 3-10-5). Tasha ultimately made landfall just east of Hong Kong.

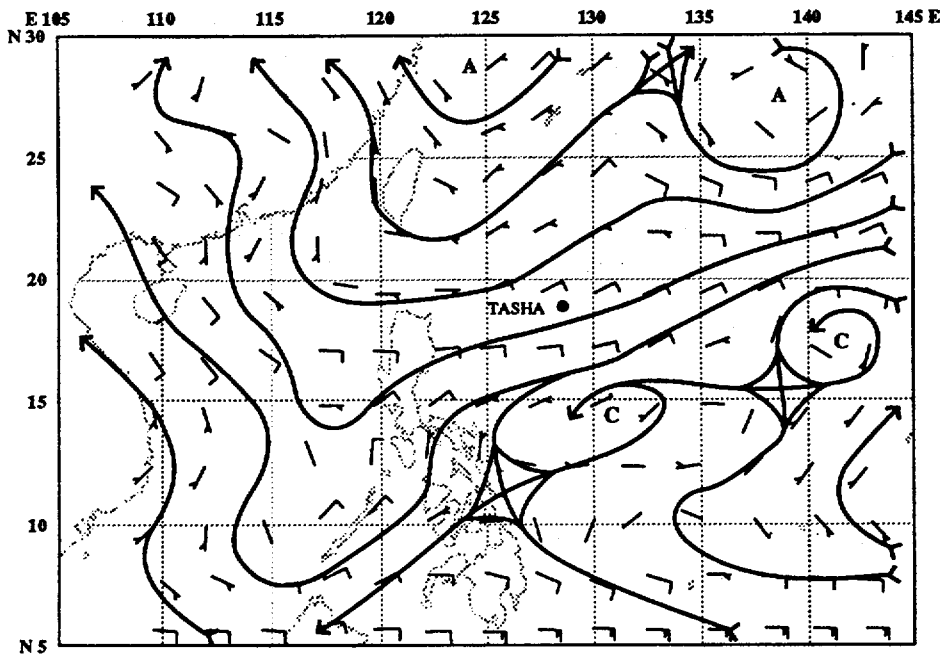
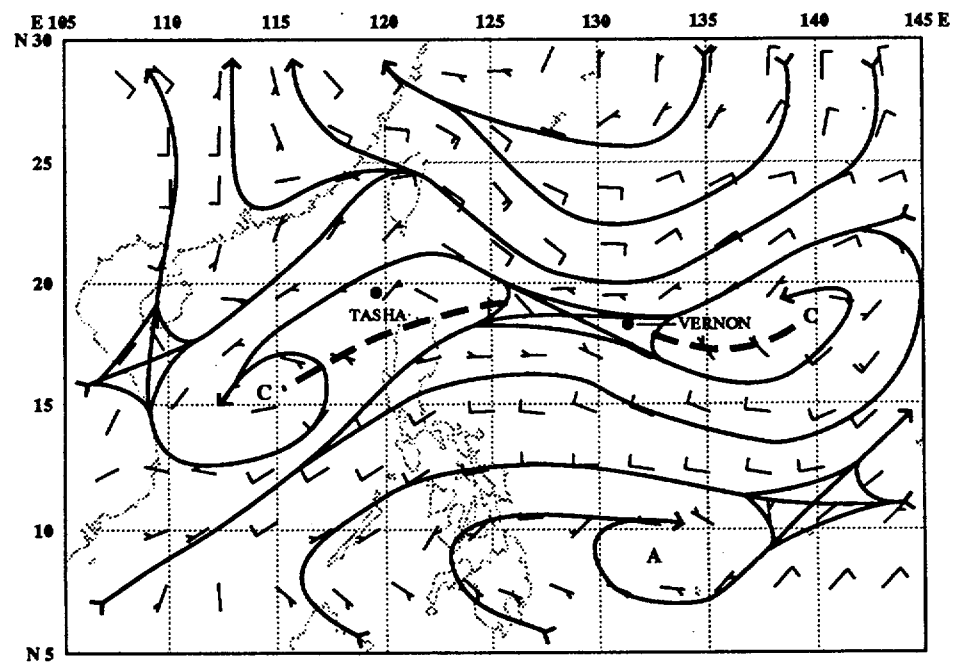


Figure 3-10-2. Deep layer mean analysis at 231200Z July, showing the mean position of the subtropical ridge over the East China Sea and weak easterly steering flow over Tasha.

Figure 3-10-3. Deep layer mean analysis at 261200Z July with Tasha beginning to interact with the large cyclonic circulation to the southwest.



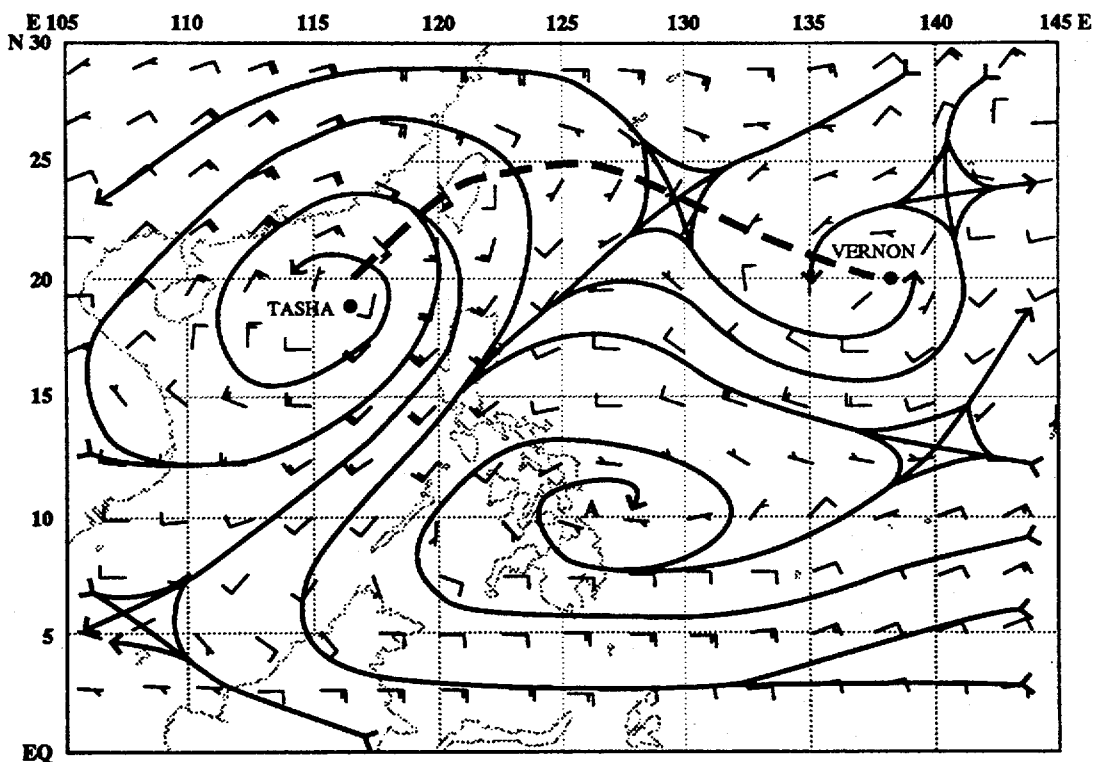


Figure 3-10-4. Deep layer mean analysis at 301200Z July depicting the moderate monsoon surge to the south of Tasha (compare with Figure 3-10-3).

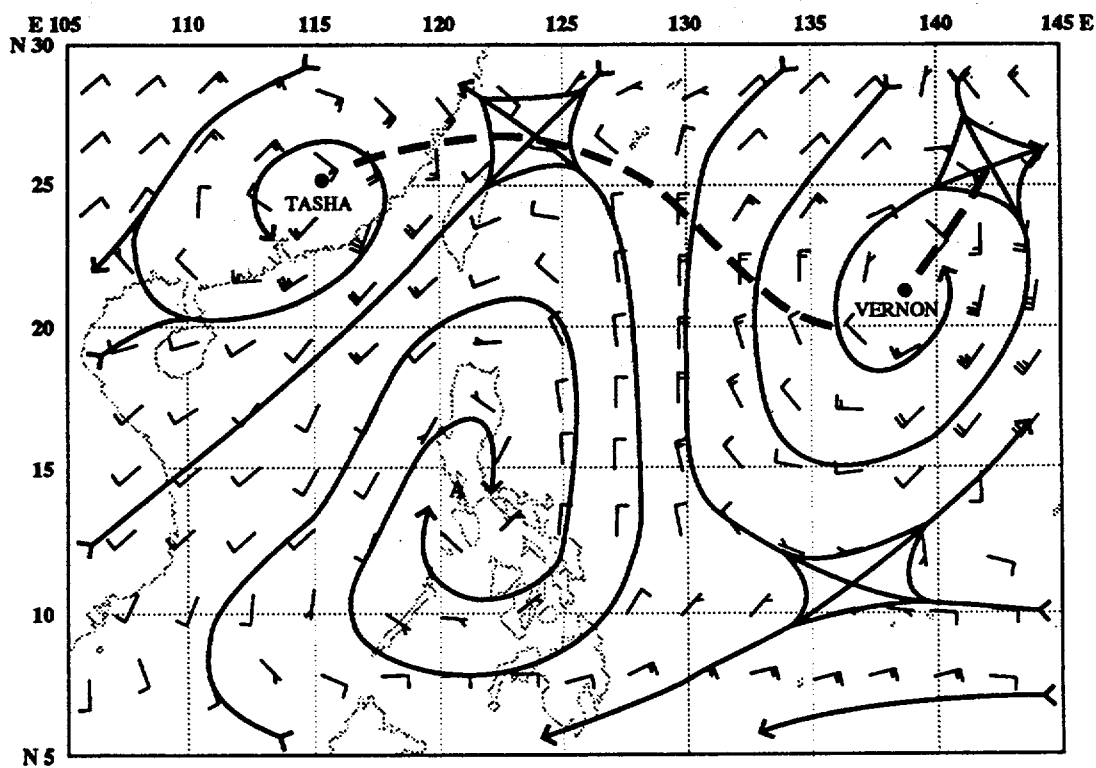


Figure 3-10-5. Deep layer mean analysis at 311200Z July, when compared with Figure 3-10-3, reveals that the axis of the monsoon trough with Tasha embedded has shifted northward over China.

IV. INTENSITY

For several days before significant development occurred, the persistent, but poorly organized, convection remained embedded in the monsoon trough, undergoing large diurnal fluctuations. During this time, the upper-level winds over the system were in excess of 30 kt (15 m/sec). However, after passing through the Luzon Strait, the tropical disturbance moved into a more favorable environment with less vertical shear near the eastern end of the tropical easterly jet. The cyclone reached peak intensity on 30 July, just prior to landfall (Figure 3-10-6). Once inland, the system dissipated due to the influence of rugged terrain in southeastern China and the loss of its oceanic source of heat and moisture.

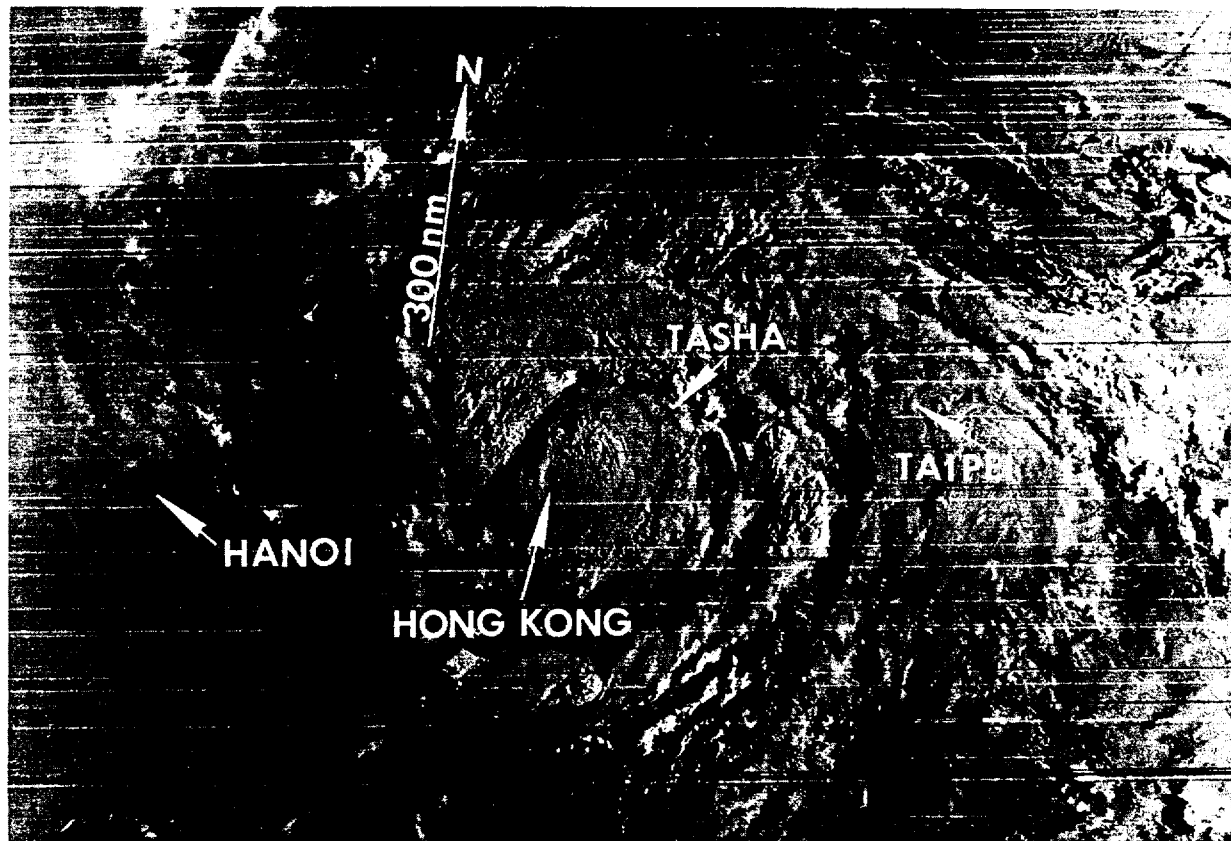


Figure 3-10-6. Tropical Storm Tasha at peak intensity moves into southern China (302210Z July DMSP visual imagery).

V. FORECASTING PERFORMANCE

Tasha's eventual northward track was not forecast initially (Figure 3-10-7). The NOGAPS prognoses maintained a weak subtropical ridge over southern China, which was expected to steer the system west-northwestward between Hainan Dao and Hong Kong. However, due to the weak steering flow depicted by the models, an alternate scenario for erratic motion was developed. On 29 July, after satellite imagery indicated that the previously mentioned monsoonal surge was beginning, the track was modified to initial northeastward movement followed by a turn to the north. If the surge turned out to be weaker than anticipated, an alternate scenario of steady northward movement was included. The alternate scenario turned out to be correct. Throughout Tasha's life, the guidance provided by the numerical forecast aids was practically useless. The major northward shift of the axis of the monsoon trough was not depicted well in the NOGAPS prognoses, and the complex and rapidly changing synoptic environment was not amenable to subjective analysis by the forecaster.

VI. IMPACT

Tasha landed 75 nm (140 km) east of Hong Kong at 312100Z and caused widespread damage due to torrential rains and flooding in Fujian and Guangdong provinces. In Fujian Province in southeastern China, 69 people were killed and 10,000 houses destroyed. Irrigation facilities were damaged, and approximately 5 million acres of farmland were flooded, with rainfall amounts reported in excess of 12 inches (305 mm). In Guangdong Province in southern China, 39 people died, 335 were injured and 25,200 houses were destroyed. Rainfall in some areas exceeded 14 inches (355 mm) with 5.3 million acres of farmland flooded. In contrast, damage in Hong Kong was relatively minor. Ferries to outlying islands, Macau and many parts of Guangdong were suspended or canceled. Seven emergency shelters were opened and many social activities were disrupted, but no serious flooding or landslides occurred.

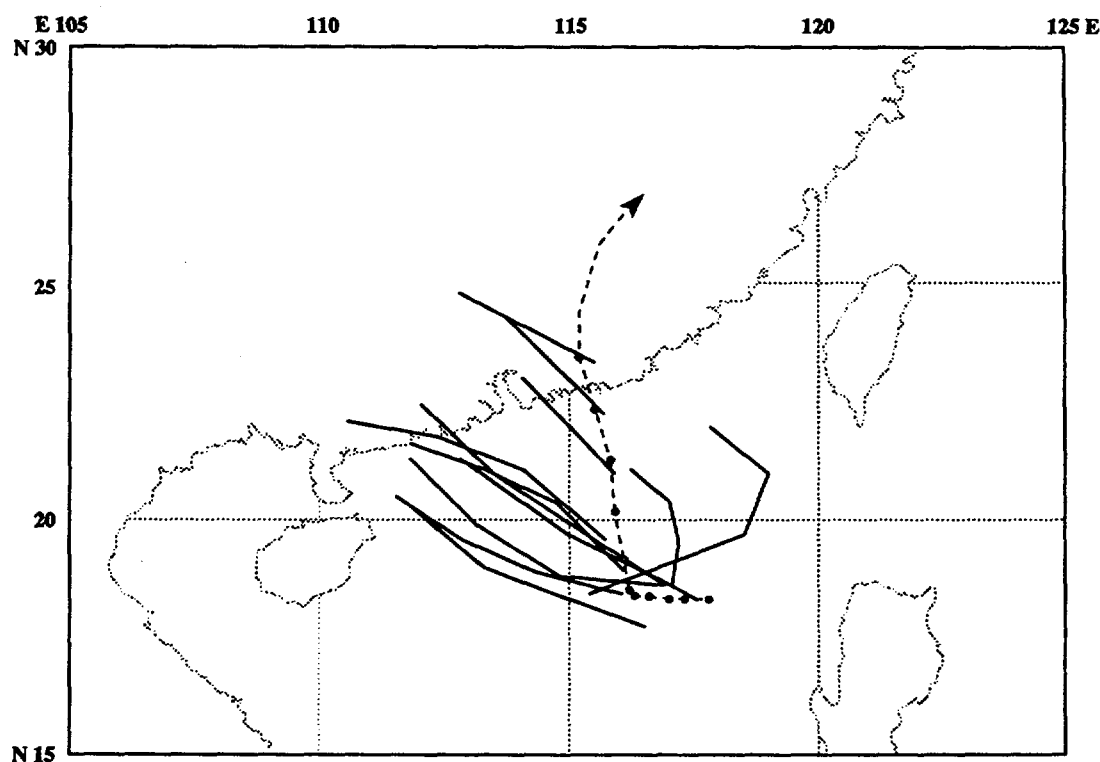
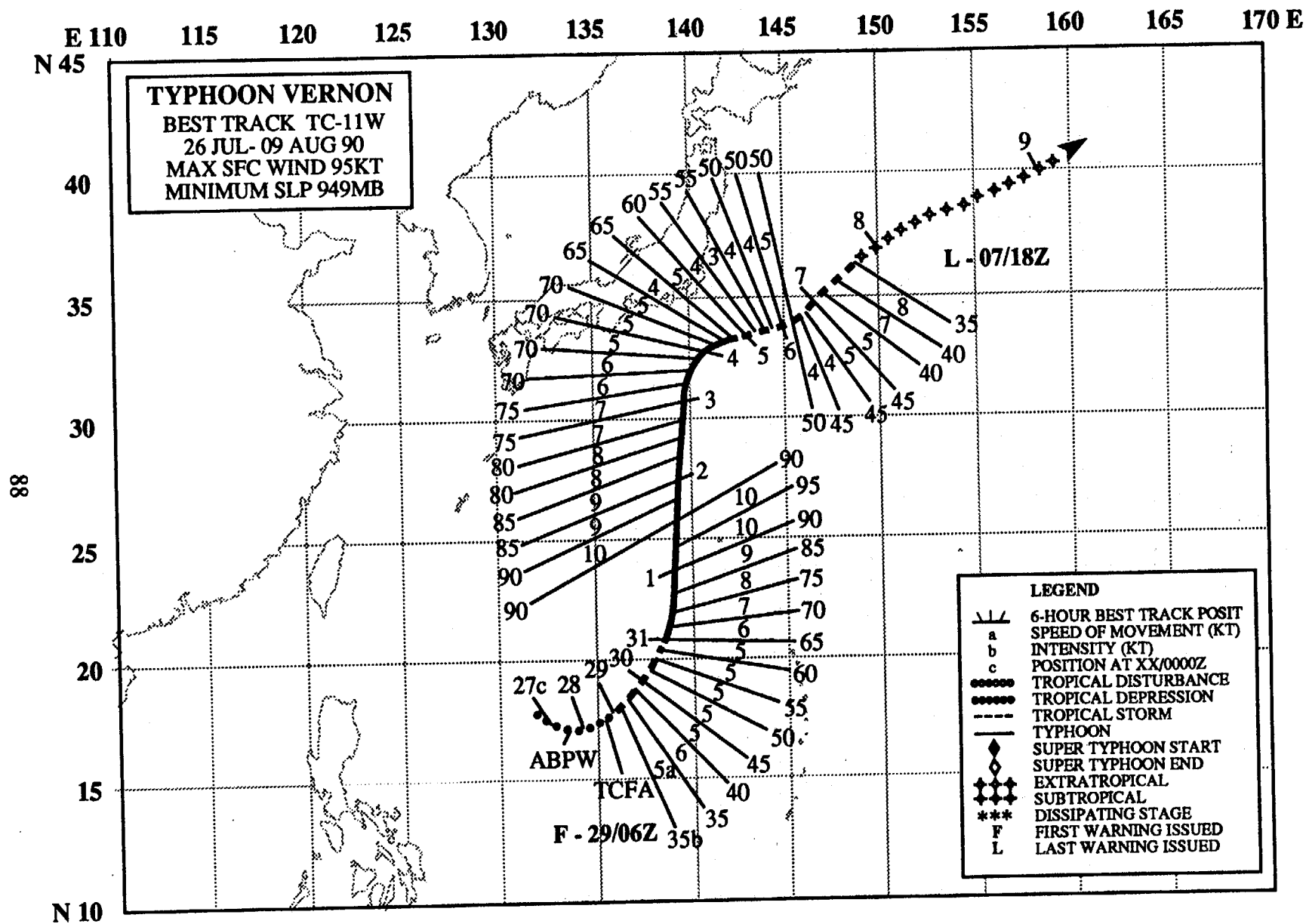


Figure 3-10-7. Summary of JTWC forecasts (solid lines) for Tasha is superimposed on the final best track (dashed line).



TYPHOON VERNON (11W)

I. HIGHLIGHTS

Vernon, the last of four tropical cyclones to develop during July, was the last of a series of storms that included Tropical Storm Tasha (10W) and Typhoon Steve (09W) to form the only three-storm outbreak in the western North Pacific during 1990. Vernon followed Steve's northward-oriented track, as the monsoon trough underwent a major displacement to the north.

II. CHRONOLOGY OF EVENTS

- 271700Z - Significant Tropical Weather Advisory reissued to include a low-level cyclonic circulation in the monsoon trough with persistent convection and an estimated minimum sea-level pressure of 1004 mb.
- 282200Z - Tropical Cyclone Formation Alert issued for increased outflow and improved convective curvature.
- 290600Z - First warning followed consolidation of convection into two interlocking cloud bands.
- 291200Z - Upgraded to tropical storm after appearance of a ragged central dense overcast.
- 310000Z - Upgraded to typhoon based on eye development.
- 010600Z - Peak intensity - 95 kt (48 m/sec) - based on intensity estimate of CI 5.5 at 010300Z.
- 050000Z - Downgraded to tropical storm intensity due to the loss of central convection.
- 071800Z - Final warning - extratropical - issued as Vernon continued to lose its supporting convection.

III. TRACK AND MOTION

During the last week of July, the western portion of the active monsoon trough was anchored in Asia by Tasha (10W). The trough extended eastward across the Philippine Sea through Steve (09W) and north-northeastward to a mid-level cyclonic circulation east of Honshu (Figure 3-11-1). Vernon developed in the monsoon trough between these two tropical cyclones and moved slowly eastward along the trough axis on the edge of the deep southwesterly flow. The eastward track along the trough axis became more northward as the entire monsoon trough shifted northward throughout the week. As Vernon approached Japan, the Asian High persisted across Honshu, and Vernon was forced to slow and

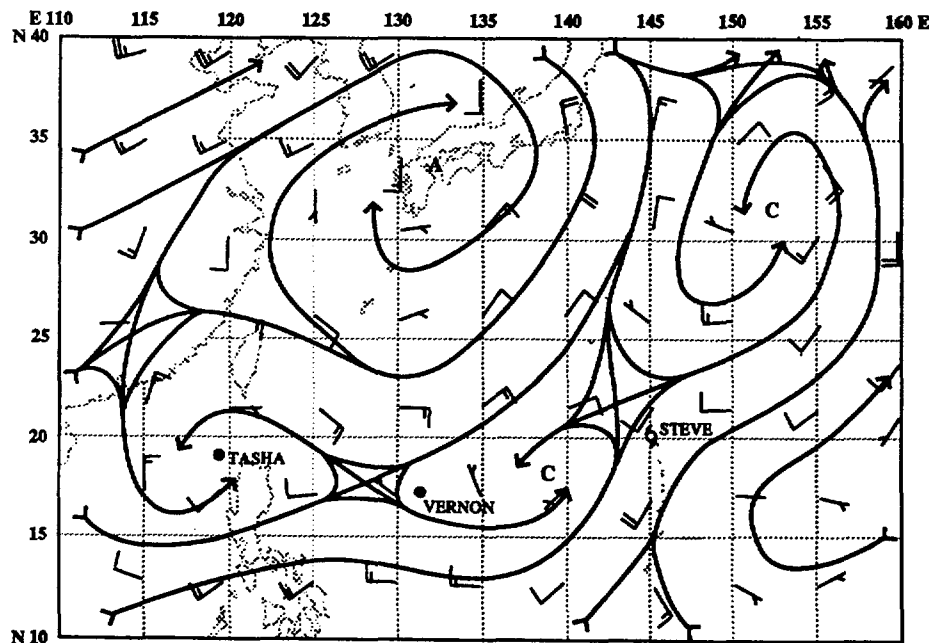


Figure 3-11-1. The 270000Z July NOGAPS 500-mb analysis shows the monsoon trough, extending eastward from Tasha (10W) through Steve (09W) to a low east of Honshu. Equatorward of the trough axis the deep southwesterly flow extends up through the middle troposphere. 84

track around the southern portion of the High. Vernon and Steve (09W) moved in a similar manner and maintained a separation of approximately 800 nm (1480 km) until Steve accelerated northeastward.

IV. INTENSIFICATION

Typhoon Vernon (Figure 3-11-2) intensified steadily despite the proximity to both Steve (09W) and Tasha (10W). The upper-level outflow from Tropical Storm Tasha, however, disrupted Vernon's vertical alignment. Only after Tasha dissipated over China on 31 July was Vernon able to develop into a typhoon (Figure 3-11-3). Approaching Honshu on 3 August, the eye of the typhoon became elongated along an east-west axis and lost much of its definition. After turning northeastward, Vernon (Figure 3-11-4) began a slow extratropical transition.

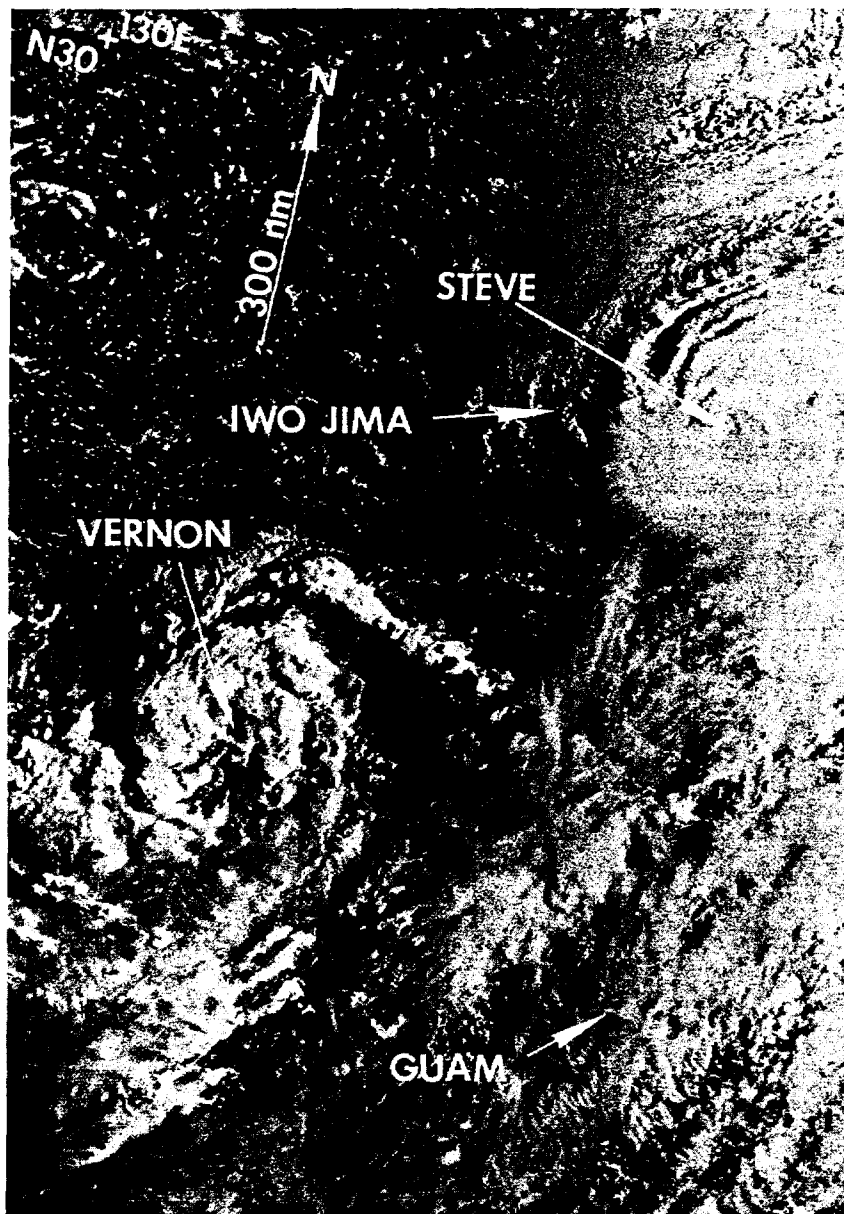


Figure 3-11-2. The tropical disturbance which became Typhoon Vernon develops approximately 700 nm (1300 km) southwest of Typhoon Steve (09W). The curved convective bands indicate the system is developing. A Tropical Cyclone Formation Alert was issued one hour after this image was received (282055Z July DMSP visual imagery).

V. FORECASTING PERFORMANCE

Due to their close proximity, forecasters initially considered the possibility of binary interaction between Vernon and Steve (09W). However, after rotation around a common midpoint was not observed, the binary interaction scenario was discarded in favor of a north-oriented forecast track similar to that taken by Steve three days earlier. The first nine forecasts (Figure 3-11-5) using this scenario were extremely accurate and had 72-hour mean forecast errors of less than 100 nm (185 km). As Vernon moved further north, forecasters experienced the same dilemma as with Steve (09W). The NOGAPS prognostic series indicated the subtropical ridge would build from the east, displacing the cyclone further west with landfall in the heavily populated areas of Japan. The forecasts were based on this guidance. As Vernon moved northward, the ridge built in from the east as forecast, but further to the south. At 040600Z August, it became evident that the mid-level ridge would hold across Honshu, and the forecast track was changed from northward to northeastward and away from Japan.

As a point of interest, the NOGAPS and JMA models had totally different prognostic solutions for the ridge motion across Japan. JMA retained the ridge and let Vernon move north-northeastward. NOGAPS linked the ridge over Japan with the maritime subtropical ridge to the east, and then moved the ridge off the island and over the Pacific. The NOGAPS guidance was used for the forecasts and the JMA guidance became the alternate scenario. In retrospect, the alternate scenario proved to be correct.

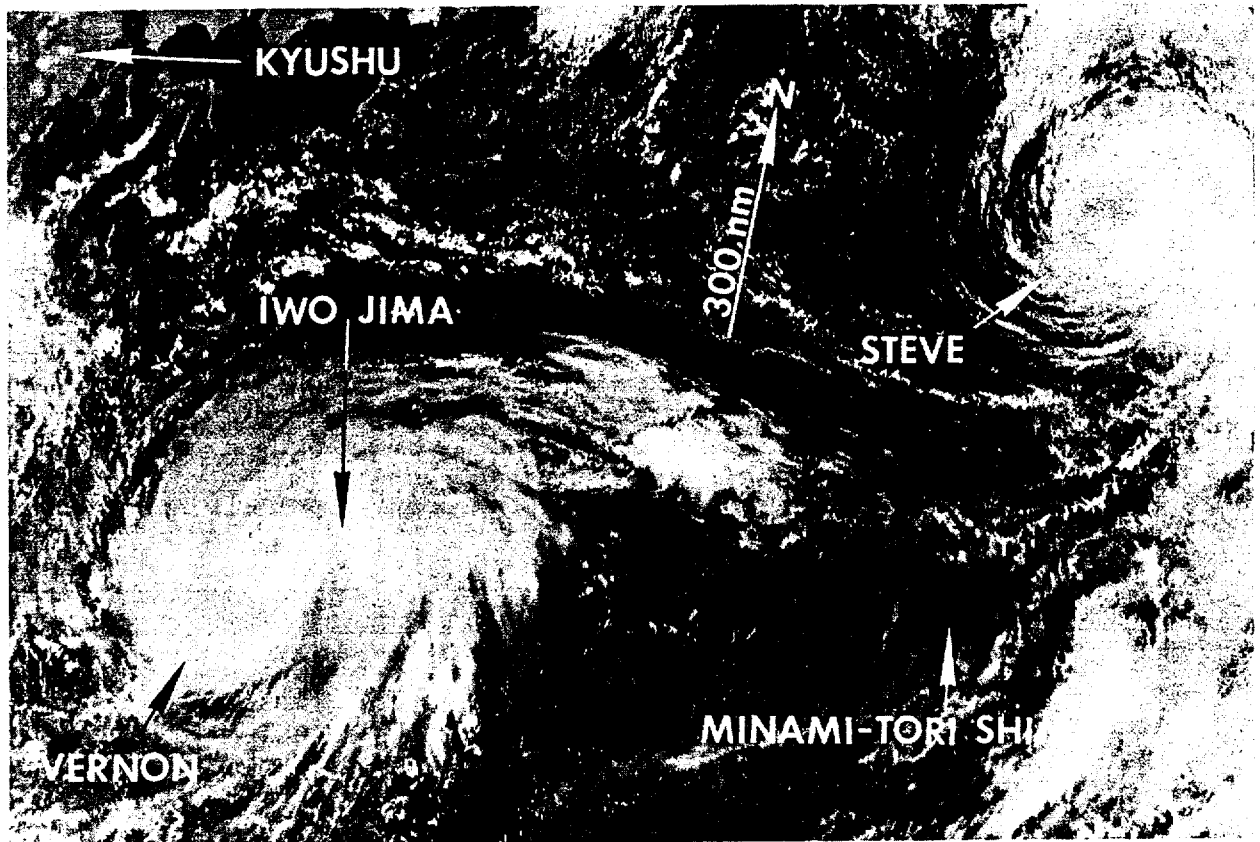


Figure 3-11-3. Typhoon Vernon near maximum intensity of 95 kt (48 m/sec). At this point, Vernon has a well-defined, but cloud-filled eye, and Typhoon Steve is weakening over water (312316Z July DMSP visual imagery).

VI. IMPACT

Although Typhoon Vernon threatened the Tokyo metropolitan area, it veered northeastward, passing within 120 nm (220 km) of the Japanese coast. There were no deaths or significant damage reports related to Vernon.

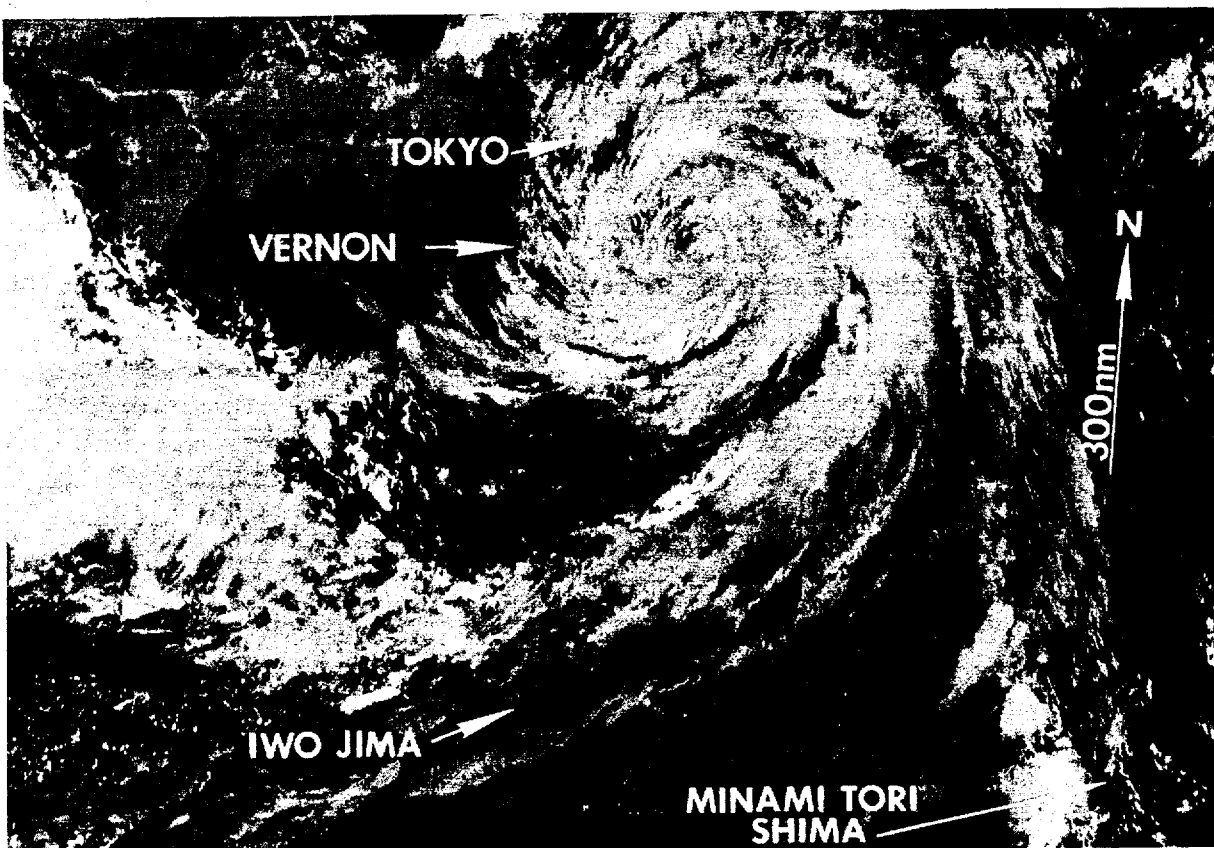


Figure 3-11-4. Vernon as it was downgraded to a tropical storm. Most of the deep central convection has diminished, leaving a well-defined low level circulation of stratocumulus and cumulus clouds (042333Z August DMSP visual imagery).

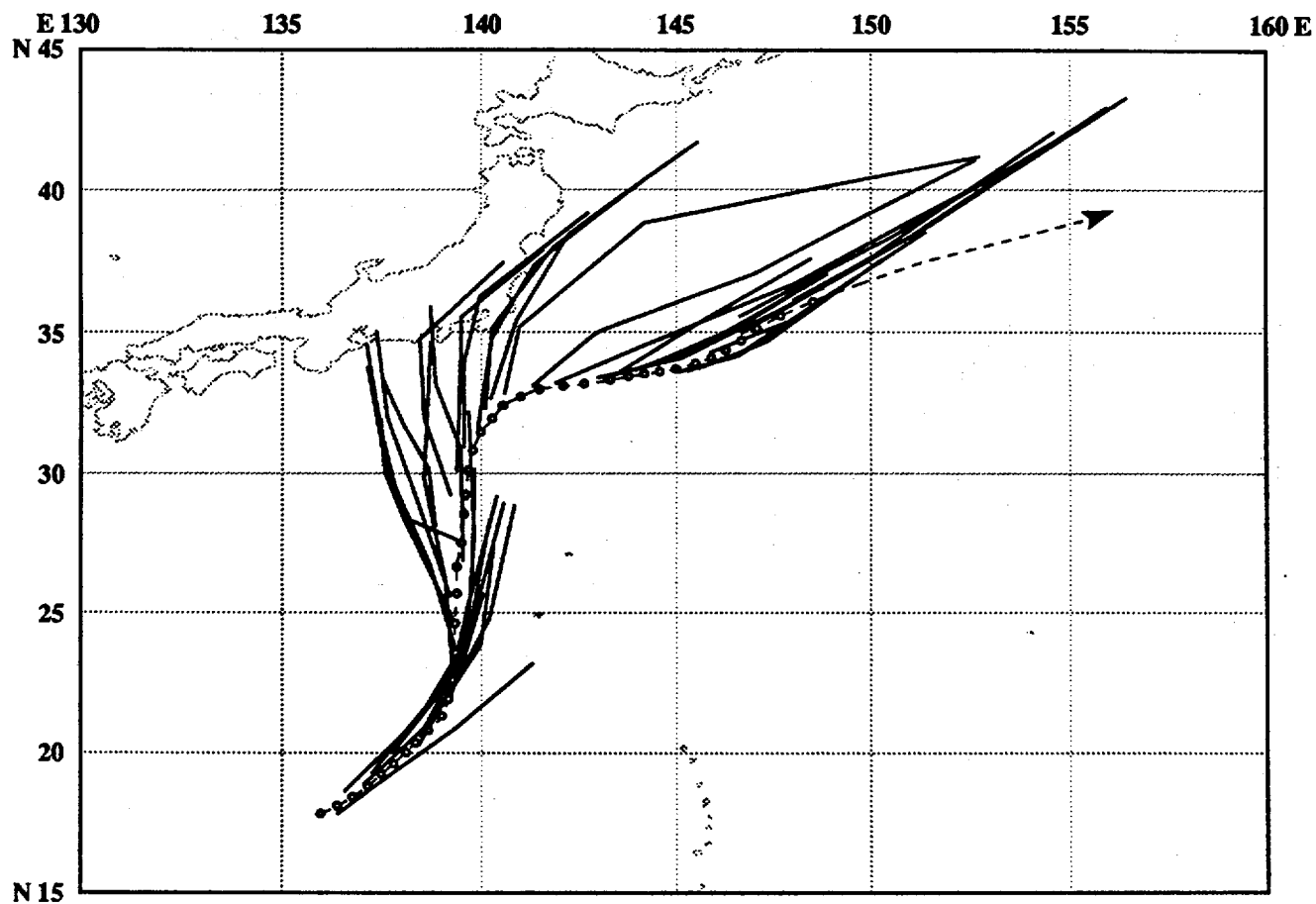
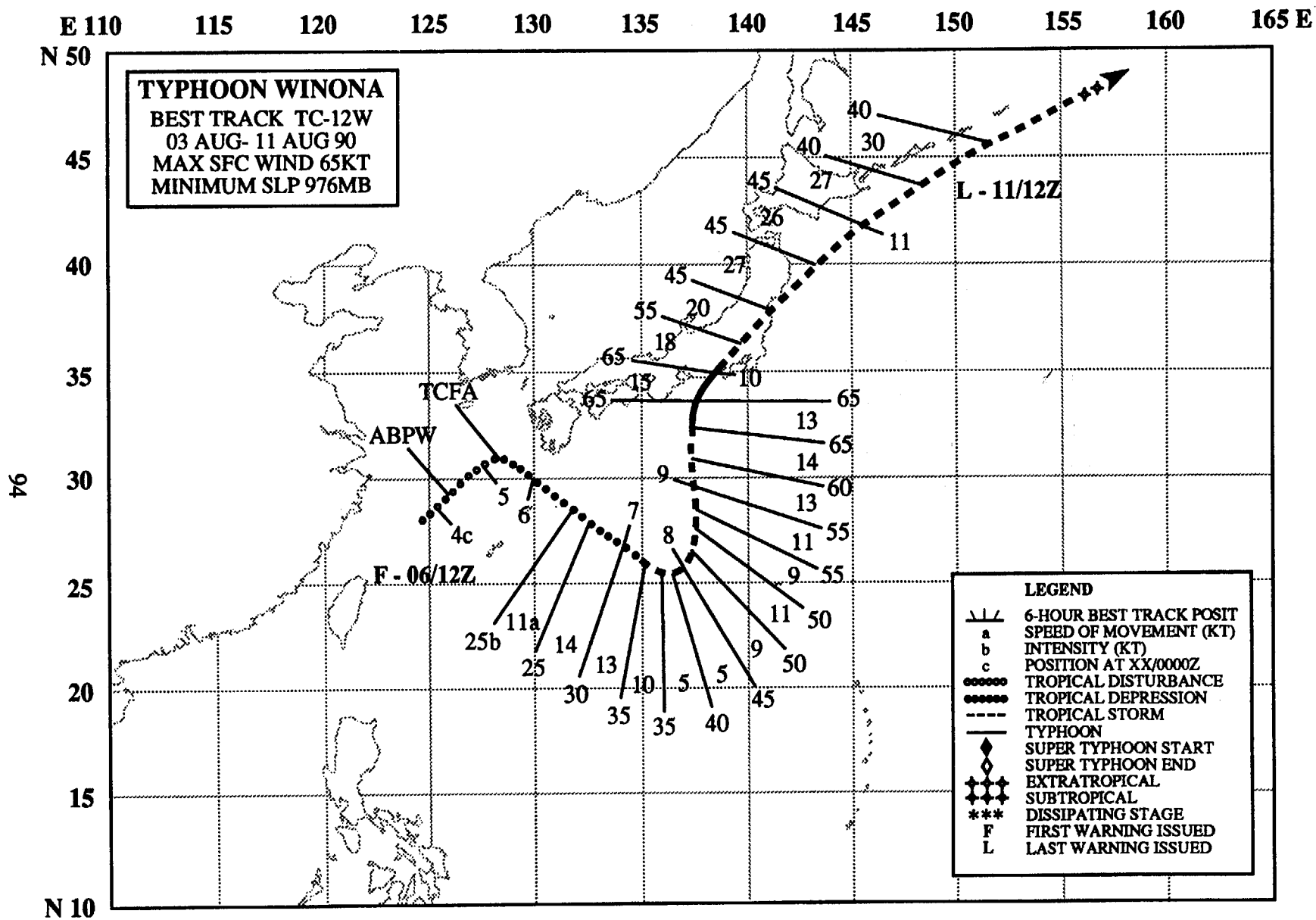


Figure 3-11-5. Summary of JTWC forecasts (solid lines) for Vernon (11W) is superimposed on the final best track (dashed lines).



TYPHOON WINONA (12W)

I. HIGHLIGHTS

Winona was the first typhoon of 1990 to hit Japan and the only tropical cyclone to form poleward of 25 degrees north latitude. It formed from the remnants of Tropical Storm Tasha (10W) in a monsoon trough displaced northward of its normal location. Winona tracked across the southern portion of the Kanto Plain, was caught in the westerlies, and completed extratropical transition as it swept just south of the Kurils.

II. CHRONOLOGY OF EVENTS

- 040600Z - First mentioned on the Significant Tropical Weather Advisory as remnants of Tropical Storm Tasha moving off China and reforming as a weak circulation in the East China Sea.
- 051100Z - Tropical Cyclone Formation Alert issued based on improved convective organization and Dvorak analysis of CI 1.0.
- 061200Z - First warning issued as a tropical depression. Although both convection and organization had improved, vertical shear from the northwest inhibited further development.
- 070600Z - Upgraded to tropical storm as vertical shear decreased and circulation center and convection became better aligned.
- 091200Z - Upgrade to typhoon and peak intensity - 65 kt (33 m/sec)- based on a ragged eye and first intensity estimate of CI 4.0.
- 100000Z - Landfall on Japan 20 nm (35 km) east of Hamamatsu, a city 110 nm (205 km) southwest of Tokyo. Downgraded to tropical storm.
- 111200Z - Final warning - extratropical - issued as Winona became embedded in mid-latitude westerlies.

III. TRACK AND MOTION

Winona was unique in regard to both its genesis and its movement. The system formed in the monsoon trough, which was displaced 300 nm (555 km) north of its normal location. The initial southeastward movement almost directly opposed the expected climatological track. Winona typified the complex interaction that can occur among tropical cyclones, the deep monsoon southwesterlies, and the subtropical ridge. Winona later moved north, then northeast, in response to a well-developed mid-latitude trough.

Enhanced convection became prevalent in the East China Sea as the low pressure area associated with the remnants of Tropical Storm Tasha (10W) moved out to sea by 040000Z August. This area of enhanced convection developed into Winona. The system tracked northeastward initially, then southeastward along the edge of the deep monsoon westerlies. The 500-mb analysis at 070000Z (Figure 3-12-1) shows Winona embedded in a complex flow pattern with Tropical Storm Vernon (11W) to the northeast. The subtropical ridge had split, with one cell centered in the Luzon Strait, and the other south of Vernon. Winona tracked toward the neutral point between the two cells.

By 080000Z, Vernon (11W) had tracked northeastward and become extratropical. At the same time, Winona slowed to 4 kt (7 km/hr) and turned sharply northward as the ridge to the southeast built poleward. As Vernon (11W) completed its extratropical transition at 090000Z near the Kamchatka Peninsula, the ridge strengthened north and northeast of Winona in response to the extratropical cyclone's rapid deepening. In response, Winona maintained a northward track until it made landfall near Hamamatsu, Japan. After landfall, it began to accelerate northeastward, and by 101200Z, Winona was embedded in the mid-latitude westerlies, beginning its extratropical transition. Winona finished its extratropical transition by 111200Z as it skirted south of the Kuril Islands.

IV. INTENSITY

Winona developed as the remnants of Tasha (10W) moved off the coast of China into the East China Sea. The disturbance generated persistent convection, but it was subject to strong upper-level northerly flow (Figure 3-12-2). The strong vertical wind shear left Winona's circulation center exposed north of the deep convection.

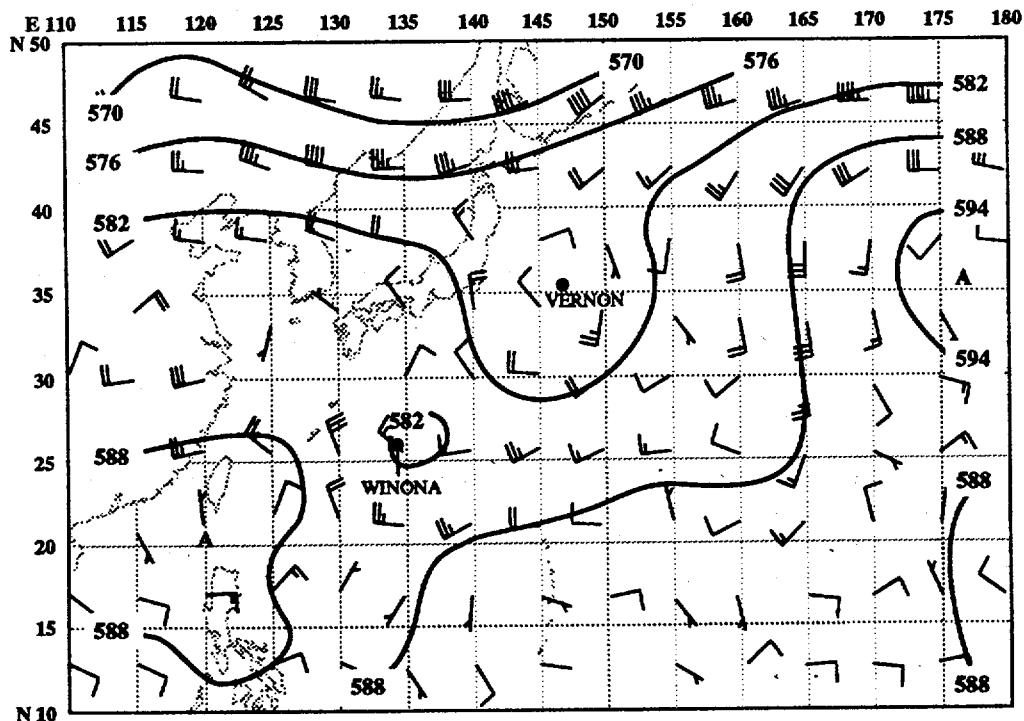
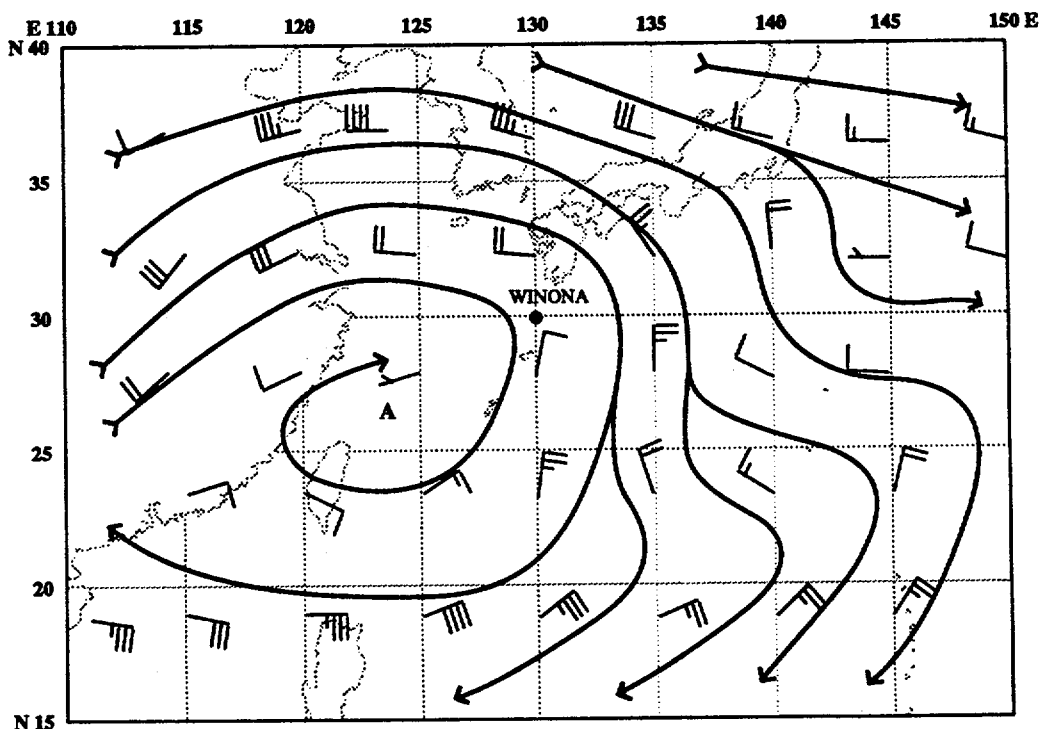


Figure 3-12-1. The 500-mb NOGAPS analysis for 070000Z August depicts Winona tracking between the two subtropical highs. Tropical Storm Vernon (11W) northeast of Winona. Note: heights are in decameters.

Figure 3-12-2. The 200-mb analysis for 060000Z August shows Winona is under the influence of unidirectional flow from the north.



As the shear decreased, the LLCC moved under the deep convection, and the system began to intensify. On 070000Z, a ship (call sign JFYD) approximately 215 nm (400 km) south of the center reported 35 kt (18 m/sec) southwesterly winds. At 070448Z (Figure 3-12-3), satellite analysts provided the first CI 2.5 Dvorak analysis, and the system was upgraded to a tropical storm. Winona continued to move southwestward toward a col and away from the shear, as it intensified. As Winona tracked northward, after an abrupt turn, it intensified further and developed dual upper-level outflow channels: one to the northeast and southwest. By 091200Z, Winona reached its maximum intensity of 65 kt (33 m/sec) and maintained it until making landfall 12 hours later. Winona weakened but managed to retain some strength and organization throughout its track over land. Yokota Air Base (WMO 47642) received peak winds of 40 kt (21 m/sec) with gusts to 57 kt (29 m/sec) at 100322Z and nearby Camp Zama had gusts up to 63 kt (32 m/sec) recorded at 100250Z. Winona got caught up in the westerlies as it reentered the water east of the Kanto Plain and became extratropical.

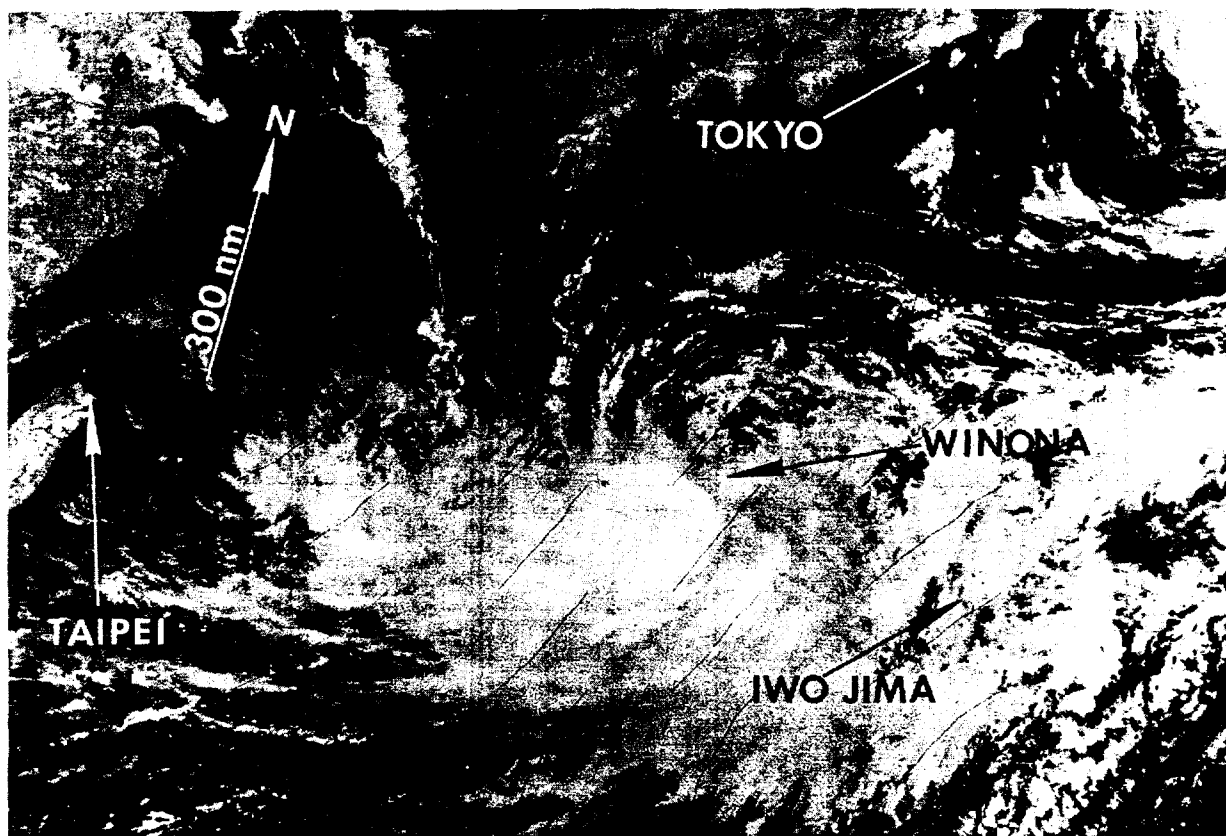


Figure 3-12-3. Winona's low-level circulation center is moving underneath the deep convection. This was the first good indication that Winona would intensify. Vernon (11W) is northeast of Winona (070448Z August NOAA visual imagery).

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance is shown in Figure 3-12-4. The initial forecasts on Winona predicted a weak, disorganized system that would be short lived. Reorganization of the system as the center moved under the deep convection caused the track to be relocated on the third warning. By then, JTWC had a much better handle on the system and correctly forecast the sharp 120 degree turn to the north. This was 12 hours ahead of other agencies. Forecasters were slow in developing the system until it made the turn. After the LLCC moved under the deep convection, JTWC correctly predicted the effect that dual outflow channels would have in rapidly deepening the system. The objective aids FBAM and CSUM had problems with Winona's track. FBAM continued to move Winona south around the ridge until the system made the turn, then it caught on and went due north. CSUM started the turn too early and made it too tight, coming in west of the actual track. NOGAPS correctly built the ridge northward, which caused the push to the north. In addition to the accurate northward forecast, JTWC accurately forecast landfall. Forecasters then expected Winona to track northeastward north of the subtropical high and get caught up in the westerlies. A big decision centered around which way the storm would track around Mt. Fuji-san. JTWC did not predict the ridge flattening overnight and opted for an initial track through central Honshu west of Mt. Fuji-san, then skirting northern Honshu just off the coast in the Sea of Japan. As a short wave trough passed to the north, the ridge damped, and Winona turned sooner than forecast. Both JTWC and the Japan Meteorological Agency brought their tracks further south once it was obvious that Winona would track south of Mt. Fuji-san. Both agencies also kept their forecast tracks over the northeastern edge of Japan, skirting along or just south of the Kurils. Winona tracked just south of these forecasts. Both agencies, however, correctly forecast the acceleration of the system as it became embedded in the westerlies and subsequently became extratropical.

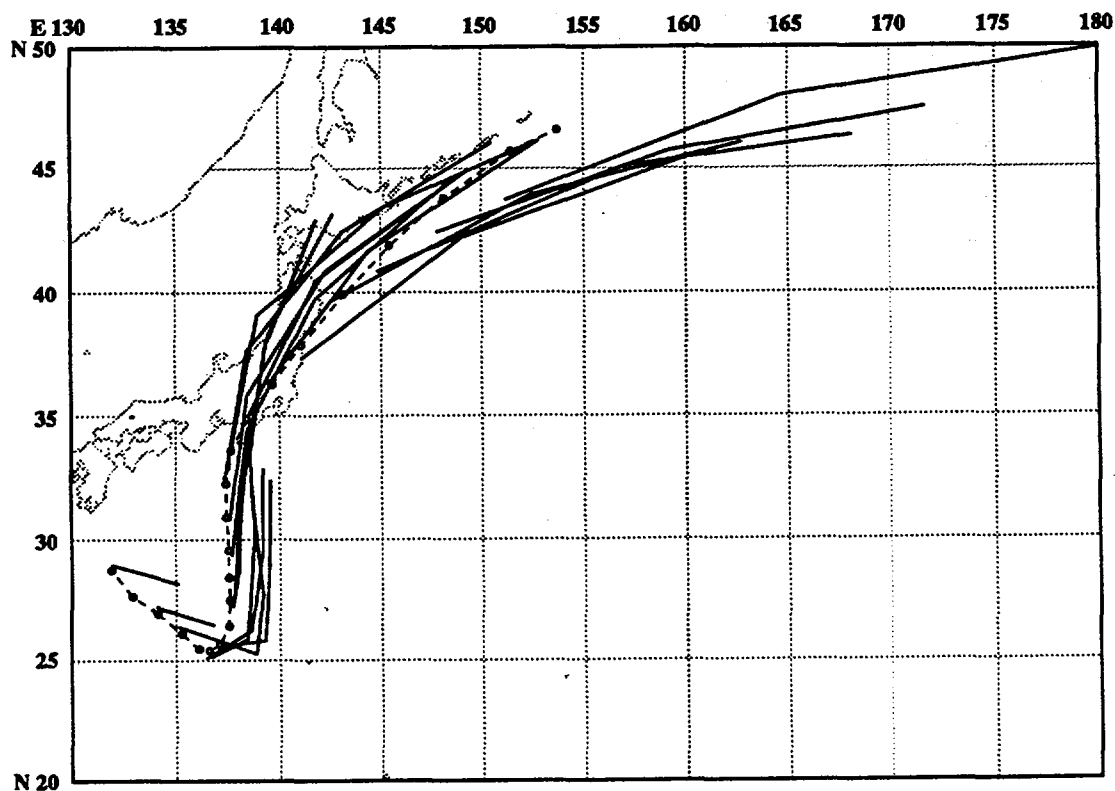


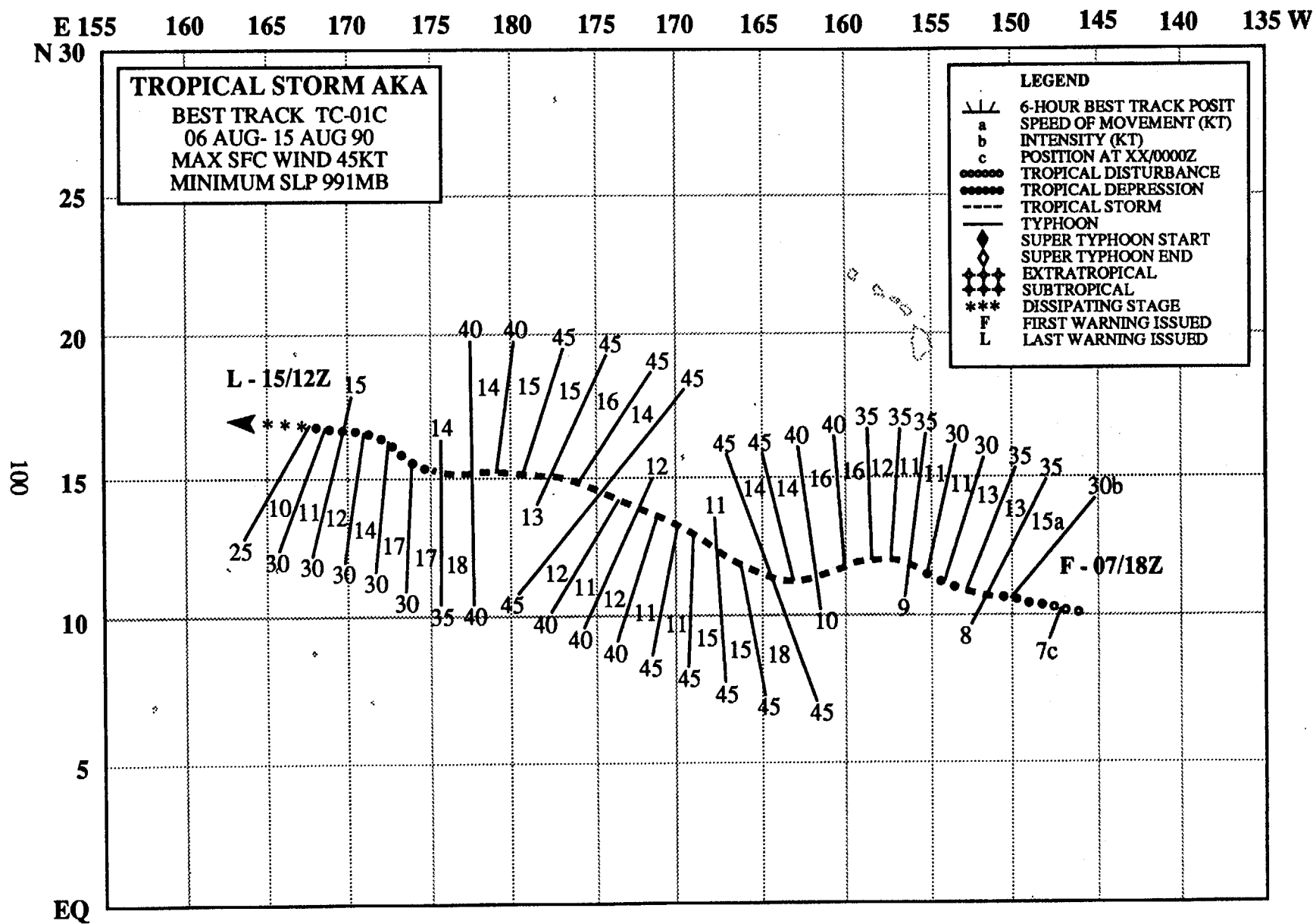
Figure 3-12-4. Summary of JTWC forecasts (solid lines) for Winona is superimposed on the final best track (dashed lines).

VI. IMPACT

Damages to U.S. military installations in Japan were minimal. Trees on bases were uprooted, tiles were blown off roofs, and there were isolated power outages.

The rest of Japan did not fare as well. According to reports from various Japanese newspapers, there were 13 typhoon related injuries but no deaths. In southeast Japan over 7000 homes in the Shizuoka Prefecture cities of Yaizu, Hamamatsu and Shimoda lost electricity as trees fell on the power lines. There were 686 homes flooded and 43 landslides. Transportation was disrupted, and over 500,000 travelers were affected by either the cancellation of 62 domestic flights from Tokyo's Haneda Airport or the many trains that were halted or delayed. All Tokaido Shinkansen bullet trains between Tokyo and Osaka were stopped. The teams scheduled to play in the Tokyo Dome could not find transportation, resulting in the first "rained out" game since the stadium was built in 1987.

The much needed rainfall poured more than 20 million tons of water into Japanese reservoirs, pushing them up to 36 percent of their total capacity. This allowed the lifting of water restrictions which had gone into effect earlier in the year.



TROPICAL STORM AKA (01C)

I. HIGHLIGHTS

Aka was the only tropical cyclone of 1990 to be in warning status when it crossed the date line from the central into the western North Pacific Ocean. It remained embedded in the trade wind trough, tracked steadily west-northwestward and never developed beyond tropical storm intensity.

II. CHRONOLOGY OF EVENTS

- 071800Z - First advisory issued by the Central Pacific Hurricane Center (CPHC) due to increased organization and amount of deep central convection.
- 090000Z - Upgraded to tropical storm intensity after convective organization improved and the first Dvorak intensity estimate of 2.5.
- 131500Z - Final advisory issued by CPHC and responsibility for Aka passed to the Joint Typhoon Warning Center (JTWC).
- 131800Z - First warning on Aka issued by JTWC.
- 140600Z - Downgraded to a tropical depression due to the loss of central convection resulting from persistent vertical wind shear.
- 151200Z - Final warning (dissipating over water) followed further weakening from vertical shear associated with a vigorous TUTT low to the northwest.

III. TRACK AND MOTION

Aka formed in the trade wind trough southeast of Hawaii (Figure 3-01C-1), remained embedded in the broad low-latitude easterlies and tracked steadily west-northwestward.

IV. INTENSITY

Although Aka persisted for nine days, its convection never became well organized. The system was maintained by low-level easterlies converging into the trade wind trough. However, the upper-level outflow pattern was continually disrupted by vertical wind shear. On 15 August, the low-level flow carried Aka westward under a vigorous TUTT low near the dateline. The upper level sheared away, the low level circulation dissipated and only the TUTT low remained (Figure 3-01C-2).

V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-01C-3. The NOGAPS prognostic series correctly maintained a mid- and low-level ridge north of Aka. Forecasters were uncertain about how long the tropical cyclone would persist as it approached the TUTT low. When dissipation became obvious, the forecast period was truncated and the final warning issued.

VI. IMPACT

No information available.

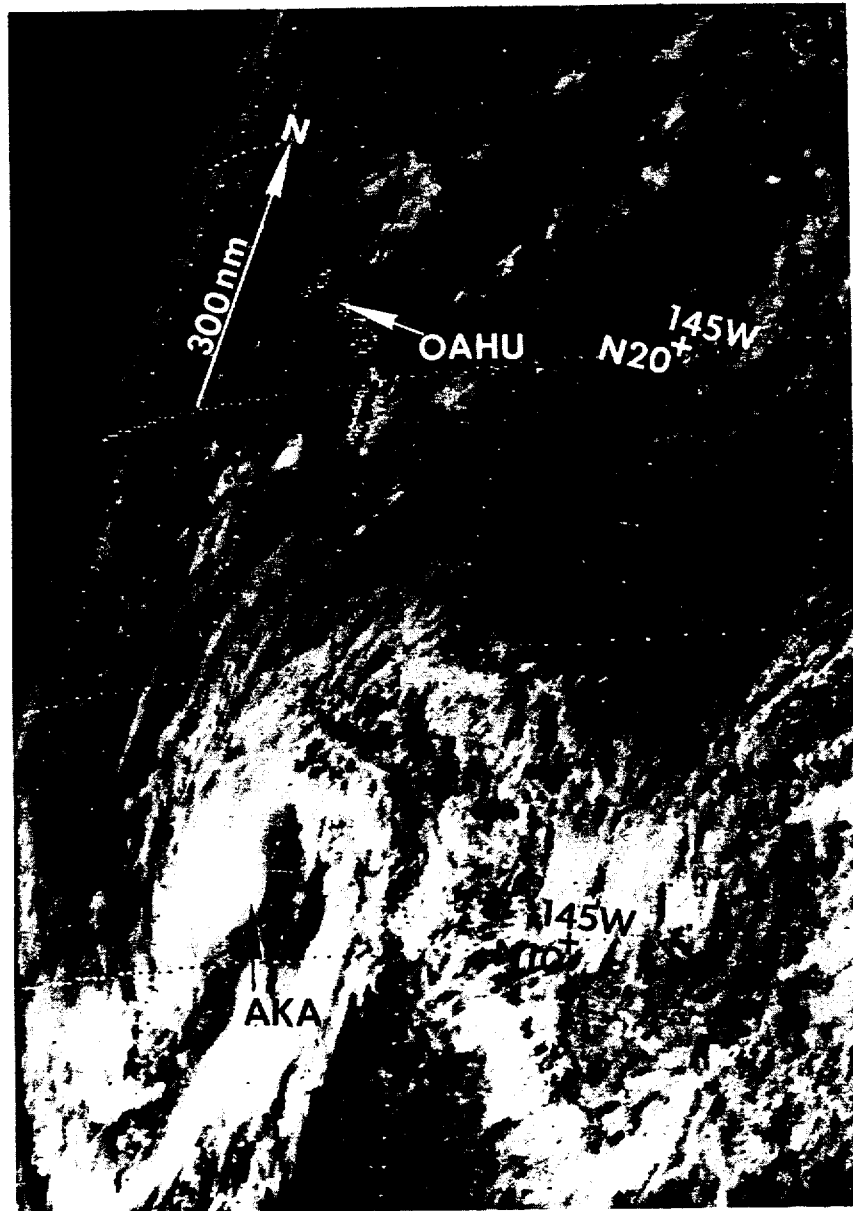


Figure 3-01C-1. Aka reaches tropical storm intensity south of the Hawaiian Islands (090101Z August GOES Central visual imagery - photo courtesy of the National Weather Service Forecast Office, Honolulu, Hawaii).

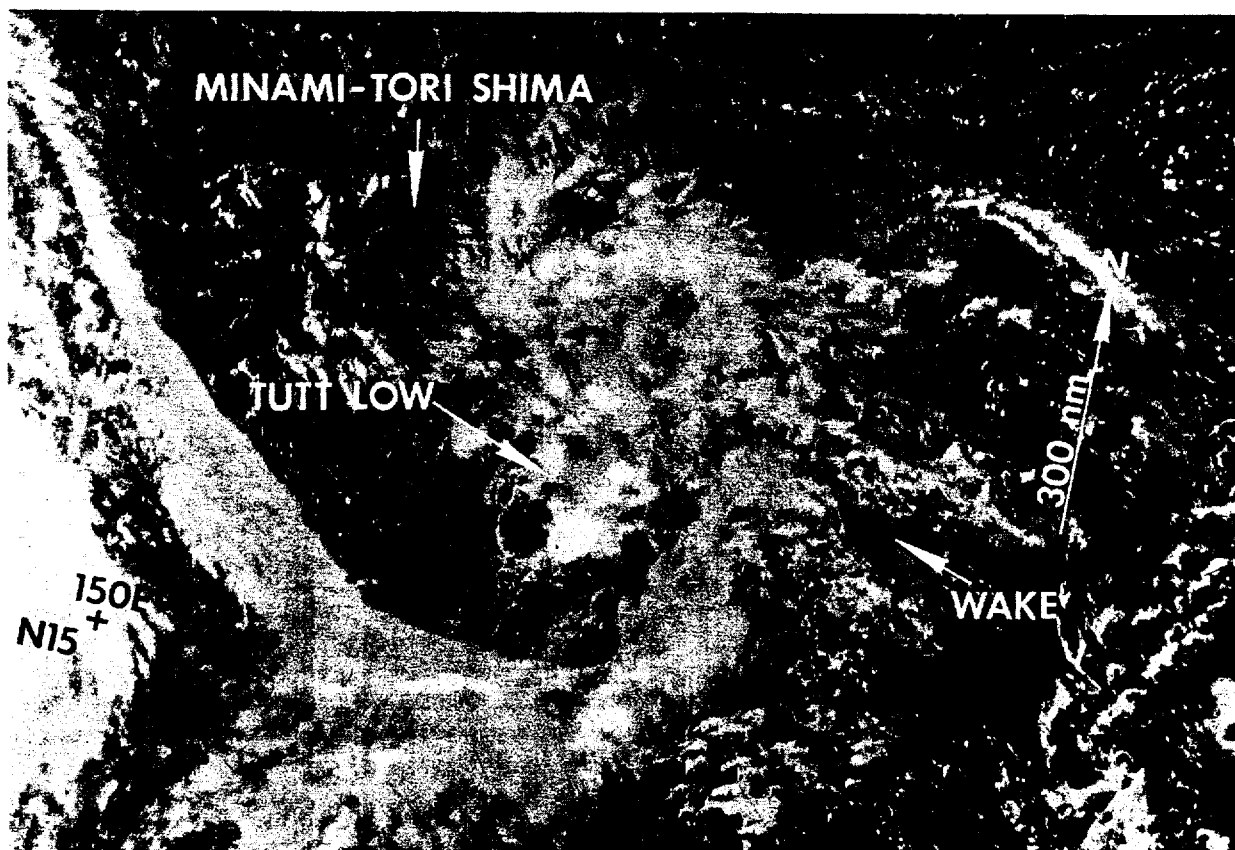


Figure 3-01C-2. The vigorous TUTT low with its random convective elements dominate the area where Aka dissipated 15 hours before (160308Z August NOAA visual imagery).

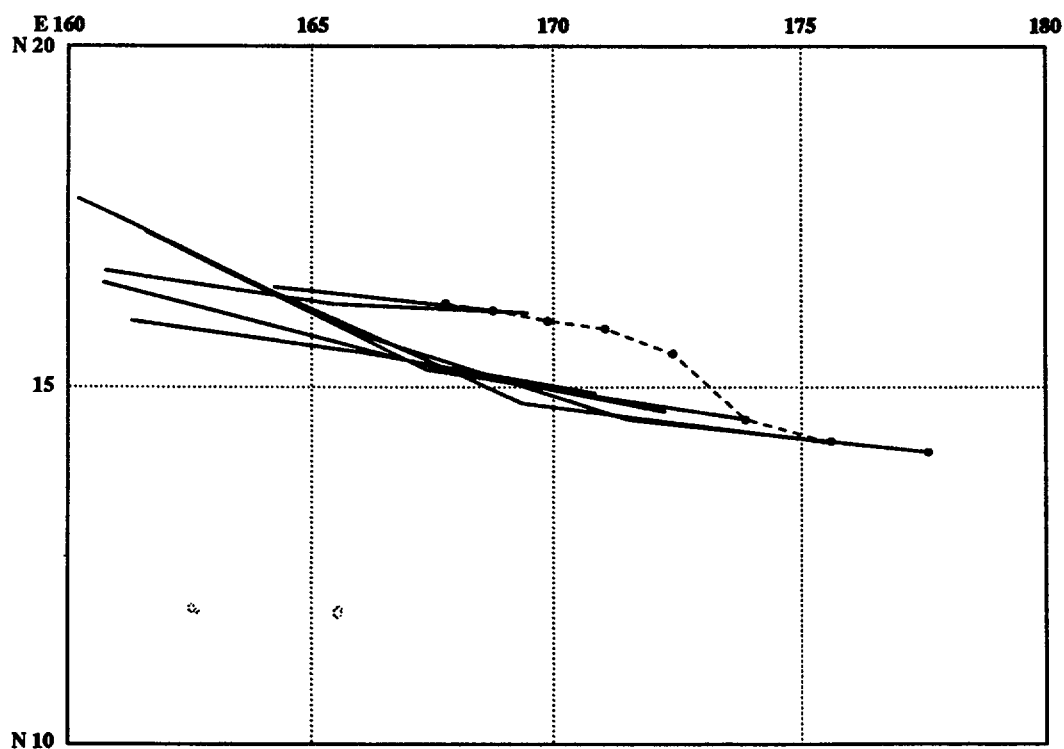
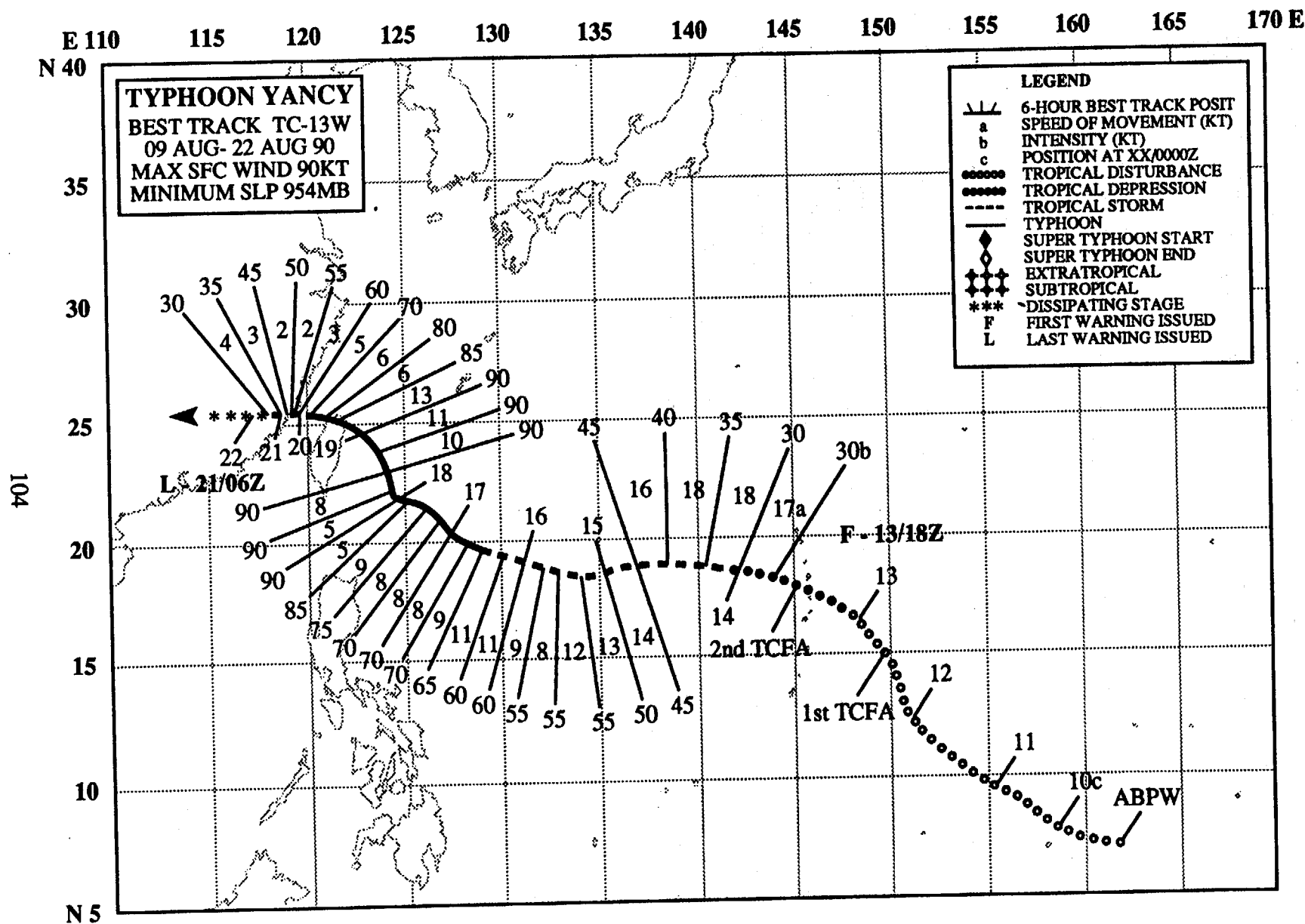


Figure 3-01C-3. Summary of JTWC forecasts (solid lines) for Aka is superimposed on the final best track (dashed line).



TYPHOON YANCY (13W)

I. HIGHLIGHTS

Yancy, JTWC's best forecast tropical cyclone of the year, was one of a series of August storms that generated in the monsoon trough. Although the track was generally toward the northwest, it contained several interesting features, including interaction with a strengthening subtropical ridge, the effects of a passing mid-latitude shortwave trough and land interaction with the mountainous terrain of Taiwan.

II. CHRONOLOGY OF EVENTS

- 090600Z - First mentioned on the Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1008 mb.
- 112100Z - First Tropical Cyclone Formation Alert based on an increase in central convection, more pronounced upper-level outflow and surface pressure decreases at several nearby land stations.
- 121400Z - Second Tropical Cyclone Formation Alert based on a northward shift of a consolidating low-level center and continued drops in surface pressure at several nearby land stations.
- 131400Z - Third Tropical Cyclone Formation Alert based on a continued increase in organization, deep central convection and an approaching surge in southwest monsoon flow.
- 131800Z - First warning due to increased consolidation of central convection and improvements in the upper-level outflow.
- 141200Z - Upgrade to tropical storm prompted by increased convective curvature, consolidation of the cyclonic center and the first intensity estimate of CI 2.5.
- 161200Z - Upgrade to typhoon prompted by a decrease in vertical wind shear, improved organization in the deep central convection, improved upper-level outflow and intensity estimates of CI 4.0.
- 180000Z - Peak intensity - 90 kt (46 m/sec) - based on intensity estimate of CI 5.0.
- 200000Z - Downgraded to tropical storm based on radar reports, synoptic reports and satellite imagery which indicated significant weakening due to land interaction as the system crossed Taiwan.
- 210600Z - Downgraded to tropical depression due to the effects of land interaction and increased vertical wind shear.
- 210600Z - Final warning - dissipated - based on a combination of land interaction and increased vertical wind shear as the system moved into mainland China.

III. TRACK AND MOTION

The LLCC which developed into Yancy generated on the eastern side of a broad monsoon depression. A series of vortices persisted at low latitudes for four days before consolidating into Yancy. In its formative stages, Yancy moved erratically as mesoscale convective elements developed, decayed, and were replaced by new elements. The resulting large monsoon depression moved generally westward at 8 to 10 kt (15 to 20 km/hr) until 13 August. A 48-hour period of rapid westward movement followed as Yancy moved into an area dominated by a strengthening subtropical ridge to the north (Figure 3-13-1). This westward track continued until a mid-latitude shortwave trough moving off the coast of China weakened the subtropical ridge over the East China Sea (Figure 3-13-2), resulting in an 18-hour period of north-northwestward movement. The system resumed its westward track across Taiwan as the subtropical ridge reestablished itself. Yancy executed a mesoscale trochoidal oscillation (wobble) about a smoothed track as it moved past Taiwan as depicted by radar position reports from Hualein (WMO 46699), Taiwan in Figure 3-13-3.

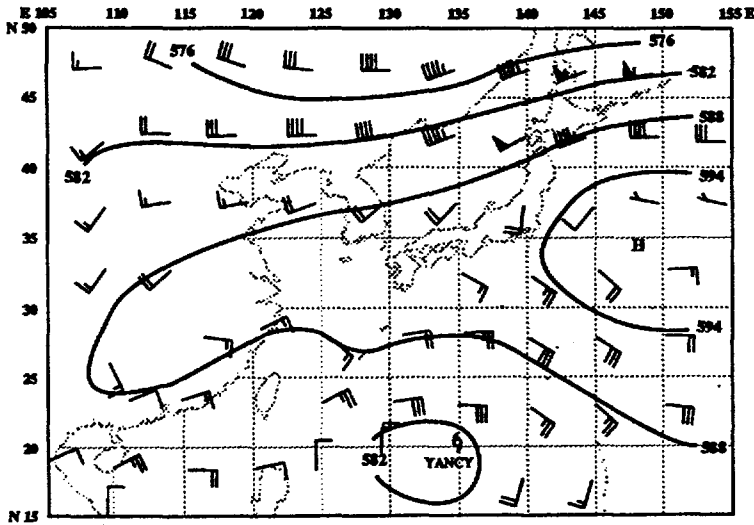


Figure 3-13-1. 500-mb NOGAPS analysis from 150000Z August, showing the strengthening of the mid-level ridge north of Yancy which resulted in the westward track.

Figure 3-13-2. 500-mb NOGAPS analysis from 170000Z August, showing a passing shortwave trough weakening the mid-level subtropical ridge, which resulted in a jog in the track to the north-northwestward.

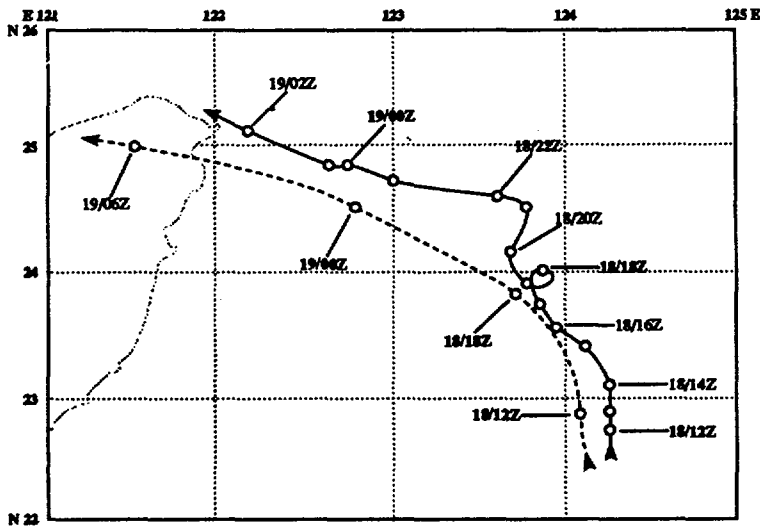
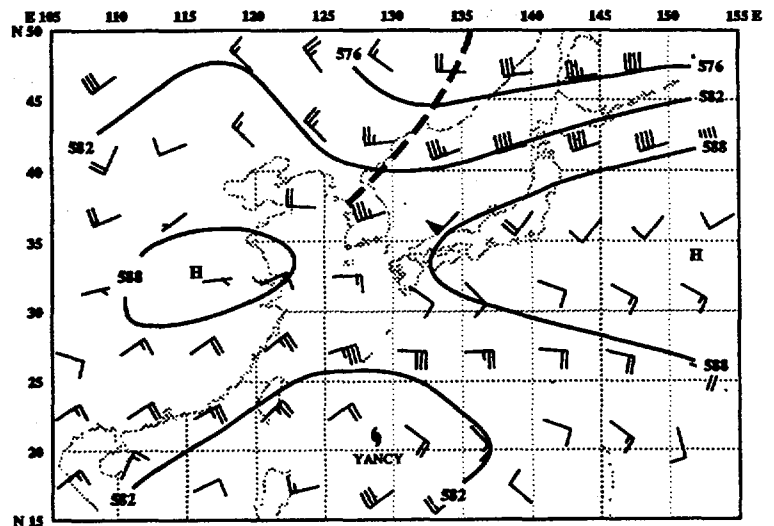


Figure 3-13-3. Plots of hourly radar positions from Hualein (WMO 46699), Taiwan compared to the smoothed best track (dashed line) show Yancy's wobble.

IV. INTENSITY

At 150000Z August, Yancy had a distinct low-level circulation center on the poleward side of the monsoon cloud mass (Figure 3-13-4). The poleward dislocation was attributed to strong upper-level flow from the north and east that apparently inhibited rapid development. A strongly divergent flow became established over the system on August 17, with outflow branches into the equatorial easterlies and into the major TUTT cell to the east-northeast (Figure 3-13-5). Fairly slow deepening to maximum intensity followed and Yancy developed an eye on 18 August (Figure 3-13-6). Weakening and decay were directly attributable to the close approach to the Taiwan mountains, followed by landfall on mainland China.

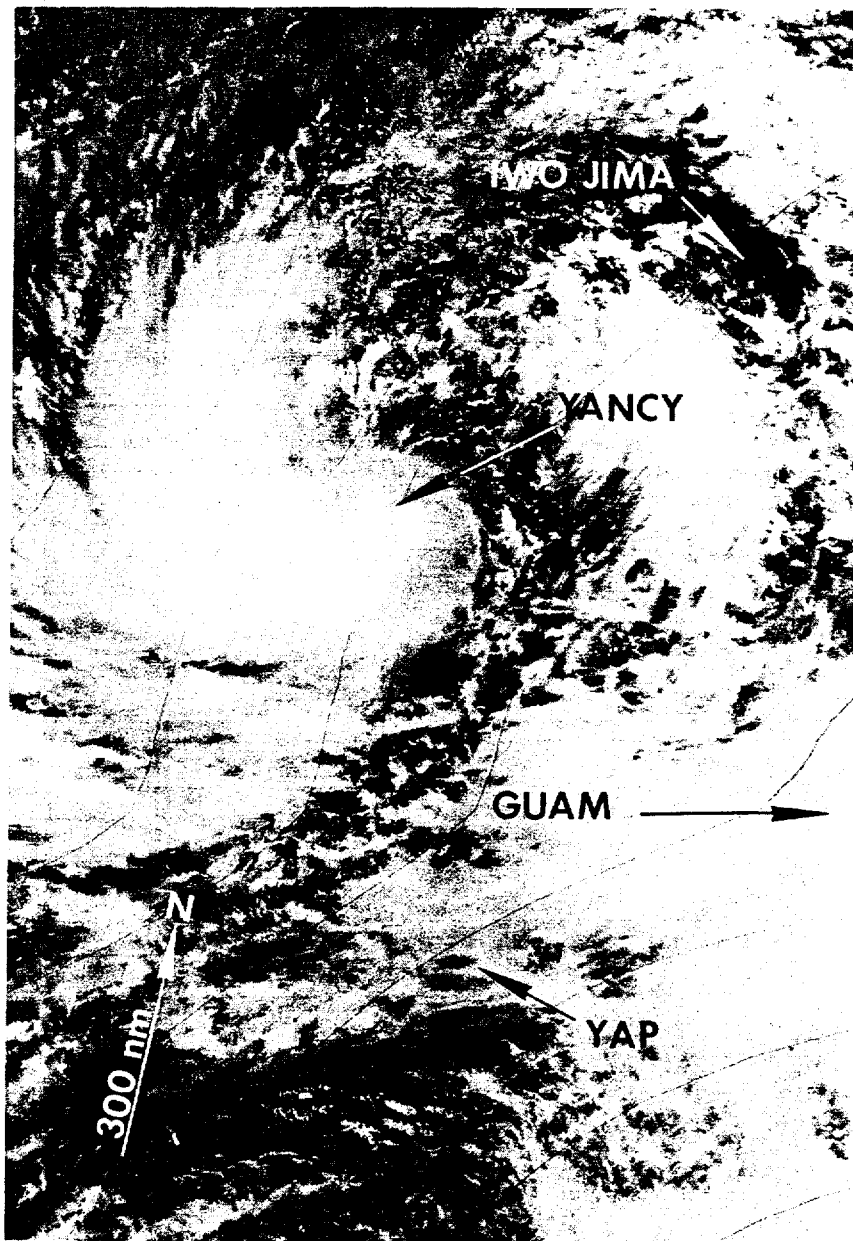


Figure 3-13-4. Tropical Storm Yancy (13W) as it separates from the convection associated with the monsoon trough. Note the area of strong low-level convergence southeast of the system. This area was associated with a strong surge in the monsoon flow from which Yancy separated (150504Z August NOAA visual imagery).

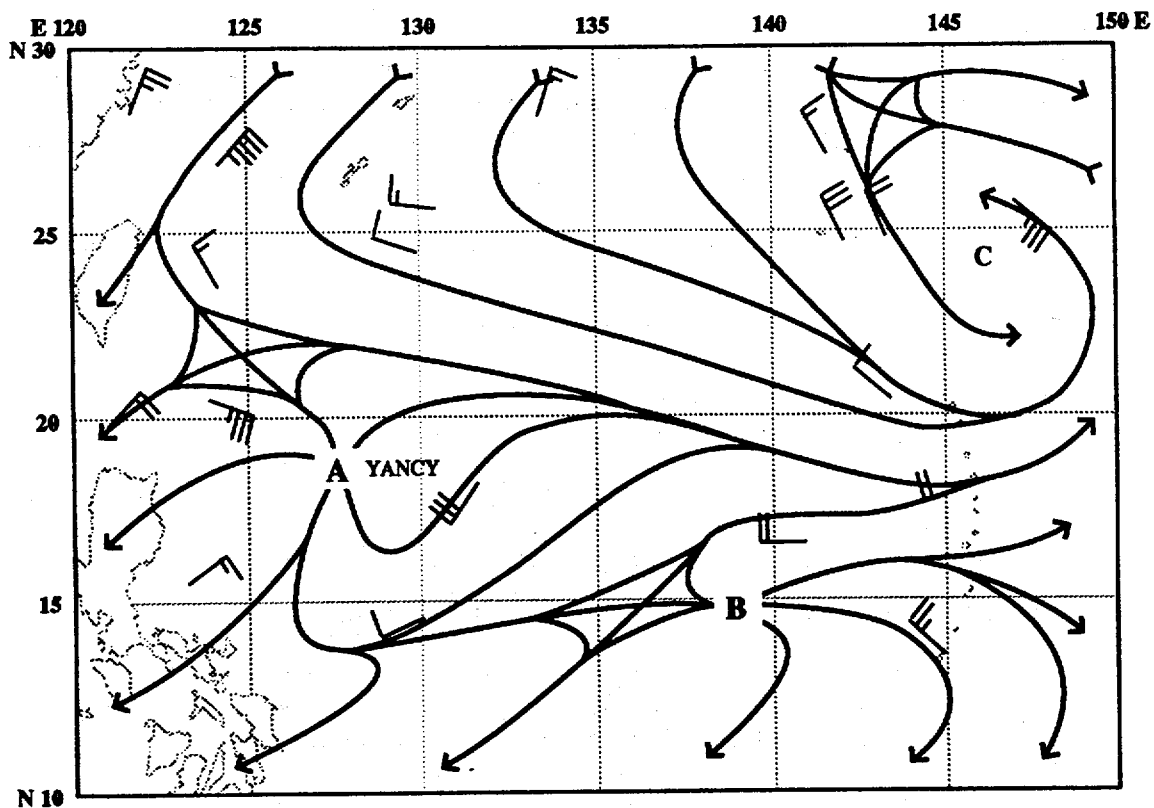


Figure 3-13-5. The 170000z August 200-mb analysis, with Yancy at point A, showing the upper-level outflow channel to the southwest and eastward into the large TUTT cell at point C. The outdraft at point B is over deep convection associated with the formation of Zola (14W).

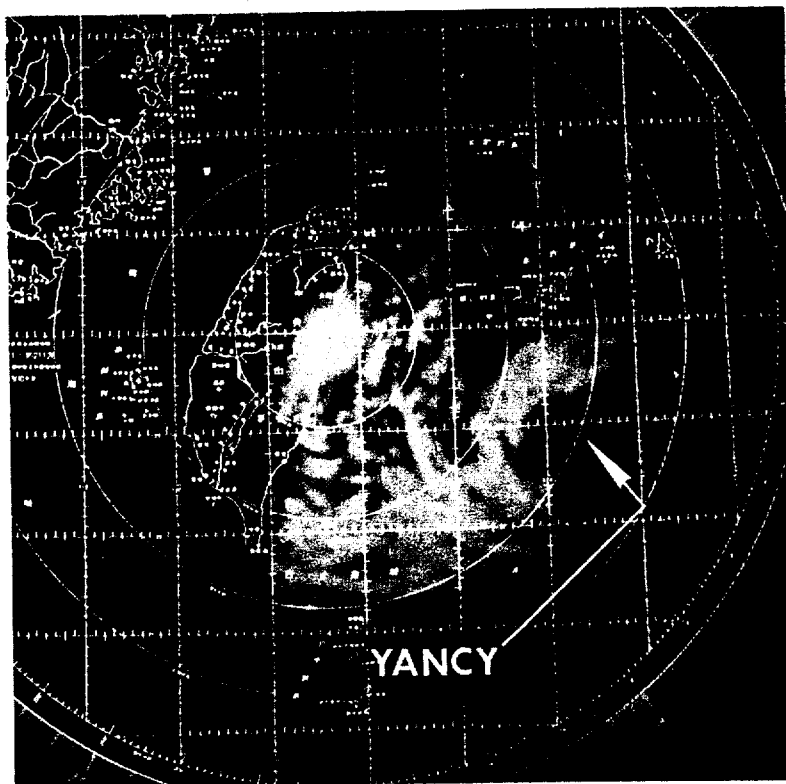


Figure 3-13-6. Yancy's eye appears at the edge of the radar scope at Hualein (WMO 46699), Taiwan (181200Z August photo courtesy of the Central Weather Bureau, Taipei, Taiwan).

V. FORECASTING PERFORMANCE

Although Yancy followed what might appear to be a simple northwestward track, it proved to be a difficult forecast scenario. At two key points during the forecast cycle, the forecast aids and NOGAPS prognostic charts were not in agreement as to the final storm track. As the tropical storm passed through the northern Mariana Islands, the statistical forecast aids indicated that the storm would recurve, while the dynamic forecast aids and the NOGAPS model indicated the system would move westward in response to a building mid-level ridge to the north. The second difficult forecast decision came as the storm approached Taiwan. The statistical forecast aids, the ECMWF model, the NMC model and the Japanese model all called for the system to recurve in response to a passing mid-latitude shortwave trough. The dynamic forecast aids and the NOGAPS model forecast the system to track westward toward a col in the mid-level subtropical ridge over eastern China. As Figure 3-13-7 indicates, JTWC chose the correct forecast at each of these key forecast points. Yancy proved to be JTWC's best forecast storm of the year, with errors of 97nm (180km) at 24 hours, 98nm (182km) at 48 hours and 108nm (200km) at 72 hours.

VI. IMPACT

Yancy passed through the northern Philippine Sea, triggering a deep monsoon surge that resulted in heavy rains and flooding on northern Luzon, leaving at least six people dead and more than 60,000 people fleeing to evacuation centers. Yancy's next impact was felt on Taiwan as it brought heavy winds and torrential rains to the northern half of the island before moving into mainland China. There, the death toll climbed to 216 people, with an additional 59 reported missing and an estimated economic loss of approximately 170 million dollars.

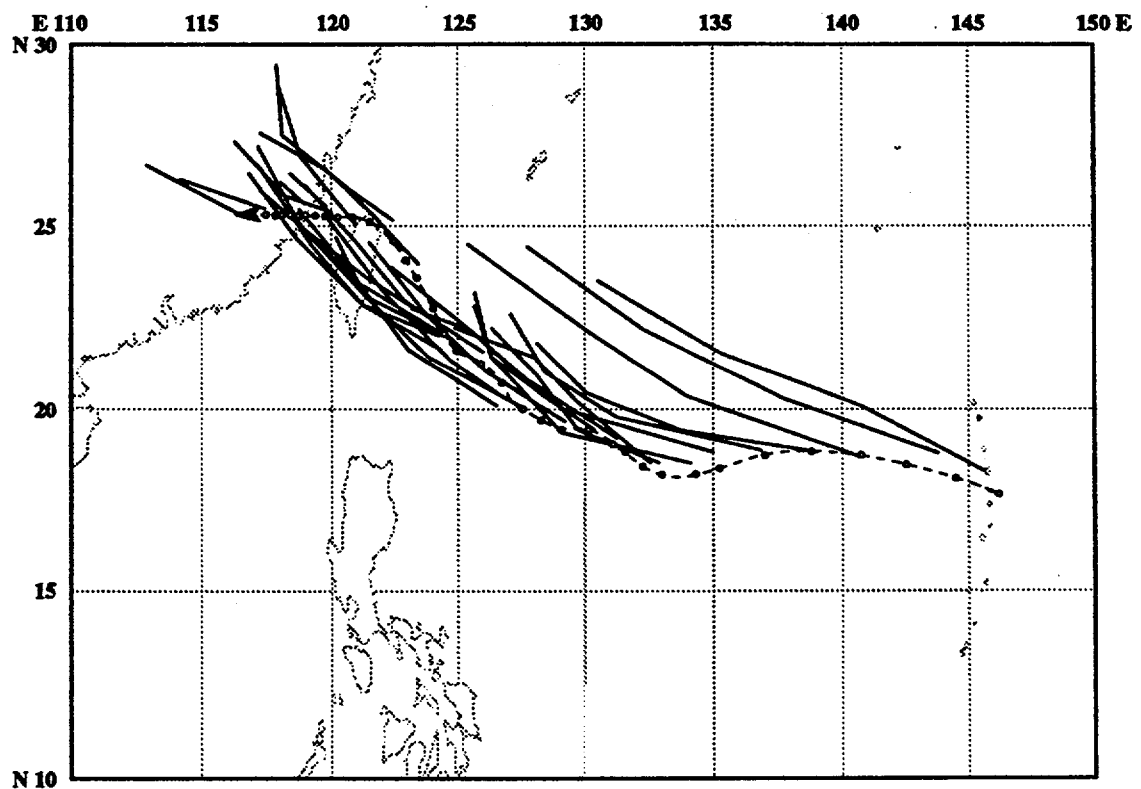
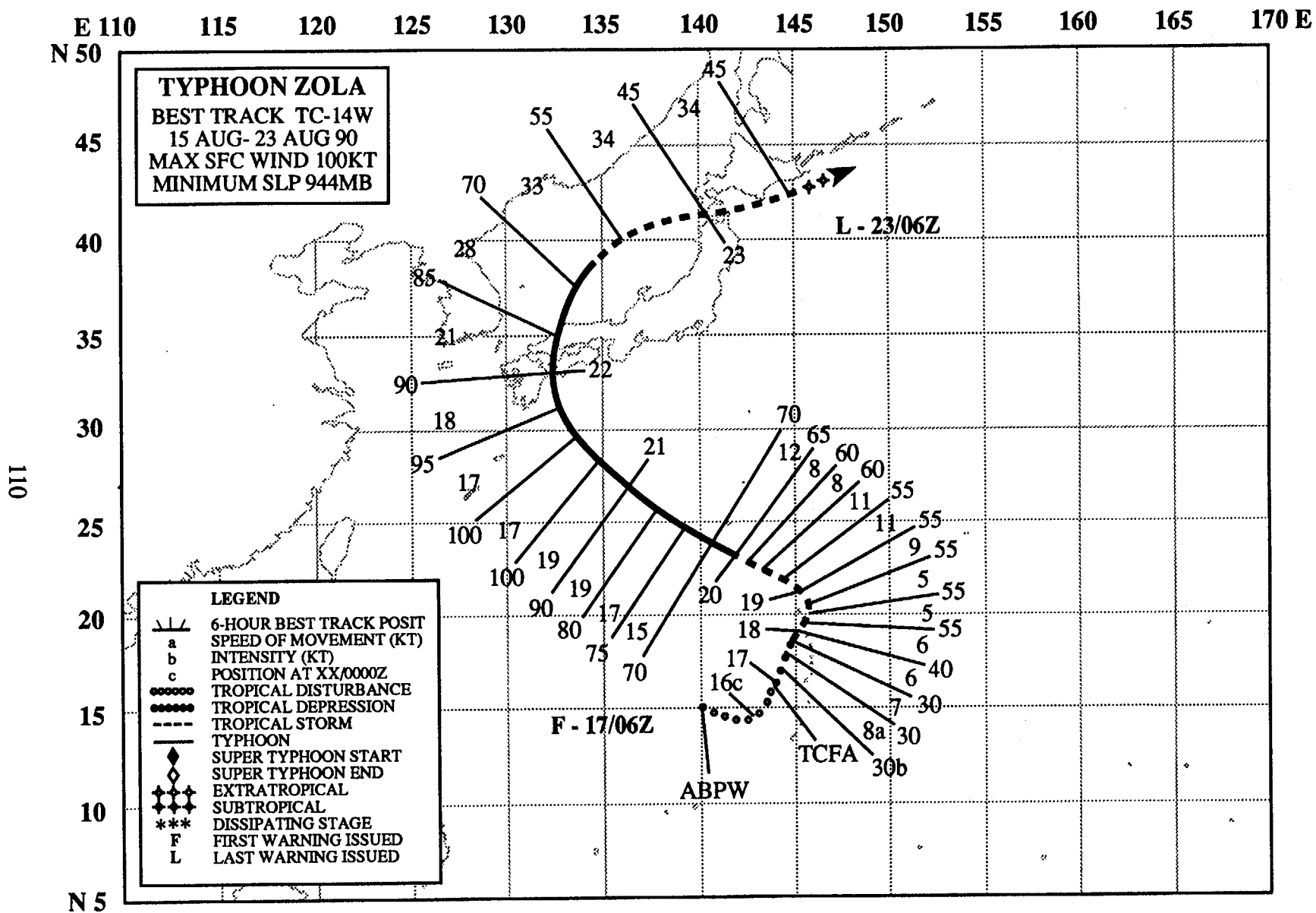


Figure 3-13-7. Summary of JTWC forecasts (solid lines) for Yancy is superimposed on the final best track (dashed line).



TYPHOON ZOLA (14W)

I. HIGHLIGHTS

In the wake of Typhoon Yancy (13W), a surge in the southwesterly monsoon flow developed and Zola formed west of Guam in the monsoon trough. The depression initially tracked northeastward with the movement of the monsoon surge and slowly intensified. Then, Zola broke away from the monsoon trough and intensified into a typhoon. The typhoon recurved over western Honshu into the Sea of Japan and accelerated to the east-northeastward.

II. CHRONOLOGY OF EVENTS

- 150600Z - First mentioned on the Significant Tropical Weather Advisory as a persistent area of convection which extended eastward from Tropical Storm Yancy (13W). The estimated minimum sea-level pressure was 1000 mb.
- 162200Z - Tropical Cyclone Formation Alert based on a transient band of convection wrapping around the low-level circulation center.
- 170600Z - First Tropical Depression Warning prompted by the persistence of deep convection associated with a surge in the monsoonal flow just to the south of the circulation center.
- 180600Z - Upgraded to a tropical storm after receipt of a ship report of 55 kt (27 m/sec) and a 998 mb sea-level pressure indicating increased periphery winds and a tightened pressure gradient to the south and east.
- 200000Z - Upgraded to typhoon based on the appearance of a 25 nm (45 km) diameter ragged eye and the first CI 4.0.
- 210600Z - Peak intensity - 100 kt (51 m/sec) - followed an increase in organization, outflow, and intensity estimate of CI 5.5.
- 221800Z - Downgraded to tropical storm due to increased vertical wind shear and the start of extratropical transition.
- 230600Z - Final warning - extratropical - issued as Zola transitioned into a mid-latitude low due to strong vertical shear associated with the mid-latitude westerlies.

III. TRACK AND MOTION

After briefly tracking eastward during its formative stages, Typhoon Zola tracked north-northeastward just west of the Northern Mariana Islands along the western side of the subtropical high to the northeast. The tropical cyclone continued to track towards the north-northeast for the next three days as a short wave trough tracked slowly eastward, north of the system. Once the shortwave passed, the subtropical high built westward and combined with a dynamic high that moved off the coast of China to Japan reestablishing the subtropical ridge over Japan (Figure 3-14-1 through Figure 3-14-3). As this happened, Zola turned sharply and started tracking northwestward around the ridge. It recurved over southern Honshu and accelerated northeastward into the Sea of Japan.

IV. INTENSITY

Starting in the monsoon trough, Zola spun up as a result of a surge in the southwest monsoon associated with Typhoon Yancy (13W). For the first several days, the tropical cyclone developed slowly, remaining a tropical depression, primarily due to vertical wind shear. As Zola continued to track northeastward, the system intensified to 55 kt (28 m/sec) as it moved into an area of upper-level divergence southeast of a Tropical Upper Tropospheric Trough (TUTT) low. However, it remained a tropical storm until breaking away from the monsoon trough (Figure 3-14-4). Once separated from the monsoon trough, Zola intensified as it developed an outflow channel to the south. Intensification continued due to enhanced outflow to the north associated with a TUTT low to the northwest. The tropical cyclone reached a peak intensity of 100 kt (51 m/sec) on 21 August (Figure 3-14-5). At

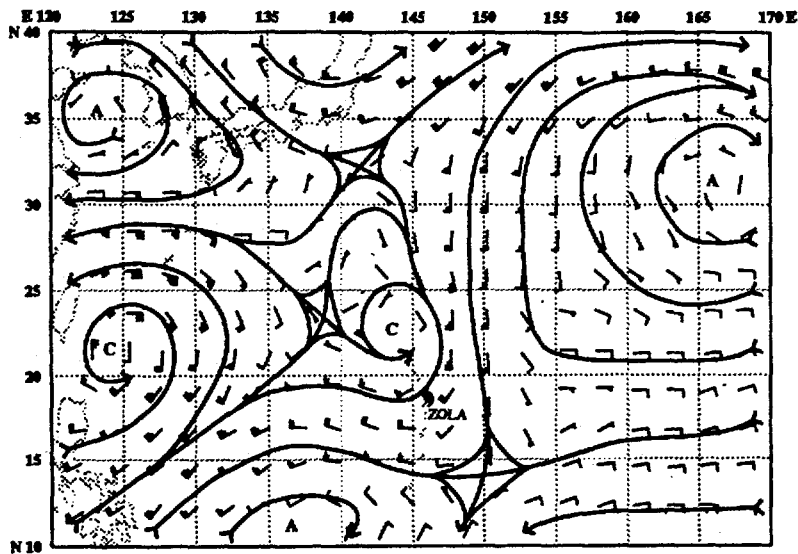


Figure 3-14-1. The 180000Z August NOGAPS deep layer mean analysis shows Zola's surface position, the subtropical high to the northeast, the mid-latitude trough to the north-northwest, and the dynamic high over the Yellow Sea. At this time, Zola was moving north-northeastward along the southwest side of the subtropical high

Figure 3-14-2. The 190000Z August NOGAPS deep layer mean circulation analysis depicts the mid-latitude trough in a position north-northeast of Zola. The dynamic high has moved to a location over southern Japan.

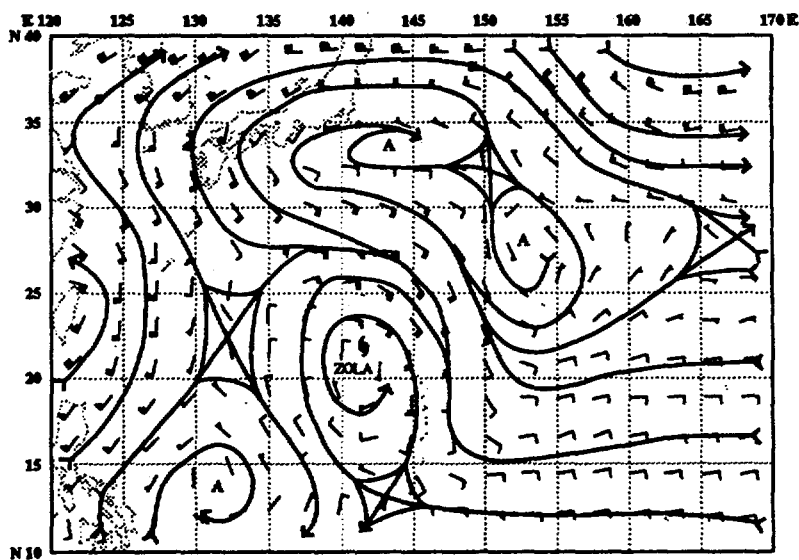
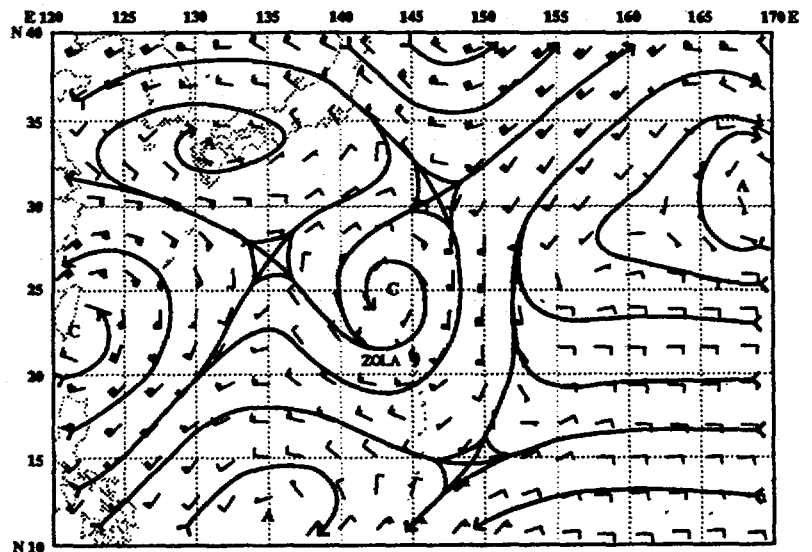


Figure 3-14-3. The 200000Z August NOGAPS deep layer mean circulation analysis shows the dynamic high and the subtropical high merging north of Zola.

211800Z, the typhoon started to weaken due to increasing vertical wind shear associated with the mid-latitude westerlies and land interaction with Japan. After recurving Zola quickly transitioned into an extratropical system.

V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-14-6. The initial warnings did not forecast Zola's sharp turn and track to the northwest. The NOGAPS prognostic series did not indicate a dynamic high moving off the coast of eastern Asia and combining with the subtropical high, reestablishing the ridge further to the west. JTWC also forecast Zola to recurve further to the east. The recurvature farther to the west may have been caused by the advection of warm, moist air from the tropics which strengthened the subtropical high to the tropical cyclone's northeast.

VI. IMPACT

No information received.

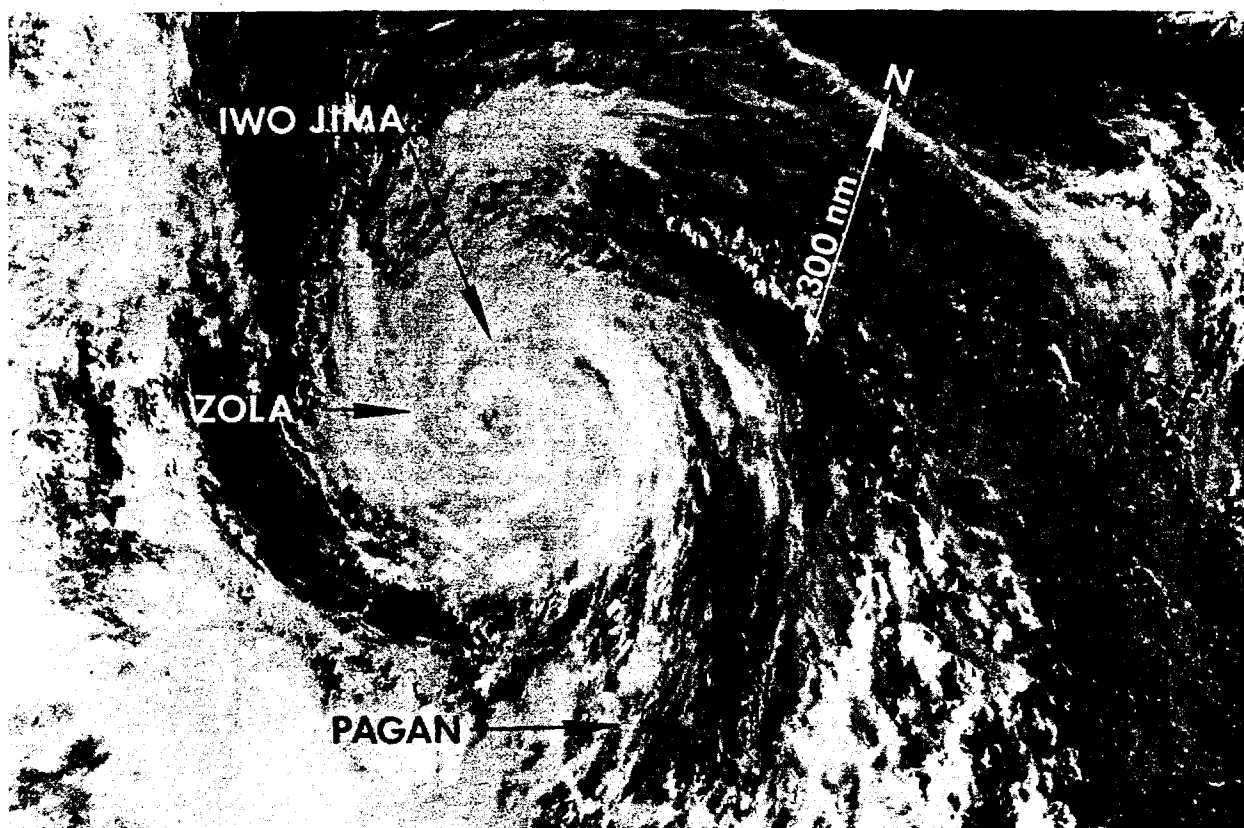


Figure 3-14-4. Zola just after breaking away from the monsoon trough. A distinct separation can be seen between Zola and the cloud mass to its south (192322Z August DMSP visual imagery).

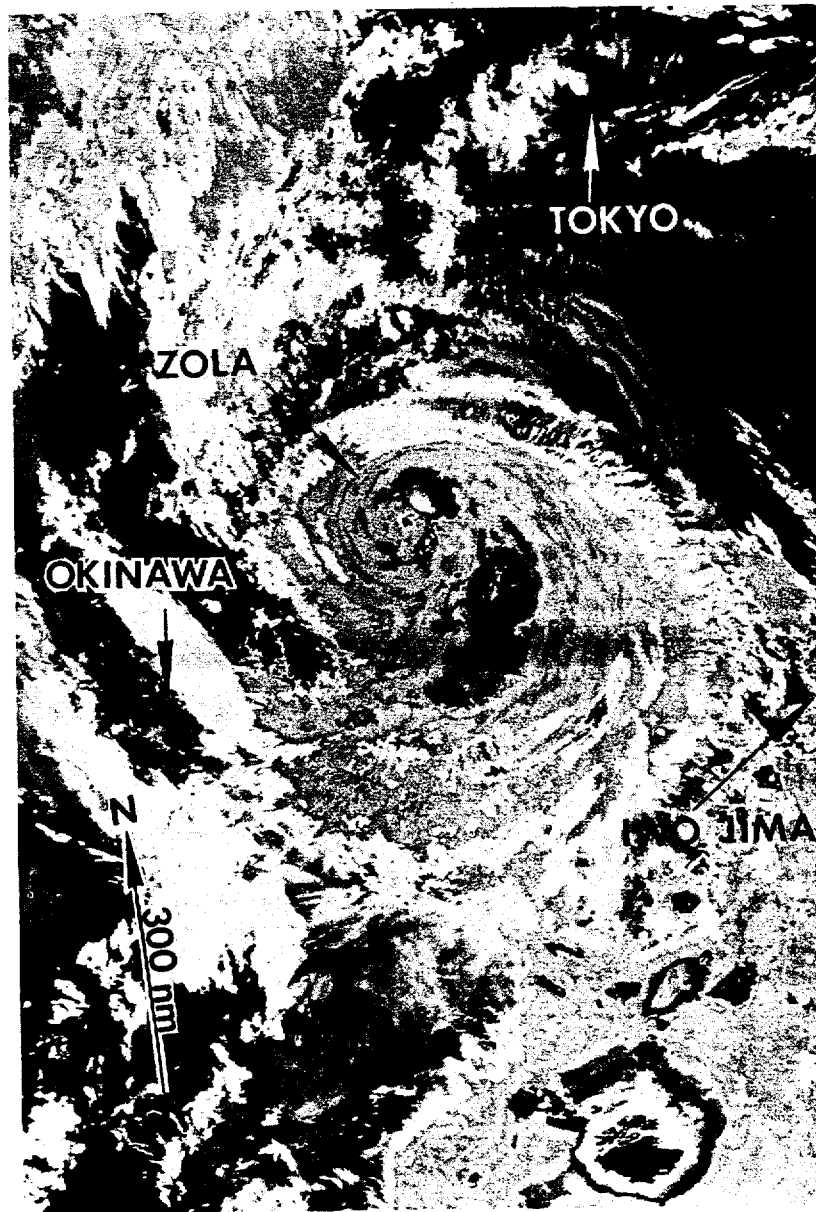


Figure 3-14-5. Zola, with a small eye and at maximum intensity, is moving northwestward towards southern Japan (210933Z August DMSP enhanced infrared imagery).

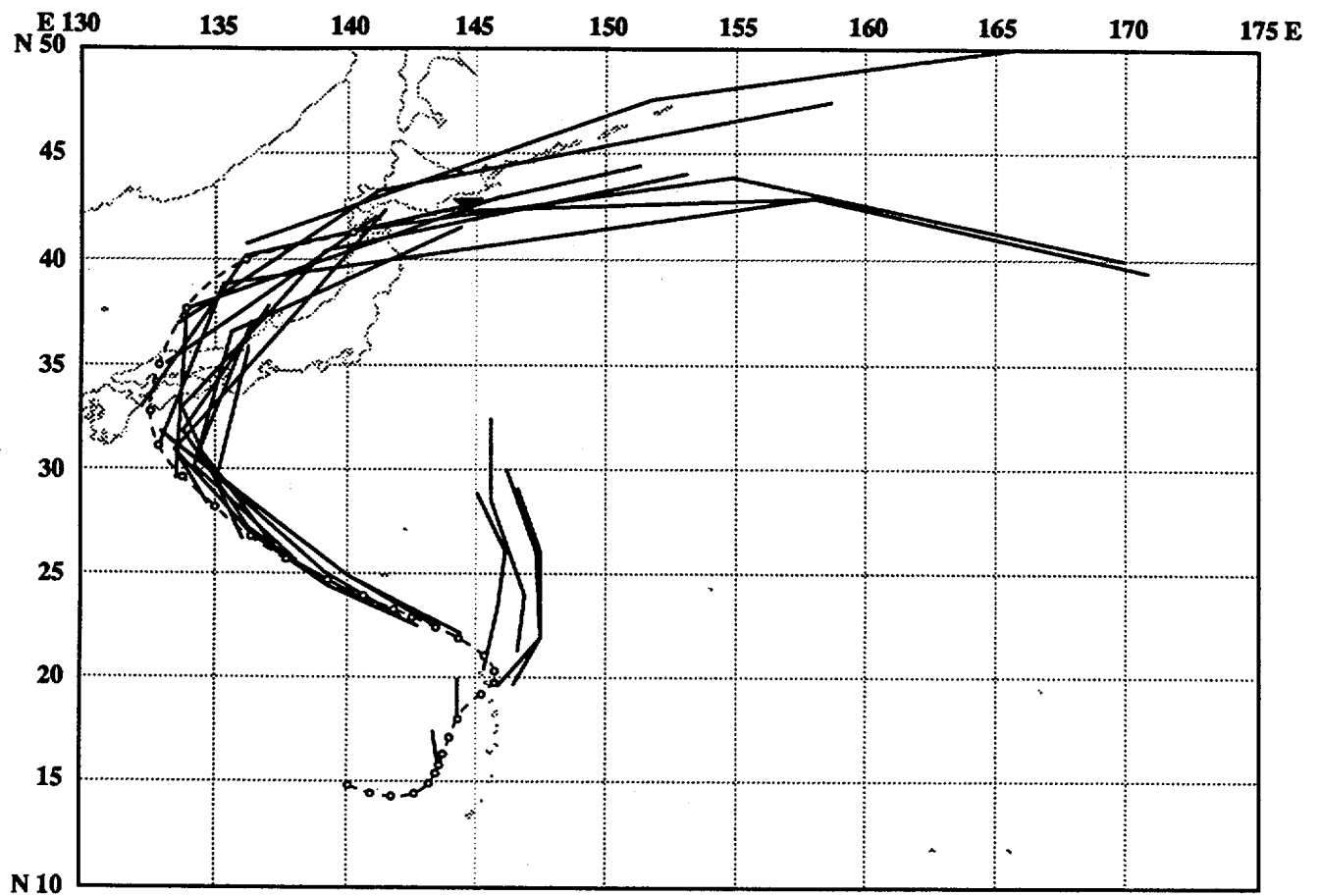
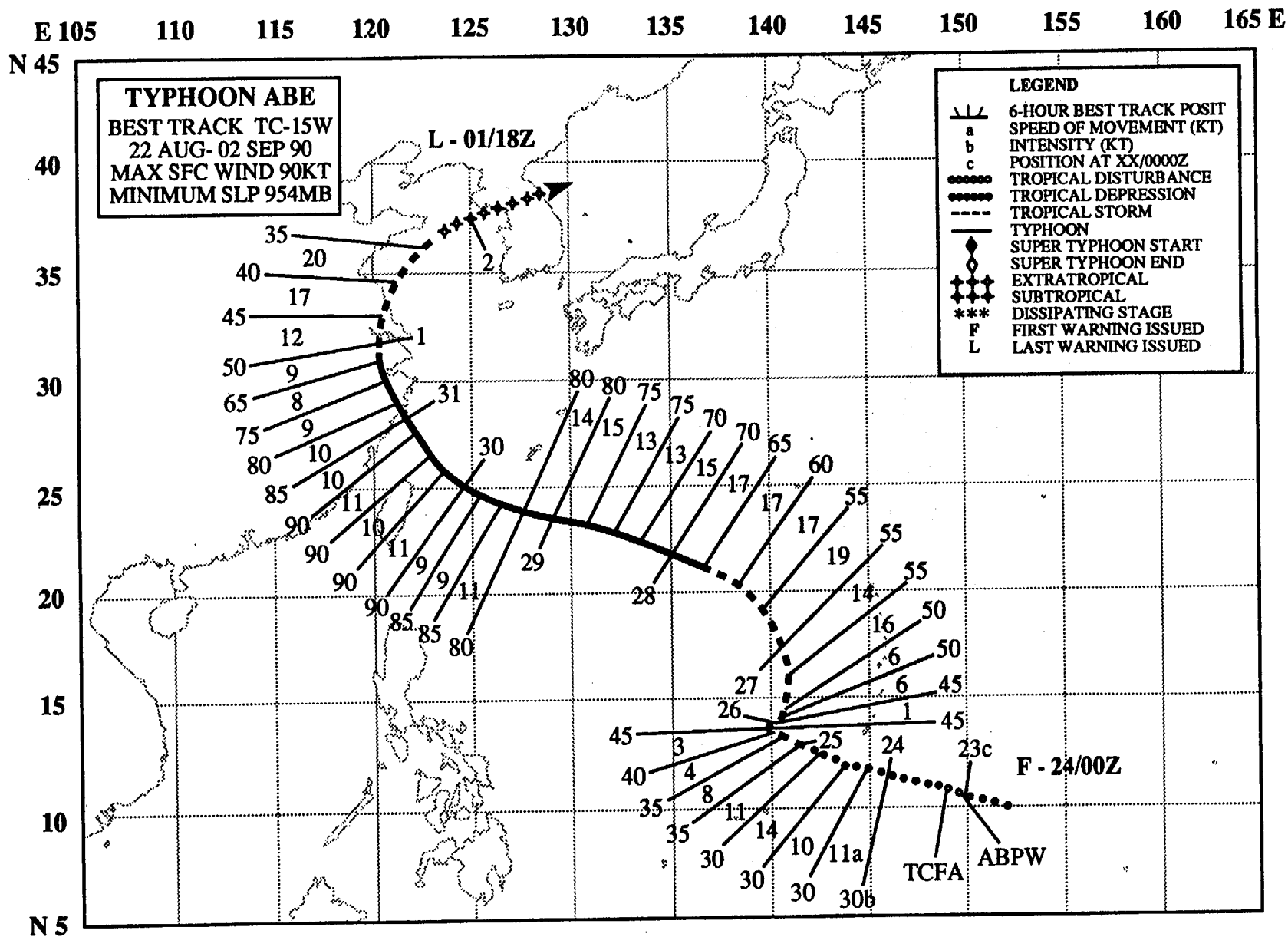


Figure 3-14-6. Summary of JTWC forecasts (solid lines) for Zola is superimposed on the final best track (dashed line).



TYPHOON ABE (15W)

I. HIGHLIGHTS

Typhoon Abe, the fourth of five tropical cyclones in August, caused extensive damage from the Republic of the Philippines through eastern China during its nine day life. Abe was also noteworthy as a classic example of the erratic motion and rapid reorganization that can occur in association with an intense monsoon surge.

II. CHRONOLOGY OF EVENTS

- 230100Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection at the end of an active monsoon trough. Minimum sea-level pressure estimated to be 1007 mb.
- 230600Z - First Tropical Cyclone Formation Alert based on increased convection, organization, and outflow aloft.
- 240000Z - First warning issued due to continued development.
- 250000Z - Upgraded to a tropical storm based on increased central convection.
- 271200Z - Upgraded to a typhoon after detection of a ragged eye.
- 300000Z - Peak intensity - 90 knots (46 m/sec) - based on intensity estimate of CI 5.0
- 311200Z - Downgraded to tropical storm as convection decreased due to land interaction.
- 011800Z - Final warning issued due to extratropical transition.

III. TRACK AND MOTION

From its initial mention on the Significant Tropical Weather Advisory until 250000Z, Abe tracked steadily west-northwestward under a well-developed subtropical ridge. By 251200Z, an intense, deep surge in the monsoon westerlies began to develop south of Abe, arresting its westward motion. The enhanced convection associated with the surge (Figures 3-15-1a, 3-15-1b and 3-15-1c) initially formed east of Abe's convective cloud mass and grew as it wrapped around to the north. Eventually, Abe's circulation center reorganized to the north, between the competing convective masses. The intensity and horizontal extent of the monsoon surge is illustrated by the time sequence of gradient level winds recorded at the National Weather Service Observatory at Taguac, Guam (WMO 91217) and shown in Figure 3-15-2. During the timeframe of the figure, Abe was located between 270 and 540 nm (500 to 1000 km) from Guam. Following the monsoon surge event that pushed the system on a brief eastward then northward track, Abe resumed a west-northwestward track along the periphery of the subtropical ridge. The typhoon eventually recurved through a weakness in the subtropical ridge associated with a passing short-wave trough. The recurvature track took Abe along the coasts of the Zhejtang and Jiangsu Provinces of China, into the Yellow Sea, and across the middle portion of South Korea.

IV. INTENSITY

From the initial warning at 240000Z until 270600Z, Abe intensified by only 25 kt (13 m/sec) due to the disruptive shearing effects of the monsoon surge. The subsequent three days of intensification to its peak of 90 kt (46 m/sec) at 300000Z was also slower than normal. The slow intensification may be attributed to some restriction of Abe's outflow into the tropical upper-level easterlies caused by the outflow of Typhoon Becky (16W). Any additional intensification that might have resulted from the eventual establishment of good outflow into the midlatitude westerlies at 310000Z was negated by the terrain effects as Abe approached China.

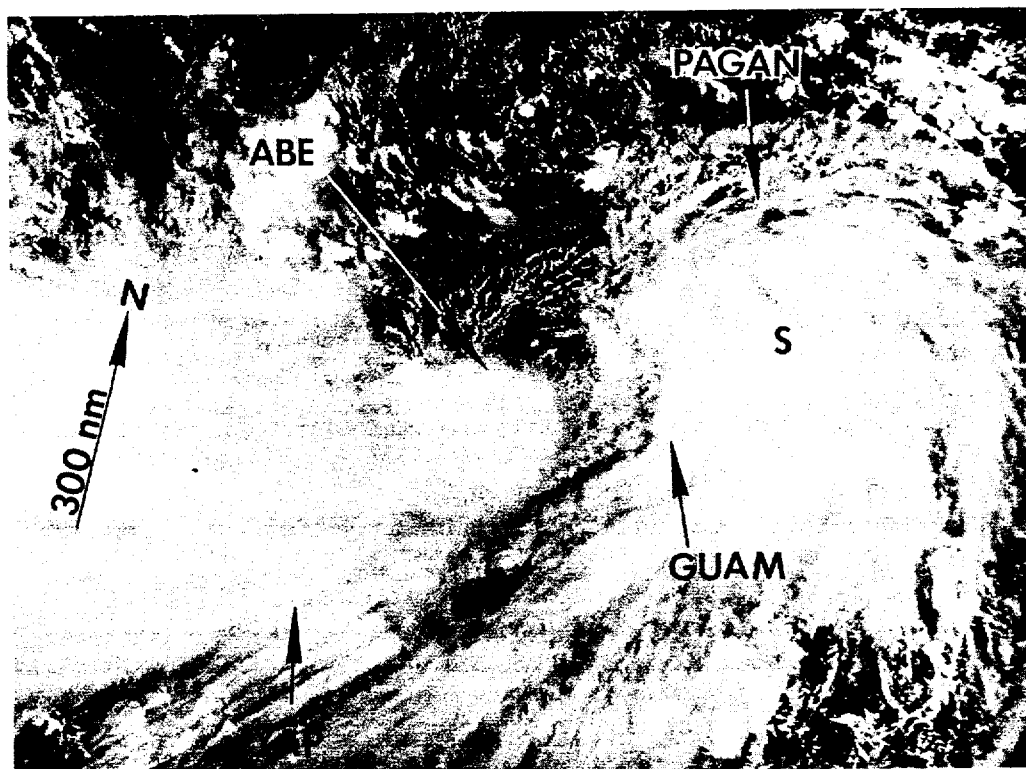


Figure 3-15-1a. The comma-shaped cloudiness (at Point S) to the northeast of Guam is associated with a monsoon surge that is wrapping around Abe's center (260441Z August NOAA visual imagery).

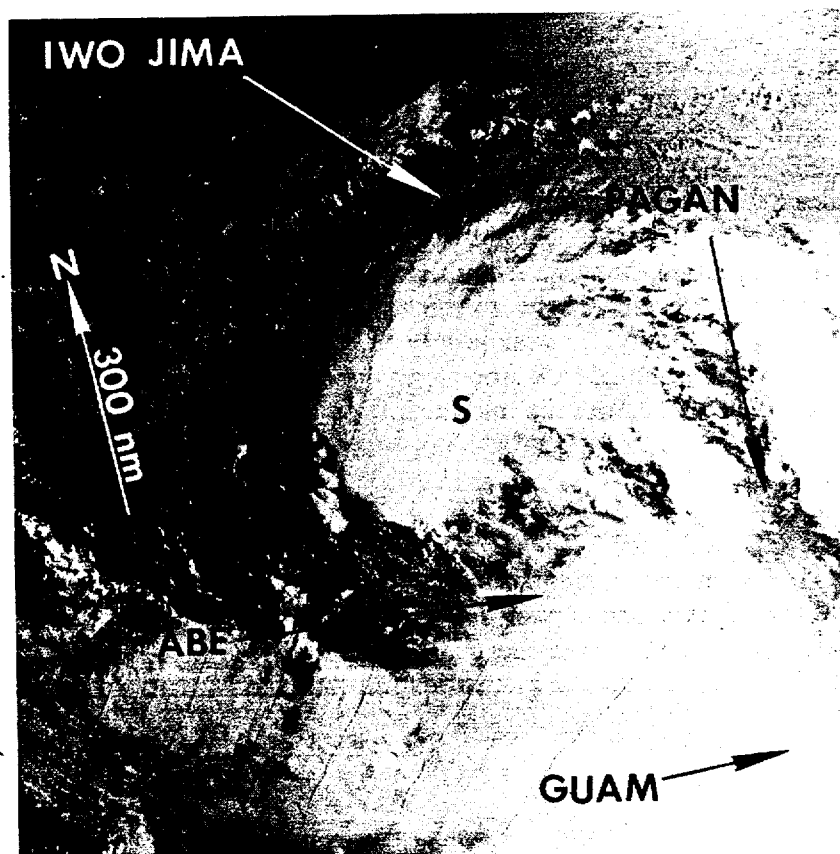


Figure 3-15-1b. The comma-shaped cloudiness (at Point S) has rotated counterclockwise around Abe's center during the past 18 hours, and is to the north (262238Z August NOAA visual imagery).



Figure 3-15-1c. The area of broken cloudiness (at Point S) which has rotated around to the west of Abel's center in the past 6 hours is associated with the monsoon surge mentioned in Figures 3-15-1a and 3-15-1b (270430Z August; NOAA visual imagery).

V. FORECASTING PERFORMANCE

As illustrated by Figure 3-15-3, the overall forecast performance of JTWC for Abe was quite good with the exception of the period when Abe made the sharp turn northward due to the monsoon surge-induced reorganization. JTWC has no objective guidance that can reliably forecast the onset of deep monsoon surges or the associated track changes that might be induced. Theoretically-based synoptic reasoning that can assist forecasters in subjectively anticipating either the onset of the monsoon surge or its effects is limited. The best tool for short-range forecast intelligence is meteorological satellite imagery.

VI. IMPACT

The impact from Abe was extensive. Monsoon rains from the surge feeding into Abe caused extensive flooding in Luzon, killing 12 people in Manila. Landslides from the heavy rains resulted in 32 deaths in the provinces of Benguet, Nueva Viscaya and Nueva Ecija to the north of Manila. According to the Red Cross, the death toll in the Philippines due to the combined effects of Abe and Becky (16W) was 85. Okinawa experienced winds as high as 60 kt (31 m/sec), and high surf conditions there swept one person out to sea. Flooding in Taiwan resulted in one death and six injuries, and landfall in China resulted in 51 deaths and 250 injuries near Shanghai.

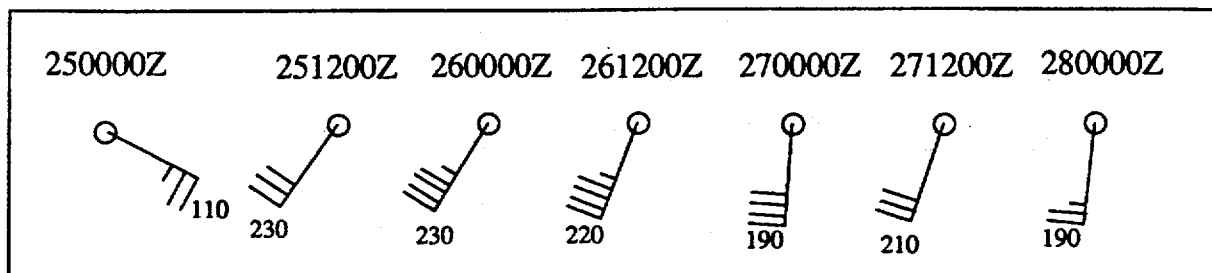


Figure 3-15-2. Gradient level winds recorded at Guam (WMO 91212) during monsoon surge associated with Abe.

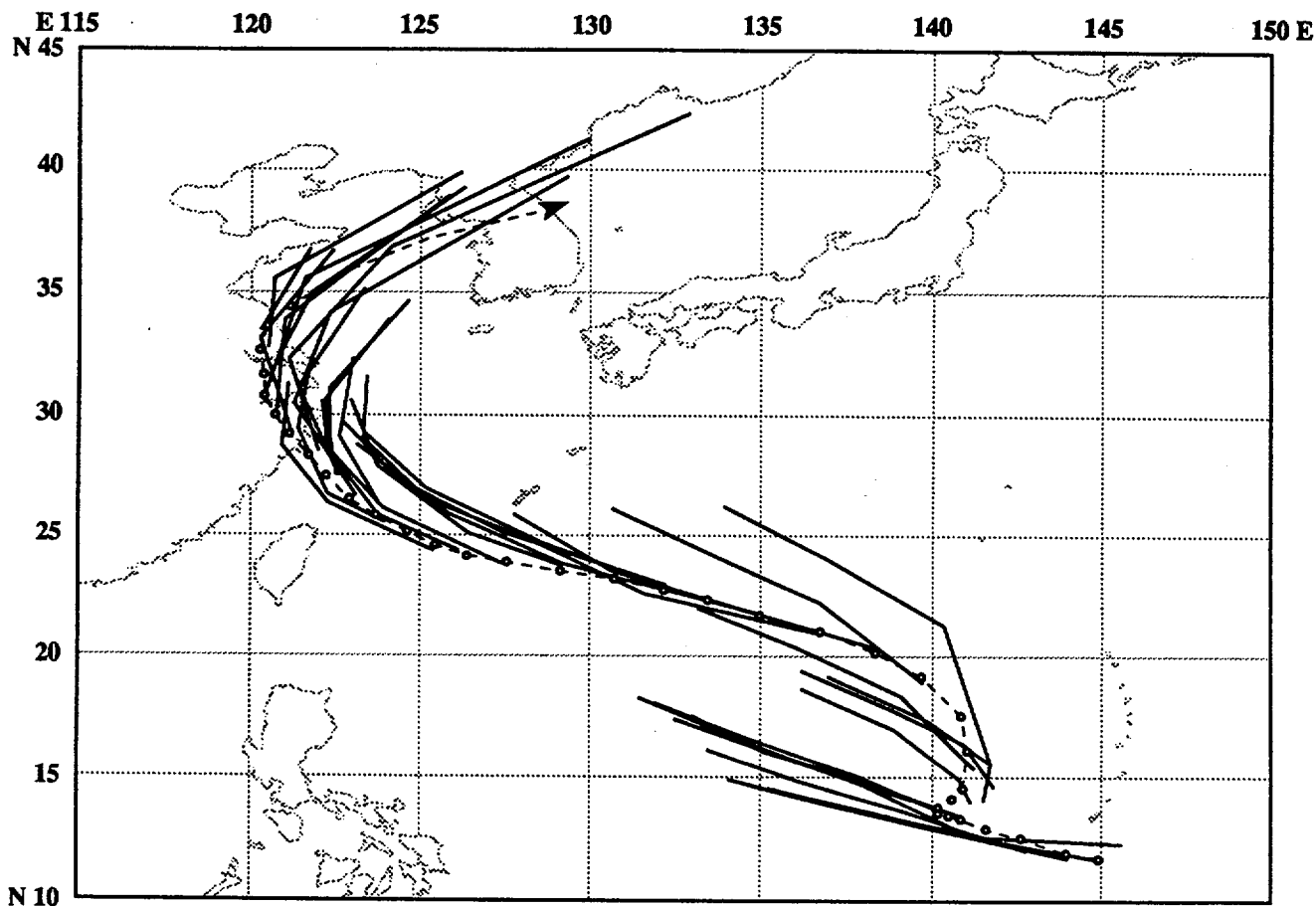
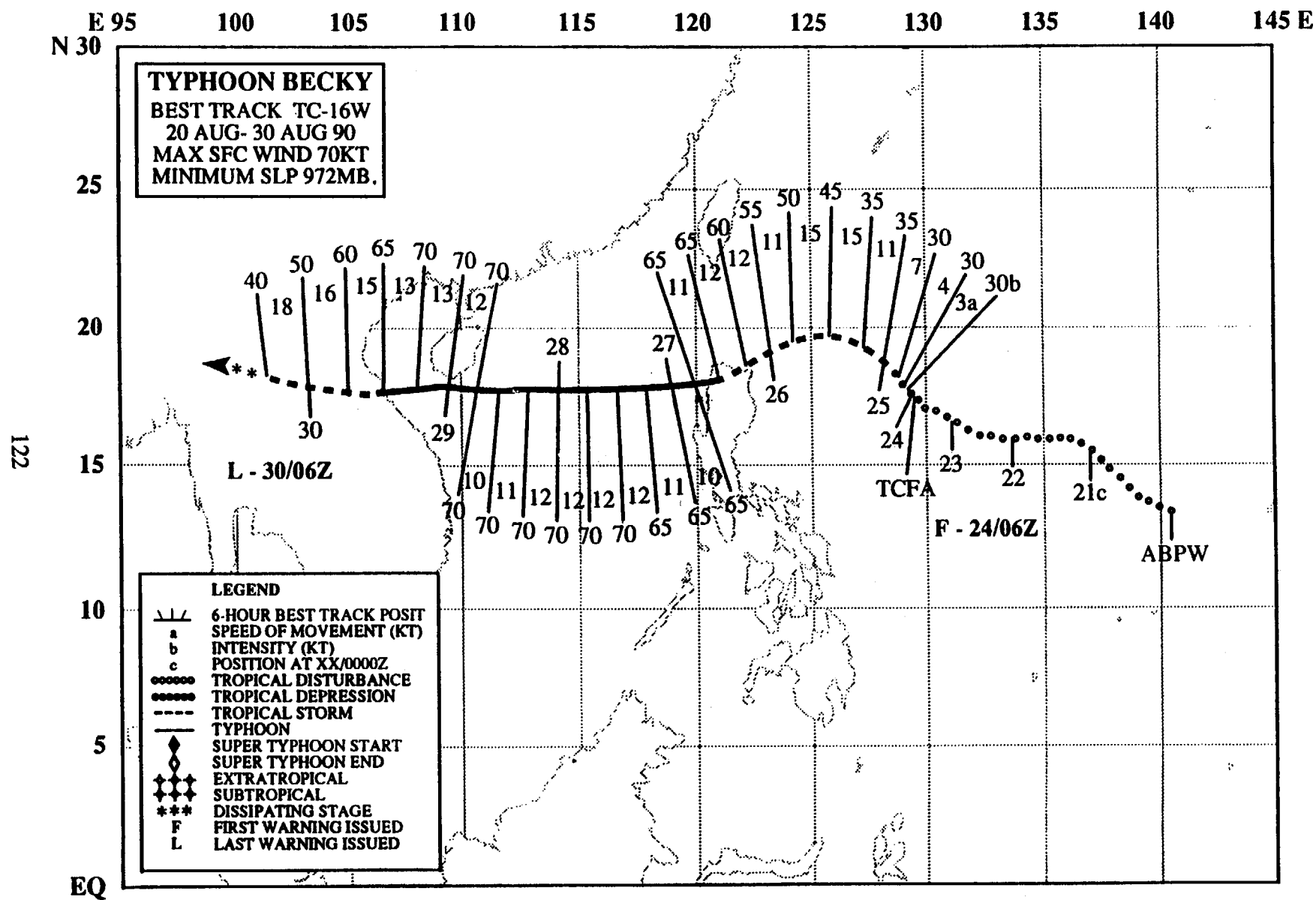


Figure 3-15-3. Summary of JTWC forecasts (solid lines) for Abe superimposed on the final best track (dashed line).



TYPHOON BECKY (16W)

I. HIGHLIGHTS

Becky, a midget typhoon and the eleventh typhoon of 1990, generated in the monsoon trough and tracked south of the subtropical ridge throughout its existence. After initially moving west-northwestward, the storm took a southwestward track across the northwestern tip of Luzon before heading westward across the South China Sea. Becky hit northern Luzon with typhoon-force winds and later slammed into northern Vietnam as a severe tropical storm.

II. CHRONOLOGY OF EVENTS

- 200600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with a minimum sea-level pressure of 1007 mb.
- 232200Z - Tropical Cyclone Formation Alert based on increased convective organization, a steady drop in sea-level pressure, and a corresponding increase in surface winds.
- 240600Z - First warning based on appearance of a well-developed low-level circulation center on the edge of the deep convection.
- 250000Z - Upgrade to tropical storm based on tighter spiral band curvature and first intensity estimate of CI 2.5.
- 261200Z - Upgraded to typhoon after appearance of a 10 nm (19 km) diameter eye and the first CI 4.0 satellite signature.
- 271200Z - Peak intensity - 70 kt (36 m/sec) - accompanied the reappearance of a small 8 nm (15 km) diameter ragged eye.
- 291800Z - Downgraded to tropical storm intensity after the central dense overcast degenerated into a poorly defined spiral cloud band.
- 300600Z - Final warning - dissipated over land.

III. TRACK AND MOTION

After forming 275 nm (510 km) west of Guam, Becky tracked slowly west-northwestward under the influence of the subtropical ridge (Figure 3-16-1) that was building westward across the wake of Typhoon Zola (14W) which was moving through the Sea of Japan. While Becky approached northern

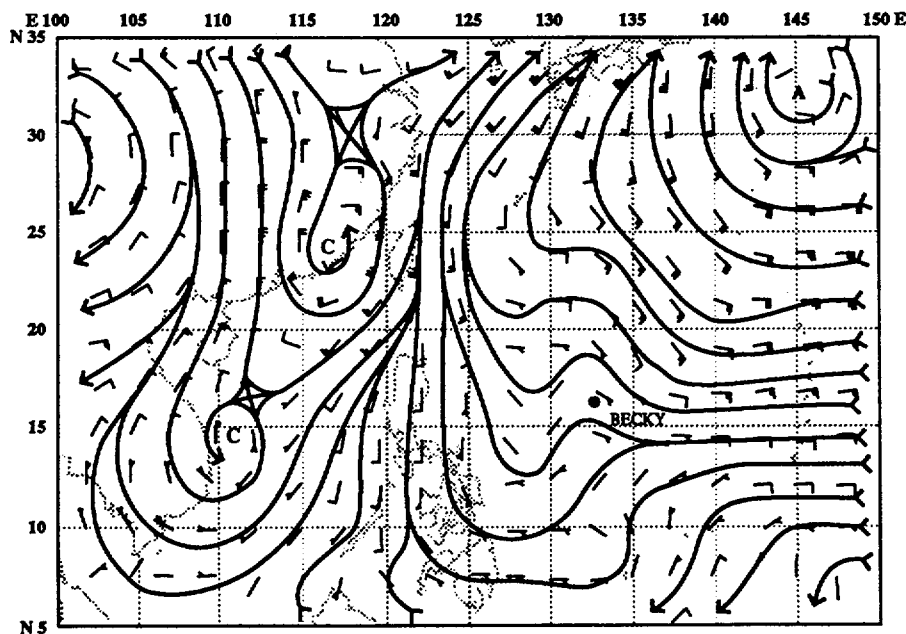


Figure 3-16-1. Deep layer mean circulation analysis from 221200Z August shows ridging north of Becky and troughing over the east coast of China.

Luzon, the trough shown over eastern China in Figure 3-16-1 moved eastward and filled (Figure 3-16-2). Subsequently, Becky accelerated as the steering flow strengthened and tracked to the west-southwest for the next day-and-a-half. With the high established to the north, the typhoon tracked due west and made landfall in northern Vietnam.

IV. INTENSITY

The disturbance that developed into Typhoon Becky originated in the low-level monsoon trough and the cloudiness left behind after Typhoon Zola (14W) separated from the trough. Strong northerly upper-level outflow from Zola slowed early development of Becky. Reestablishment of the TUTT to the north of the tropical cyclone effectively reduced the vertical shear and allowed the tropical cyclone to reach tropical storm intensity on 25 August. Becky attained minimal typhoon intensity and exhibited a 10 nm (19 km) diameter eye just as it crossed the northwestern tip of Luzon (Figure 3-16-3). After entering the South China Sea, Becky (Figure 3-16-4) maintained minimum typhoon intensity until it made landfall in northern Vietnam and rapidly dissipated.

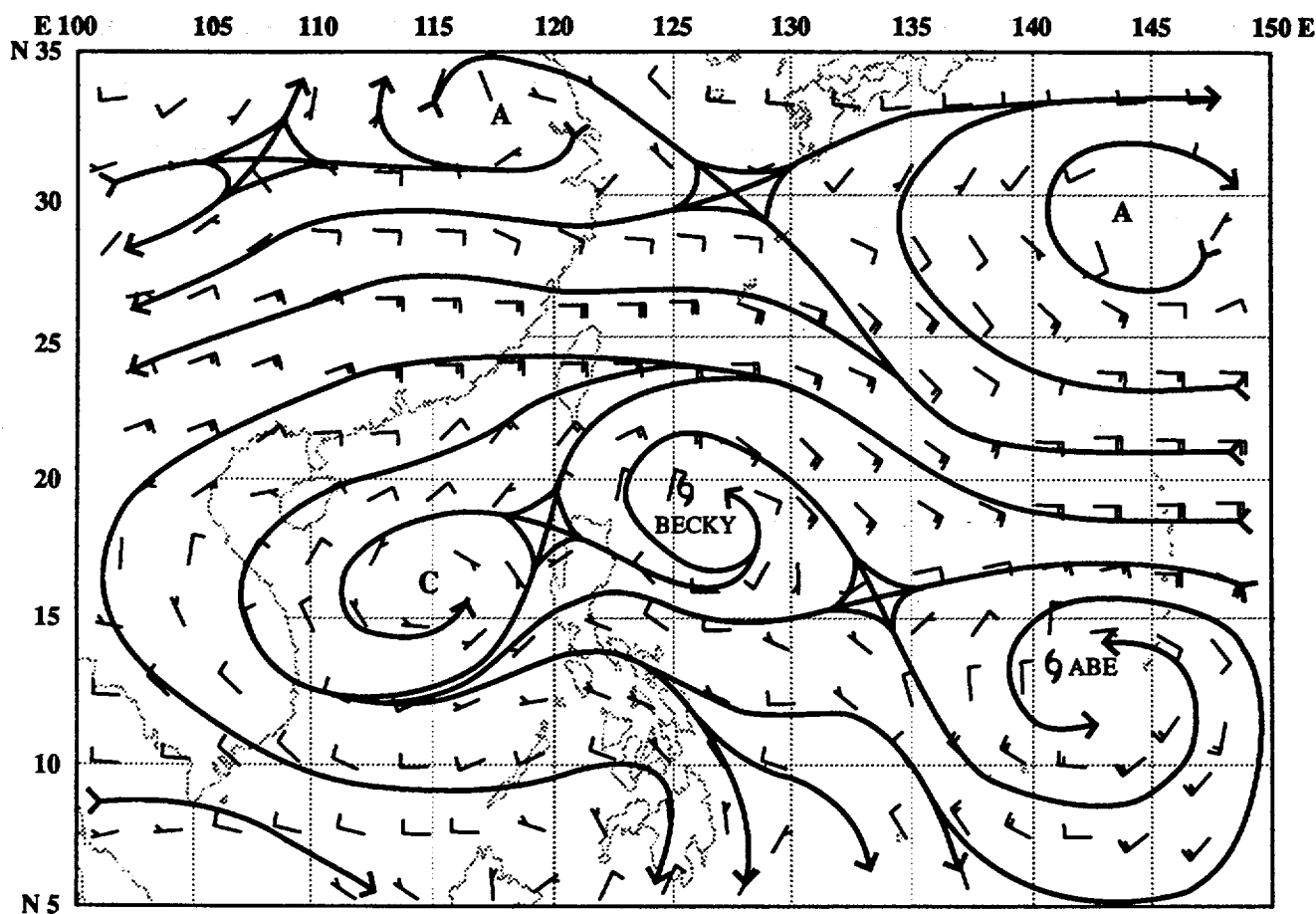


Figure 3-16-2. Deep layer mean circulation analysis for 251200Z August shows ridging over northeast China.



Figure 3-16-3. Becky reaches minimum typhoon intensity just as it hits northern Luzon (260039Z August DMSP visual imagery).

V. FORECASTING PERFORMANCE

Except for the first two warnings, JTWC correctly anticipated that Becky would turn and accelerate onto a more west-northwestward heading as it passed northern Luzon in response to the building ridge over eastern Asia (Figure 3-16-5). However, the strength of the ridge development was underestimated, resulting in a delay in forecasting the west-southwest portion of Becky's track.

VI. IMPACT

Becky crossed northern Luzon as it reached typhoon intensity, killing 32 people and forcing the evacuation of thousands due to heavy flooding. News reports from Vietnam stated that the northern province of Nghe Tinh experienced winds greater than 60 kt (30 m/sec) which severely damaged 400,000 acres of rice paddy and many homes. Three boats with a total of 20 fishermen aboard were reported missing.

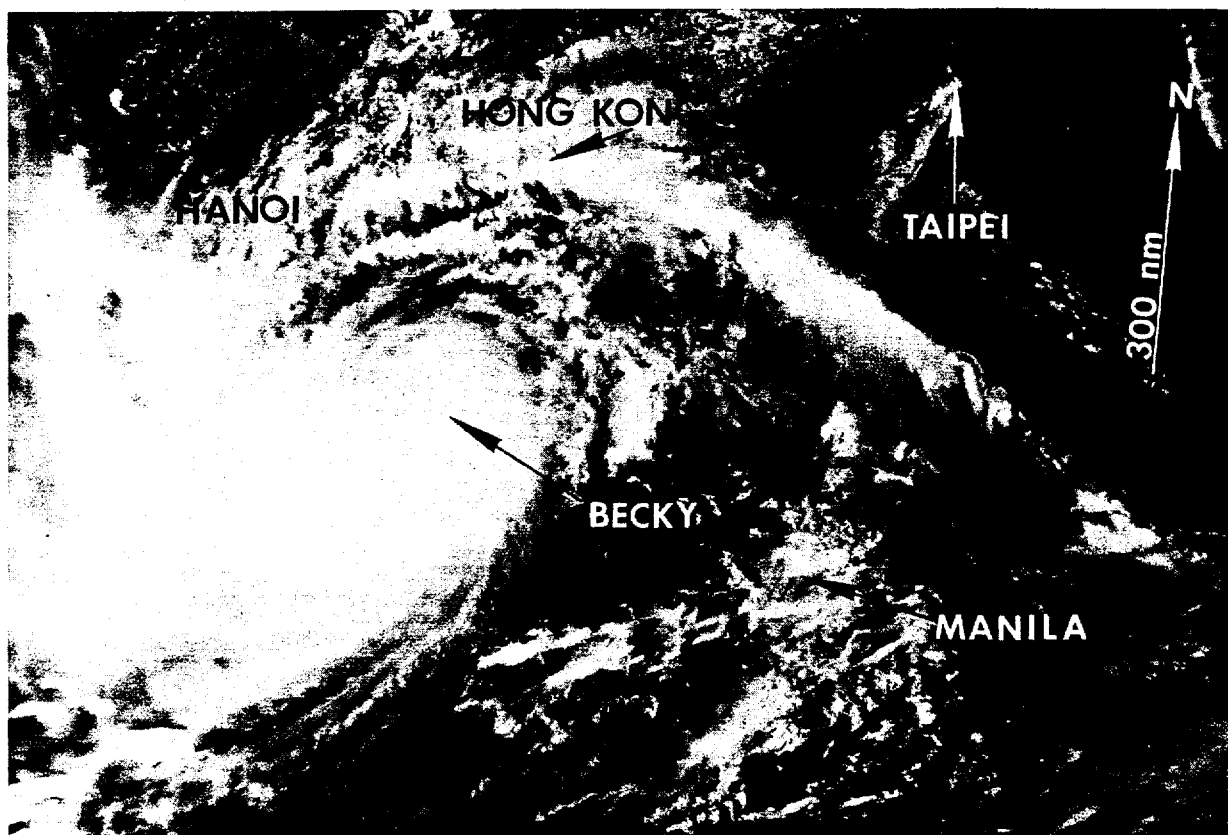


Figure 3-16-4. Becky at peak intensity of 70 kt (36 m/sec) before making landfall in northern Vietnam (280600Z August NOAA visual imagery).

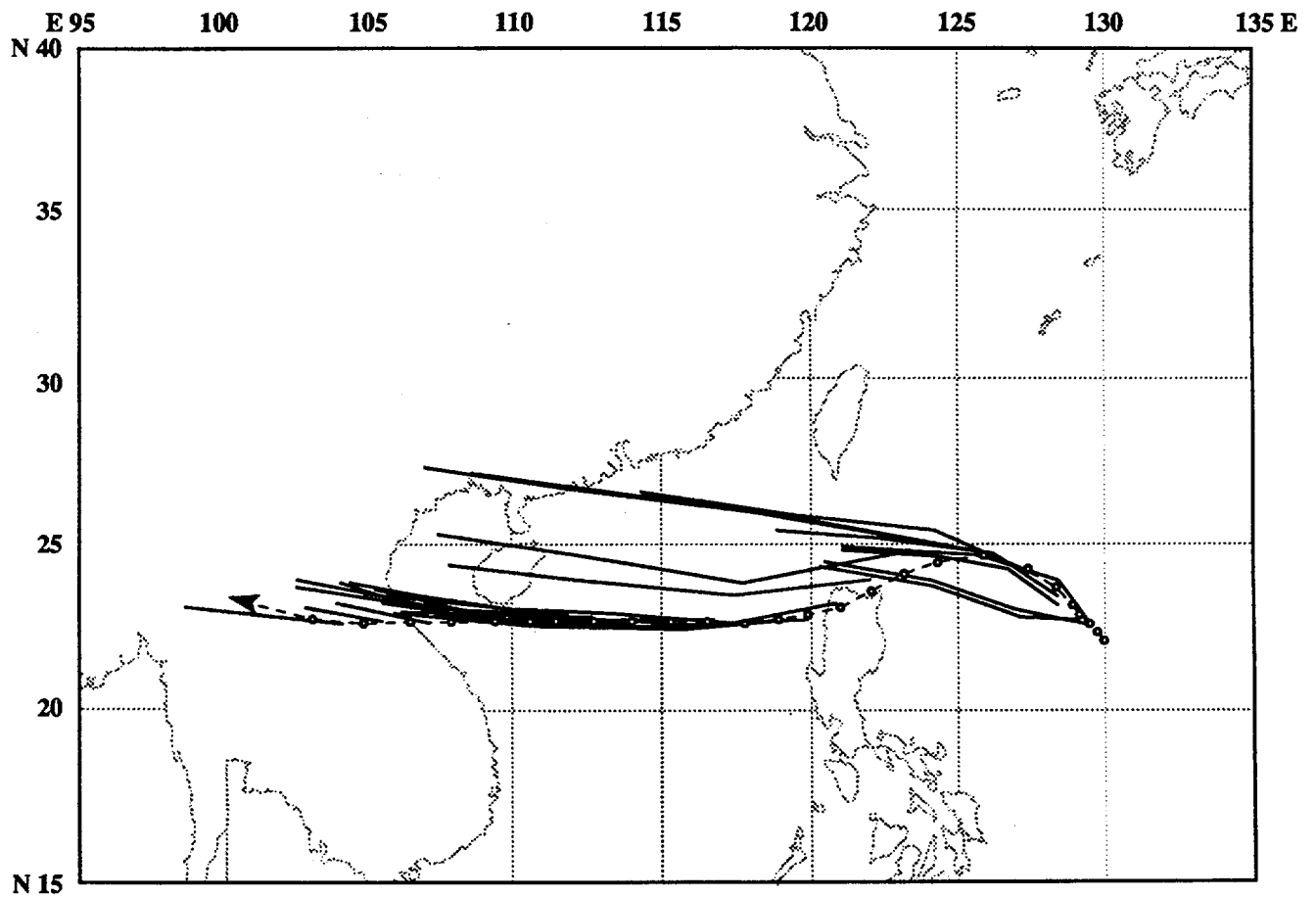
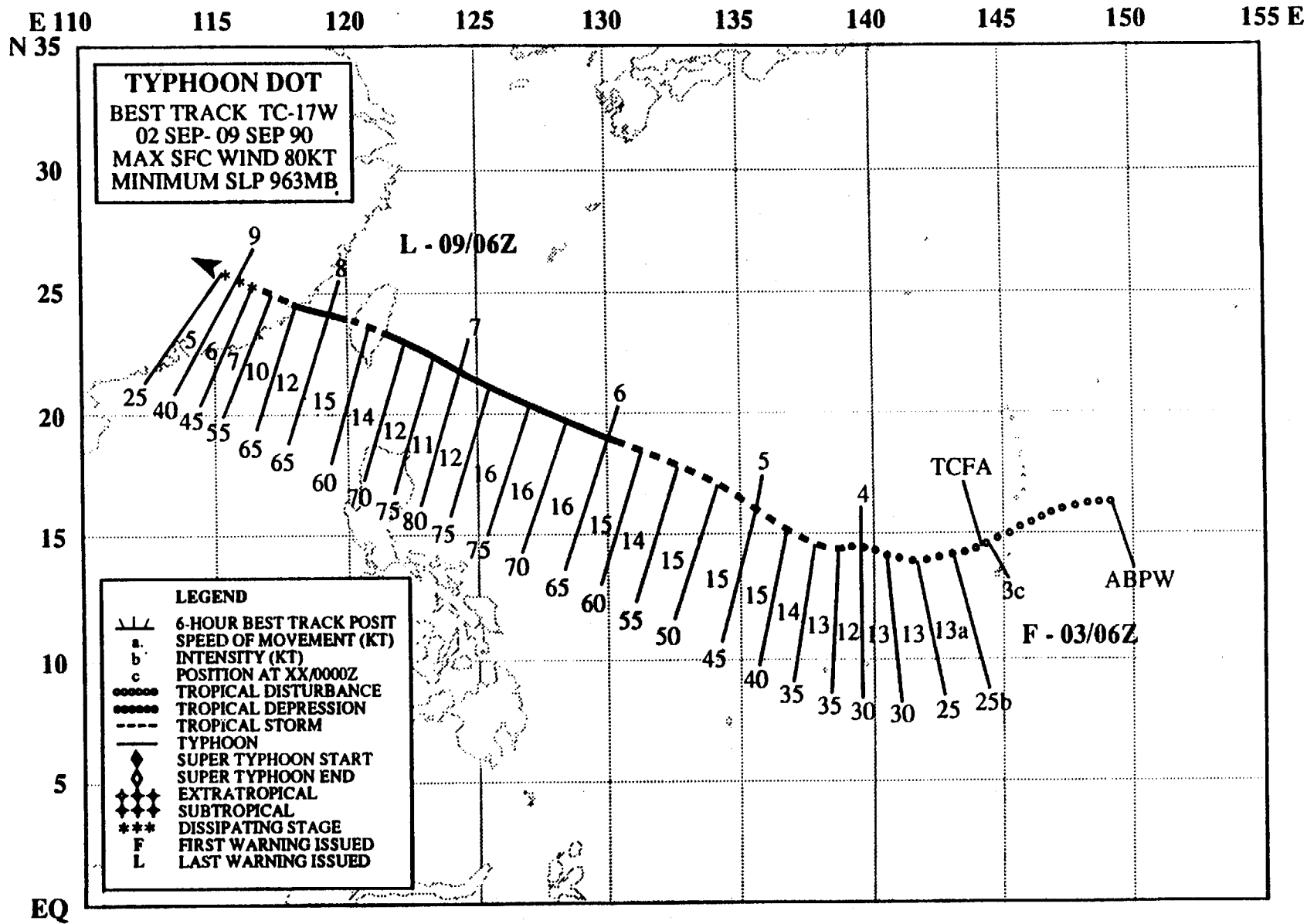


Figure 3-16-5. Summary of JTWC forecasts (solid lines) for Becky superimposed on the final best track (dashed line).



TYPHOON DOT (17W)

I. HIGHLIGHTS

Dot developed in the monsoon trough at the same time as Tropical Storm Cecil (18W) and brought enhanced southwesterly wind flow and heavy rains across Guam. Later, as Dot crossed central Taiwan, torrential monsoon rains from the associated monsoon surge caused extensive flooding in northern Luzon. During its passage across Taiwan and the Fujian Province of China, surface winds in the Formosa Strait exceeded 50 kt (26 m/sec) for 30 hours.

II. CHRONOLOGY OF EVENTS

- 300600Z - (August) First mentioned on the Significant Tropical Weather Advisory as a weak cyclonic circulation.
- 030100Z - (September) Tropical Cyclone Formation Alert issued due to improved vertical alignment between the low level circulation and the convection.
- 030600Z - First warning prompted by increased convective cloud organization.
- 040600Z - Upgrade to tropical storm based on consolidation of central cloud mass.
- 060000Z - Upgrade to typhoon based on formation of a banding-type eye.
- 070000Z - Peak intensity - 80 kt (41 m/sec) - as deep convection around a ragged eye increased.
- 071800Z - Downgraded to tropical storm due to the effects of mountainous terrain in central Taiwan.
- 080000Z - Upgraded to typhoon as eye redeveloped over the Formosa Strait.
- 081200Z - Final warning - dissipated - after Dot moved over land.

III. TRACK AND MOTION

The disturbance which later became Dot generated in the eastern extension of an active monsoon trough. Initially, Dot's cloud system center remained poorly organized and difficult to position. Consequently, six of the first seven warnings on the tropical cyclone were relocated as the convection fluctuated between the multiple circulation centers in the broad monsoon trough (Figure 3-17-1). After

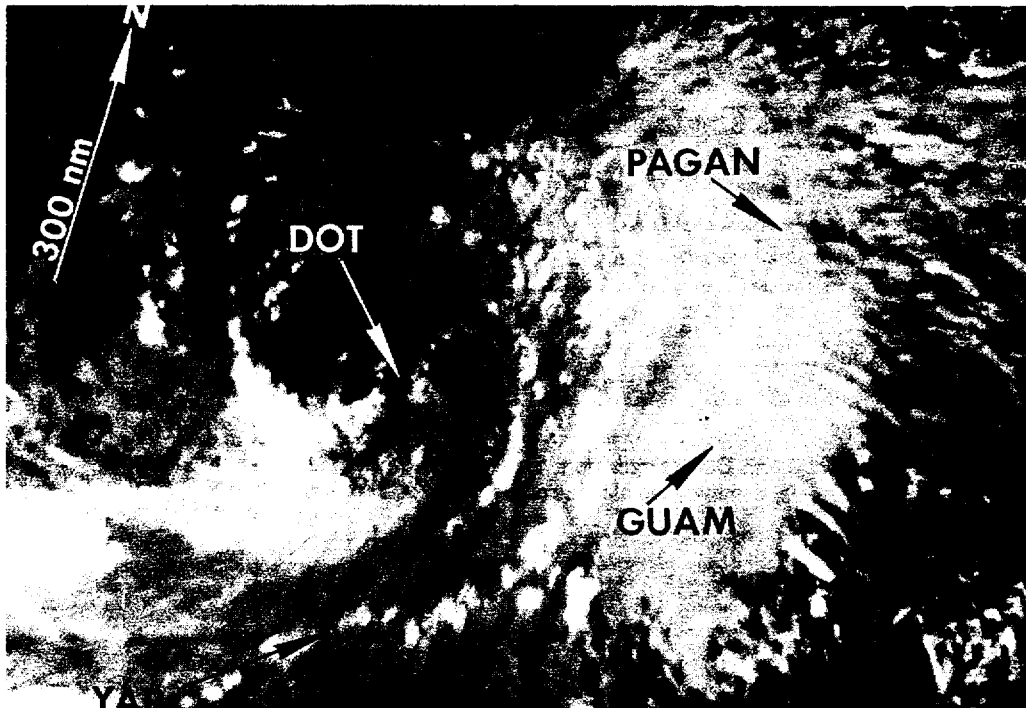


Figure 3-17-1. The broad circulation associated with Tropical Depression 17W extends over 300 nm (483 km) from its poorly defined circulation center (040423Z September NOAA visual imagery).

consolidation took place on 5 September (Figure 3-17-2), Dot tracked steadily west-northwestward south of the subtropical ridge, crossed central Taiwan and dissipated over Fujian Province in southeastern China.



Figure 3-17-2. Tropical Storm Dot emerges from the monsoon trough and begins to consolidate around a single, dominant circulation center (042232Z September NOAA visual imagery).

IV. INTENSITY

As a broad monsoon depression, Dot intensified at a rate of only 5 kt (3 m/sec) per day in its early stage of development. As the upper-level shear across the system diminished, convection increased around the circulation center, and a faster rate of intensification commenced. After becoming a tropical, Dot intensified at a steady rate of 20 kt (10 m/sec) per day prior to landfall in Taiwan. At maximum intensity, Typhoon Dot had a ragged eye approximately 25 nm (40 km) in diameter (Figure 3-17-3).

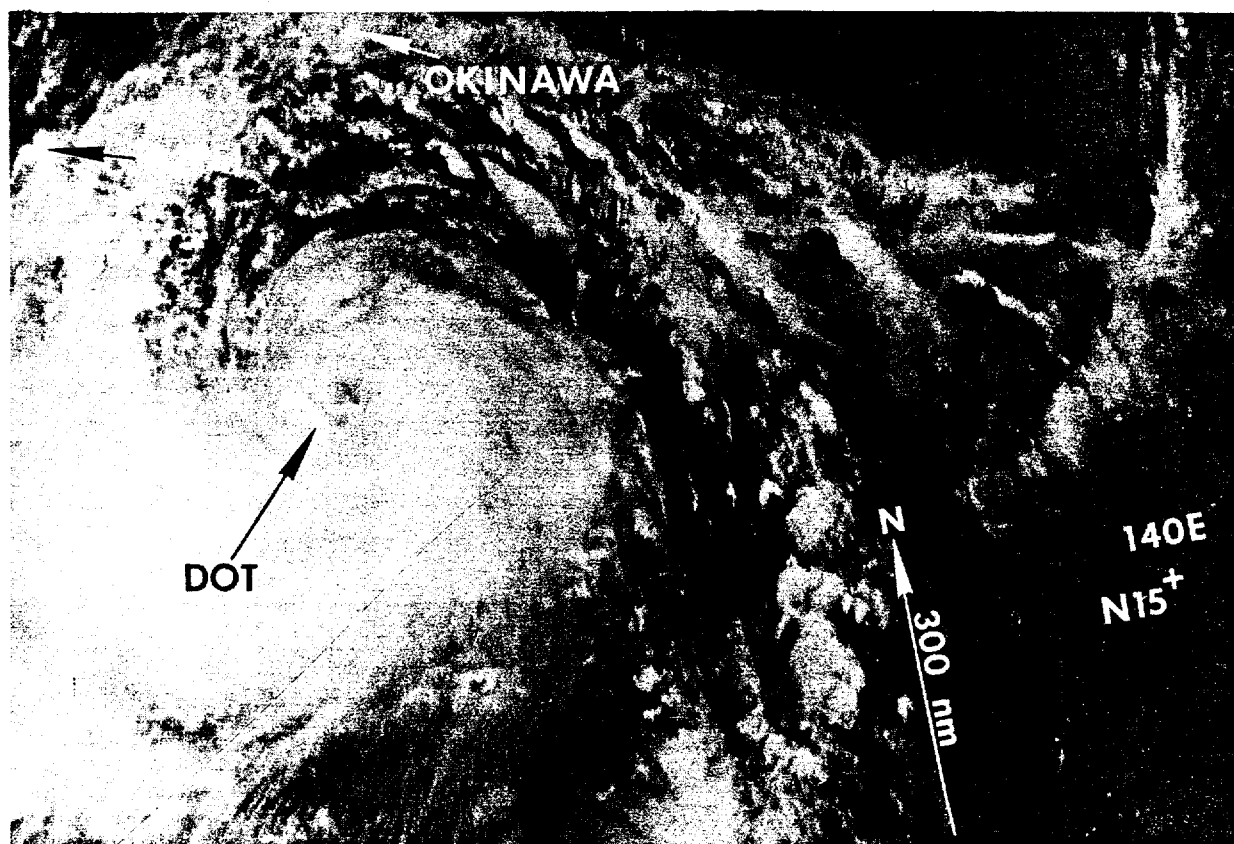


Figure 3-17-3. Typhoon Dot 11 hours prior to reaching maximum intensity east of Taiwan (061251Z September DMSP visual imagery)

Dot weakened significantly over the mountainous terrain of central Taiwan, then reintensified in the Formosa Strait. Dot's ragged eye was visible on radar (Figure 3-17-4) prior to landfall south of Zhangzhou in southern China.

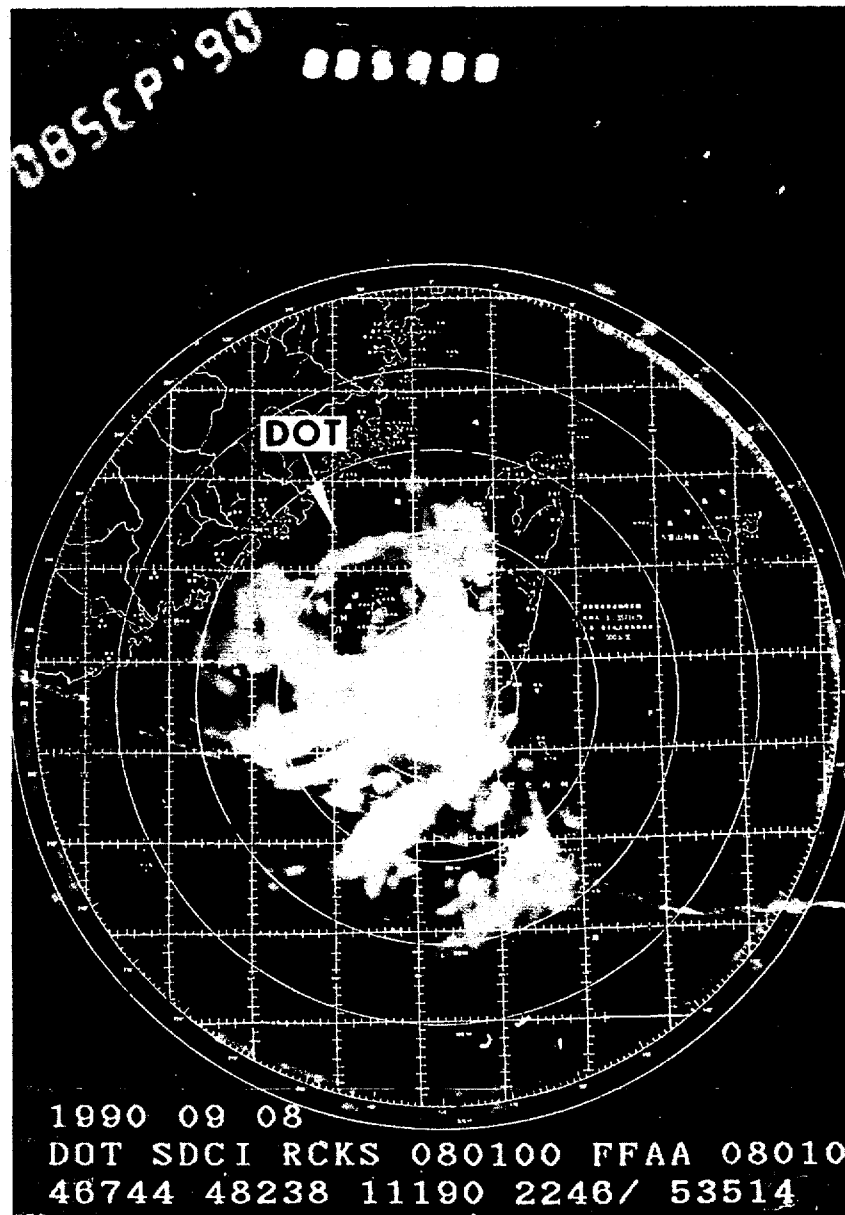


Figure 3-17-4. Evidence of redevelopment of an eye after Typhoon Dot passed across central Taiwan as seen by radar at Kaohsiung (WMO 46744) at 080100Z September (Photograph courtesy of Central Weather Bureau, Taipei, Taiwan).

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance is shown in Figure 3-17-5. Uncertainty about Dot's motion on 4 September resulted in larger forecast errors, but once its motion was more clearly established, JTWC forecast a west-northwestward track south of the subtropical ridge.

VI. IMPACT

Heavy rains from convergent low-level wind flow into Dot caused flooding on Guam, Luzon and Taiwan. The floods in northern Luzon caused the deaths of four people and the evacuation of an estimated 65,000 more. At least three deaths were reported in Taiwan.

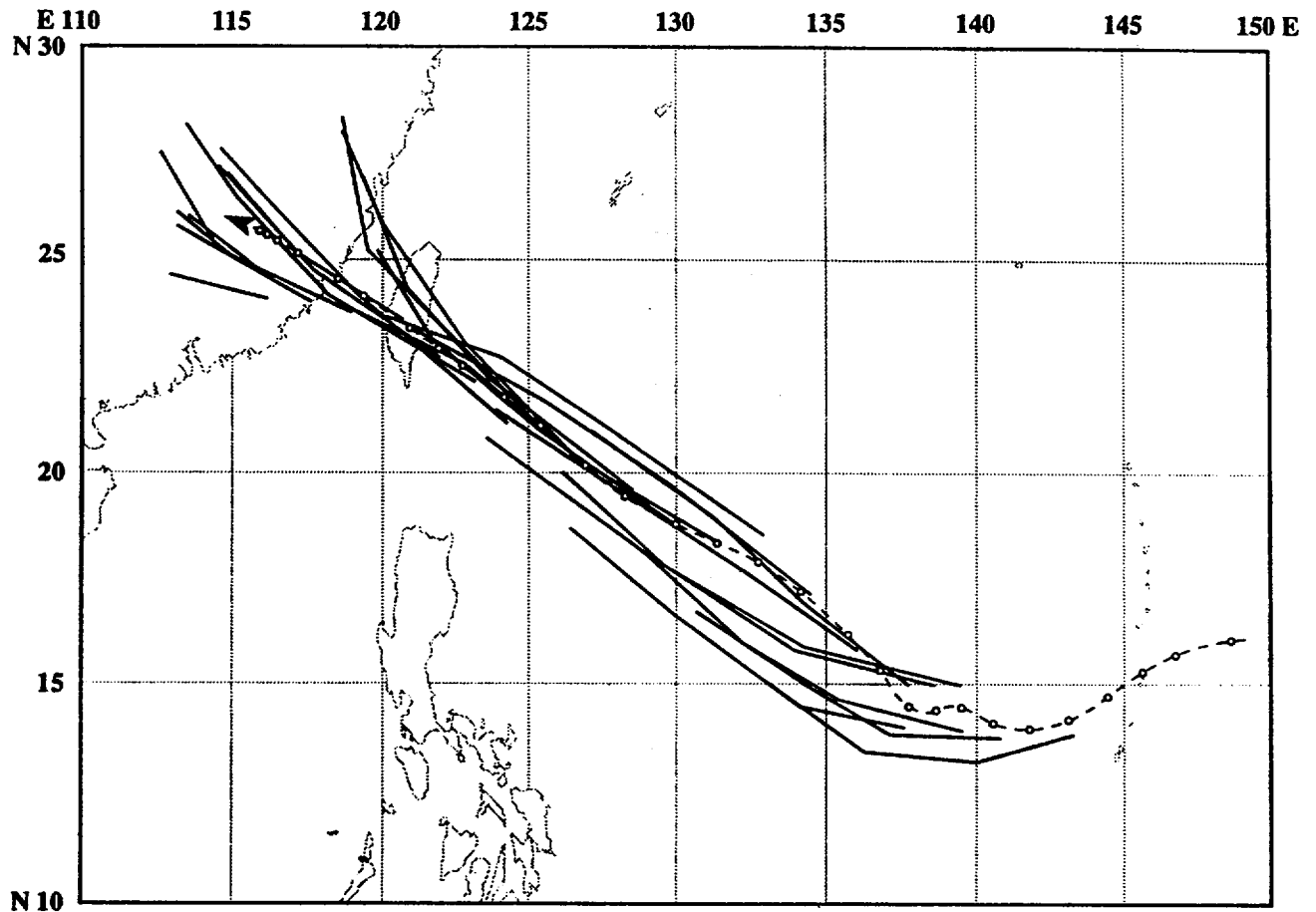
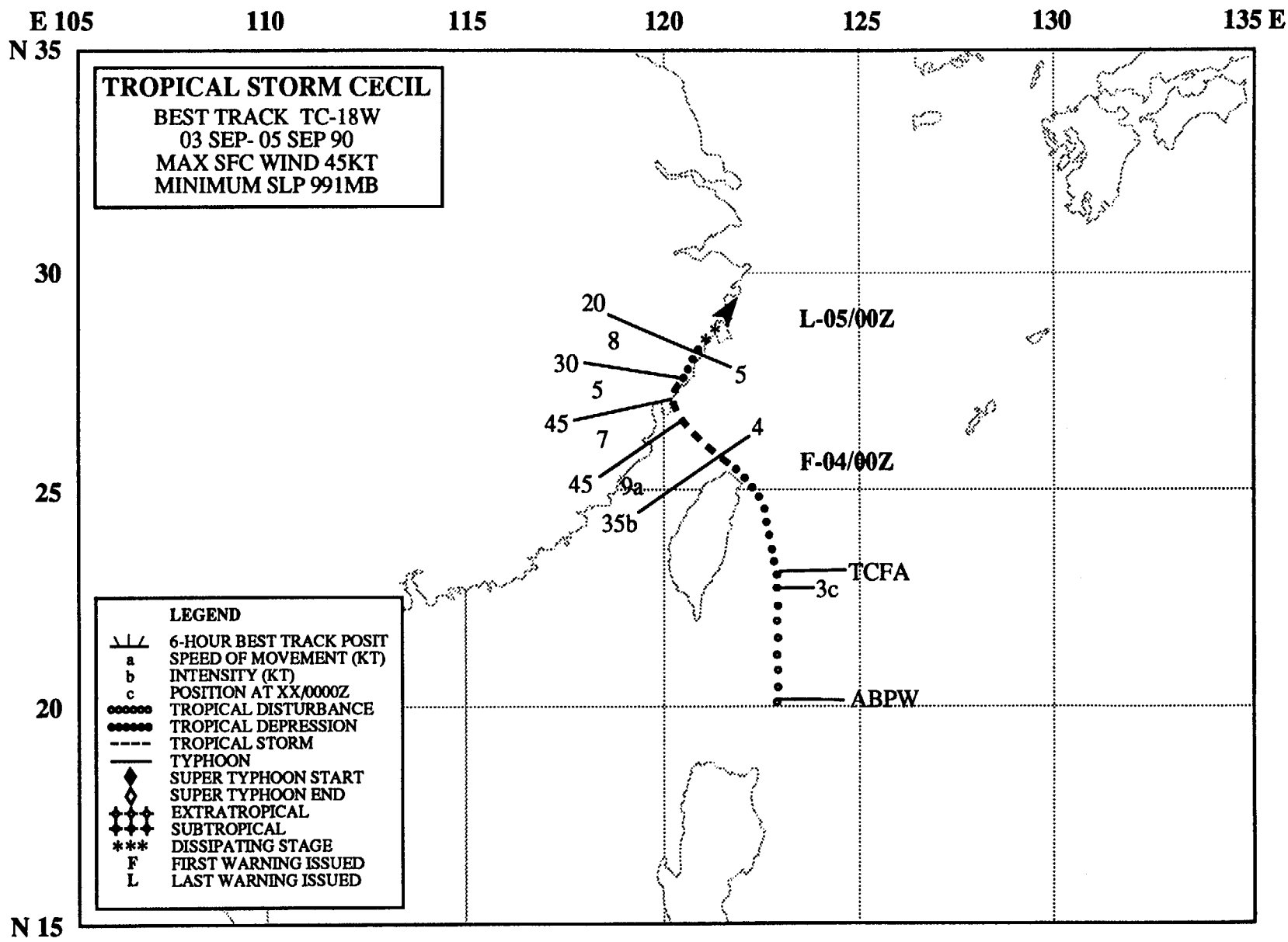


Figure 3-17-5. Summary of JTWC forecasts (solid lines) for Dot superimposed on the final best track (dashed line).



TROPICAL STORM CECIL (18W)

I. HIGHLIGHTS

Tropical Storm Cecil was a short-lived, midget tropical cyclone that formed in the wake of Typhoon Abe (15W). As Abe raced poleward, the monsoon trough reestablished itself over northern Luzon, and Cecil formed at the northeast end of the trough. Cecil tracked northward and skirted the northern coast of Taiwan before making landfall in southeastern China.

II. CHRONOLOGY OF EVENTS

020600Z - First mentioned on Significant Tropical Weather Advisory as a low level cyclonic circulation in the monsoon trough.

030400Z - Tropical Cyclone Formation Alert issued due to Dvorak intensity estimate of CI 1.5.

040000Z - First tropical storm warning issued after Dvorak analysis jumped up to CI 2.5.

040600Z - Peak intensity - 45 kt (23 m/sec) - based on synoptic report.

050000Z - Final warning issued as Cecil moved over land and dissipated.

III. TRACK AND MOTION

After Abe (15W) moved inland and northward over China, Cecil formed northeast of Luzon in association with a surge in the monsoon. Cecil tracked northward around the western periphery of the maritime subtropical ridge. As the tropical cyclone approached northern Taiwan, it turned northwestward and made landfall over southeastern China.

IV. INTENSITY

The mountainous terrain of Taiwan inhibited Cecil's development, however the tropical cyclone consolidated on a smaller than normal scale to become a midget tropical storm. The peak intensity of 45 kt (23 m/sec) was attained while crossing the Formosa Strait. Cecil (Figure 3-18-1) dissipated rapidly after making landfall over mainland China.

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance is shown in Figure 3-18-2. Forecasters initially expected Cecil to continue to track northward around the western side of the subtropical ridge. Subsequent forecasts reflected the northwestward track across the Formosa Strait and dissipation over land.

VI. IMPACT

No information was received.

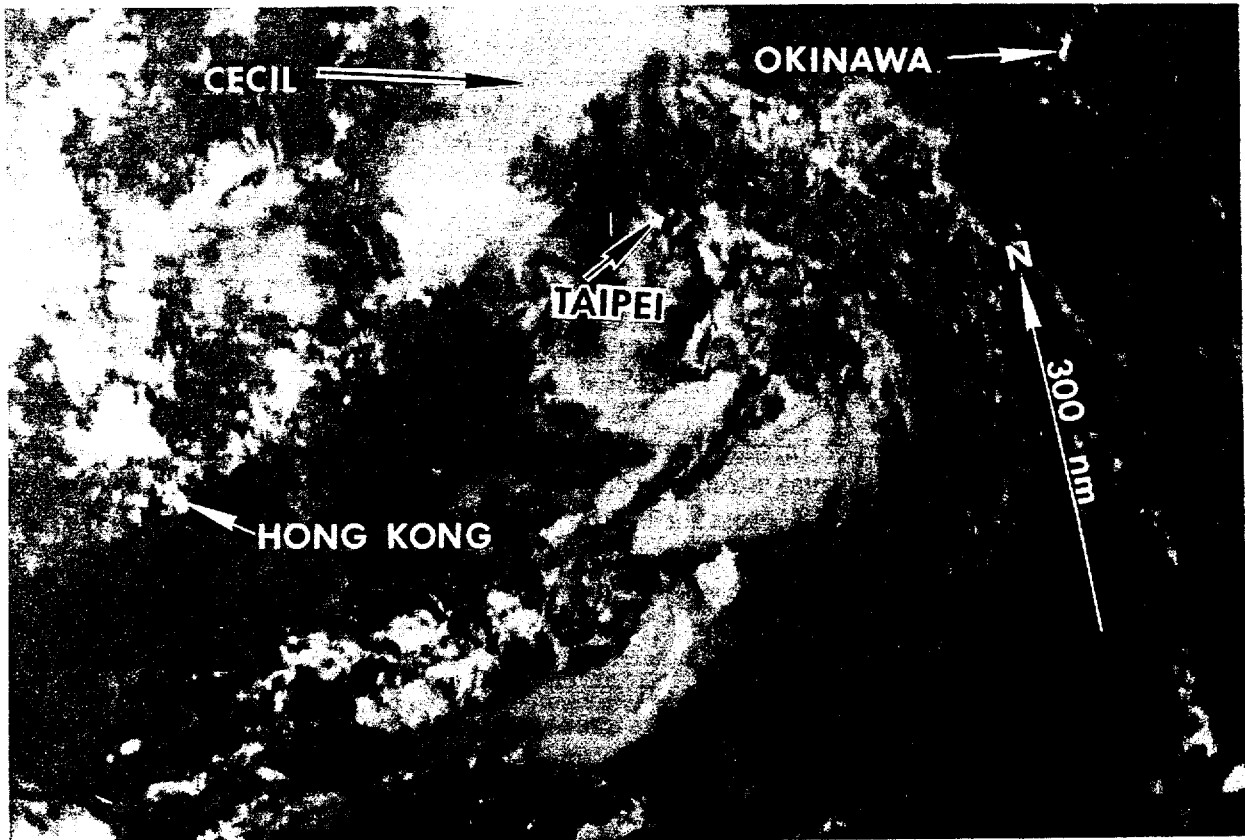


Figure 3-18-1. Moonlight photo of Cecil's small central dense overcast just on the coast of southeastern China. There is another tropical disturbance just southeast of Taipei. Note the bright city lights of Hong Kong, Taipei and Okinawa (041333Z September DMSP visual imagery).

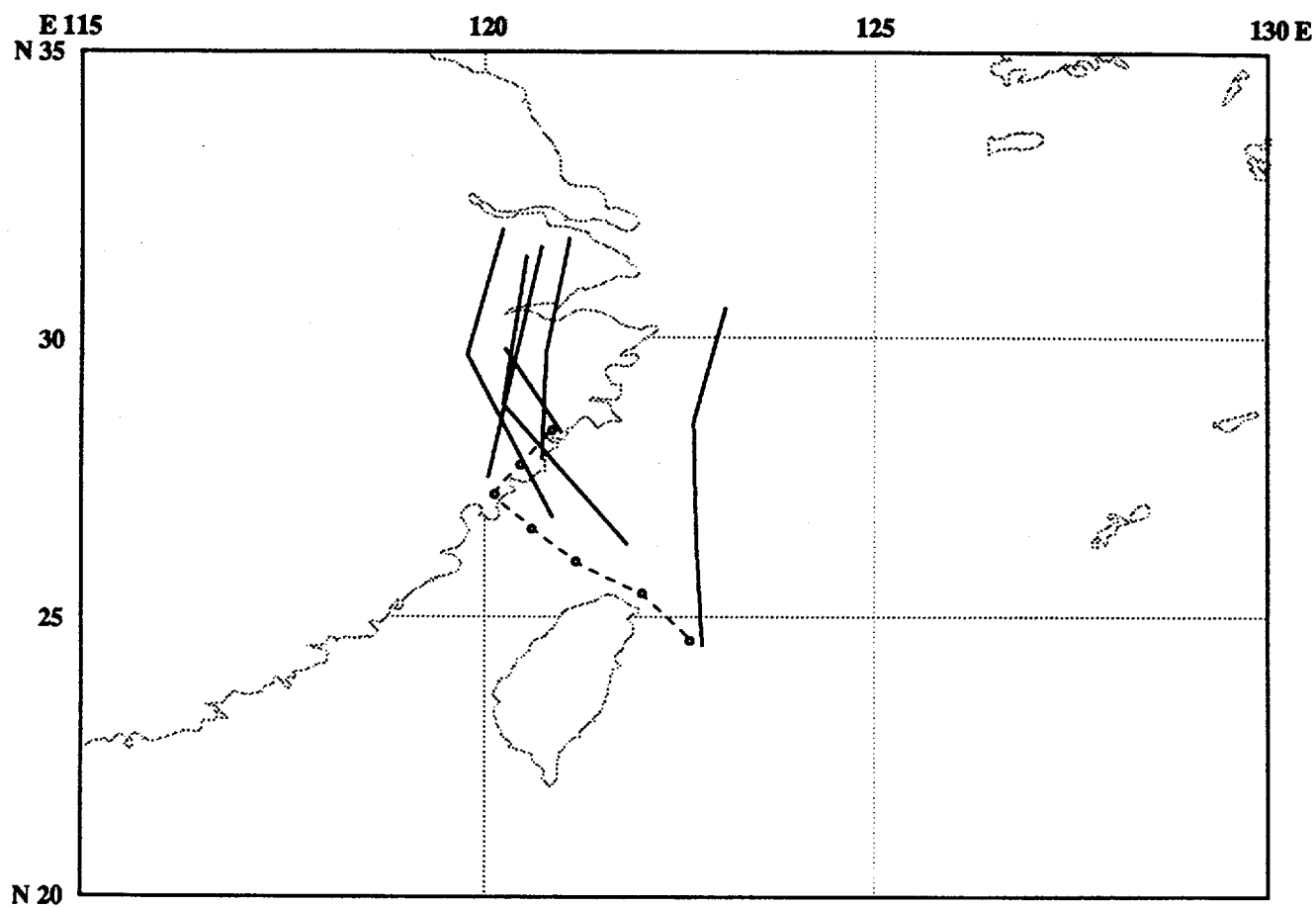
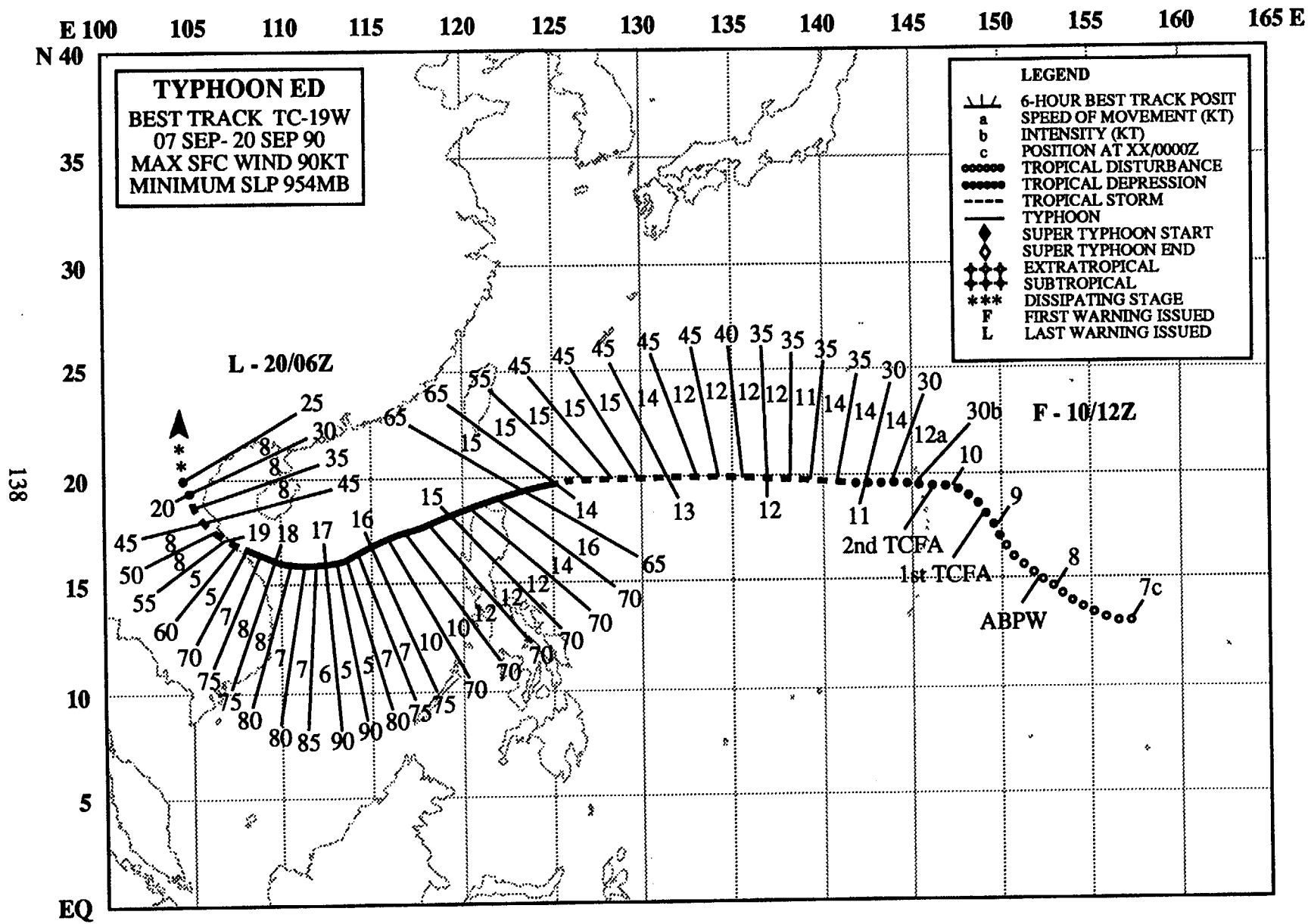


Figure 3-18-2. Summary of JTWC forecasts (solid lines) for Cecil are superimposed on the best track (dashed line).



TYPHOON ED (19W)

I. HIGHLIGHTS

Ed, which had the second longest track (3150 nm (5835 km)) of any "straight runner" in 1990, formed in the Marshall Islands and continued westward for nearly two weeks before finally making landfall in northern Vietnam. It was the third of six tropical cyclones to form in September.

II. CHRONOLOGY OF EVENTS

- 080600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1008 mb.
- 090730Z - First Tropical Cyclone Formation Alert based on increased in convective organization; synoptic data in the area indicating a small compact surface circulation; and favorable outflow conditions aloft.
- 100600Z - Second Tropical Cyclone Formation Alert based on persistent, well developed low-level circulation indicated by synoptic data.
- 101200Z - First warning followed improved organization in the convection, fair upper-level outflow in all quadrants and the first intensity estimate of CI 2.0.
- 121200Z - Upgrade to a tropical storm based on synoptic data, consolidation of the convection, improved upper-level outflow and the first intensity estimate of CI 2.5.
- 140000Z - Upgraded to typhoon due to better definition in the spiral banding, development of a partial eye wall and the first intensity estimate of CI 4.0.
- 181800Z - Downgraded to tropical storm based on a decrease in central convection, and an intensity estimate of CI 3.5.
- 200000Z - Downgraded to a tropical depression after a decrease in organization, land interaction and an intensity estimate of CI 2.0.
- 200600Z - Final warning - dissipated over land - issued as Ed moved inland over northern Vietnam.

III. TRACK AND MOTION

As a disturbance, Ed initially tracked northwestward in response to the deep layer flow around the subtropical anticyclone to the northeast (Figure 3-19-1). The tropical cyclone became involved in a

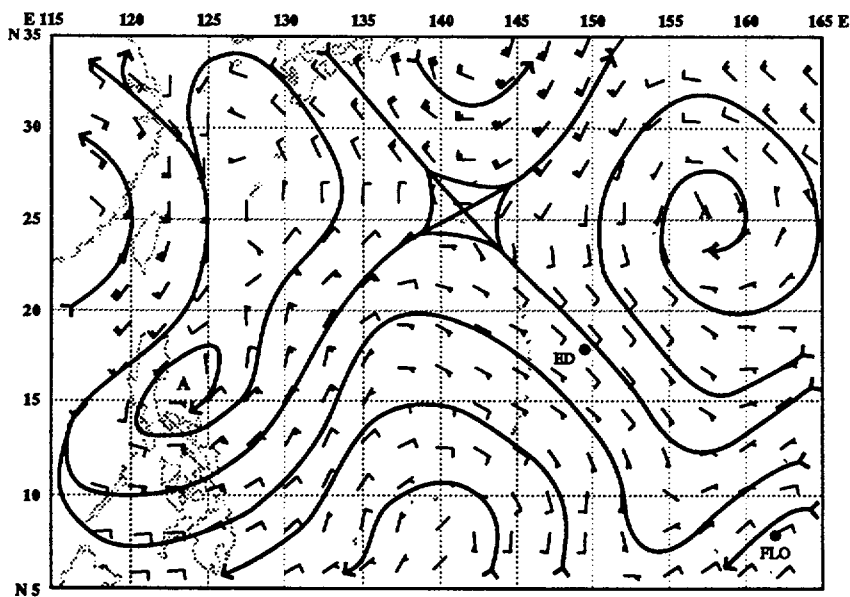


Figure 3-19-1. NOGAPS deep layer mean analysis from 090000Z September showing Ed embedded in southeasterly flow associated with the anticyclone to the northeast. Flo (20W) is located to the southeast of Ed.

ridge building process to the north and took a more westerly track on 10 September. Then the mid-level ridge strengthened to the north (Figure 3-19-2) and the typhoon turned west-southwestward at 140000Z. For four days Ed continued to track to the west-southwest before turning northwestward along the coast of Vietnam. The northwestward turn appeared to be the combined result of the steering flow becoming southerly when a mid-level ridge formed between Ed and Flo (20W)(Figure 3-19-3), and the barrier

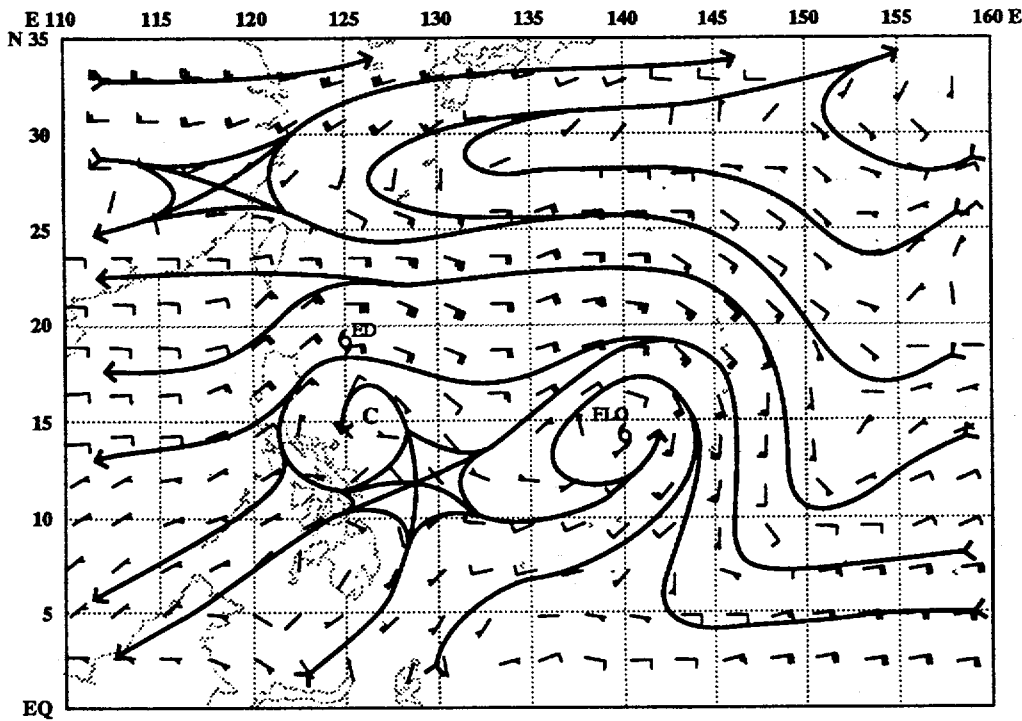
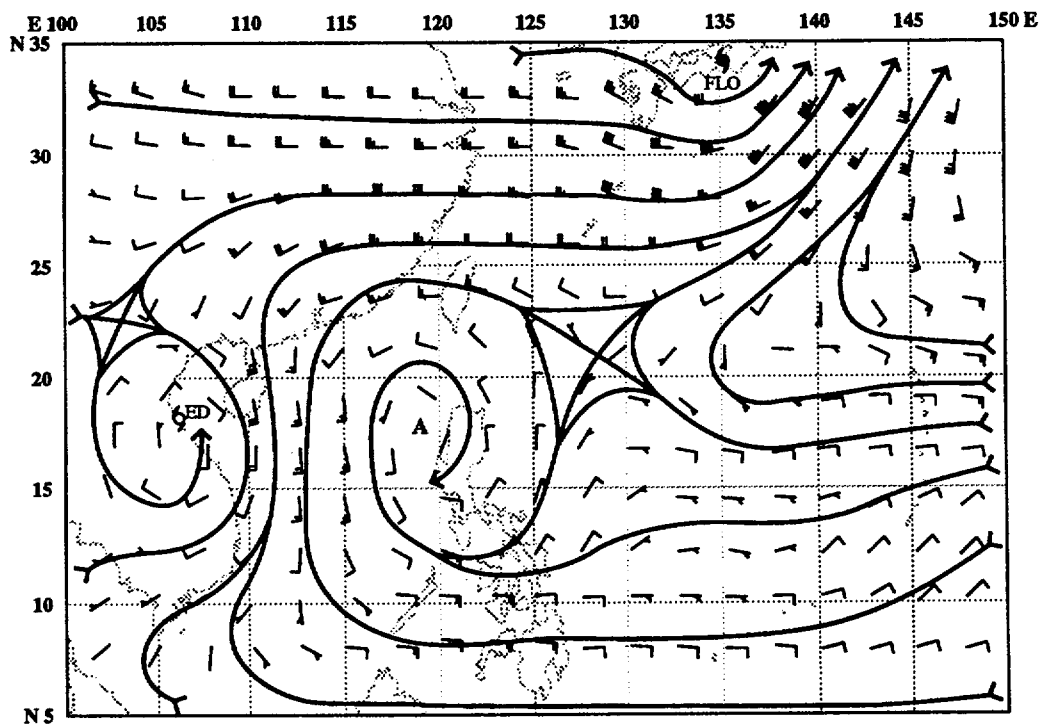


Figure 3-19-2. NOGAPS deep layer mean analysis from 140000Z September shows the increased wind flow between Ed and the building subtropical ridge to the north. Flo (20W) is located to the east-southeast of Ed.

Figure 3-19-3. NOGAPS deep layer mean analysis from 191200Z September shows a mid-level anticyclone over the Philippine Islands producing southerly flow across the South China Sea.



effect of the coastal mountains of Vietnam.

IV. INTENSITY

On 10 September, Ed's compact cluster of cumulonimbus clouds moved into a more favorable upper-level environment with low vertical wind shear. At the same time, the disturbance which would later become Flo (20W), was rapidly taking shape southeast of Guam (Figure 3-19-4). The 20-30 kt (10-15 m/sec) low-level monsoonal southwesterlies to the south aided the development of both systems. Intensification continued at a normal rate of one T-number per day, and Ed became a typhoon on 14 September (Figure 3-19-5). The tropical cyclone maintained typhoon intensity until it struck the Vietnamese coast on 18 September.

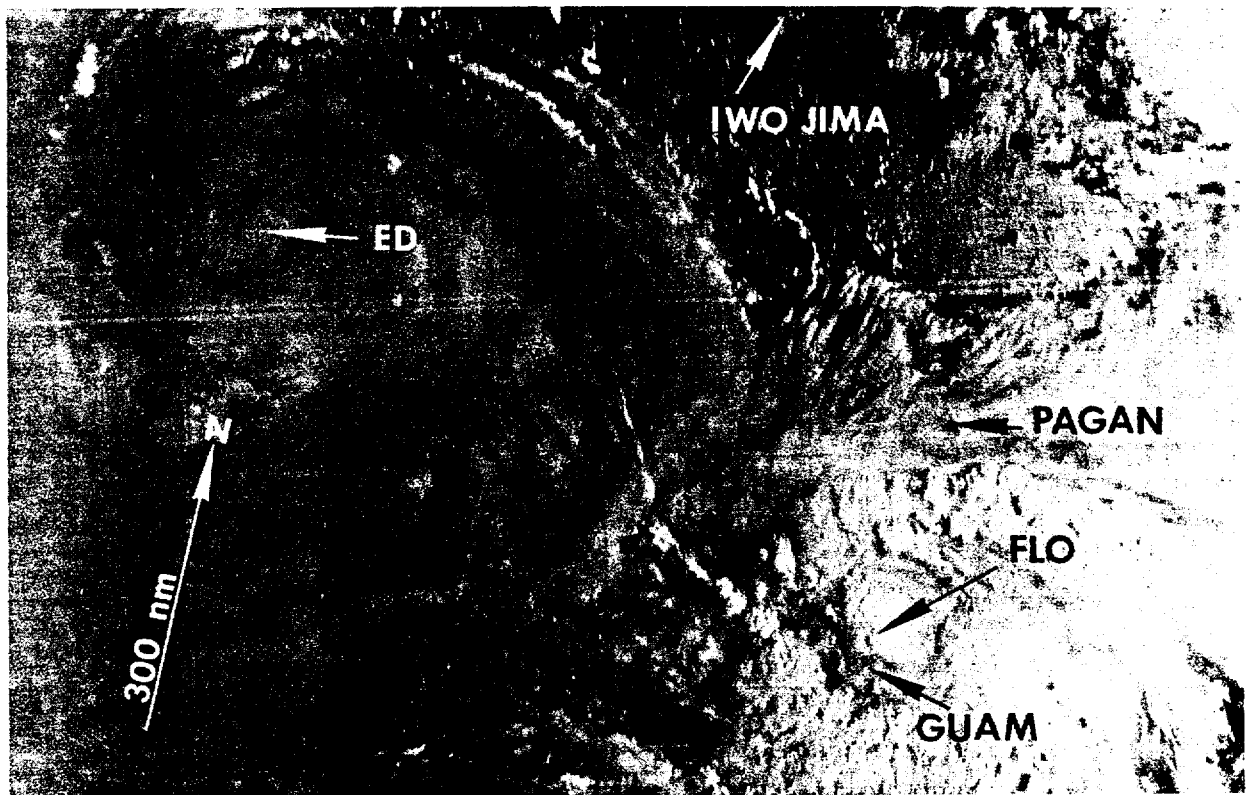


Figure 3-19-4. Tropical Storm Ed as it moves westward across the Philippine Sea. Tropical Storm Flo is rapidly developing to the southeast near Guam (122042Z September DMSP visual imagery).

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance is shown in Figure 3-19-6. The initial forecasts were to the right of Ed's westward track and were influenced by the NOGAPS 500-mb prognostic series, which continued to forecast significant weakening of the mid-level ridge over the East China Sea. The ridge actually strengthened, keeping the tropical cyclone on a more westward track. JTWC was strongly influenced by the dynamic aid, OTCM, which forecast a west-northwestward track throughout Ed's life. Of interest, the dynamic aid FBAM, which used the smoothed deep layer mean fields for steering, correctly forecast Ed's turn to the southwest, but missed the track change off Vietnam.

VI. IMPACT

No information was received.

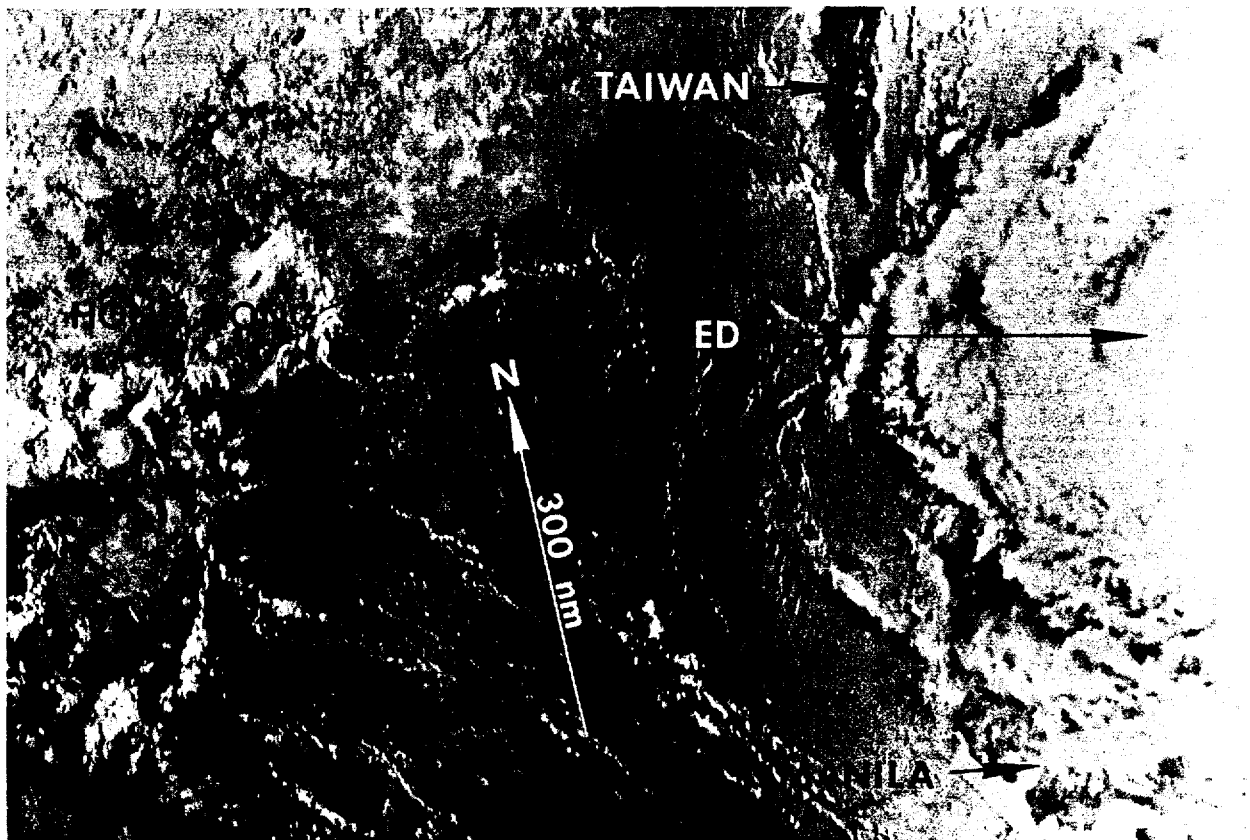


Figure 3-19-5. Typhoon Ed approaches the Straits of Luzon (140007Z September NOAA visual imagery).

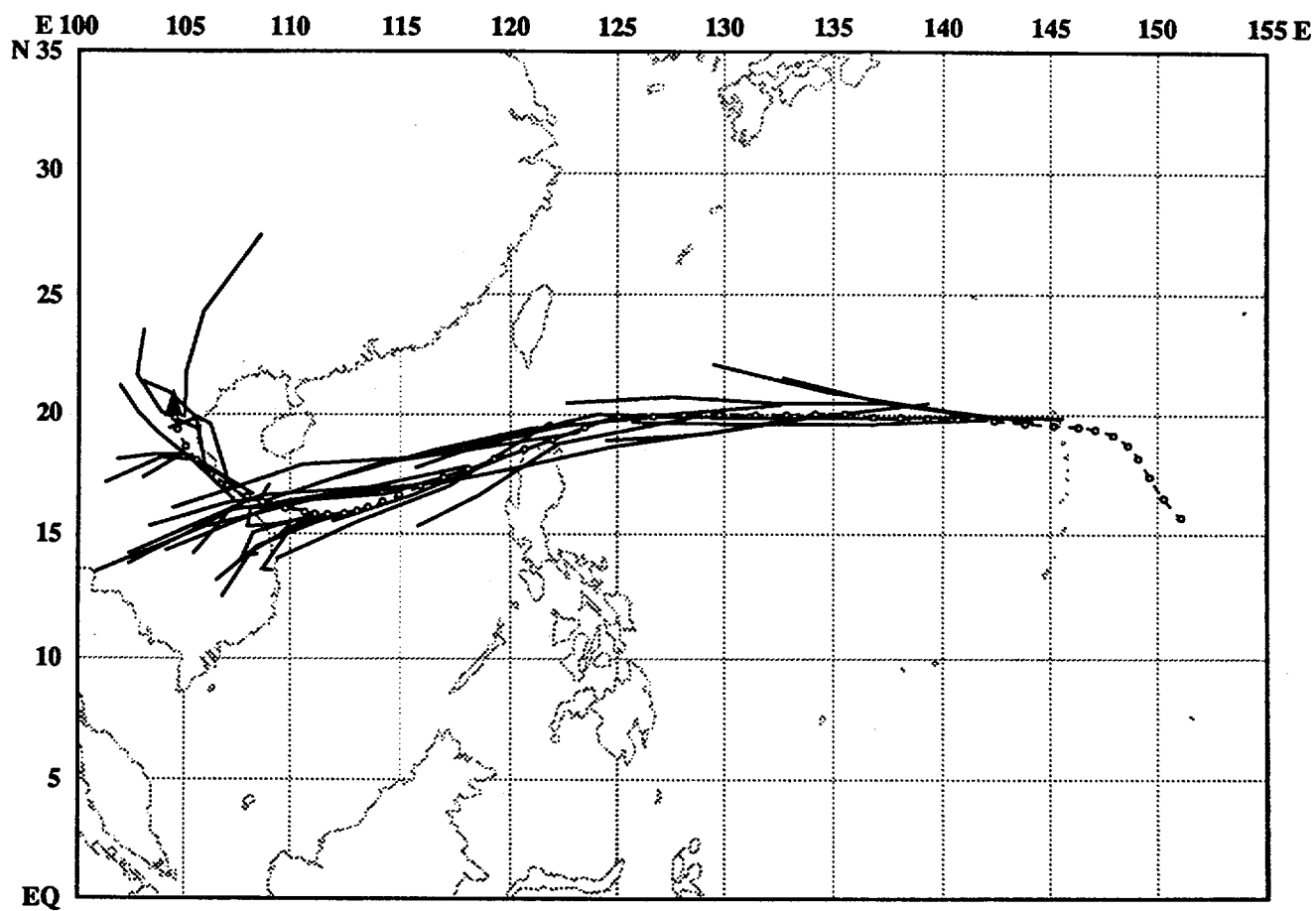
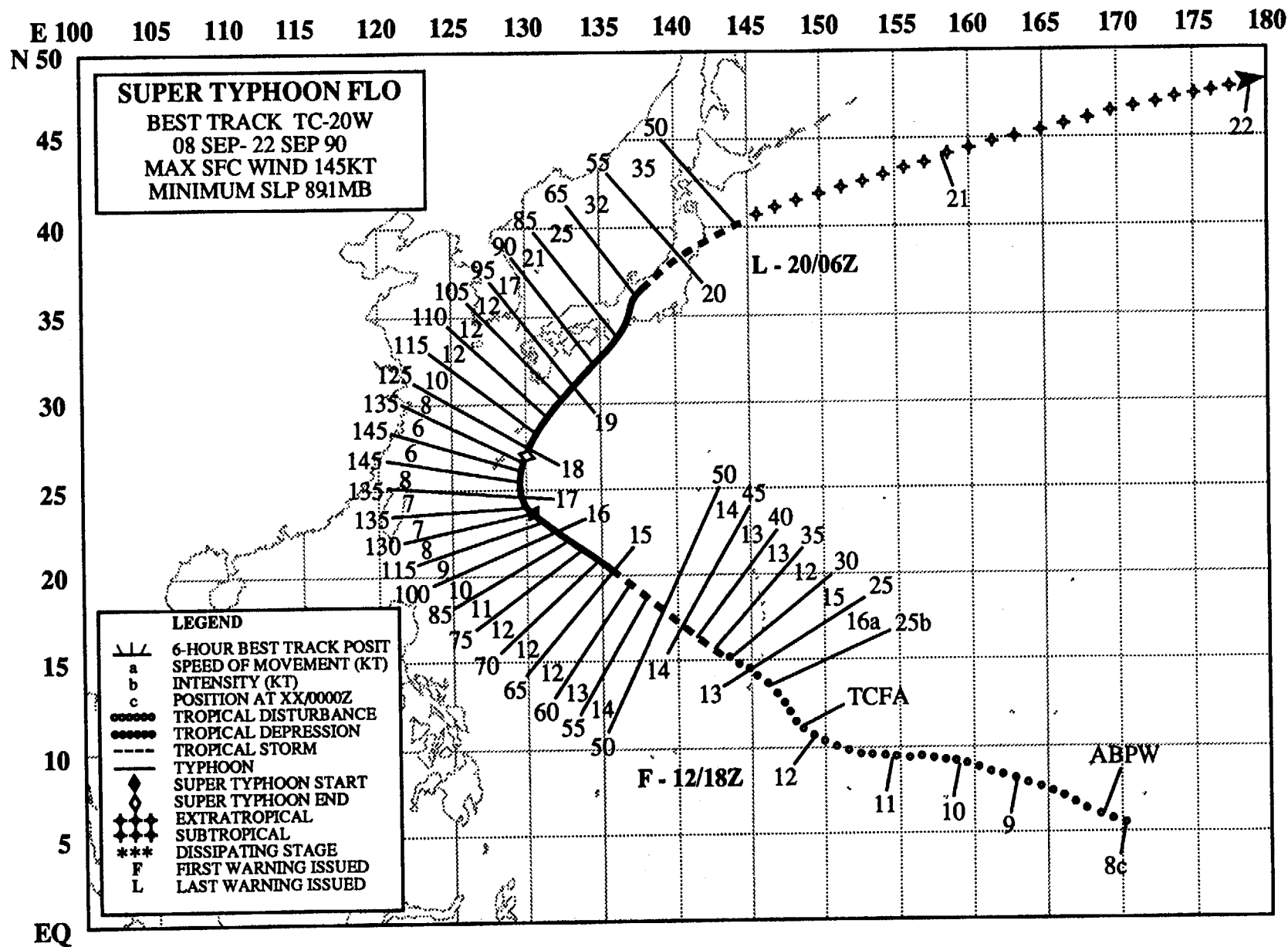


Figure 3-19-6. Summary of JTWC forecasts (solid lines) for Ed superimposed on the final best track (dashed line).



SUPER TYPHOON FLO (20W)

I. HIGHLIGHTS

Flo was the fourth of six tropical cyclones to develop in September, the first of four super typhoons this year, and the object of over three consecutive days of upper-tropospheric aircraft reconnaissance missions during the TCM-90 field experiment. Flo formed in the wake of Typhoon Ed (19W), passed close by Guam, then rapidly intensified into a super typhoon as it approached Okinawa. Recurvature was slow before the tropical cyclone accelerated northeastward towards the Japanese mainland where it was called the most powerful typhoon to hit Honshu in 19 years. At least 38 people were reported dead or missing, damage was estimated in the millions of dollars, and transportation, communications and power were disrupted.

II. CHRONOLOGY OF EVENTS

- 080600Z - First mentioned on the Significant Tropical Weather Advisory as an area of convection in the monsoon trough with an estimated sea-level pressure of 1008 mb.
- 120530Z - Tropical Cyclone Formation Alert based on increased organization associated with well defined upper-level circulation center.
- 121800Z - First warning issued due to a continued increase in convective organization.
- 131800Z - Upgraded to tropical storm after a Dvorak intensity estimate of CI 2.5.
- 150600Z - Upgrade to typhoon based on the appearance of a small circular eye and the first CI 4.0.
- 161200Z - Reached super typhoon intensity after undergoing a period of rapid deepening: intensity estimate of CI 7.0.
- 170600Z - Peak intensity - 145 kt (75 m/sec) - based on the 891 mb report from a TCM-90 aircraft reconnaissance dropsonde.
- 180000Z - Downgraded to typhoon intensity after eye became ragged and intensity estimate of CI 6.0.
- 200000Z - Downgraded to a tropical storm due to increased vertical wind shear and the start of extratropical transition.
- 200600Z - Final warning issued as Flo transformed into an extratropical low.

III. TRACK AND MOTION

Within a day after Ed (19W) began to consolidate on 7 September in the monsoon trough 750 nm (1390 km) east of Guam, a persistent area of convection that would become Flo developed farther to the east in the southern Marshall Islands. Under the steering influence of the subtropical ridge to the north, Flo drifted west-northwestward for the next eight days at approximately 12 kt (22 km/hr). As the tropical cyclone approached Okinawa on 15 September, a mid-latitude short wave trough deepened to the northwest and induced a break in the subtropical ridge. On 17 September, Flo slowed and started to recurve around the western periphery of the ridge. It slowly accelerated in response to the passing trough. Finally, on 19 September, the typhoon accelerated northeastward across Honshu, as it became embedded in the stronger mid-latitude westerlies aloft. Flo subsequently transitioned to an extratropical system east of Japan on 20 September.

IV. INTENSITY

Flo existed as a weak disturbance for four and a half days (8 - 12 September) before it started to intensify. Nearing Guam on 12 September, the disturbance's convection and low-level circulation appeared to consolidate (Figure 3-20-1). This consolidation process seemed related to the availability of deeper monsoonal southwesterly flow that was enhanced by the presence of Ed (19W) to the west. During the subsequent intensification process, the TCM-90 Doppler radar profiler on the island of

Saipan, 100 nm (185 km) to the north-northeast of Guam, recorded an interesting event. A time-height cross-section of meridional wind speed for 13 September revealed a mid-tropospheric 50 kt (25 m/sec) wind maximum (Figure 3-20-2) that extended around the eastern edge of Flo. The presence of the mid-level jet was concurrent with the intensification of Flo into a tropical storm.

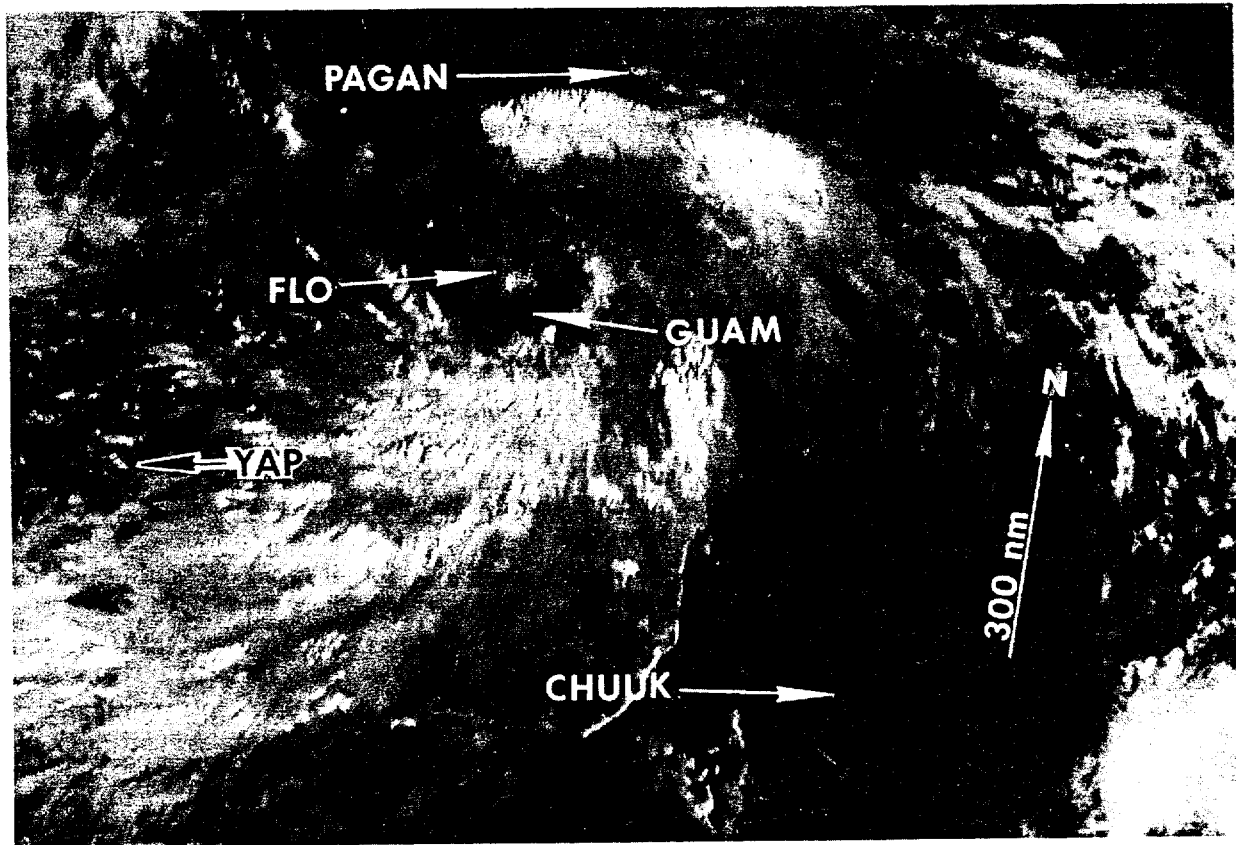


Figure 3-20-1. Flo as a tropical depression near Guam. The relatively clear area northeast of Guam is related to subsidence from Ed (19W), which is just off the top left edge of the photo (122325Z September DMSP visual imagery).

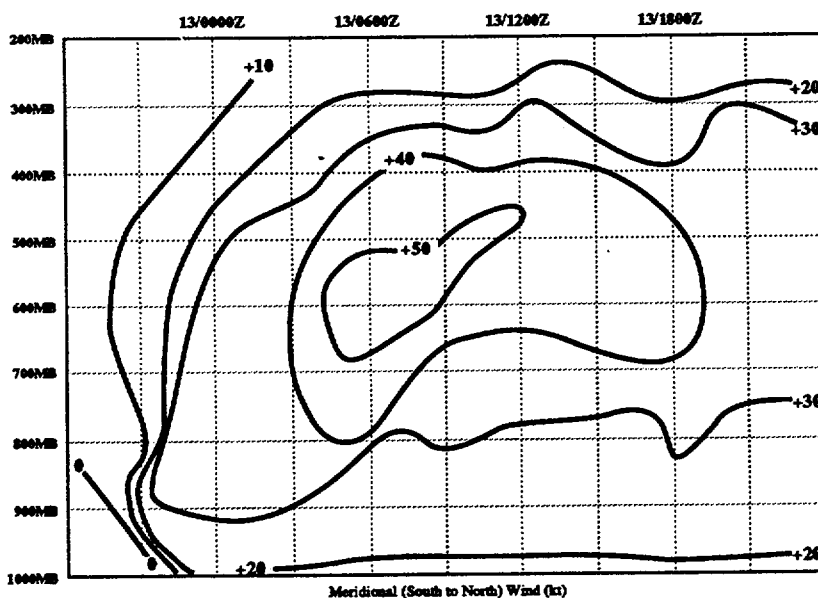


Figure 3-20-2. Time-height cross-section of the meridional wind speed for 13 September from the TCM-90 Doppler radar profiler on Saipan. The analysis shows the mid-level wind maximum that was observed on the east side of Flo.

Subsequently, the tropical cyclone intensified at a normal rate of one T-number per day until it reached typhoon intensity on 15 September (Figure 3-20-3). Then Flo rapidly intensified for the next 36 hours. On 16 September, as Flo was becoming a super typhoon, it also was the subject of an Intensive Observing Period (IOP) as part of the TCM-90 experiment. During the IOP, the NASA DC-8 reconnaissance aircraft provided JTWC with invaluable information on the location, structure, and intensity of Flo as the storm approached Okinawa (Figures 3-20-4 and 3-20-5). These data were the first-ever upper tropospheric (near 200 mb) winds from a western North Pacific tropical cyclone to be collected and used operationally. As a result of the information provided, JTWC increased the maximum winds from 135 kt (69 m/sec) to 145 kt (75 m/sec) at 170600Z September (Figure 3-20-6). The flight level for the reconnaissance missions ranged from 37,000 to 43,000 ft (11.3 km to 13.1 km), approximately the 200-mb level. These data revealed that there was intense cyclonic flow around Flo's core with what appeared to be very little direct outflow evident close to the eye. Flight-level winds of 110 kt (55 m/sec) were recorded just east of the eye on 17 September. In addition, the presence of an anticyclonic eddy to the southeast of the eye was documented. A central pressure of 891 mb obtained via the dropsonde on the same day correlated well with both the Dvorak (1984) estimates of current intensity and the Atkinson-Holliday (1977) pressure-wind relationship.

As Flo began to recurve, it remained over the warm waters of the Kuroshio Current. Vertical wind shear weakened the typhoon, but it still had 90 kt (45 m/sec) sustained surface winds when it slammed into southern Honshu on 19 September. Interaction with land further weakened the tropical cyclone, and it transitioned to an extratropical cyclone the following day.

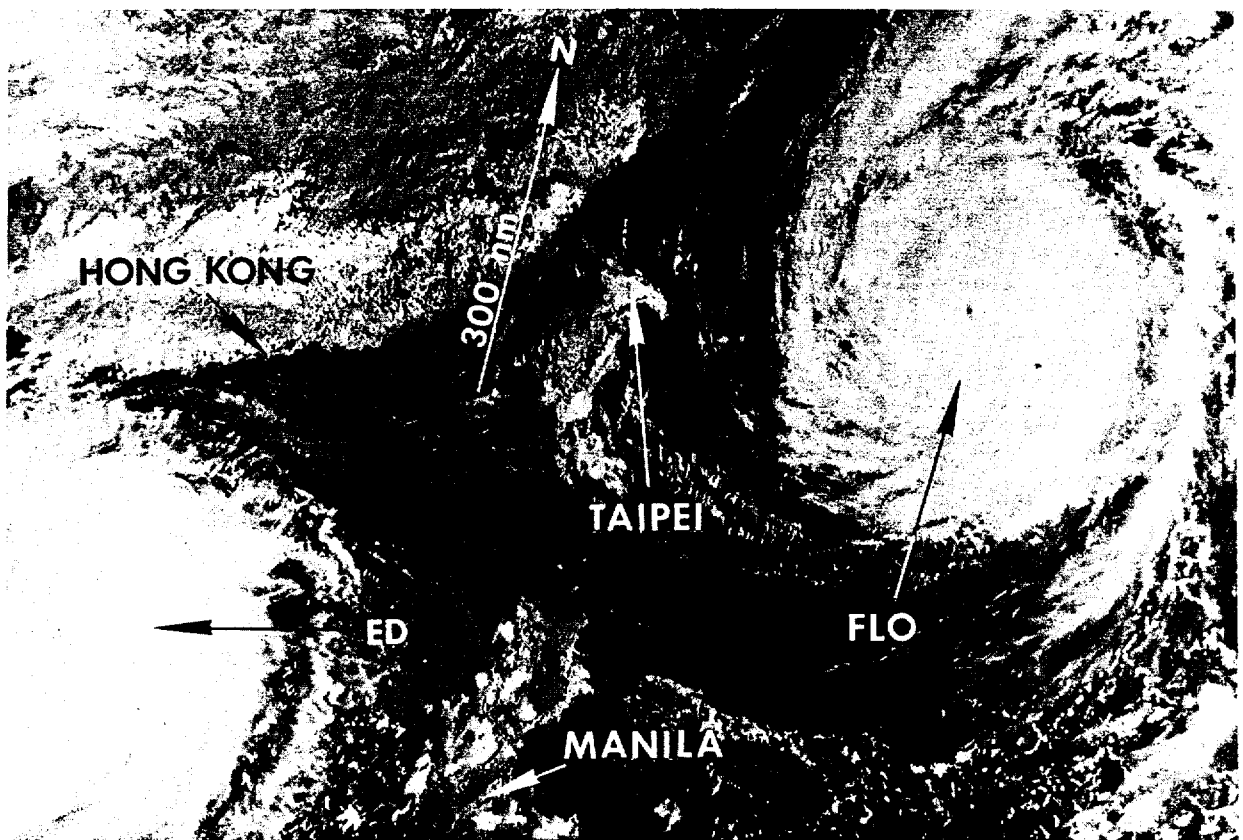


Figure 3-20-3. Flo at maximum intensity is starting to recurve just east of Okinawa while Typhoon Ed (19W) approaches Vietnam (170541Z September DMSP visual imagery).



Figure 3-20-4. Photograph of Flo from the NASA DC-8 reconnaissance aircraft flying near 200 mb on 17 September shows the top and side of the wall cloud (Photo courtesy of Mr. Franz Wen-Ching Yeh, TCM-90 experiment).

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance is shown in Figure 3-20-7. Although the overall 72-hour position error was well below average at 215 nm (395 km), some forecast problems were encountered. Twenty-four hours prior to recurvature, JTWC forecast Flo to make landfall in southern Kyushu whereas the actual landfall was farther east on southern Honshu. A mid-latitude short wave moving off the coast of Asia, deepened more rapidly than anticipated; thus, the forecasts were too far to the west.

VI. IMPACT

Flo passed only 60 nm (110 km) east of Okinawa - close enough to break the drought with 5 to 10 inches (125 to 255 mm) of rain, but just far enough away to spare the island from the most extreme winds near the eye. The maximum wind gusts reported at Naha, Okinawa were 66 kt (35 m/sec). Futenma Marine Corps Air Station and Kadena Air Base reported 64 kt (35 m/sec) and 60 kt (30 m/sec), respectively. The crew of the NASA DC-8 estimated that 100 kt (50 m/sec) winds were just off the east coast of Okinawa. Damage to Okinawa was minor; however, there were news reports that four people died and three were missing in landslides. Flo made landfall on Honshu, 60 nm (110 km) south of Osaka with an intensity of 90 kt (45 m/sec). It was the most powerful typhoon to hit Honshu in 19 years according to news releases. The typhoon brought widespread flooding and caused 115 landslides in Honshu, leaving at least 32 people dead, six missing and 90 people injured. Property and crop damage were estimated in the millions of dollars, and communications, power, and transportation systems were interrupted. A tornado also occurred, injuring 12 people, damaging or destroying 200 homes and other buildings, and downing power lines.



Figure 3-20-5. Photograph of Flo from NASA DC-8 reconnaissance aircraft flying near 200mb on 17 September showing the stratocumulus cloud spirals that define the low-level center (Photo courtesy of Mr. Franz Wen-Ching Yeh, TCM-90 experiment).

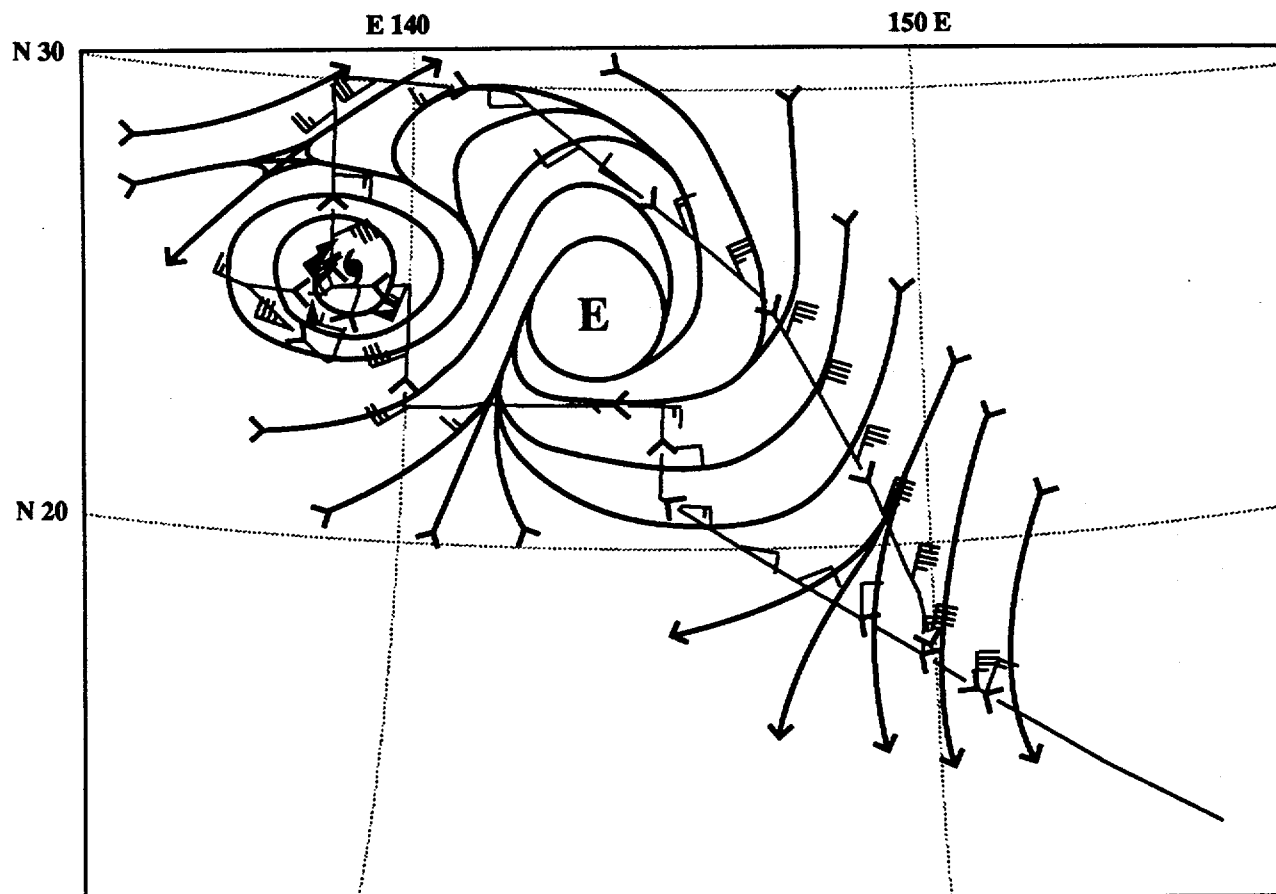


Figure 3-20-6. Flight-level winds reports from aircraft reconnaissance at 37,000-43,000ft for the period 170204Z to 170904Z September show the intense cyclonic circulation near Flo's eye and the anticyclonic eddy to the southeast.

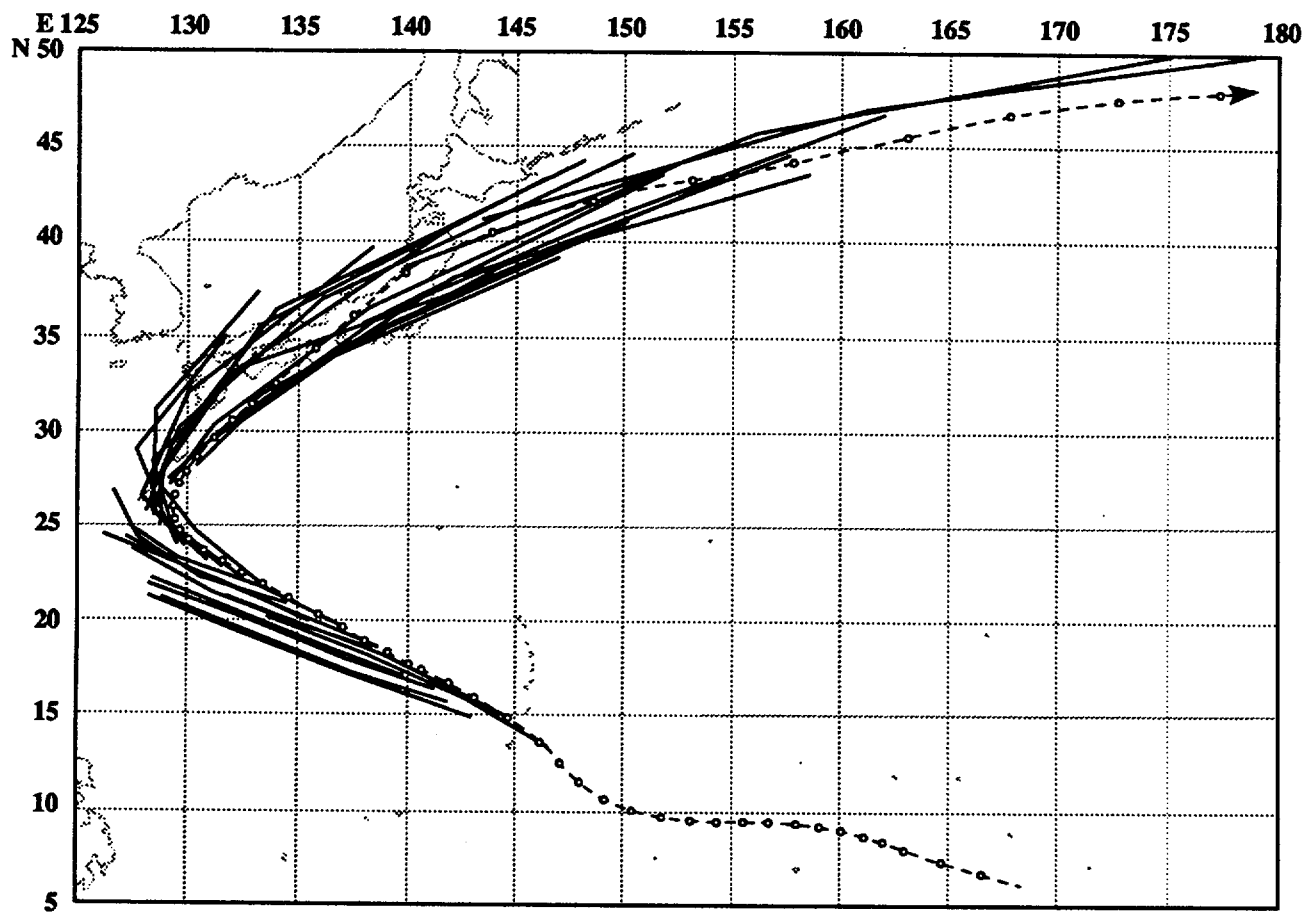
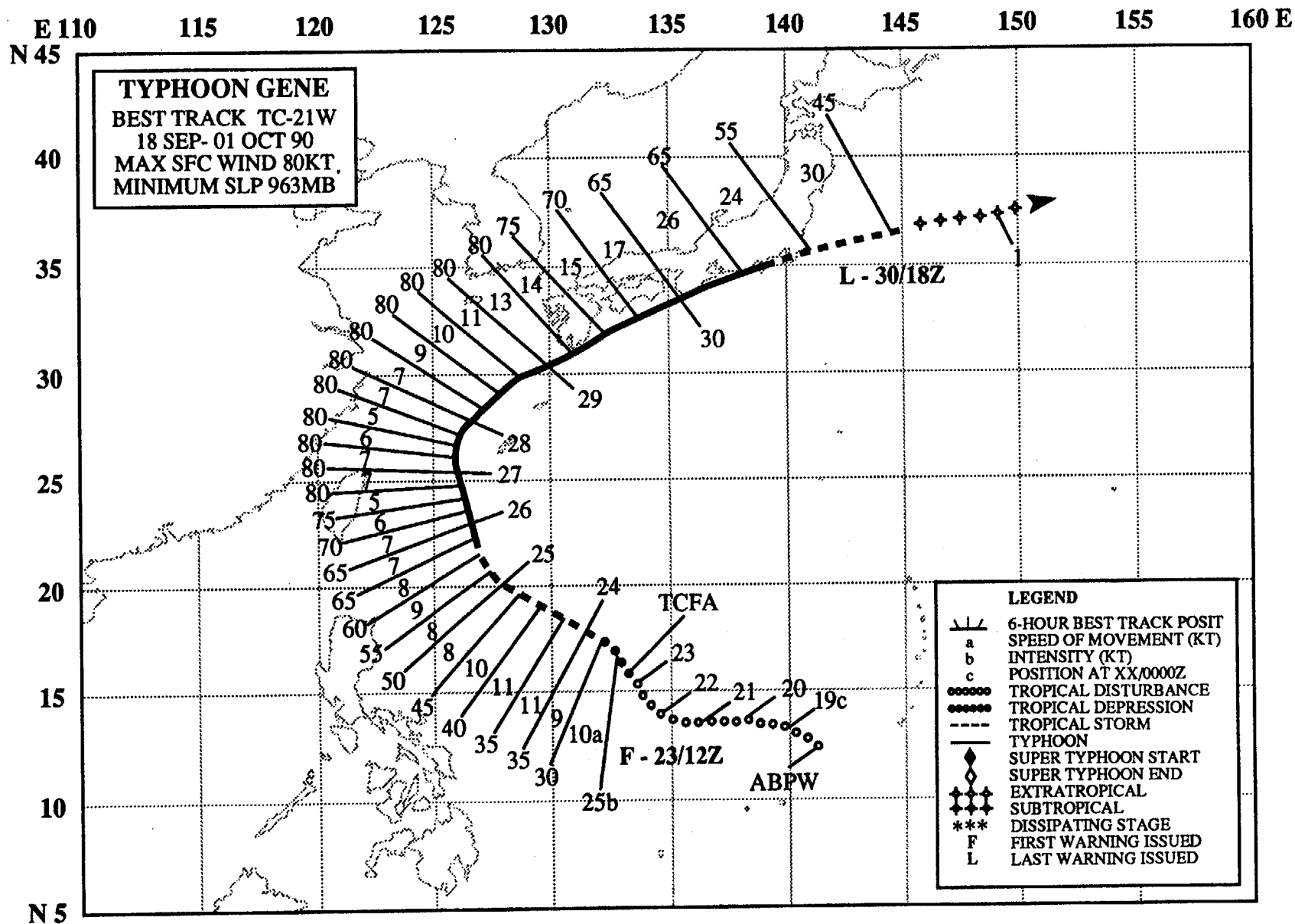


Figure 3-20-7. Summary of JTWC forecasts (solid lines) for Flo is superimposed on the final best track (dashed line).



TYPHOON GENE (21W)

I. HIGHLIGHTS

Gene was the fifth significant tropical cyclone to form in September and the fifteenth of the year to reach typhoon intensity. The initial disturbance formed 250 nm (465 km) west-southwest of Guam and tracked westward for three days before turning northwestward. Gene followed a classic recurvature pattern, passing west of Okinawa and skirting southern Japan. The orientation of Gene's recurvature track resulted in sustained radar contact from 251400Z to 300400Z and an excellent, high quality set of 250 position reports from land radar sites in the islands nearby.

II. CHRONOLOGY OF EVENTS

- 180600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1008 mb.
- 230600Z - First Tropical Cyclone Formation Alert issued in response to increased organization induced by TUTT cell to the west.
- 231200Z - First warning prompted by a Dvorak current intensity estimate of 2.0 and an increase in total convection.
- 240600Z - Upgrade to tropical storm based on improved organization and enhanced outflow.
- 251800Z - Upgrade to typhoon based on a CI 4.0.
- 271200Z - Peak intensity of 80 kt (40 m/sec) maintained until 290600Z.
- 300600Z - Downgraded to tropical storm based on synoptic data.
- 301800Z - Final warning issued due to extratropical transition.

III. TRACK AND MOTION

Gene followed a typical recurvature track. The tropical cyclone initially tracked along the equatorward side of the mid-level subtropical ridge, then turned northwestward to approach a break in the axis in the ridge. Recurvature occurred on 27 October 100 nm (185 km) west of Okinawa in conjunction with a passing short-wave trough. Now under the influence of stronger westerly winds aloft, Gene accelerated east-northeastward and changed into an extratropical low 300 nm (555 km) east of Tokyo.

IV. INTENSITY

For five days Gene's winds remained less than 25 kt (13 m/sec). However, on 23 September, assisted by a TUTT cell to the west, normal intensification of one T-number per day started. Although the track followed the warm Kuroshio ocean current, restricted outflow aloft limited development. Nevertheless, after attaining peak intensity, Gene (Figure 3-21-1) maintained 80 kt (40 m/sec) for two and a half days before slowly weakening.

V. FORECASTING PERFORMANCE

Figure 3-21-2, provides an overview of the forecasts. It illustrates two points: first, when NOGAPS prognoses are slow to weaken the mid-level subtropical ridge in response to the passing short-wave, JTWC's dynamic aids were slow (Figure 3-21-3); and second, if the initial forecast philosophy is for a "straight runner," there is a reluctance to shift to recurvature at the first indication of a change.

VI. IMPACT

There were no reports of damage on Okinawa, but as Gene moved along the southern coastlines of the Kyushu and Honshu, it caused four deaths, 12 injuries, and localized flooding. Wind speeds of 70 kt (36 m/sec) were measured on Kyushu, but weakened to 38 kt (20 m/sec) as Gene brushed by Tokyo.

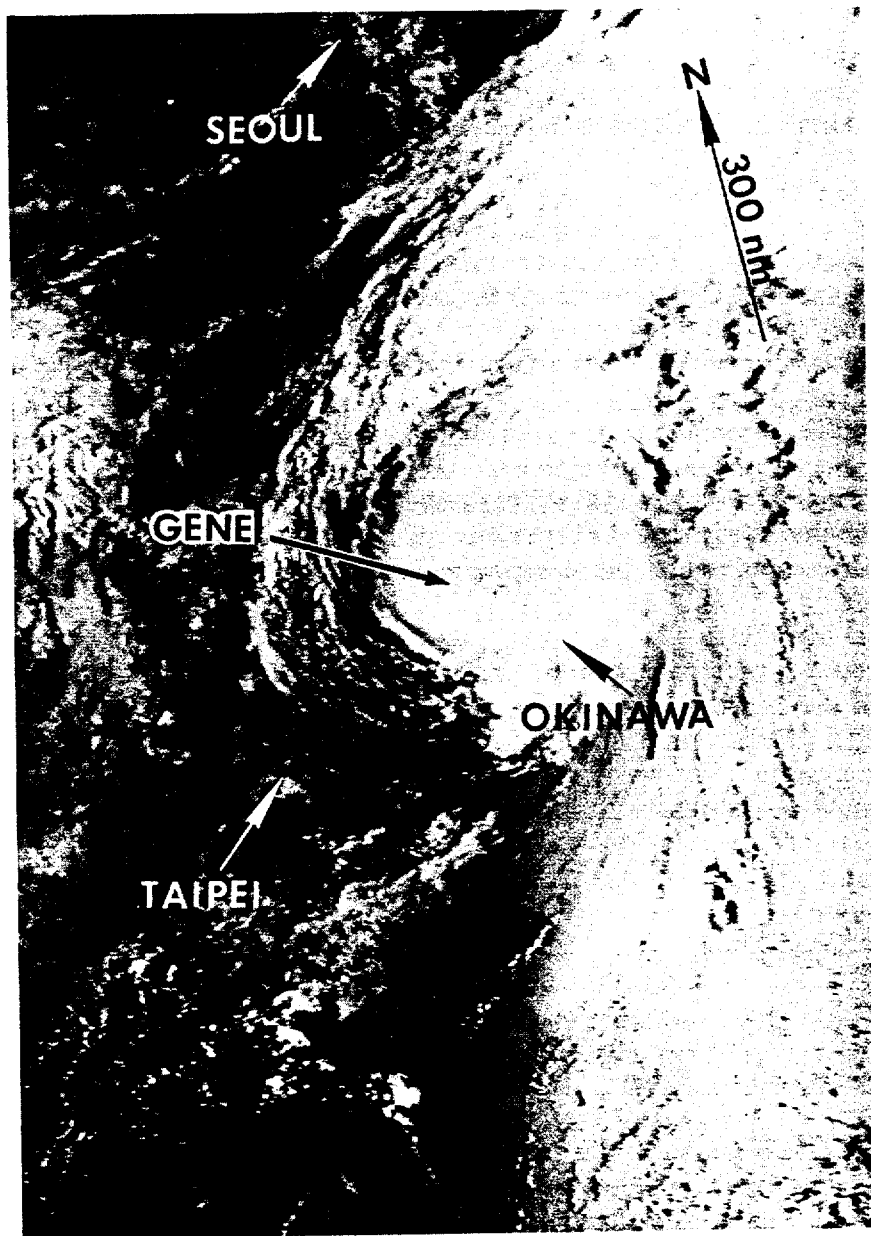


Figure 3-21-1. Gene at peak intensity (272346Z September NOAA visual imagery).

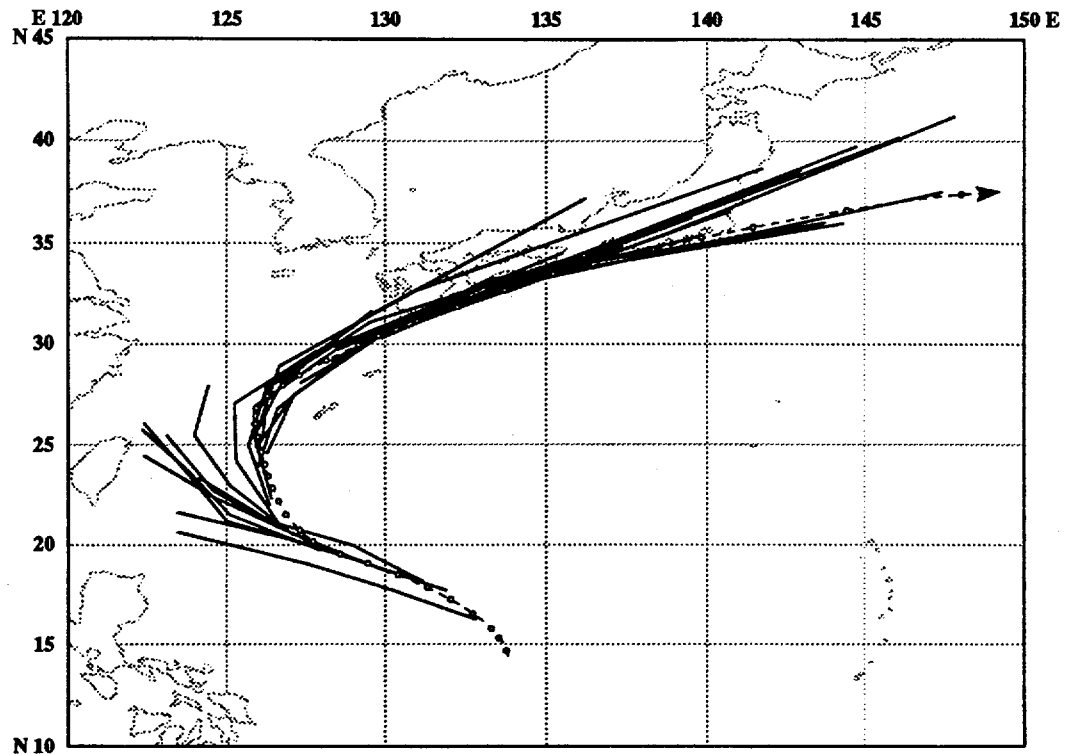


Figure 3-21-2. Summary of JTWC forecasts (solid lines) for Gene superimposed on the final best track (dashed line).

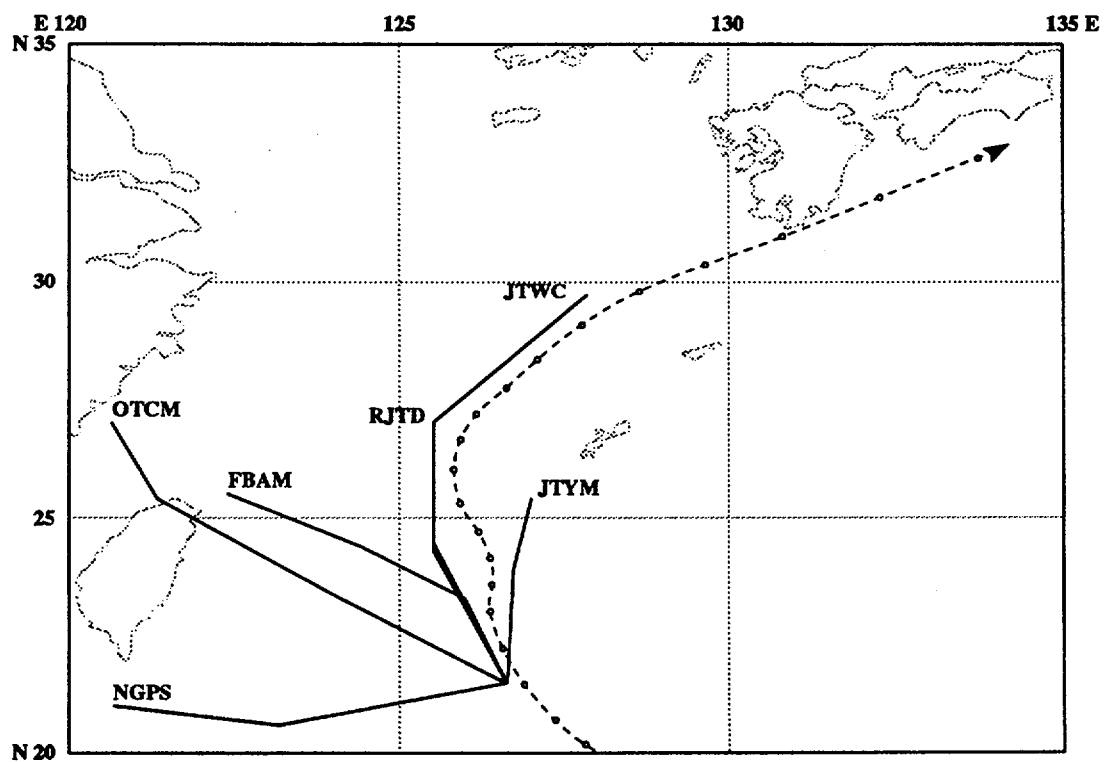
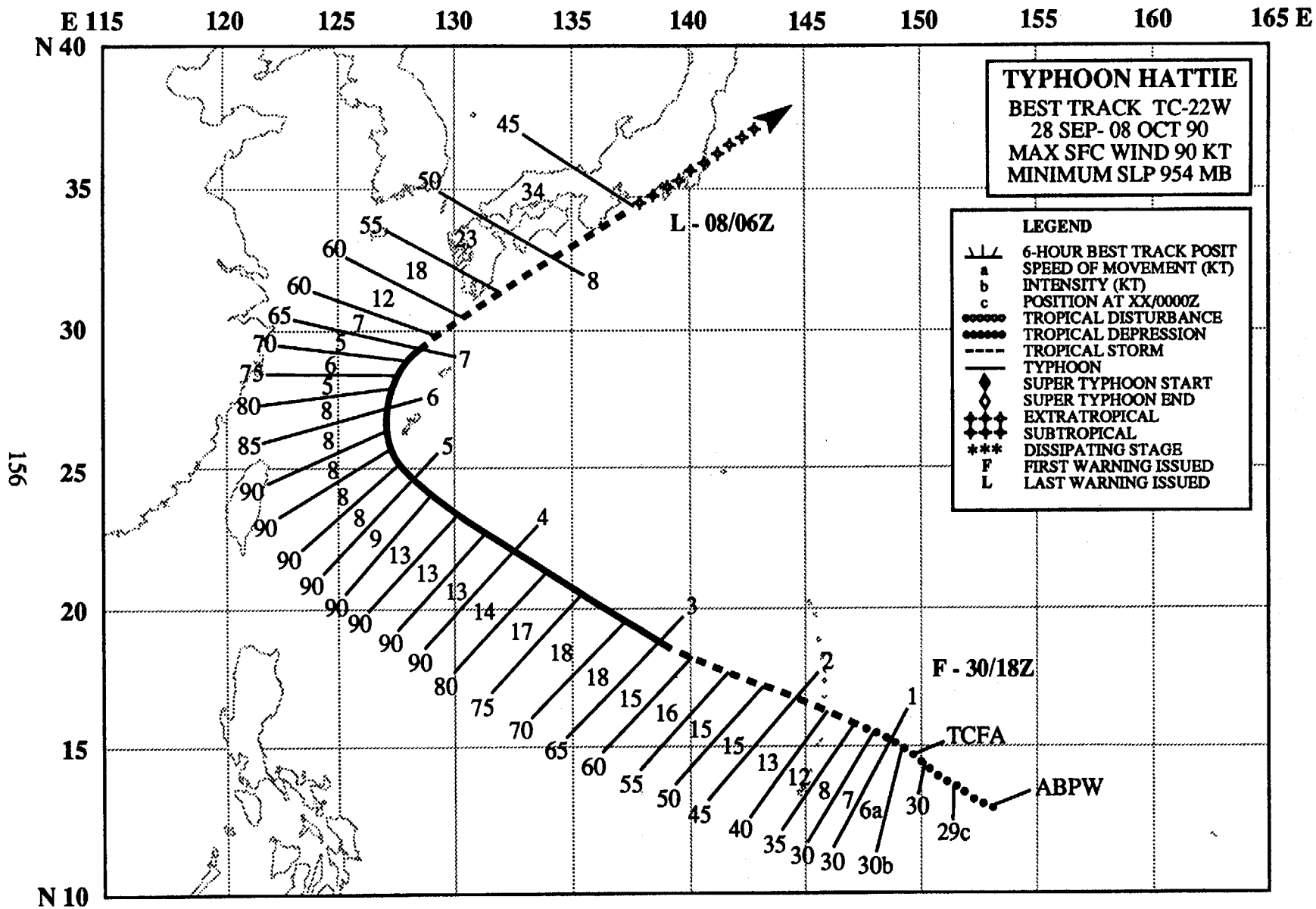


Figure 3-21-3. Comparison of 251200Z September forecasts by JTWC and supporting dynamic aids NGPS, FBAM and OTCM. Also shown are forecasts by the Japan Meteorological Agency (RJTD) and the Japanese Typhoon Model (JTYM).



TYPHOON HATTIE (22W)

I. HIGHLIGHTS

Hattie, the last of six tropical cyclones to form in September, was the fourth tropical cyclone in a six-week period to affect Okinawa and southern Japan. Its track was a classic example of recurvature.

II. CHRONOLOGY OF EVENTS

- 280600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1010 mb.
- 300730Z - Tropical Cyclone Formation Alert followed a flare up of deep convection and the first Dvorak intensity estimate of 1.0.
- 301800Z - First warning based on intensification manifested by the cirrus outflow layer showing signs of rapid growth, appearance of overshooting cumulonimbus tops, and a CI 2.0.
- 011200Z - Upgraded to tropical storm due to increase in convective extent and organization.
- 030000Z - Upgrade to typhoon based on the formation of a small eye on satellite imagery and a CI 4.0.
- 040000Z - Peak intensity - 90 kt (46 m/sec) - based on a decrease in eyewall cloud top temperatures and a CI 5.0.
- 071200Z - Downgraded to a tropical storm due to increased upper-level westerly wind shear and loss of central convection.
- 080600Z - Final warning issued following Hattie's transformation into an extratropical cyclone.

III. TRACK AND MOTION

As Typhoon Gene (21W), which had just recurved, accelerated towards the main islands of Japan on 28 September, Hattie formed in the monsoon trough 100 nm (185 km) east of Guam. Hattie followed a smooth track west-northwestward, slowed late on 4 October as it approached the lighter winds near the axis of the subtropical ridge, and recurved just to the west of Okinawa late on 5 October. Then on 7 October, Hattie accelerated northeastward in the strong southwesterly flow and churned by Tokyo, Japan on 8 October.

IV. INTENSITY

For a three-day period, 29 September to 1 October, Hattie's intensification was arrested by westerly winds aloft and to the north. On 1 October, anticyclonically curved cirrus outflow was observed to push northward from the depression's cloud system center (Figure 3-22-1), and Hattie began to intensify at a normal rate of one T-number per day. This steady intensification continued until 4 October, when the typhoon peaked at 90 kt (45 m/sec)(Figure 3-22-2). Hattie maintained its peak intensity for almost two days before moving into the strong vertical shear region north of the subtropical ridge axis. The typhoon weakened as it lost central convection and transitioned to an extratropical system on 8 October.

V. FORECASTING PERFORMANCE

The first three warnings issued by JTWC were based on Hattie's poorly defined cloud system center and verified significantly south of track (Fig 3-22-3). Then warning 04 relocated Hattie's center to the north as the convection consolidated the low-level circulation. All subsequent track forecasts verified well. In particular, three consecutive warnings beginning 48 hours prior to the recurvature point achieved exceptionally low 72-hour forecast errors near 90 nm (165 km).

VI. IMPACT

Typhoon Hattie passed 30 nm (55 km) west of Okinawa, causing damage in excess of \$1.7 million to U.S. military bases. Roof damage and beach erosion were extensive. Maximum wind gusts as high as 75 kt (38 m/sec) were recorded on Okinawa. On a positive note, the water rationing in since mid-September was lifted. Total rainfall from Flo (20W), Gene (21W) and Hattie provided 15-20 inches (380-510 mm) to fill up the almost empty reservoirs.

After Hattie recurved, it tracked along the south coast of Japan, bringing heavy rains and strong winds. Three people in Shikoku were killed and 14 injured as the bus they were riding in was struck by a landslide.

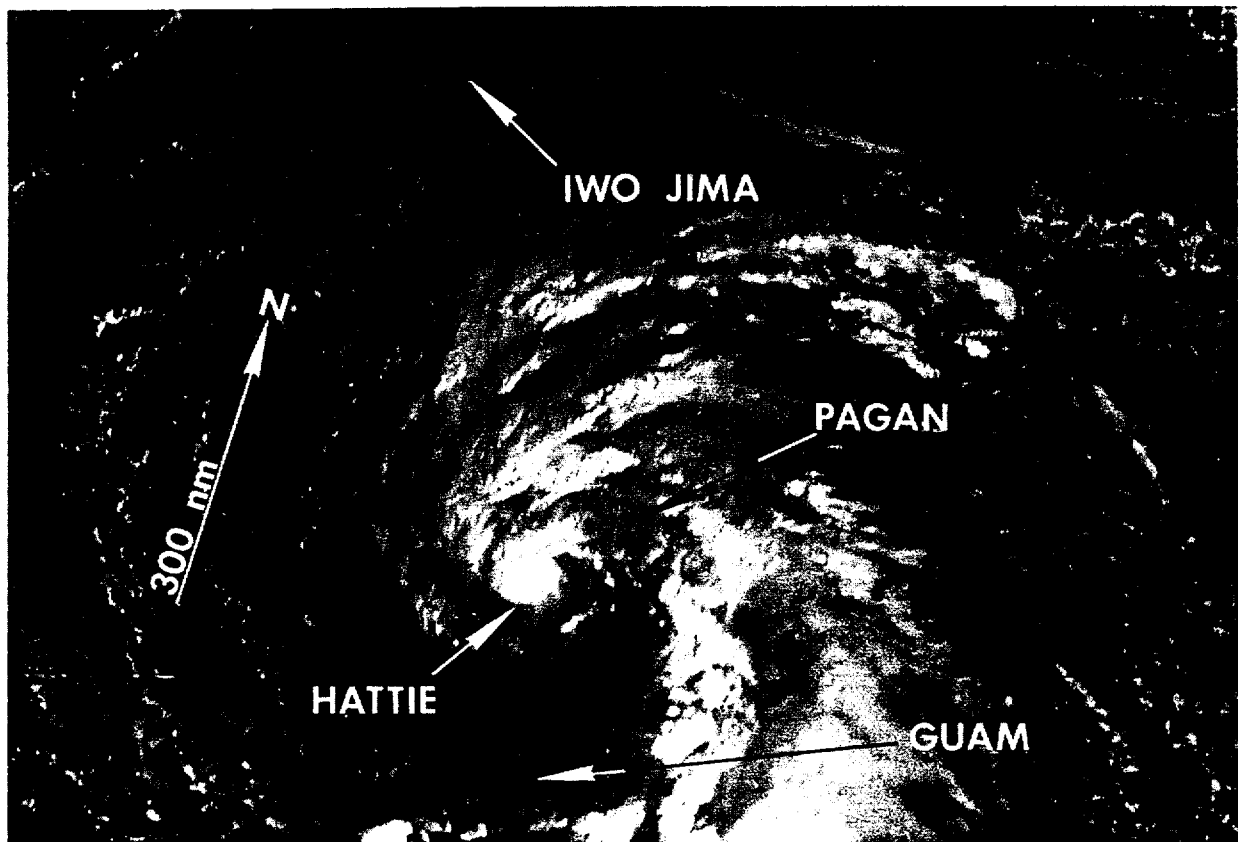


Figure 3-22-1. Hattie intensifies and its cirrus outflow pushes northward (012330Z October DMSP visual imagery).

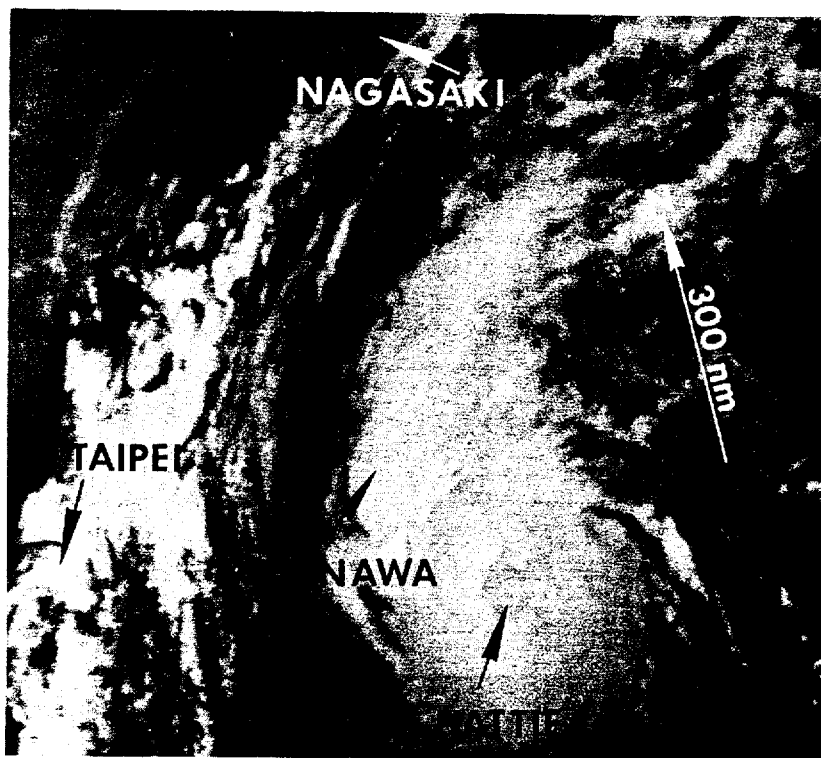


Figure 3-22-2. Typhoon Hattie at peak intensity (041307 October DMSP nighttime visual imagery).

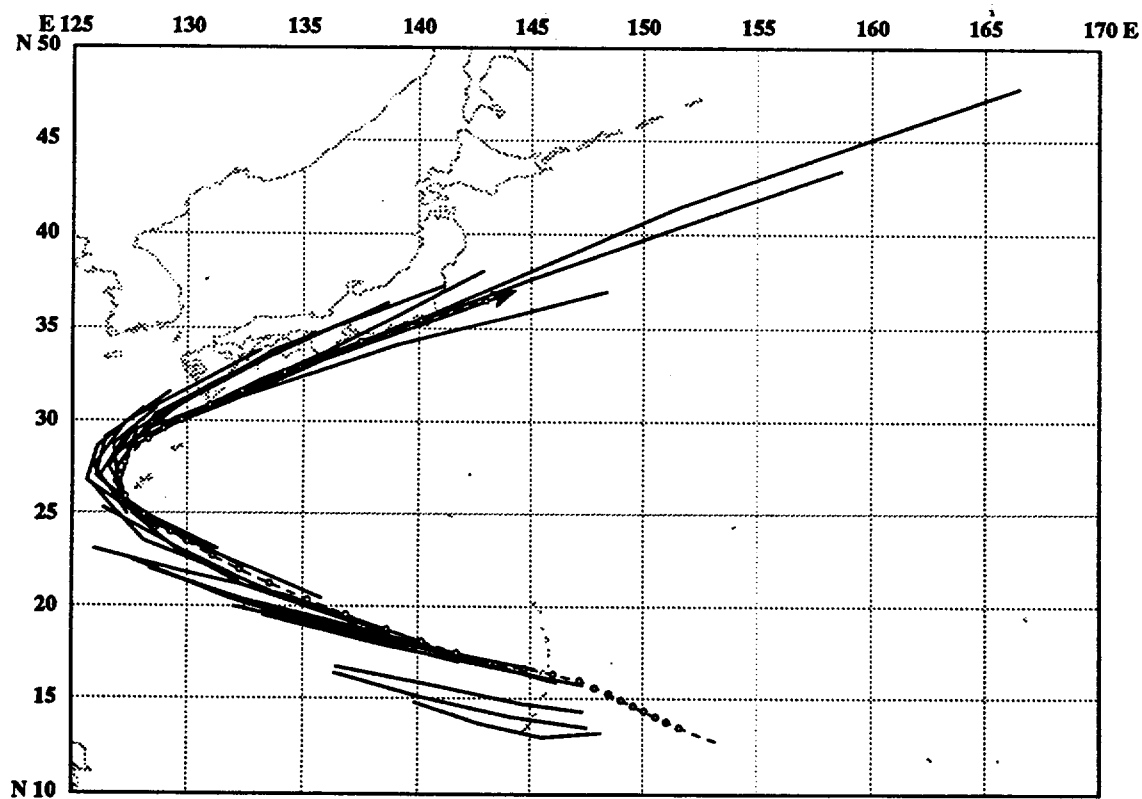
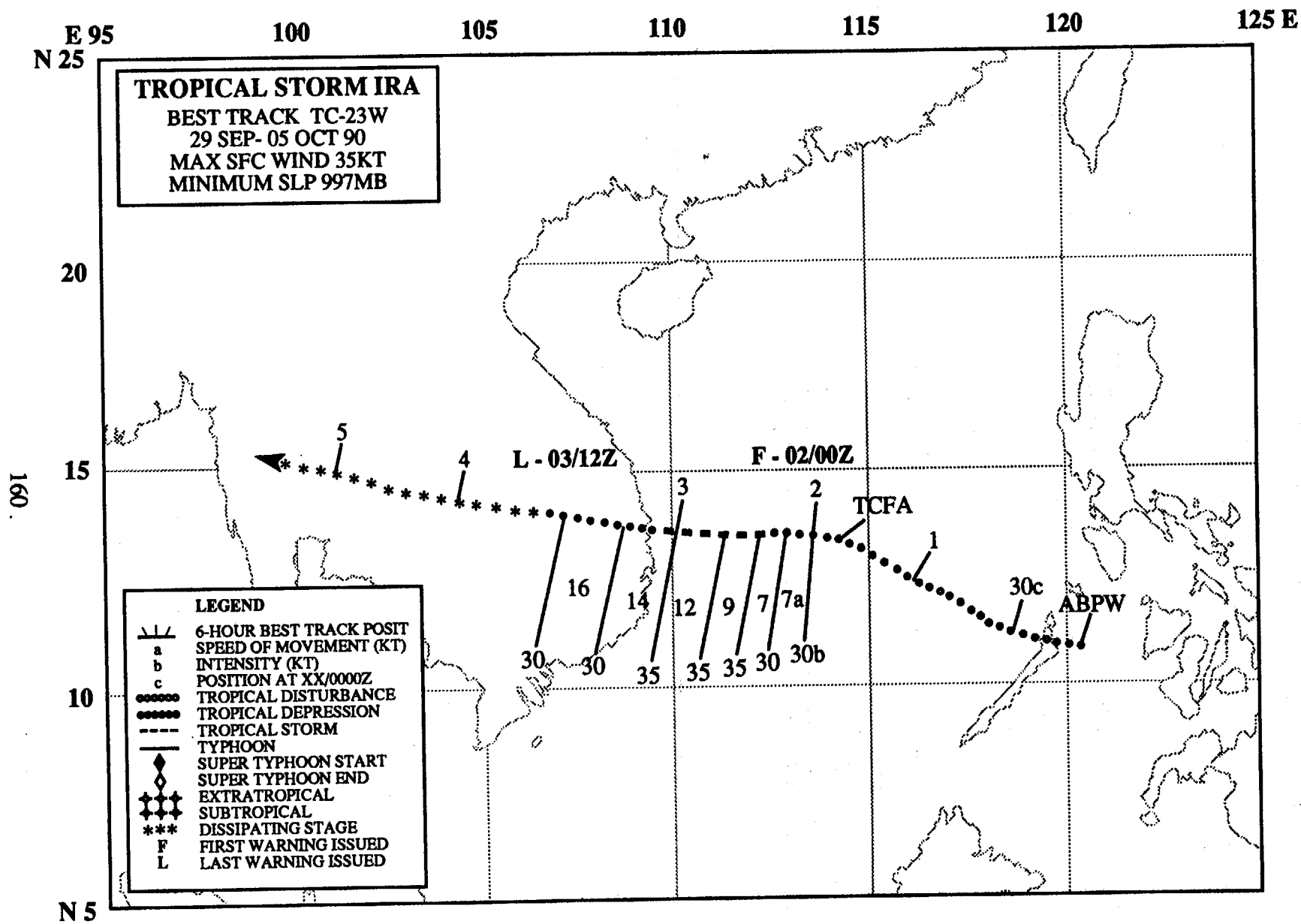


Figure 3-22-3. Summary of JTWC forecasts (solid lines) for Hattie is superimposed on the final best track (dashed line).



TROPICAL STORM IRA (23W)

I. HIGHLIGHTS

Ira, the eighth tropical cyclone to hit Vietnam in 1990 and the last in a series of weak, highly sheared tropical systems in the South China Sea, formed in a broad area of convection near Palawan Island. The convective cloud mass tracked steadily westward in the deep easterly flow and made landfall at Qui Nhon Vietnam on the third of October.

II. CHRONOLOGY OF EVENTS

- 290600Z - First mentioned on the Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1009 mb.
- 011800Z - Tropical Cyclone Formation Alert based on 12 hours of persistent cirrus outflow, the consolidation of the central convection and first CI 1.0 estimate.
- 020000Z - First warning issued due to increased deep central convection and upper-level outflow.
- 021200Z - Upgraded to tropical storm because of a continued increase in central convection.
- 030600Z - Downgraded to tropical depression based on synoptic reports along the Vietnamese coast.
- 031200Z - Final warning issued due to land interaction, and severing of the low-level overwater moisture source.

III. TRACK AND MOTION

Ira developed in the monsoon trough near the southern Philippine island of Palawan on 29 September and tracked steadily westward on the south side of a persistent mid-level ridge centered over southern China (Figure 3-23-1). On 2 October, as the tropical cyclone approached the coast of Vietnam, increased mid- and low-level easterly flow accelerated Ira on shore over Vietnam.

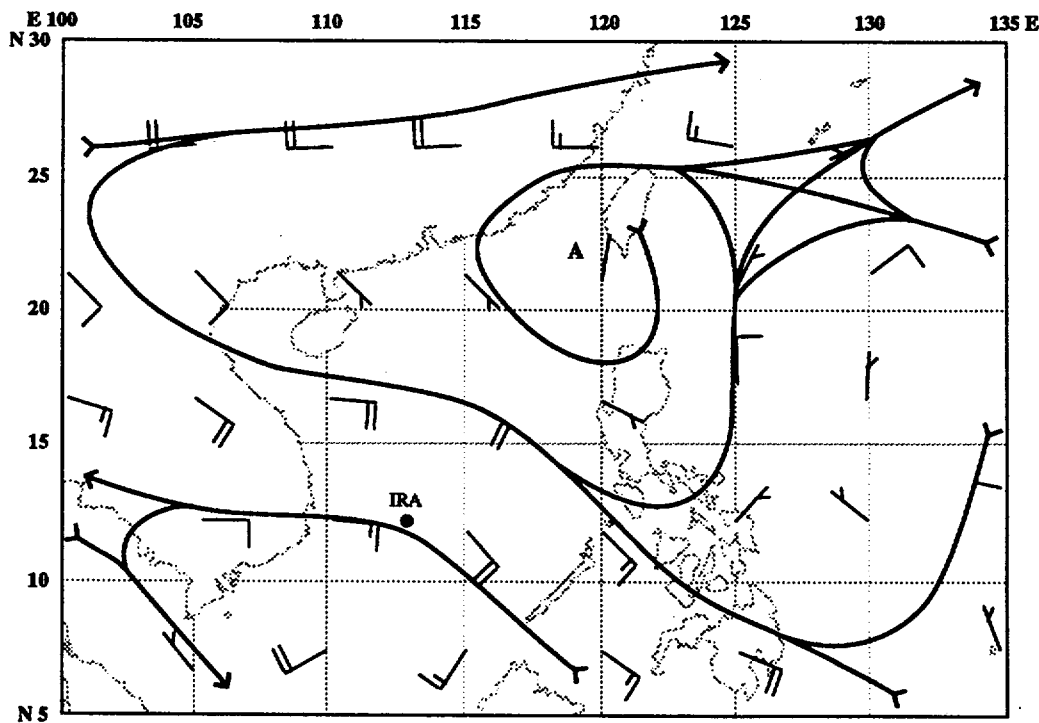


Figure 3-23-1. 500-mb NOGAPS analysis from 020000Z October, showing a persistent mid-level ridge positioned along the coast of southern China.

IV. INTENSITY

Ira's convective cloud mass was poorly organized throughout the life of the storm. The cloud system was embedded in an unfavorable environment of strong upper-level unidirectional southeasterly flow (Figure 3-23-2). Therefore, Ira was unable to develop an efficient outflow pattern during its early stages of development (Figure 3-23-3). Later, after the tropical cyclone moved over land on 3 October, its remnants tracked westward across virtually all of Indochina. There was even brief mention on the Significant Tropical Weather Advisory of a possible regeneration if the remnants moved into the Bay of Bengal.

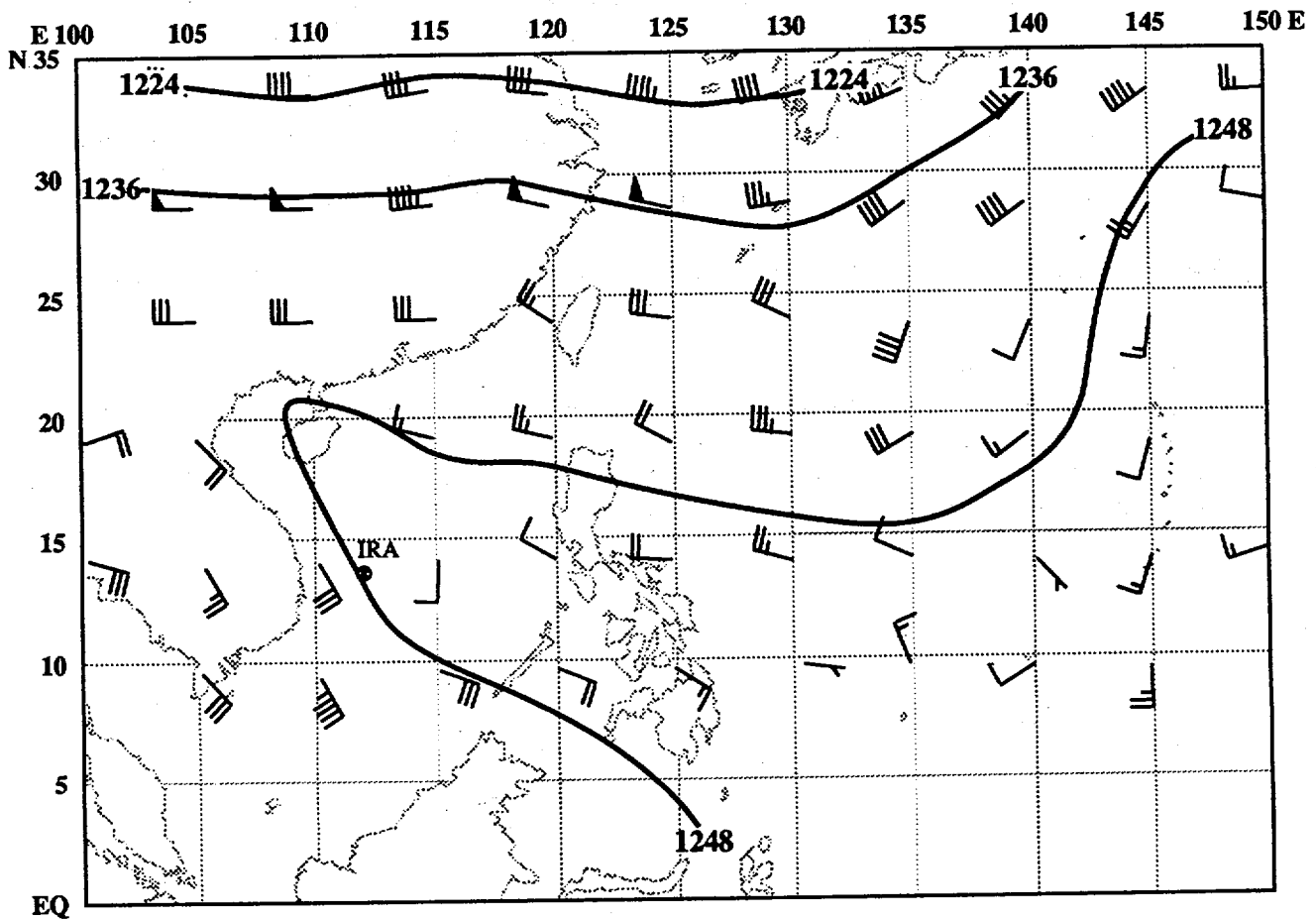


Figure 3-23-2. The 200-mb NOGAPS analysis with heights in decameters at 020000Z October, showing strong unidirectional southeasterly flow over the South China Sea which restricted the development of an efficient upper-level outflow pattern above Ira.

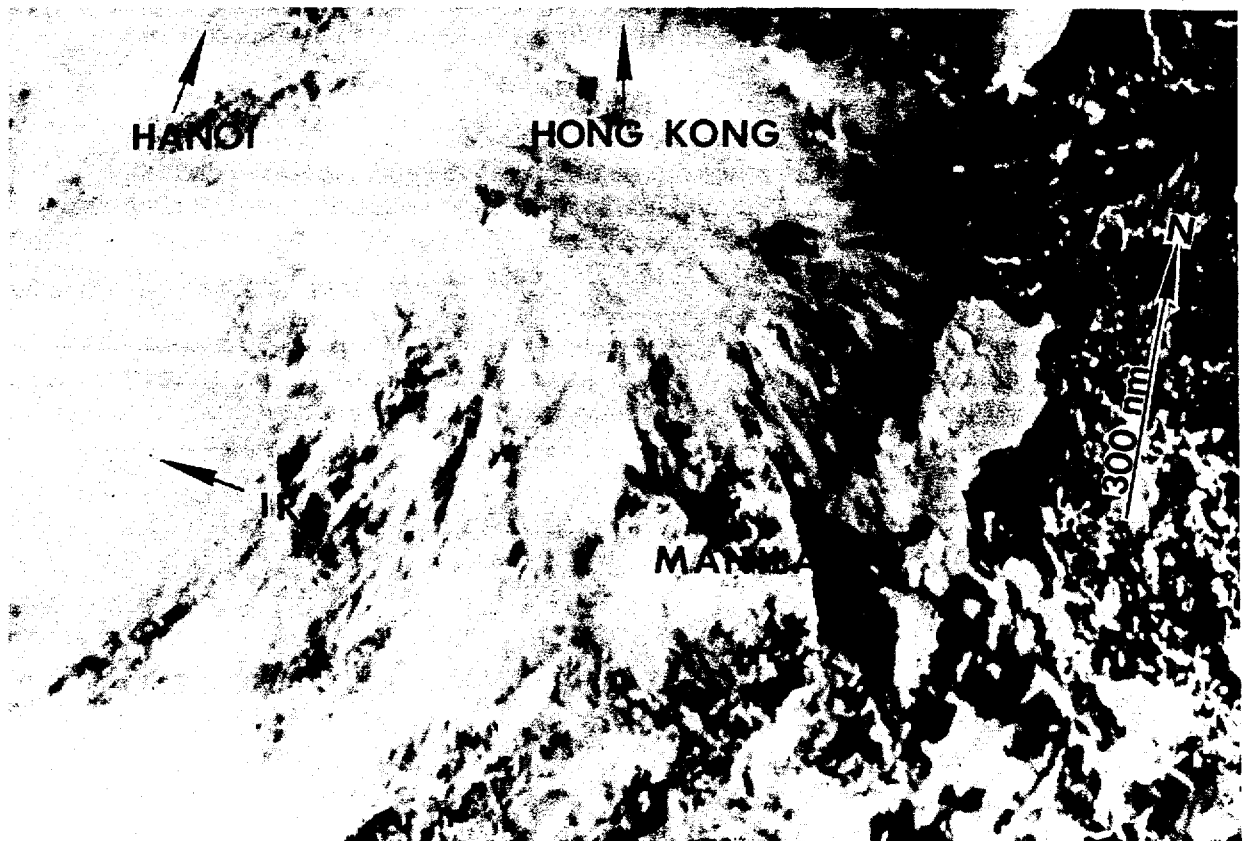


Figure 3-23-3. Ira approaches the coast of Vietnam (030607Z October NOAA visual imagery).

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance is shown in Figure 3-23-4. The only major problems encountered were the result of conflicts between synoptic reports and satellite fixes. JTWC biased the initial warning positions toward the synoptic reports after the second warning, and the accuracy of the initial motion forecasts quickly improved.

VI. IMPACT

The following report was released by the United Press International in Bangkok,

A typhoon damaged 110,000 houses and killed seven people in the coastal provinces of Vietnam, official Radio Hanoi reported.

The radio, in a broadcast Monday, said the Central Flood and Typhoon Control Committee has reported that the eighth typhoon to hit Vietnam this year caused heavy rainfall in Thua Thien-Hue Province, 320 miles south of Hanoi.

"The average rainfall was between 12 to 27.5 inches," the broadcast said, according to a translation made available Wednesday.

"Heavy rains submerged 110 of the 145 villages and more than 110,000 houses, (and) killed seven people," the radio said.

The radio did not give the exact date when the storm hit the country.

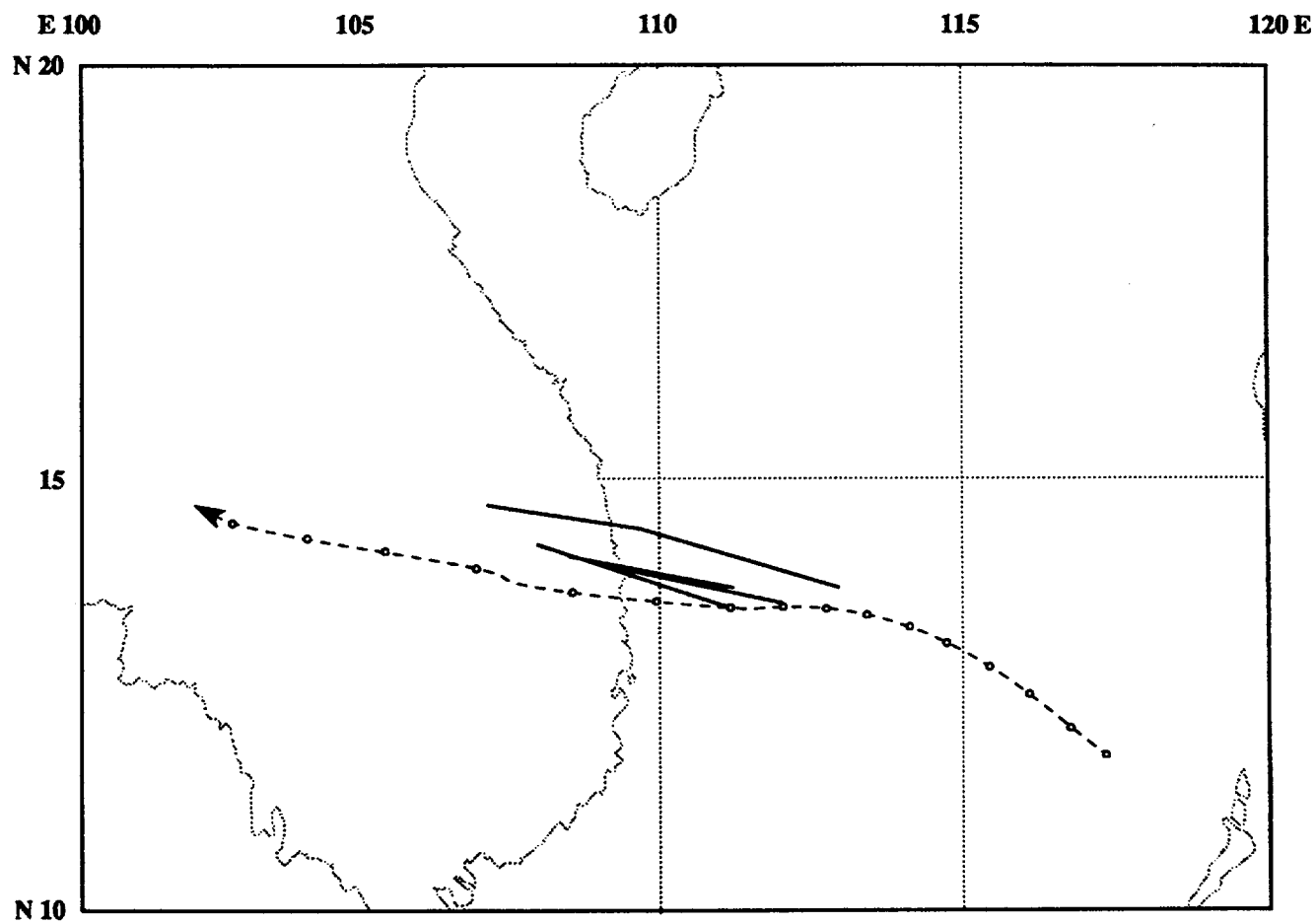
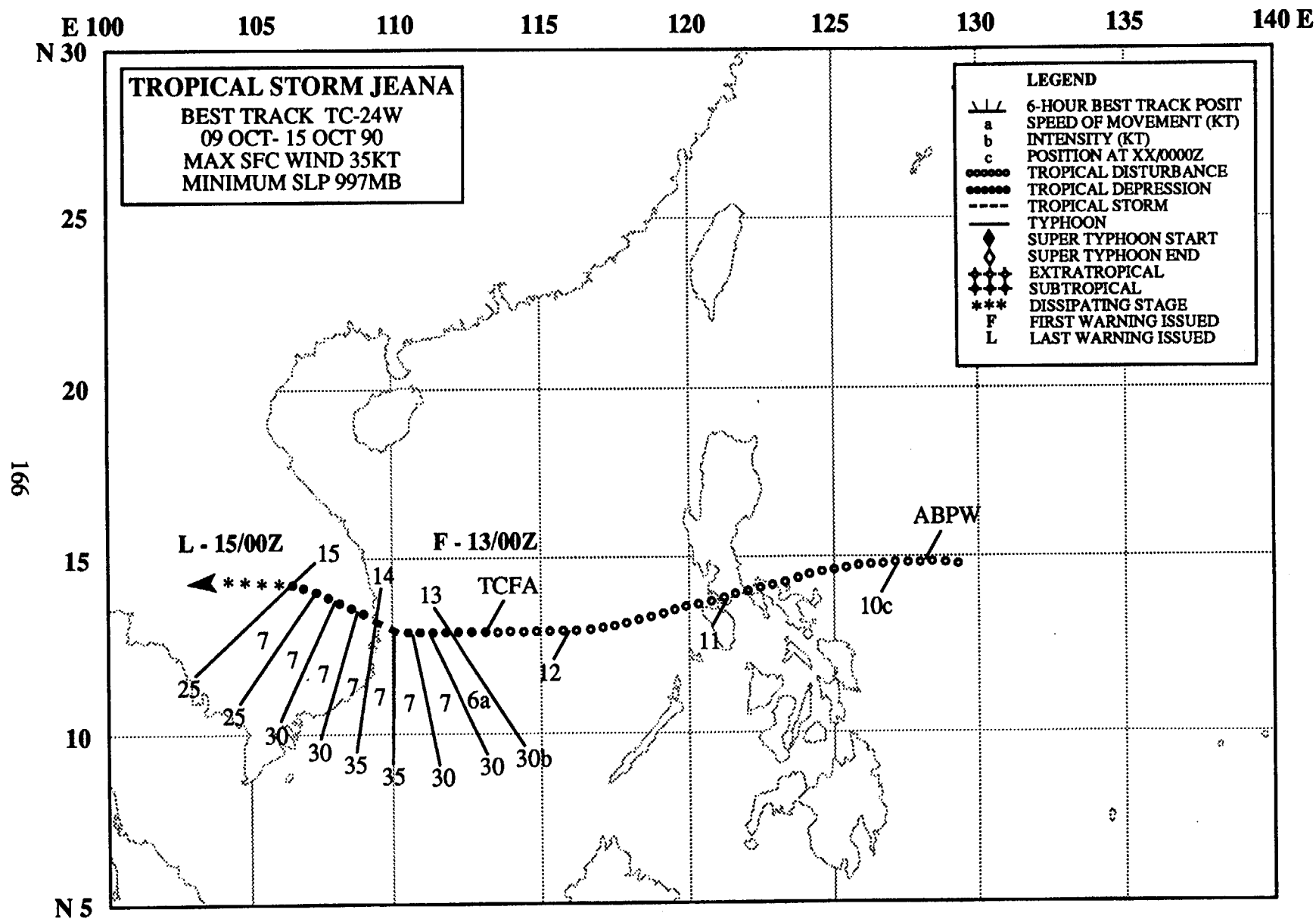


Figure 3-24-4. Summary of JTWC forecasts (solid lines) for Ira superimposed on the final best track (dashed line).



TROPICAL STORM JEANA (24W)

I. HIGHLIGHTS

Jeana, the second of four tropical cyclones to form in October, was the fifth to churn across the South China Sea in 1990. This minimal tropical storm proved as difficult to estimate intensity for, as it was to position.

II. CHRONOLOGY OF EVENTS

- 091900Z - First mentioned on the Significant Tropical Weather Advisory as a weak cyclonic circulation associated with a shear zone in the Philippine Sea.
- 121330Z - Tropical Cyclone Formation Alert issued based on the presence of a 1005 mb minimum sea-level pressure, a well defined surface cyclone, brisk northeasterly trade winds of 25-30 kt (13-15 m/sec) extending 200-400 nm (350-750 km) to the north, and a CI 1.0.
- 130000Z - First warning issued due to continued development and a CI 2.0.
- 141200Z - Upgrade to tropical storm based on synoptic ship reports of 35 to 40 kt (17-20 m/sec) and an CI 2.5.
- 141800Z - Downgraded to a tropical depression after moving over land.
- 150000Z - Final warning issued based on Jeana's movement further inland and dissipation.

III. TRACK AND MOTION

On 9 October, Jeana developed in the Philippine Sea 540 nm (1000 km) southeast of Manila. For the next six days, the tropical cyclone tracked westward, south of a narrow subtropical ridge that extended across southern China and eastward along 25° north latitude. On 13 and 14 October, because of the poorly defined cloud system center (Figure 3-24-1), the exact location of Jeana as it approached the coast of Vietnam was difficult to fix. Synoptic and satellite fixes differed, which resulted in relocations for warnings number 02 and 03 and holding Jeana quasi-stationary for number 04. Only after the 15 October data became available and the low-level circulation located well inland in Laos, was the final best track constructed.

IV. INTENSITY

On 12 October intensification started in the South China Sea. An extensive area of peripheral northeasterly gales developed to the north of Jeana because of the tightening surface pressure gradient. On 14 October, ships reported 35-40 kt (17-20 m/sec) southerly winds 80-100 nm (150-185 km) east of Jeana's apparent center. It appeared that Jeana had most likely intensified into an tropical storm just before landfall on 14 October and the gale force winds remained overwater as the center of the circulation moved inland.

V. FORECASTING PERFORMANCE

The JTWC forecasts are superimposed on the final best track in Figure 3-24-2. Jeana's westward direction was correctly forecast.

VI. IMPACT

None reported.

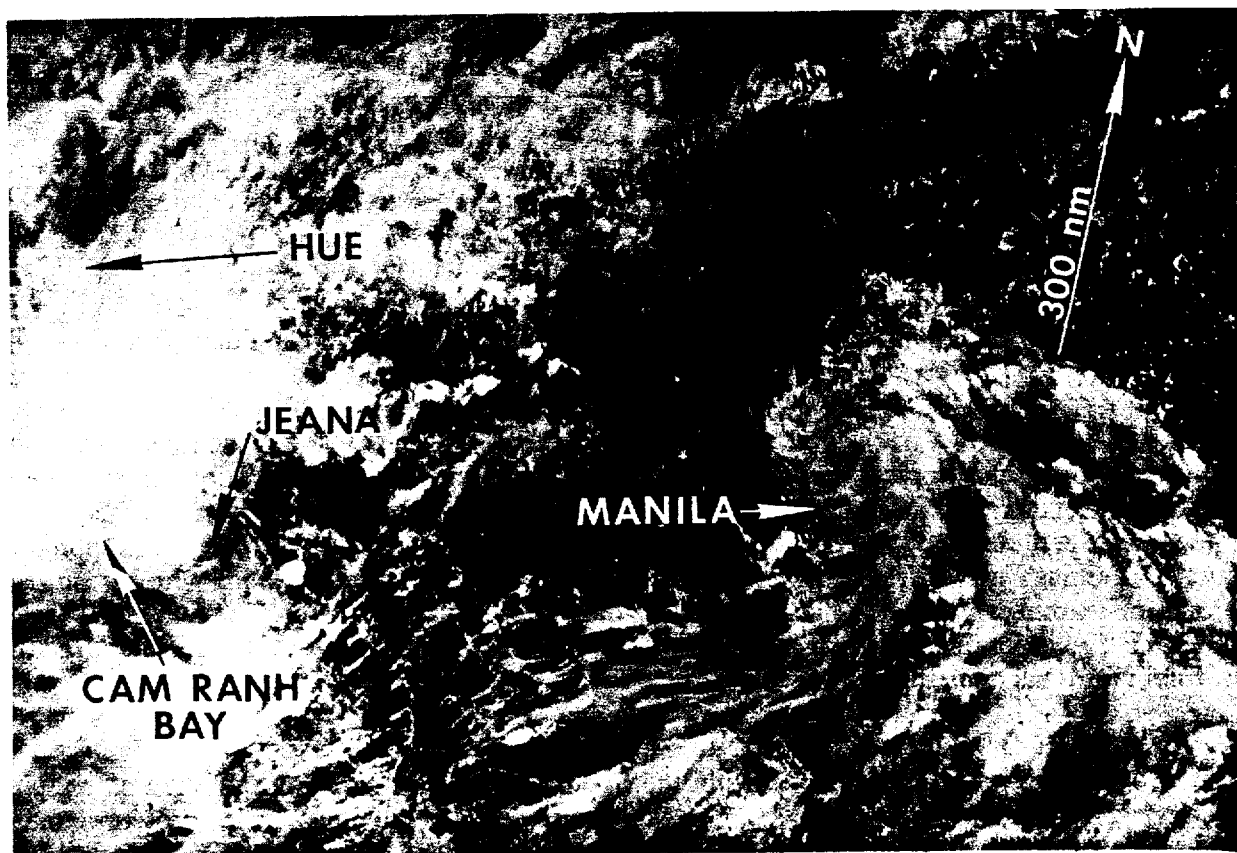


Figure 3-24-1. The partially exposed low-level circulation center of Tropical Depression 24W (Jeana) 12 hours prior to maximum intensity (130557Z October NOAA visual imagery).

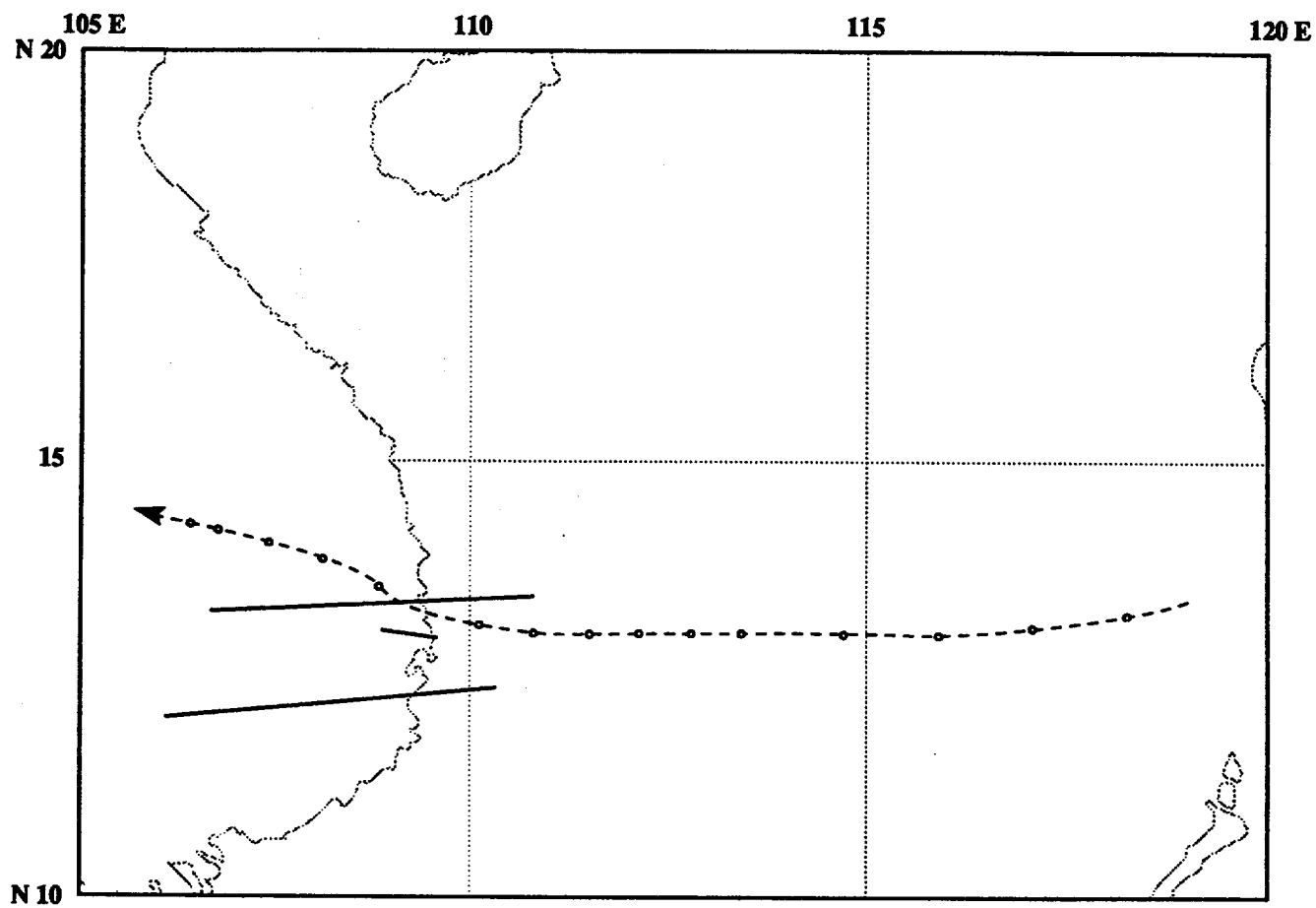
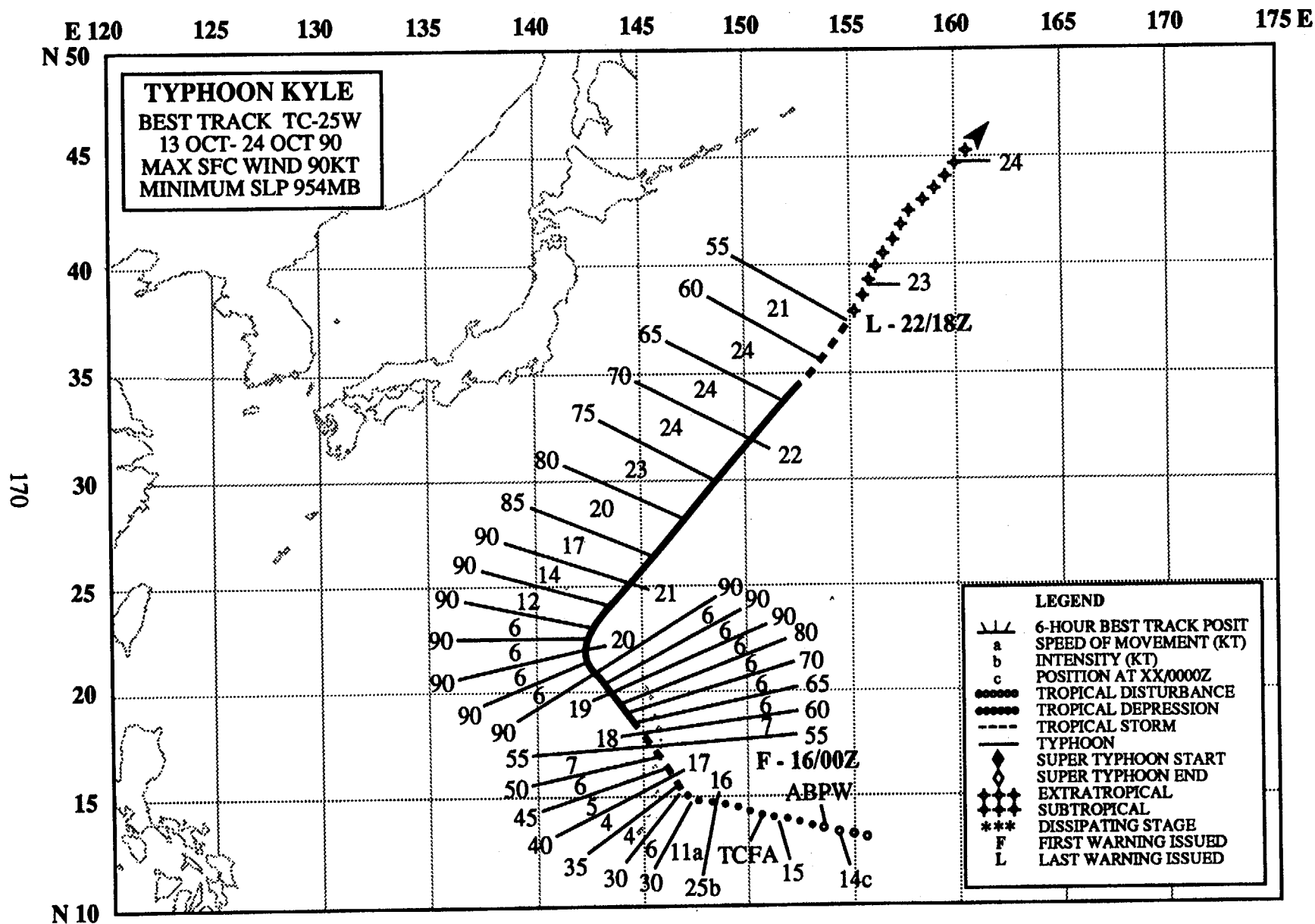


Figure 3-24-2. Summary of JTWC forecasts (solid lines) for Jeana superimposed on the final best track (dashed line).



TYPHOON KYLE (25W)

I. HIGHLIGHTS

Kyle generated from a disturbance in the monsoon trough 600 nm (1110 km) east of Guam. Separating from the trough, the cloud system gained organization and began to track along the southern edge of the subtropical ridge to its northeast. The subtropical ridge and a series of fast moving mid-latitude short-wave troughs strongly influenced Kyle's track. The tropical cyclone passed through the northern Mariana Islands, causing minimal damage, intensified into a typhoon, and recurved several hundred miles east of Japan.

II. CHRONOLOGY OF EVENTS

140600Z - First mentioned on the Significant Tropical Weather Advisory as having fair potential for development due to its favorable location east of a TUTT cyclone.

150500Z - Tropical Cyclone Formation Alert issued based on first Dvorak intensity estimate of CI 1.0.

160000Z - First warning issued following intensity estimate of CI 1.5.

161800Z - Upgraded to tropical storm due to Dvorak current intensity estimate of 2.5.

180600Z - Upgrade to typhoon based on a CI 4.0 and weaker vertical shear.

190000Z - Peak intensity - 90 kt (45 m/sec) - followed on Dvorak current intensity of 5.0.

221200Z - Downgraded to tropical storm after the eye disappeared and interaction began with mid-latitude trough to the north.

221800Z - Final warning issued as Kyle underwent extratropical transition.

III. TRACK AND MOTION

Kyle formed at the eastern end of the monsoon trough. As the circulation consolidated and separated from the trough, it began to track west-northwestward under the influence of the mid-tropospheric subtropical ridge. On 16 October Kyle was headed directly towards the island of Saipan, which is located 100 nm (185 km) north-northeast of Guam. A mid-latitude short-wave trough approaching from Asia weakened the subtropical ridge and caused the tropical cyclone to slow and turn northwestward over the northern Marianas. The tropical cyclone continued to track northwestward along the western edge of the ridge and recurved on 20 October. Kyle maintained its tropical characteristics until extratropical transition occurred on 22 October.

IV. INTENSITY

Until 16 October, Kyle encountered upper-level wind shear which restricted its outflow to the west. Then the vertical wind shear lessened, Kyle intensified and interlocking cloud bands formed. A small eye was briefly observed on the 18 October satellite images, but disappeared into the ragged central dense overcast as the short wave approached from the northwest. Twenty-four hours later, after the short wave exited the area, the eye reappeared. As Kyle moved into higher latitudes, its eye became elongated due to pressure from increasing westerly winds aloft. Gradual weakening accompanied this interaction (Figure 3-25-1) and Kyle became extratropical on 22 October.

V. FORECASTING PERFORMANCE

JTWC beat all the objective aids with overall errors of 98 nm (181 km), 166 nm (307 km), and 196 nm (363 km) at 24, 48 and 72 hours, respectively. However, JTWC forecasters missed the turning point to the northwest on 16 October (Figure 3-25-2). The half persistence/half climatology model, HPAC, suggested a sharper turn than that predicted by the dynamical models, OTCM and FBAM. JTWC relied on the guidance from OTCM, since the ridge was not anticipated to weaken as drastically as it did. Later, JTWC forecasters accurately predicted the time and point of recurvature on 20 October,

achieving 72-hour forecast errors of less than 100 nm (185 km) for 3 consecutive warnings, beginning 60 hours prior to the event.

VI. IMPACT

No information was received.

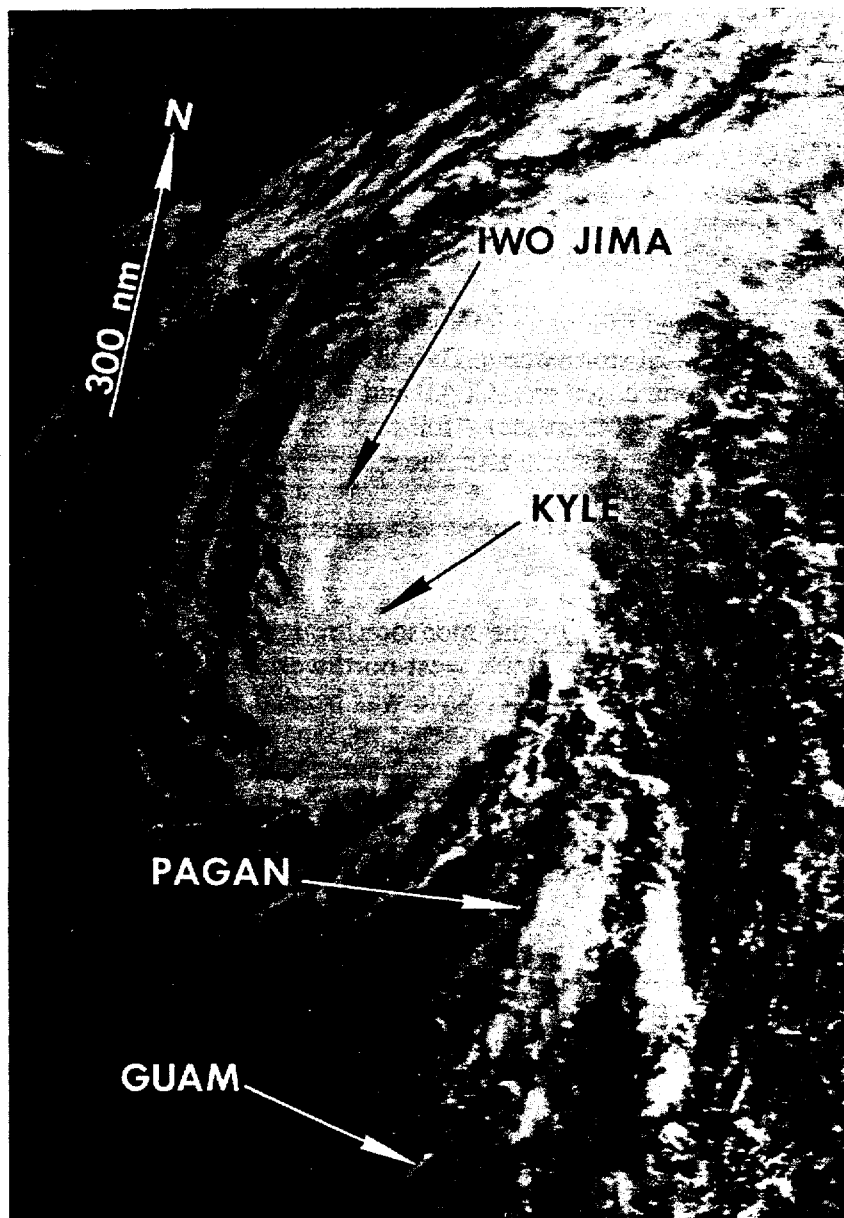


Figure 3-25-1. Typhoon Kyle with elongated eye begins to interact with a frontal system moving southeastward from Japan (200439Z October NOAA visual imagery).

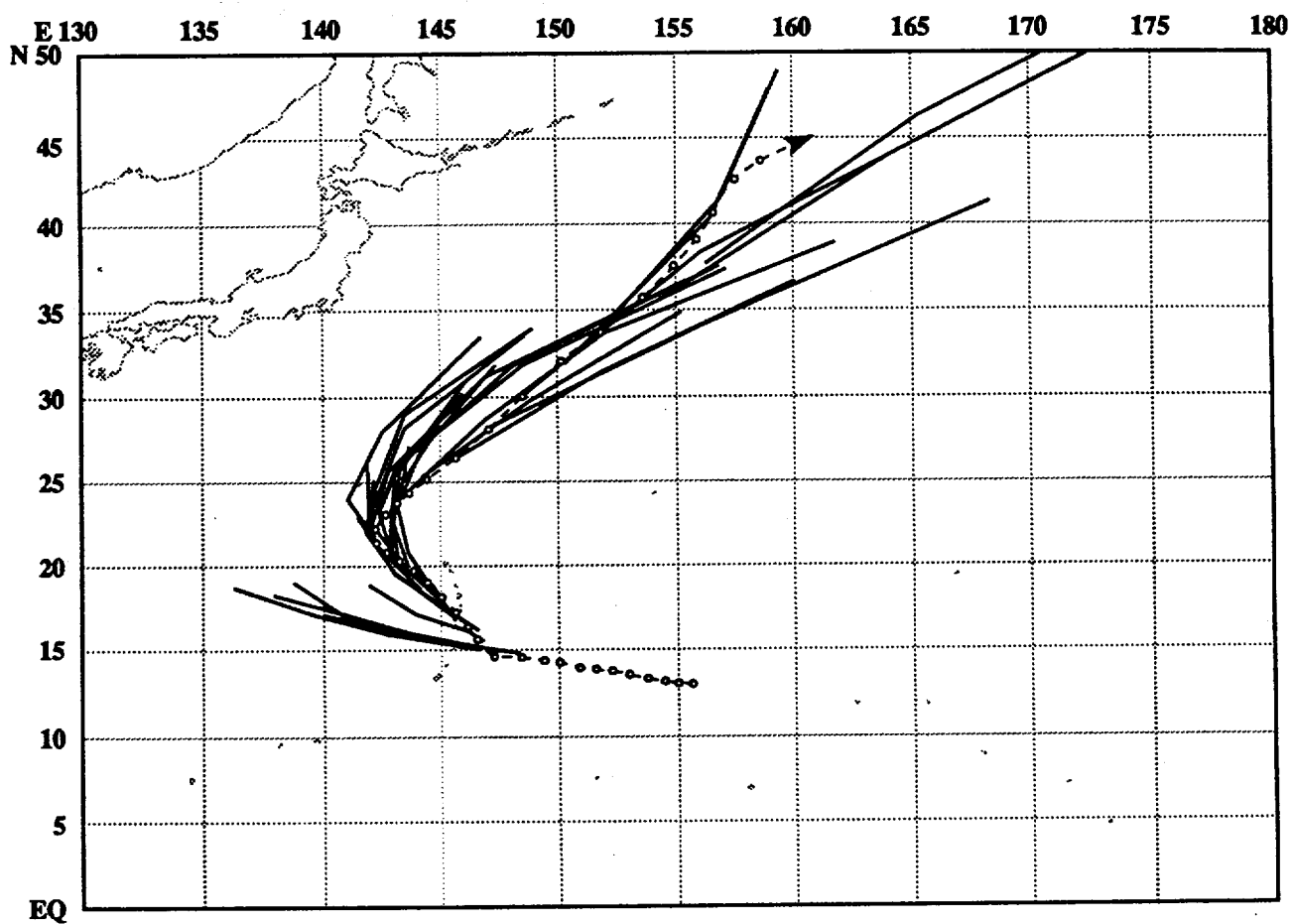
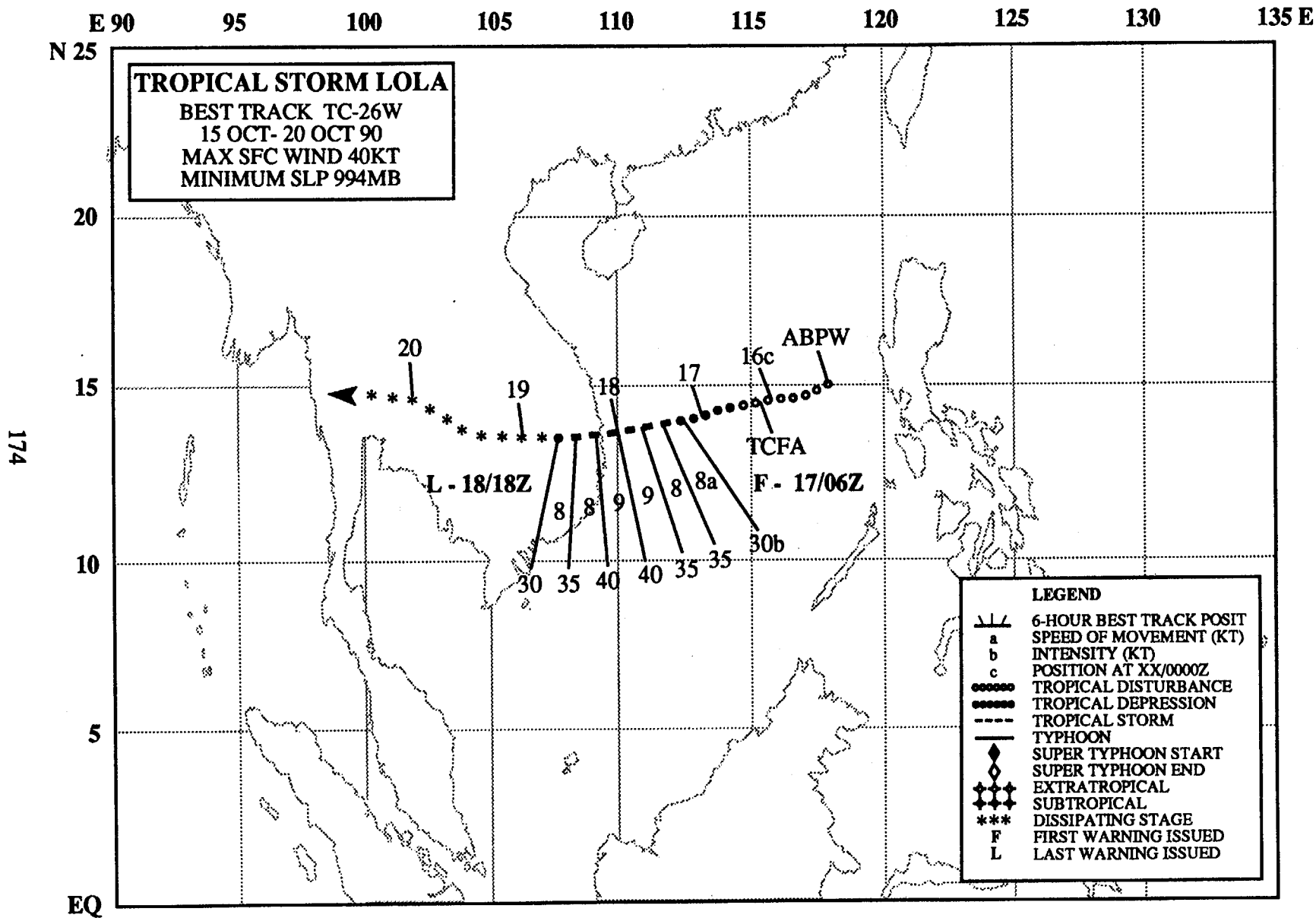


Figure 3-25-2. The overall JTWC forecast performance (solid lines) is superimposed on the final best track (dashed line).



TROPICAL STORM LOLA (26W)

I. HIGHLIGHTS

Lola, the last of four tropical cyclones to develop in October, formed in the South China Sea. It tracked westward along the same path taken by Tropical Storm Jeana (24W) four days earlier.

II. CHRONOLOGY OF EVENTS

- 150600Z - First mentioned on the Significant Tropical Weather as an area of convection associated with a weak low-level cyclonic circulation underneath a weak upper-level anticyclone.
- 160530Z - Tropical Cyclone Formation Alert based on increased organization associated with a surge of northeasterly winds coming off the coast of China.
- 170600Z - First Tropical Depression Warning issued due to a continued increase in organization as the system moved into an area of weaker vertical wind shear.
- 171200Z - Upgraded to a tropical storm due to an increase in organization associated with less vertical wind shear.
- 180000Z - Peak intensity - 40 kt (21 m/sec) - followed a small improvement in organization and an intensity estimate of 2.5.
- 181800Z - Final warning issued as Lola dissipated over land.

III. TRACK AND MOTION

Lola started in the monsoon trough on 15 October and tracked south of a narrow mid-level subtropical ridge for the next five days. Until landfall on the coast of Vietnam on 18 October, the track was consistently south of west because of the strong northeasterly low-level surge across the northern portion of the South China Sea.

IV. INTENSITY

Lola developed when a strong northeasterly surge spun up the low-level circulation. Reaching minimal tropical storm intensity on 17 October, Lola continued to intensify very slowly until it made landfall a day later (Figure 3-26-1). After landfall, the low-level cyclonic circulation persisted, as it crossed into Thailand, and dissipated just before entering the Bay of Bengal.

V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-26-2. Objective aid guidance and mid-level steering flow was used by forecasters; however, due to the strong low-level northeasterly surge Lola consistently tracked south of the forecasts.

VI. IMPACT

No information received.

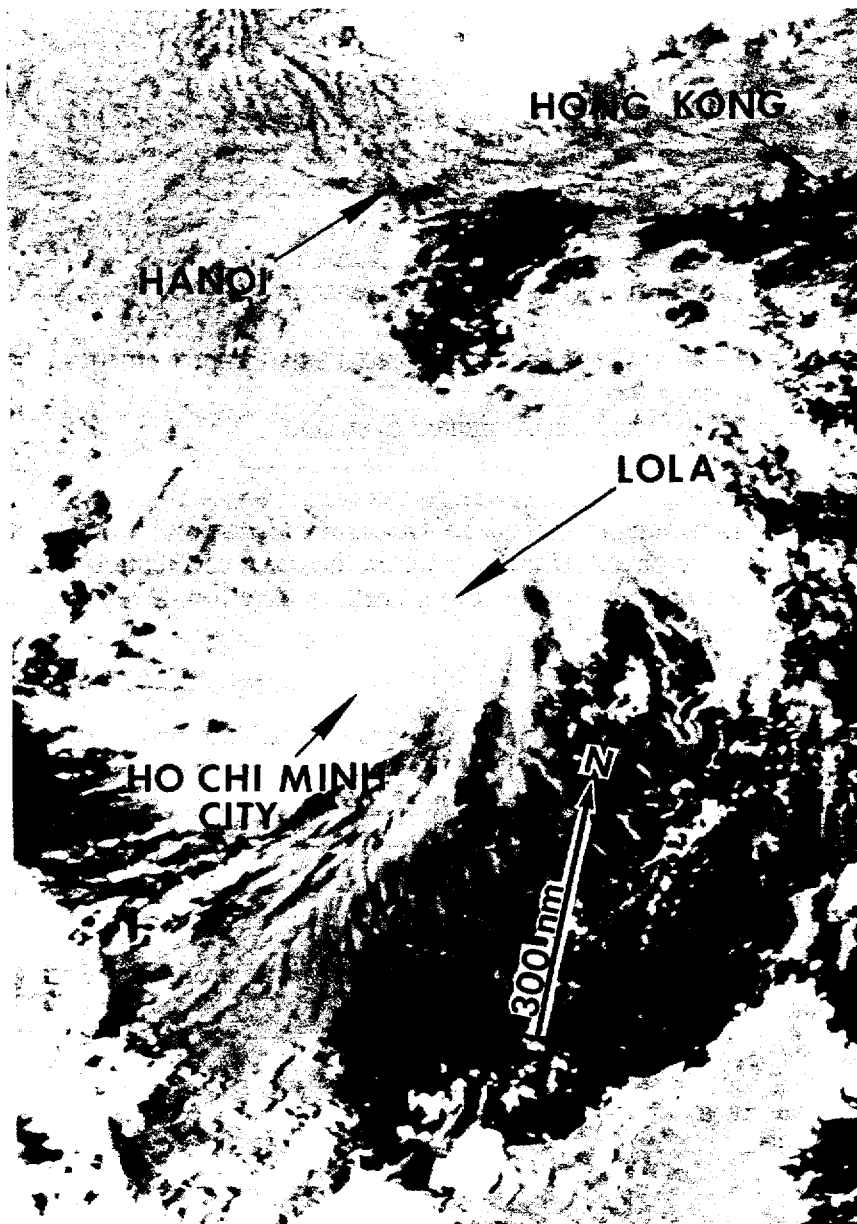


Figure 3-26-1. Lola at maximum intensity (180643Z October DMSP visual imagery).

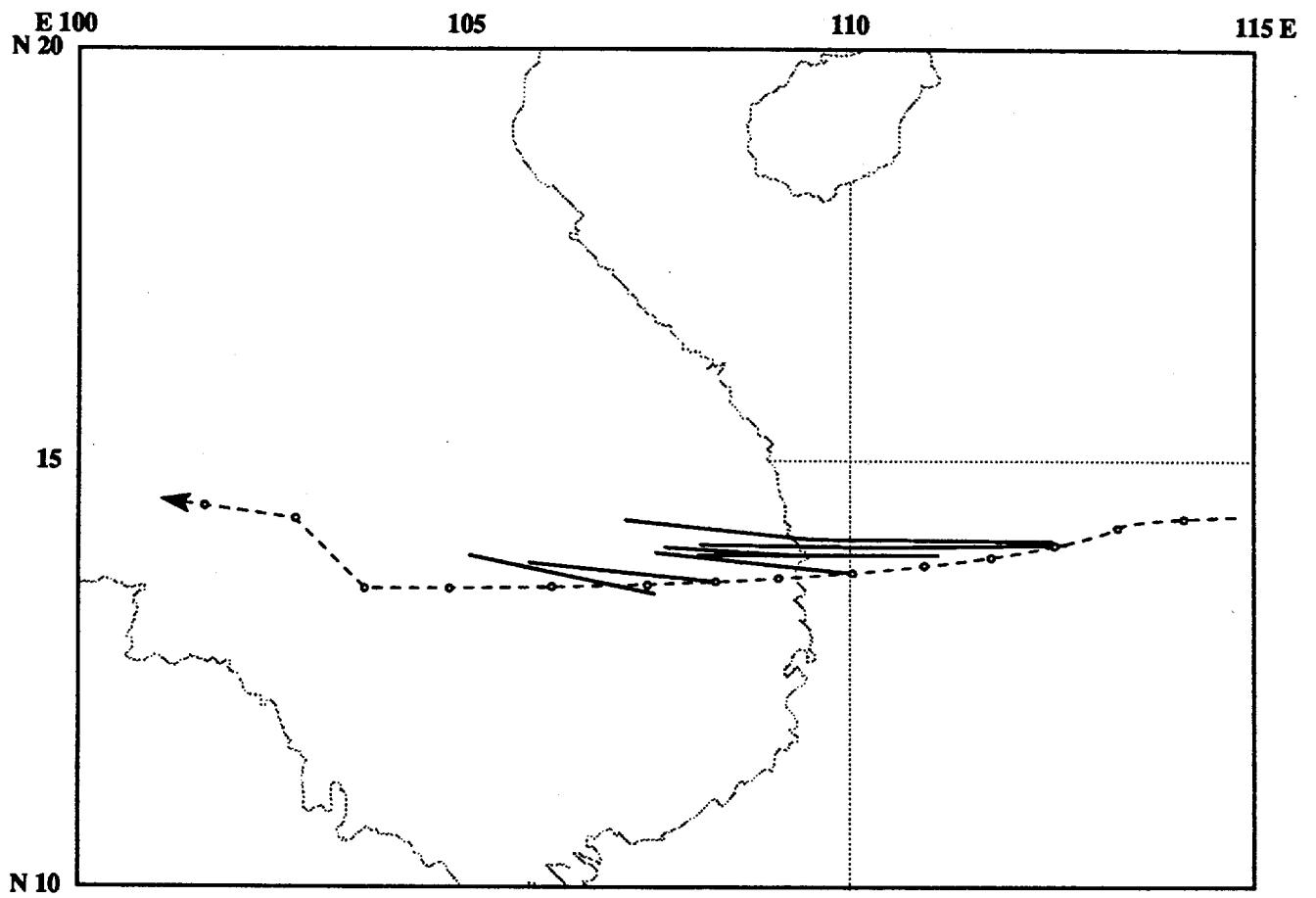
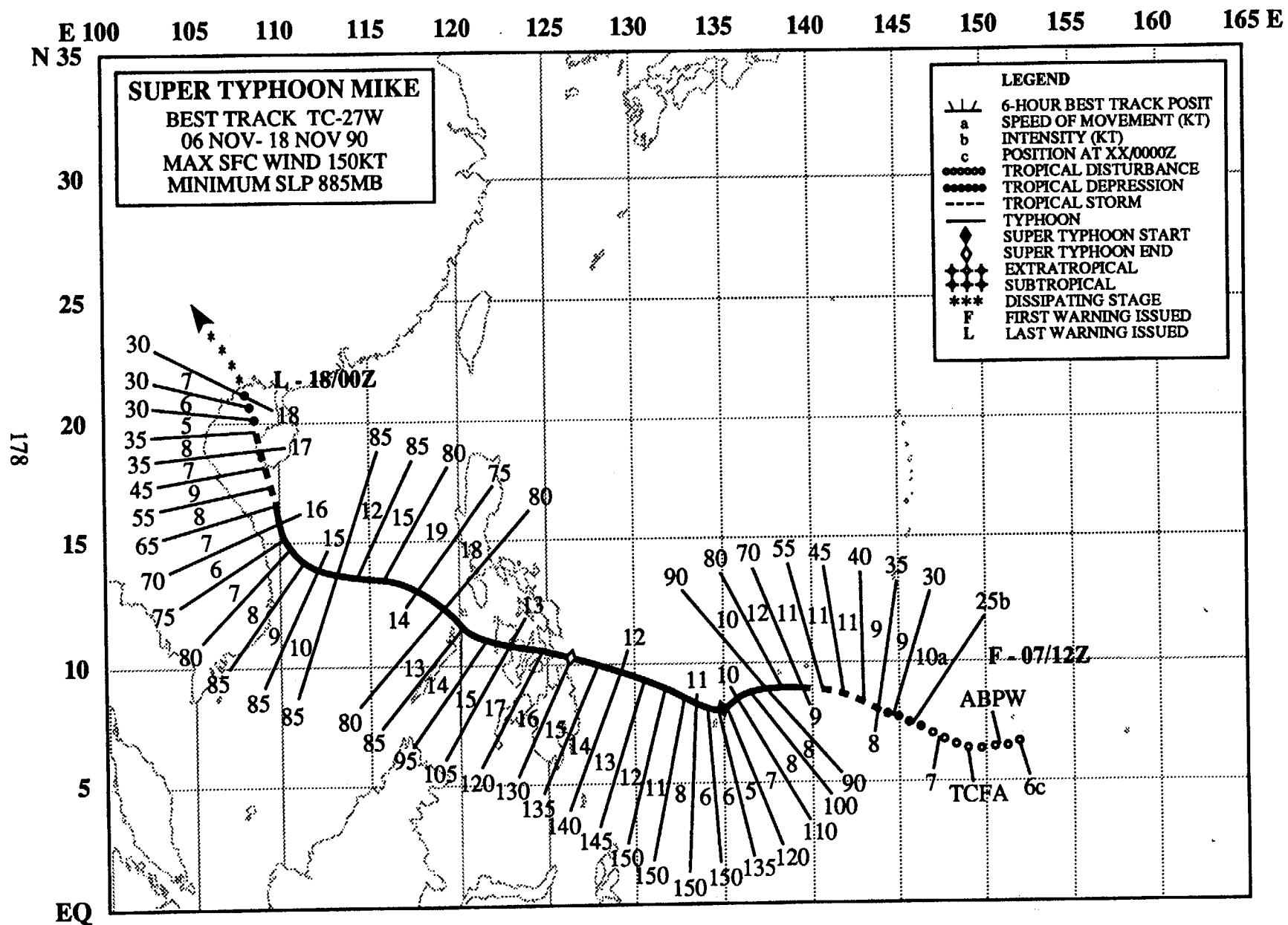


Figure 3-26-2. Summary of JTWC forecasts (solid lines) for Lola is superimposed on the final best track (dashed line).



SUPER TYPHOON MIKE (27W)

I. HIGHLIGHTS

Mike, one of the most intense and destructive tropical cyclones of 1990, caused havoc in western Carolines and in the central Philippine islands. Although basically a west-northwestward "straight runner," it posed numerous forecast challenges due to frequent direction, speed and intensity changes. As a result of the devastation and death in the Republic of the Philippines, Super Typhoon Mike's name was retired from the JTWC naming list.

II. CHRONOLOGY OF EVENTS

- 060600Z – First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1008 mb.
- 061530Z – Tropical Cyclone Formation Alert based on rapidly improving outflow and curvature, an increase in central convection, and a CI 1.0 estimate.
- 071200Z – First warning issued due to continued increase in convection and good outflow in all quadrants. Synoptic data indicated minimum sea-level pressure of 1002 mb.
- 080000Z – Upgraded to tropical storm because synoptic data indicated 35 kt (18 m/sec) around the system.
- 090000Z – Upgraded to typhoon due to formation of an eye and a CI 4.5 estimate.
- 101200Z – Upgraded to super typhoon based on a Dvorak current intensity of 7.0, a small 15 nm (24 km) diameter symmetrical eye, and good outflow in all quadrants.
- 101800Z - Peak intensity - 150 kt (77 m/sec) - established with a CI 7.5.
- 121200Z – Downgraded from super typhoon status due to interaction with land, the eye had become ragged and cloud-filled, and the temperature at the top of the convection around the eye had warmed.
- 161200Z – Downgraded to tropical storm based on interaction with Vietnam coast and degraded satellite cloud signature due to increased vertical wind shear.
- 171200Z – Downgraded to tropical depression due to synoptic reports of weak winds and rising surface pressures, and disorganized cloud signature.
- 180000Z - Final warning (dissipated over land).

III. TRACK AND MOTION

Mike initially tracked west-northwestward under the influence of the mid-tropospheric subtropical ridge to the north. While undergoing rapid intensification on 9 November, it slowed and tracked west-southwestward (Figure 3-27-1). The reason for this change was not apparent, but could be related to the temporary effect of rapid intensification on the environment, and conversely the environment's adjustment to the massive release of latent heat. At 101200Z, Mike resumed its west-northwestward track which took it across the central Philippine islands and into the South China Sea. On 15 November, it turned north-northwestward toward a weakness in the subtropical ridge (Figure 3-27-2). This track took Mike across the western side of Hainan Dao and into southern China, where it dissipated.

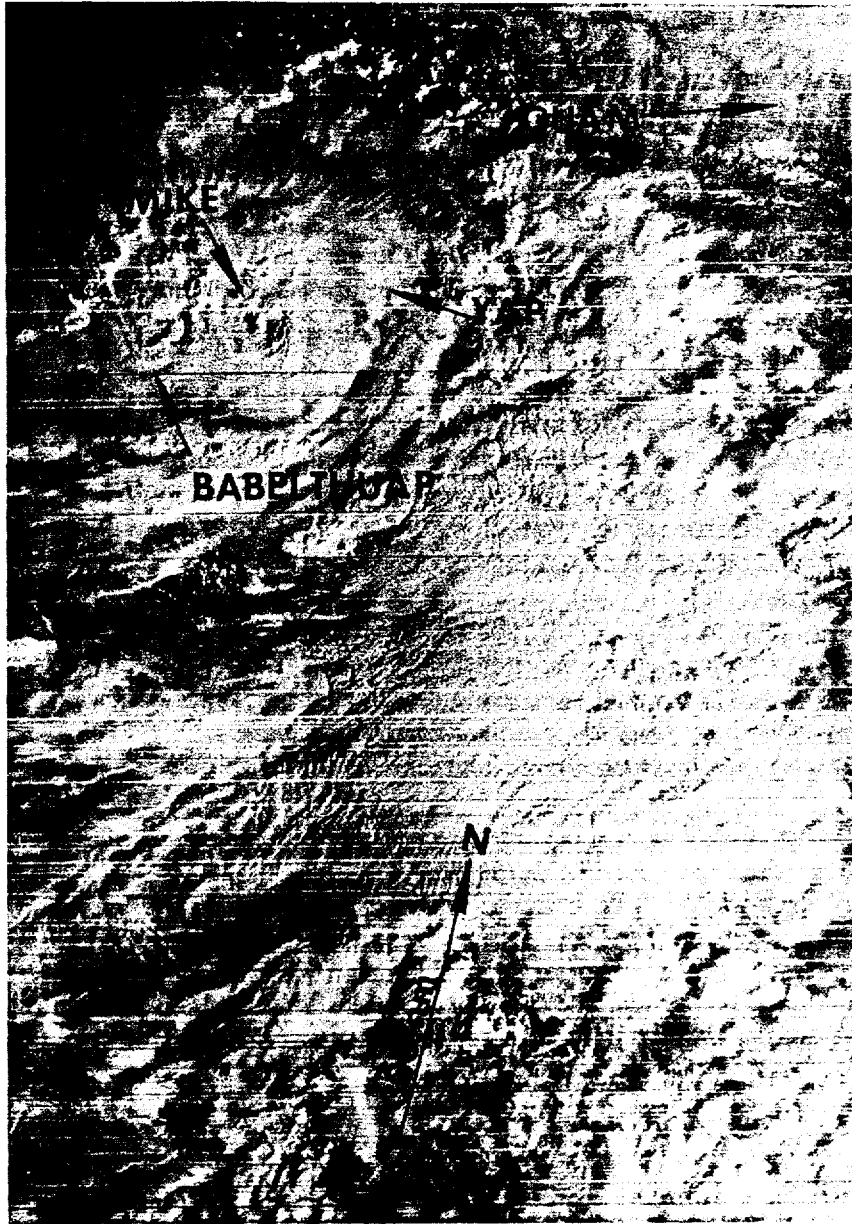


Figure 3-27-1. Mike is rapidly intensifying into a super typhoon as it passes through the western Caroline Islands (092106Z November DMSP visual imagery).

IV. INTENSITY

Mike intensified at a normal rate of T-number per day until reaching moderate tropical storm status at 081200Z. Then intensification accelerated and reached a peak of 150 kt (77 m/sec) at 101800Z. The maximum sustained surface winds increased an additional 95 kt and the estimated minimum sea-level pressure fell 99 mb to 885 mb (Figure 3-27-3) during this 48-hour period. A 200-mb trough to the northeast and broad cross equatorial flow to the south and southwest of Mike provided dual outflow channels that efficiently supported intensification. As Super Typhoon Mike approached landfall in the central Philippine Islands on 12 November, it weakened to just under super typhoon intensity at 121800Z due to the disruptive affects of the mountainous island chain across its path. After further weakening to 80 kt (40 m/sec), the typhoon reintensified to 85 kt (42 m/sec) at 141200Z over the open waters of the South China Sea. As Mike turned north-northwestward off the coast of Vietnam, increased vertical shear started the weakening process again. Dissipation followed on 18 November over southern China.

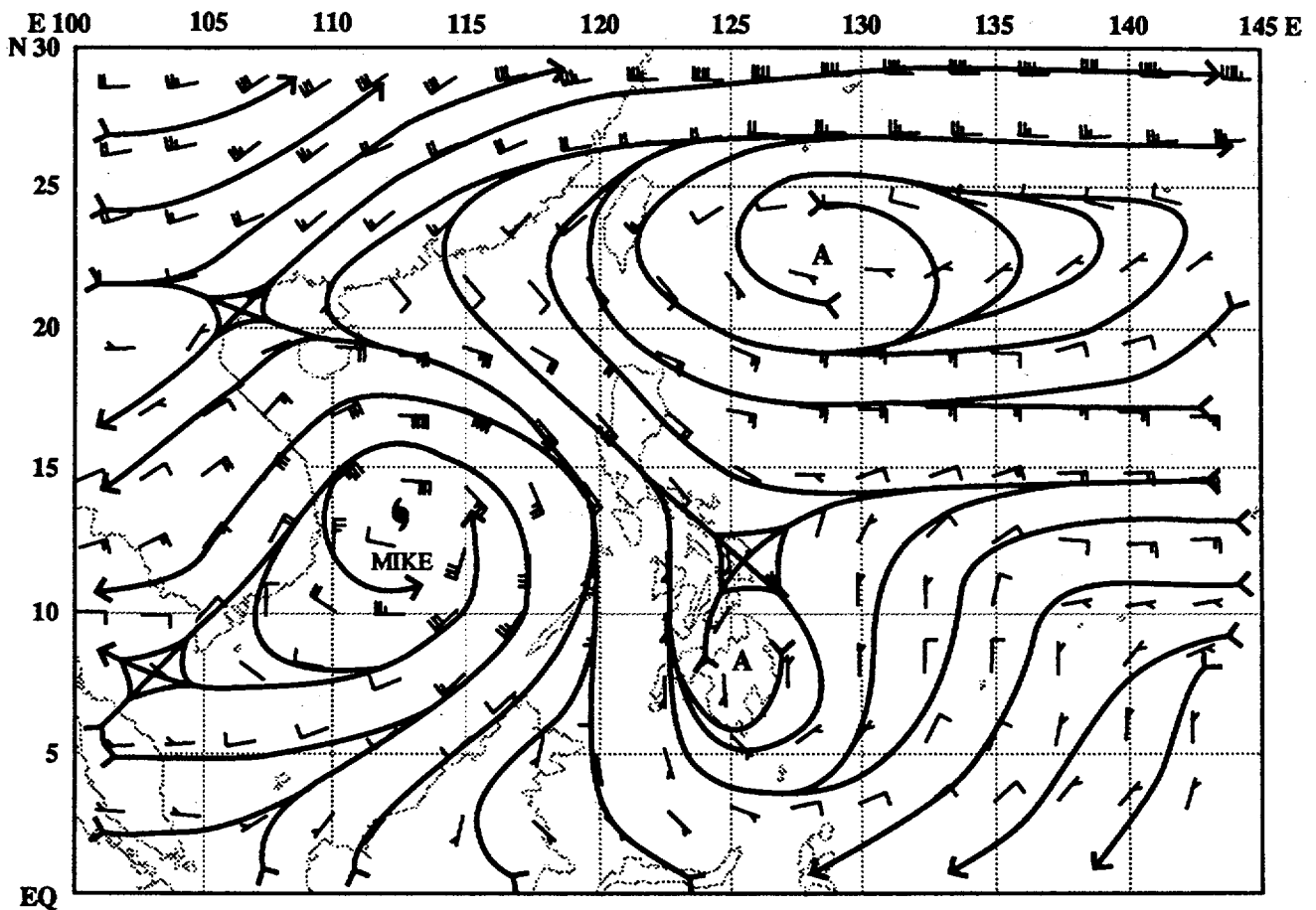


Figure 3-27-2. The deep layer mean analysis for 150000Z November shows Mike and the weakness in the subtropical ridge to the northwest.

V. FORECASTING PERFORMANCE

The overall JTWC forecast performance with respect to the best track is shown in Figure 3-27-4. Initially, JTWC forecast the tropical cyclone to move towards the northwest into the central Philippine Sea. At 091200Z, JTWC introduced a moderate probability alternate scenario of movement across the Philippine Islands, although the NOGAPS prognostic series continued to indicate that a weakness would develop in the ridge just east of Luzon. At 100000Z, it was apparent that Typhoon Mike was moving west-southwestward as the system approached Palau. Mike was expected to resume its west-northwestward track within 24 hours. The prognostic series continued to indicate a weakness in the subtropical ridge, and JTWC continued to forecast northwest motion. However, at the 120000Z, the NOGAPS prognostics changed to reflect a stronger subtropical ridge north of Mike, and subsequently JTWC forecasts reflected motion across the central Philippines, rather than up the east coast of Luzon. As Mike continued west-northwestward into the South China Sea, forecasters expected it to make landfall in Vietnam. Again, the models provided erroneous guidance. The prognostic series failed to predict the weakness that eventually developed in the subtropical ridge to the north (see again Figure 3-27-2).

Mike's favorable outflow pointed to rapid intensification, which was in fact forecast. Despite the fact that there are no objective aids, or hard and fast rules of thumb, to predict the exact rate or peak intensity, the forecast of 130 kt (67 m/sec) maximum was made 48 hours before Mike actually peaked at 150 kts (77 m/sec). Later as Mike approached the Philippine Islands, preliminary results from a climatological study of tropical cyclones crossing the Philippines correctly indicated that it would weaken to 85 kt (44 m/sec), enabling JTWC to issue a perfect 72-hour intensity forecast.

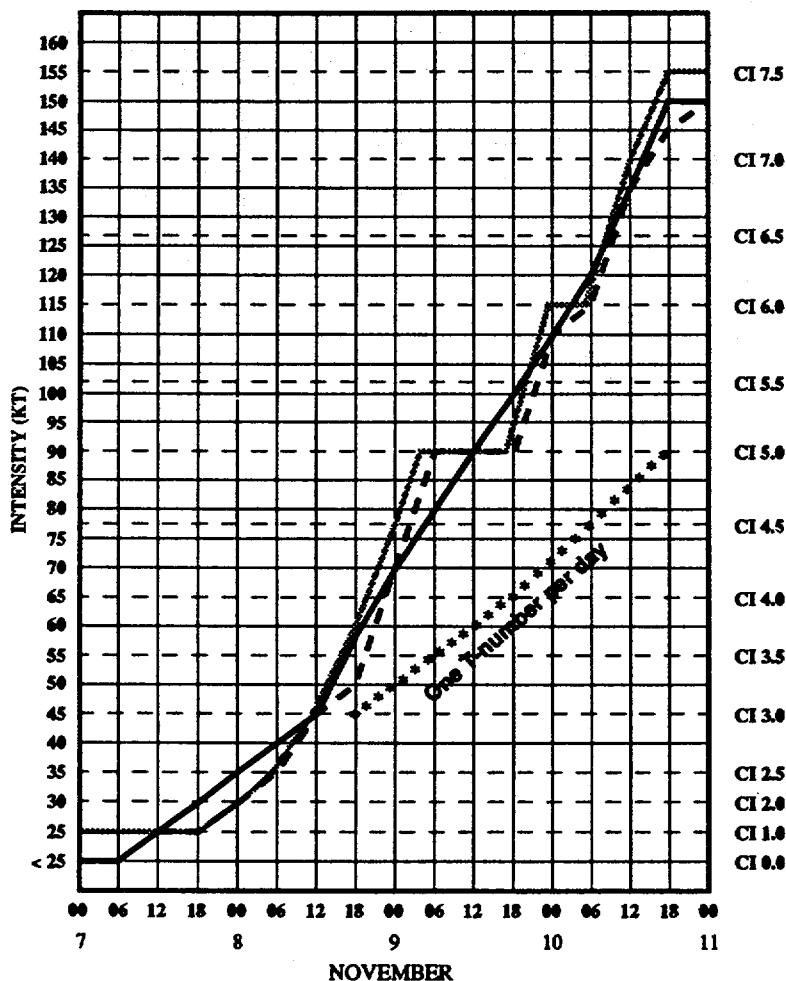


Figure 3-27-3. Plots of the satellite current intensity values (gray line), actual warning intensities (dashed line), and final best track (solid line) on a time-intensity comparison chart depict Mike's greater than normal rate of intensification after 081200Z. The normal development of one T-number per day (starred line) is included as a reference.

VI. IMPACT

Super Typhoon Mike was extremely destructive to the western Carolines and central Philippine Islands. On Koror, 45 nm (85 km) south of Mike's center, many roofs were lost and extensive damage occurred to boats, greenhouses, aquiculture projects, fruit trees and vegetable gardens. Fortunately there were no fatalities and only one serious injury was reported. Power, water and telephone services were completely out and roads were blocked by fallen trees. The National Weather Service Office at Koror (WMO 91408) recorded maximum wind gusts to 72 kt (36 m/sec), a minimum sea-level pressure of 980.5 mb and 9.8 inches (250 mm) of rain. Closer to Mike's center, where maximum wind gusts were estimated to range from 135 to 165 kt (69 to 85 m/sec), Kayangel Island just to the north of Babelthup was almost totally devastated. Many people lost everything. Most trees used for subsistence were destroyed, with some, such as breadfruit, expected to take up to ten years to replace.

Super Typhoon Mike became the most powerful typhoon to strike the Philippine Islands this year and was reported to be the most devastating to hit the country since 1981. In the central Philippine islands at least 250 people were reported dead or missing, mostly from landslides, and 2 million people were forced from their homes into temporary shelters. Over 37,000 houses were destroyed, and at least \$14 million worth of damage was recorded. Cebu city, the commercial and transportation capital of the region, was severely damaged and more than 57 water craft, mostly in the port of Cebu, sank.

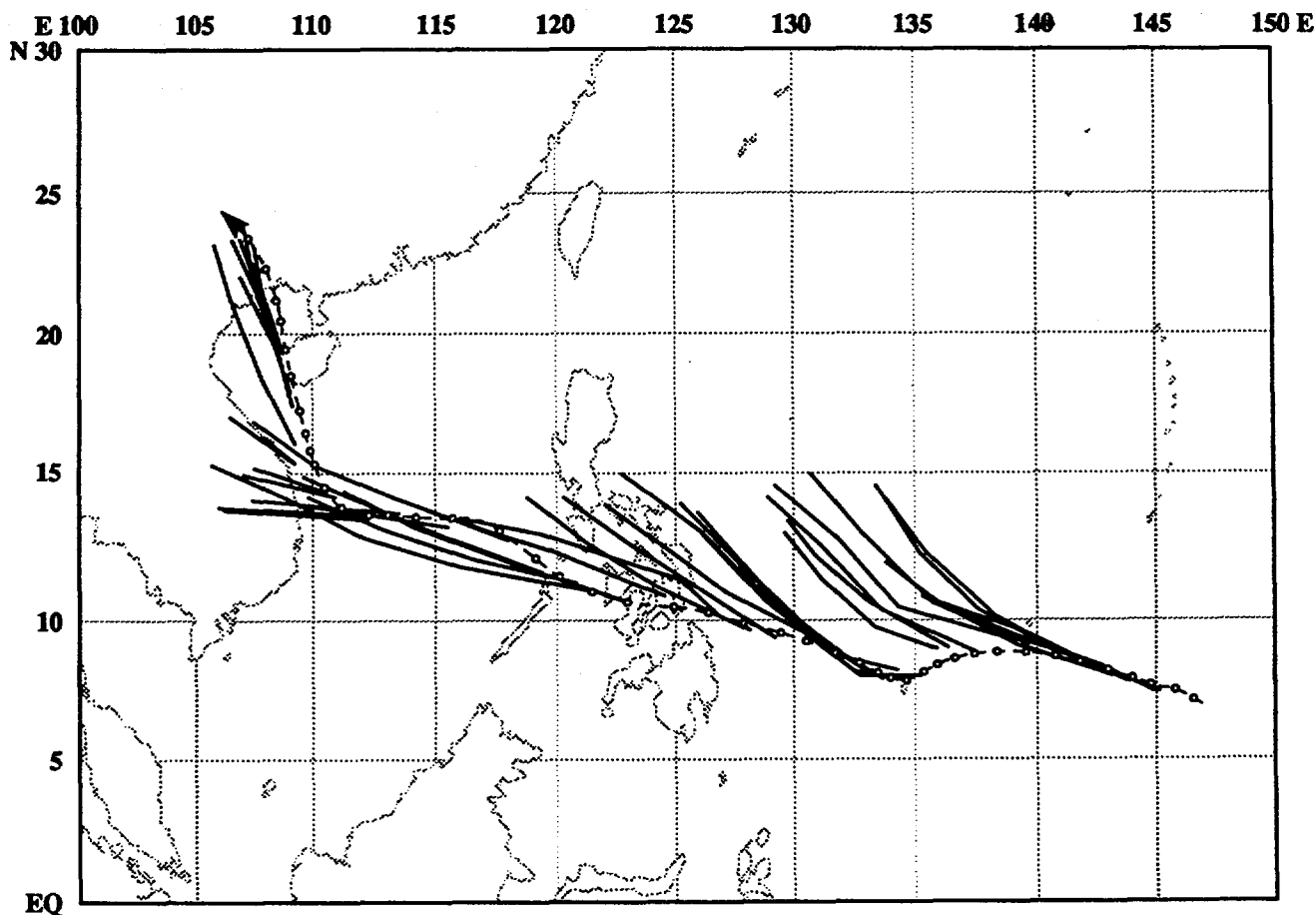
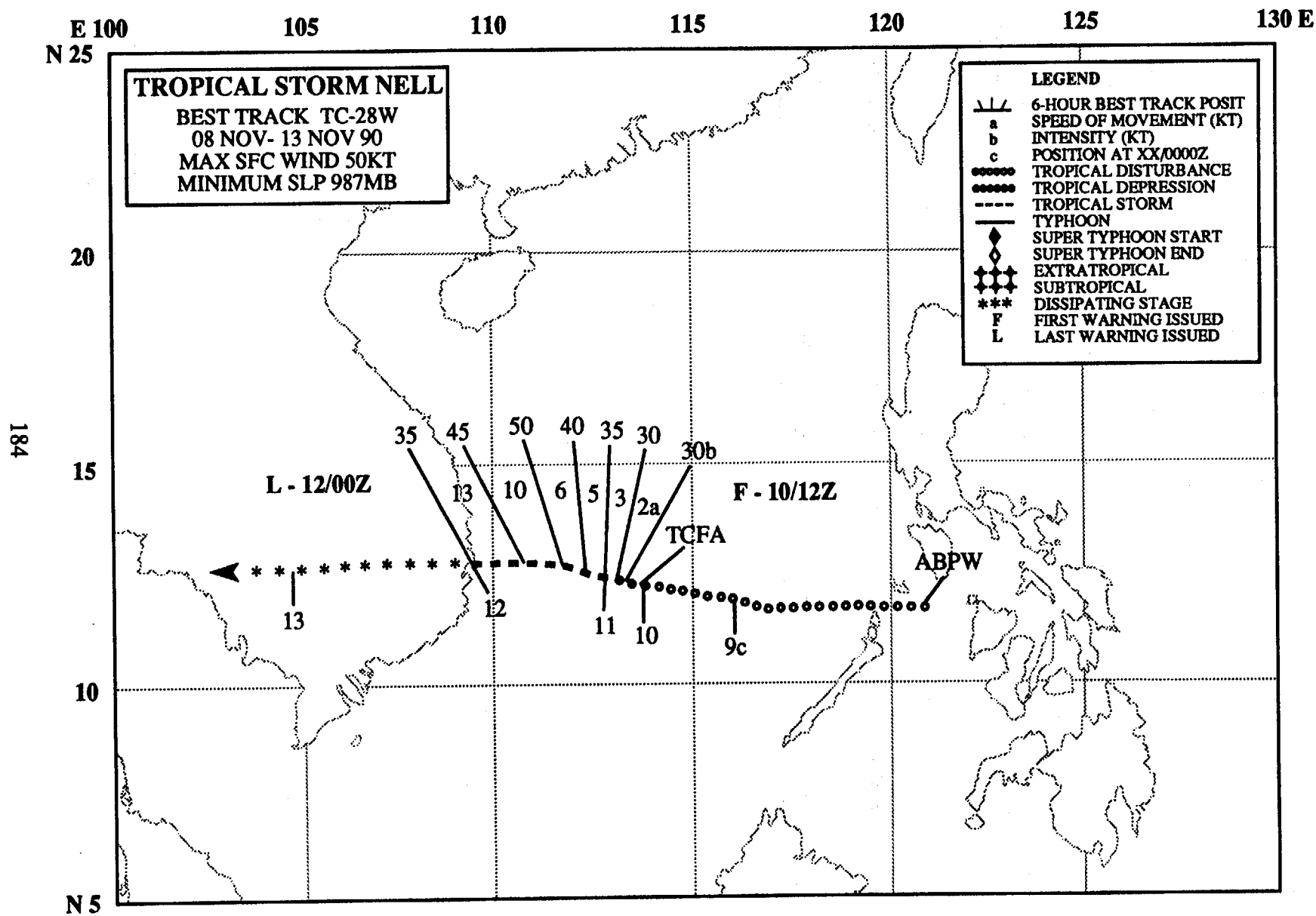


Figure 3-27-4. The JTWC forecasts (solid lines) for Mike superimposed on the final best track (dashed line).



TROPICAL STORM NELL (28W)

I. HIGHLIGHTS

Nell, the second of four November tropical cyclones, intensified in the South China Sea and tracked westward, making landfall in Vietnam.

II. CHRONOLOGY OF EVENTS

- 080600Z - First mentioned on Significant Tropical Weather Advisory as a disturbance with a persistent area of convection with estimated minimum sea-level pressure of 1005 mb.
- 100330Z - Tropical Cyclone Formation Alert issued due to increased convective organization.
- 101200Z - First warning issued due to surface synoptic reports of 25-30 kt (13-15 m/sec) winds.
- 110000Z - Upgrade to tropical storm based on a synoptic report of 996 mb sea-level pressure and 35 kt (18 m/sec) surface wind.
- 111200Z - Peak intensity of 50 kt (25 m/sec) based on ship reports.
- 120000Z - Final warning issued as Nell moved over land and began to dissipate in Vietnam.

III. TRACK AND MOTION

Nell formed west of the central Philippine Islands and tracked across the South China Sea south of the subtropical ridge which remained near 20° north latitude. After landfall, the low-level cyclonic circulation moved westward into Thailand.

IV. INTENSITY

Nell developed in association with a surge in the northeast monsoon, reached tropical storm intensity (Figure 3-28-1) on 11 November, and peaked at 50 kt (25 m/sec) despite indications of strong vertical shear.

V. FORECASTING PERFORMANCE

Initially, fix position uncertainties and the strength of the surge in the northeast monsoon led forecasters to believe Nell would move south-southwestward. Later forecasts reflected the movement to the west (Figure 3-28-2).

VI. IMPACT

No information received.

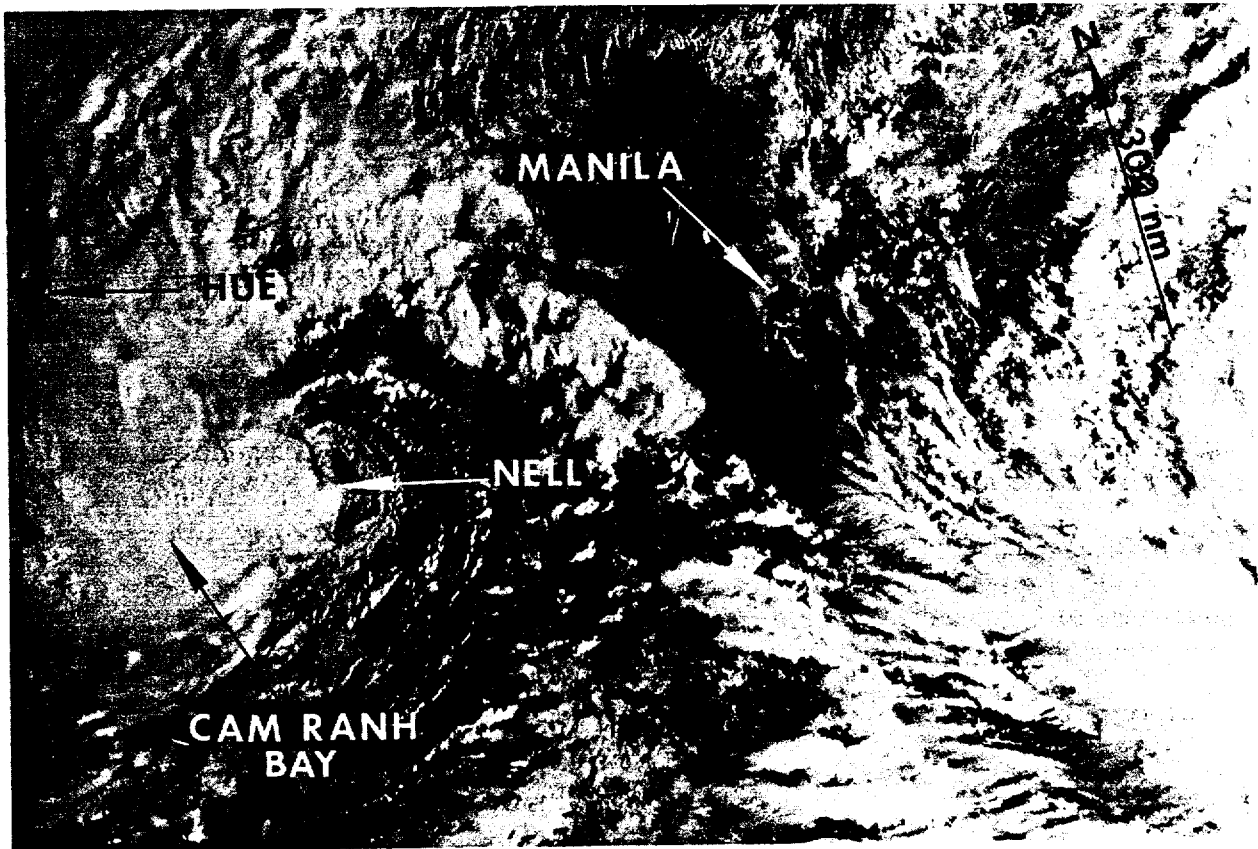


Figure 3-28-1. Just before reaching tropical storm intensity, Nell's LLCC is partially exposed (102335Z November NOAA visual imagery)

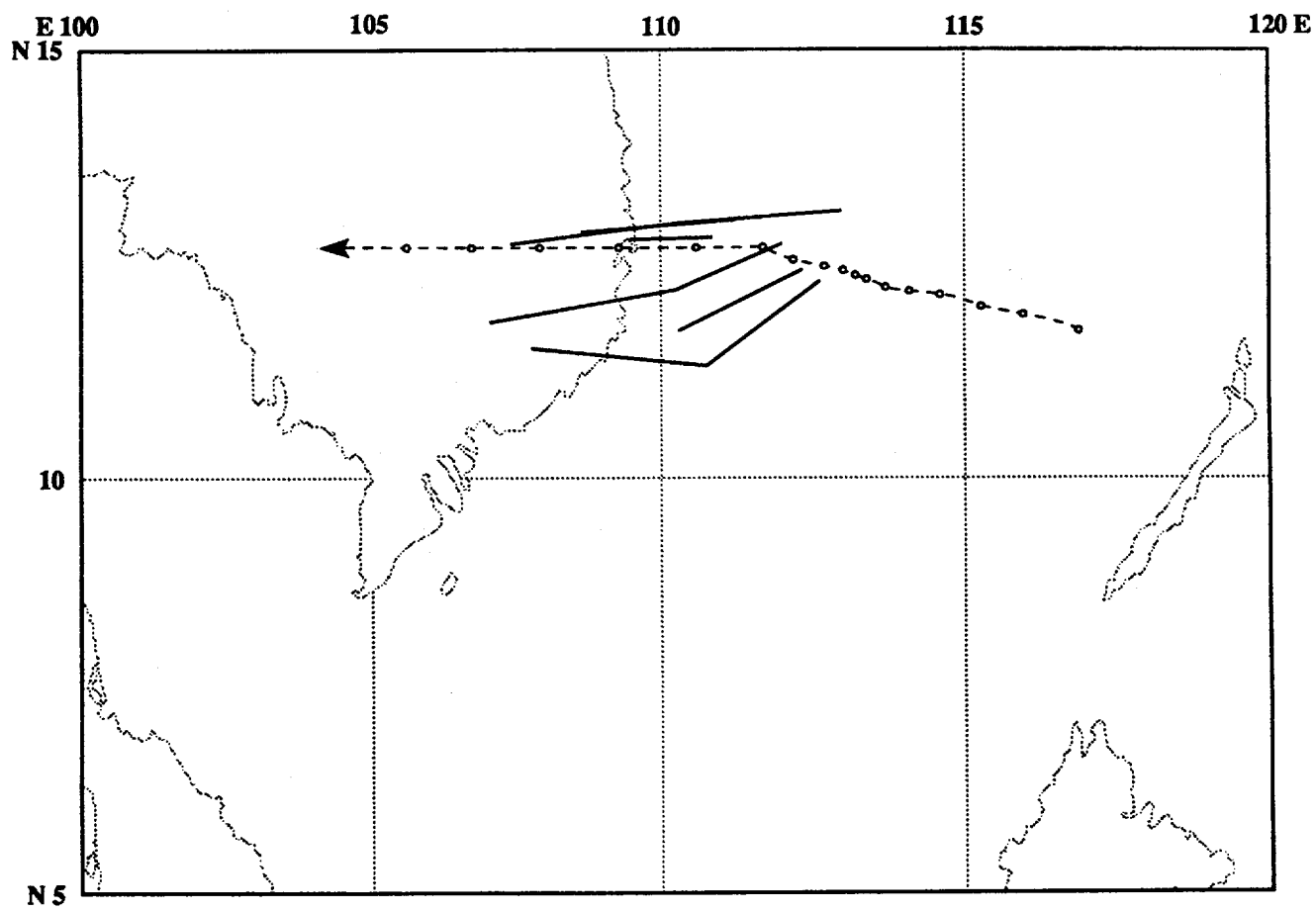


Figure 3-28-2. Summary of JTWC forecasts (solid lines) for Nell is superimposed on the final best track (dashed line).

E 110 115 120 125 130 135 140 145 150 155 160 165 170 175 180
N 45

SUPER TYPHOON PAGE

BEST TRACK TC-29W
05 NOV- 01 DEC 90
MAX SFC WIND 140KT
MINIMUM SLP 898MB

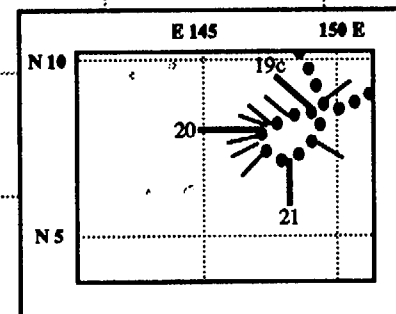
L - 30/12Z

DTG SPEED INTENSITY

19/00Z	--	30
19/06Z	10	30
19/12Z	7	25
19/18Z	1	25
20/00Z	3	30
20/06Z	2	30
20/12Z	4	30
20/18Z	5	30
21/00Z	10	25
21/06Z	10	25
21/12Z	11	25
21/18Z	14	25

LEGEND

- 6-HOUR BEST TRACK POSIT
- a SPEED OF MOVEMENT (KT)
- b INTENSITY (KT)
- c POSITION AT XX/0000Z
- TROPICAL DISTURBANCE
- TROPICAL DEPRESSION
- TROPICAL STORM
- TYPHOON
- ◆ SUPER TYPHOON START
- ◆ SUPER TYPHOON END
- ◆ EXTRATROPICAL
- ◆ SUBTROPICAL
- *** DISSIPATING STAGE
- F FIRST WARNING ISSUED
- L LAST WARNING ISSUED



F - 19/00Z

1st TCFA

2nd TCFA

ABPW

SUPER TYPHOON PAGE (29W)

I. HIGHLIGHTS

Page was the third of four tropical cyclones to form in November, the second super typhoon of the month, and part of the three-storm outbreak which included a pair of tropical cyclones near the date line: Owen (30W) in the northern hemisphere and Sina (TC 03P) in the southern hemisphere. Persisting as a discrete disturbance for nearly two weeks before the first warning was issued, Page took only three days to intensify to 140 kt (70 m/sec) once development commenced.

II. CHRONOLOGY OF EVENTS

- 050600Z - First mentioned on the Significant Tropical Weather Advisory as an area of persistent convection with an estimated minimum sea-level pressure of 1008 mb.
- 170300Z - First Tropical Cyclone Formation Alert based on better convective organization with increased low-level inflow, indications in the NOGAPS prognostic series of a decrease in vertical wind shear over the area, and a CI 1.5 estimate.
- 180300Z - Second Tropical Cyclone Formation Alert based on a broadening low-level circulation with decreasing vertical wind shear and a surge in the easterlies north of the disturbance.
- 190000Z - First warning issued due to the low-level circulation center moving under the edge of the central cloud mass, a developing upper-level anticyclone, and a current intensity estimate of CI 2.0.
- 220000Z - Upgraded to a tropical storm after convective curvature increased, upper-level outflow improved, and the first intensity estimate of CI 2.5.
- 240600Z - Upgraded to typhoon intensity after formation of an eye wall and intensity estimates of CI 4.0.
- 260600Z - Upgrade to a super typhoon followed the development of a well defined 40 nm (75 km) diameter eye and intensity estimates of CI 6.5.
- 271800Z - Downgraded to a typhoon after a decrease in central convection, visible loss of eye wall definition and an intensity estimate of CI 6.0.
- 300600Z - Downgrade to tropical storm based on increased vertical wind shear and the start of extratropical transition.
- 301200Z - Final warning based on a combination of land interaction with Honshu and extratropical transition.

III. TRACK AND MOTION

Page formed in the Marshall Islands near Kwajalein Atoll and tracked slowly westward on the south side of the subtropical ridge. As the disturbance passed south of Guam on 19 November, it interacted with enhanced low-level equatorial westerlies supporting a multiple cyclone outbreak further eastward near the date line (Figure 3-29-1). Page executed a counterclockwise loop which took two days to complete and then resumed a westward track on 22 November. As Page neared 125° east longitude, it tracked northward through a break in the subtropical ridge, recurved on 27 November, and accelerated northeastward.

IV. INTENSITY

Page's swirl of low-level cloudiness remained intact, but poorly organized, for two weeks beneath strong easterly flow aloft which restricted vertical development (Figure 3-29-2). On 23 November, the tropical cyclone began steady intensification in an area of lower vertical wind shear. Over the next three days, Page (Figure 3-29-3) underwent several periods of rapid intensification to reach a peak of 140 kt (72 m/sec) on 26 November. During this 72-hour period, the estimated sea-level pressure (Atkinson-Holliday, 1977) dropped 93 mb to a minimum of 898 mb with a subsequent 95 kt (50 m/sec) increase in the maximum winds. After maintaining peak intensity for a day, Page began to weaken due to increasing vertical wind shear as it encountered the mid-latitude westerlies. Extratropical transition occurred over Honshu on 30 November.

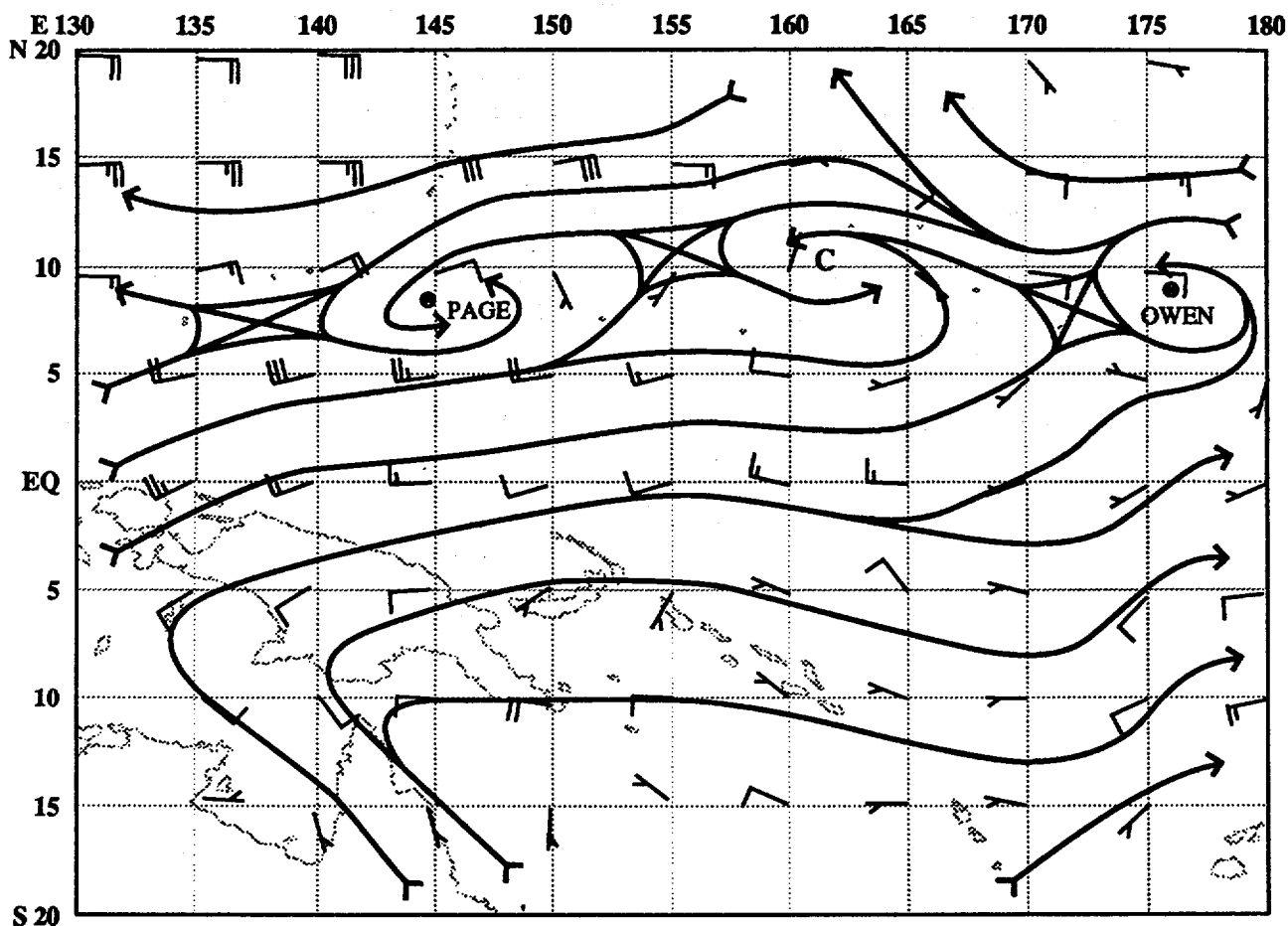


Figure 3-29-1. The 190000Z November NOGAPS 850-mb analysis shows enhanced low-latitude flow, extending eastward to the dateline where Owen (30W) was developing.

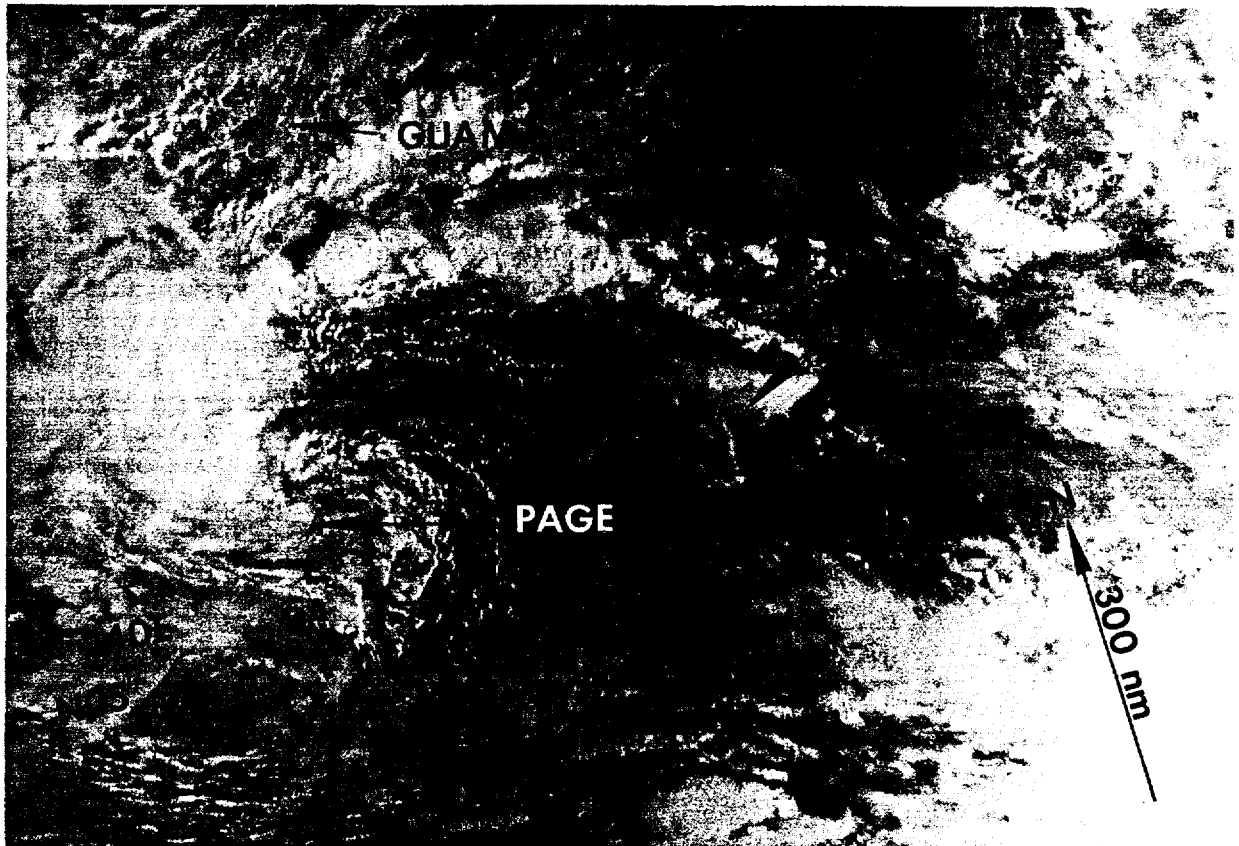


Figure 3-29-2. The exposed low-level circulation center associated with TD 29W as it loops south of Guam (202124Z November NOAA visual imagery).

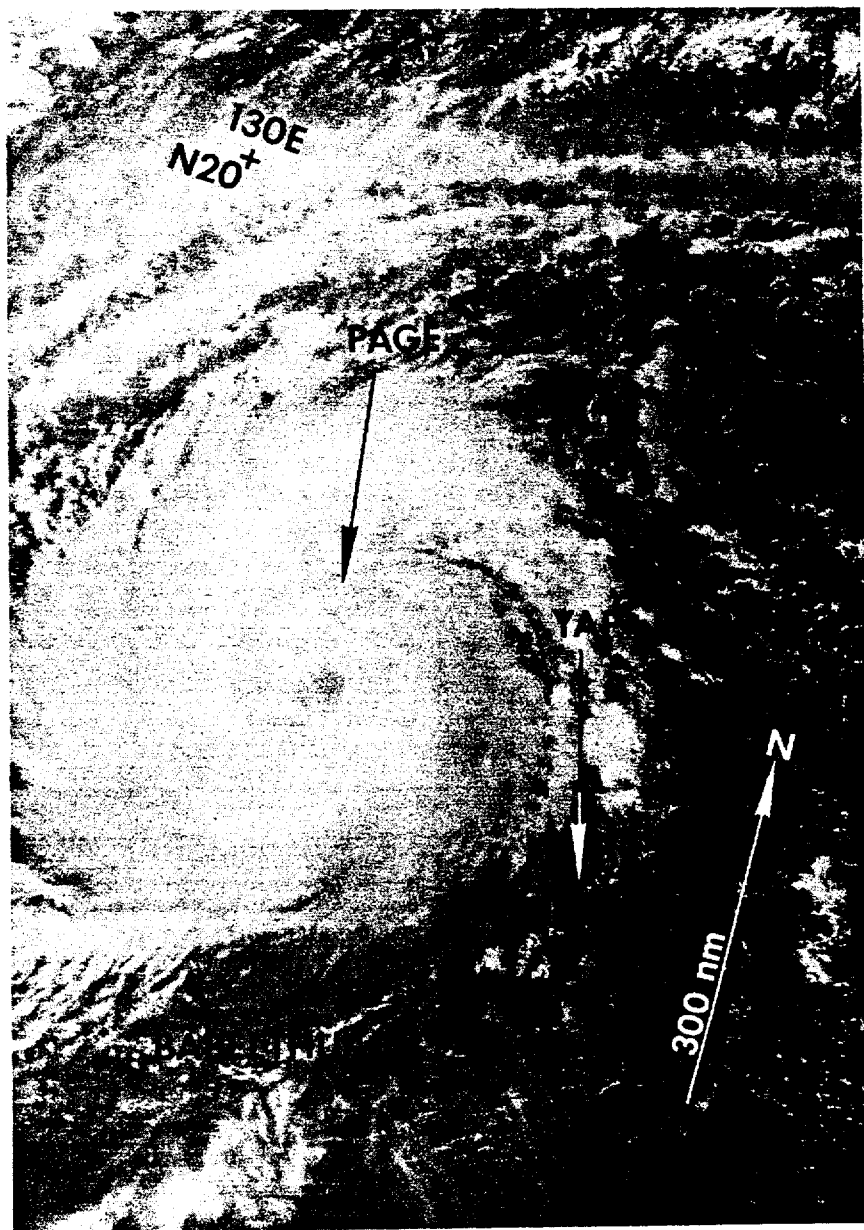


Figure 3-29-3. Super Typhoon Page near its peak intensity (250443Z November NOAA visual imagery).

V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-29-4. The difficulties came from two sources: first the loop south of Guam was unexpected, and second the NOGAPS prognostic series maintained a weak mid-level ridge over the Philippine Sea to the north of Page, supporting a west-northwestward track into the northern Philippines. At 250000Z, a moderate probability alternate scenario was formulated calling for Page to recurve east of the Philippines in response to a developing weakness in the subtropical ridge associated with a passing shortwave trough. This alternate became the primary forecast at 260000Z, as the ridge broke and recurvature followed.

VI. IMPACT

Guam received peak gusts to 46 kt (23 m/sec) at the International Airport (WMO 91212) on 23 November and over 5 inches (125 mm) of rain, which resulted in some localized flooding. No information was received about Page's passage over Honshu.

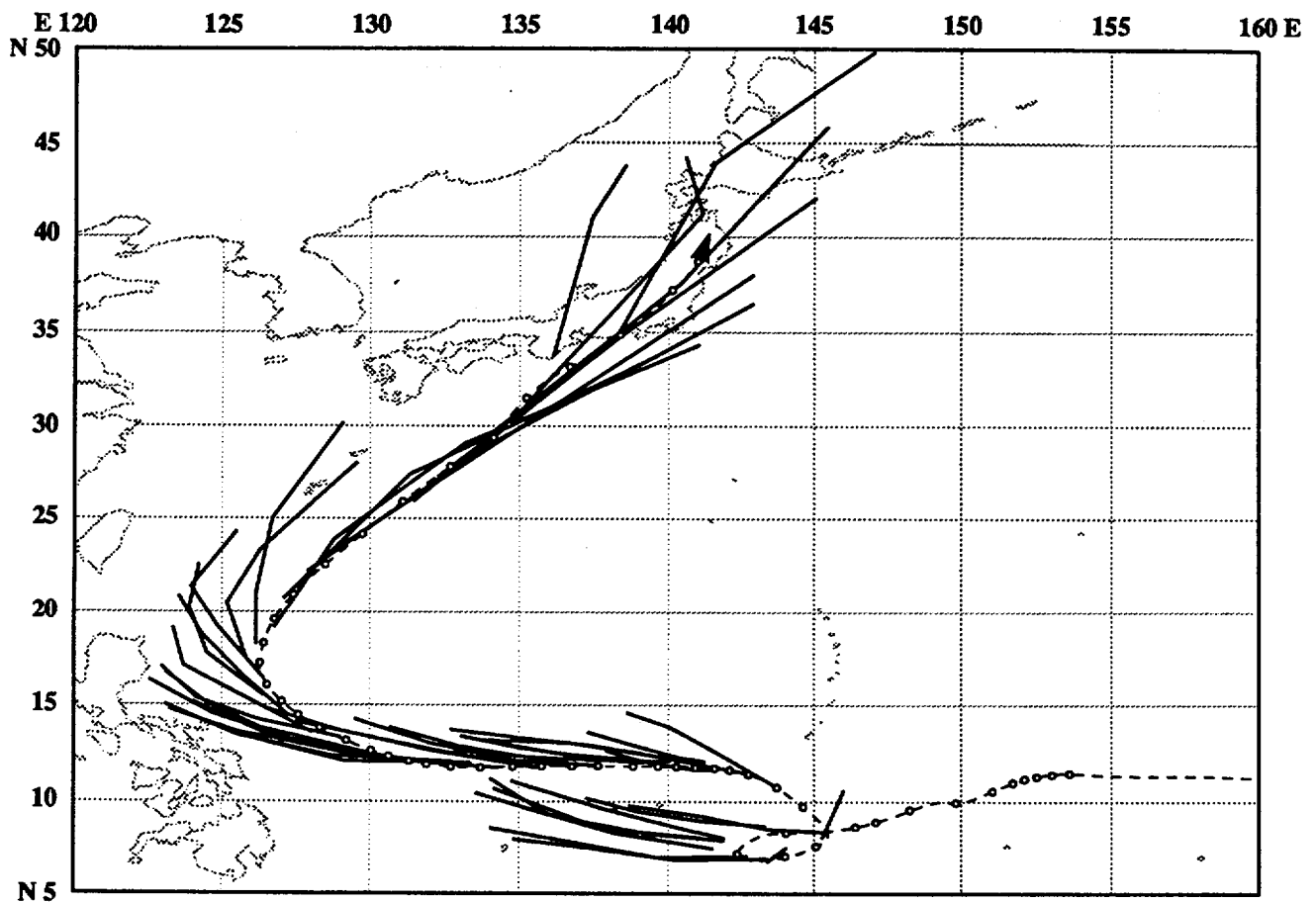
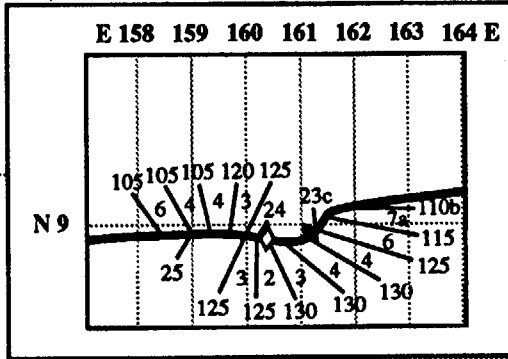


Figure 3-29-4. Summary of JTWC forecasts (solid lines) for Page superimposed on the final best track (dashed line).

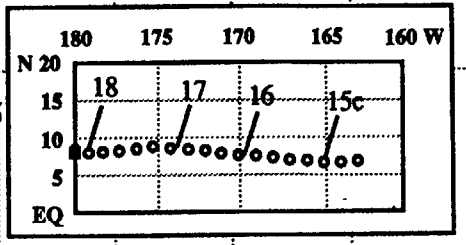
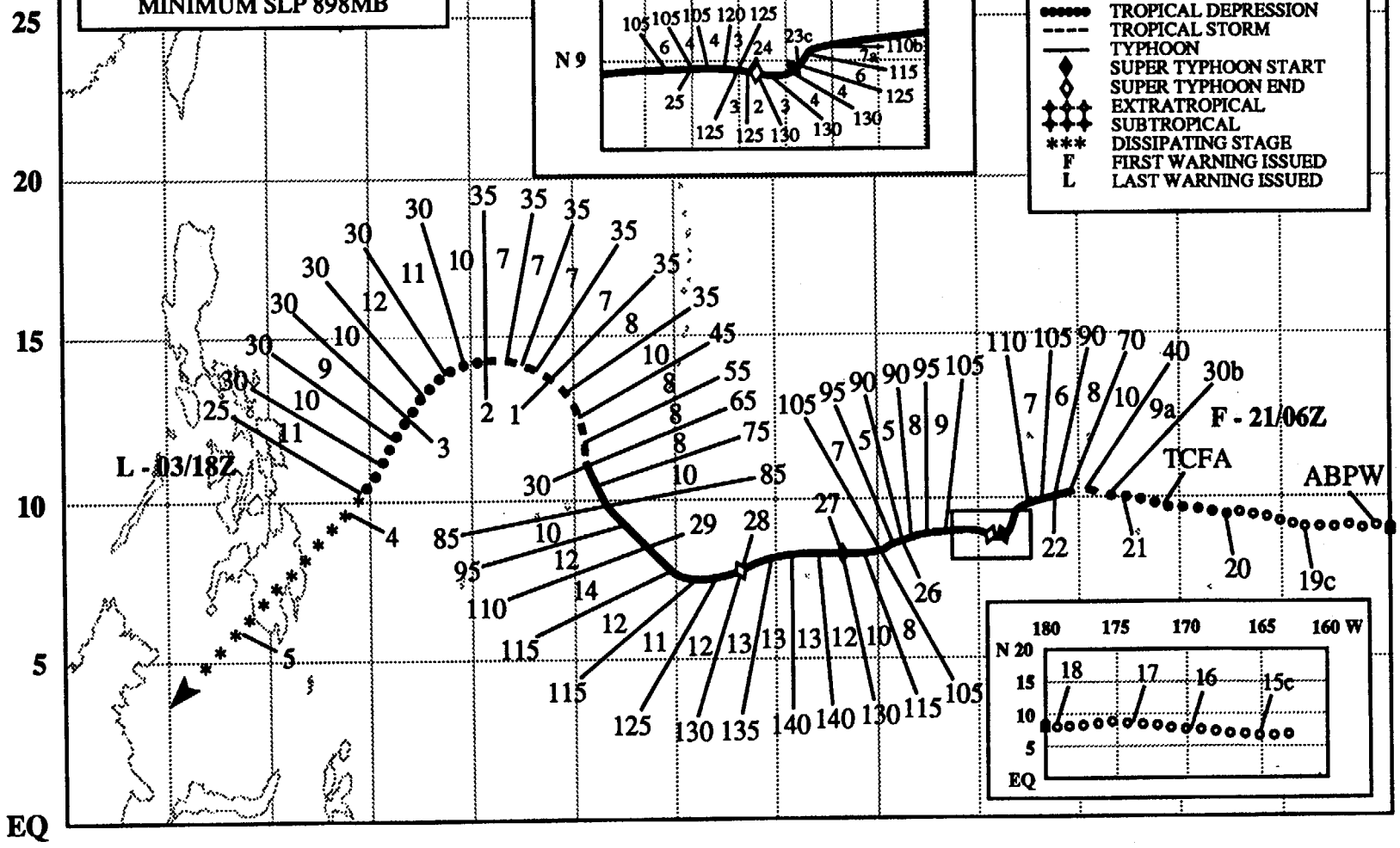
E 115 120 125 130 135 140 145 150 155 160 165 170 175 180
N 30

SUPER TYPHOON OWEN
BEST TRACK TC-30W
14 NOV- 05 DEC 90
MAX SFC WIND 140KT
MINIMUM SLP 898MB



LEGEND

- 6-HOUR BEST TRACK POSIT
- a SPEED OF MOVEMENT (KT)
- b INTENSITY (KT)
- c POSITION AT XX/0000Z
- TROPICAL DISTURBANCE
- TROPICAL DEPRESSION
- TROPICAL STORM
- TYPHOON
- ◆ SUPER TYPHOON START
- ◆ SUPER TYPHOON END
- ◆ EXTRATROPICAL
- ◆ SUBTROPICAL
- *** DISSIPATING STAGE
- F FIRST WARNING ISSUED
- L LAST WARNING ISSUED



194

SUPER TYPHOON OWEN (30W)

I. HIGHLIGHTS

Owen was both the longest lasting and one of the most interesting tropical cyclones of 1990. It started to rapidly intensify while still a tropical depression, explosively deepened to super typhoon intensity, weakened and then reintensified to a super typhoon. Owen started as a discrete cloud mass southwest of Hawaii, maintained its integrity as it tracked westward in the trade wind trough, but did not intensify until it crossed the date line and passed north of Kwajalein in the Marshall Islands. It then reached typhoon intensity in less than 18 hours and continued westward over the central Caroline Islands where it weakened and reintensified. Its deep convection sheared away southeast of Ulithi Island in the western Carolines. The exposed low-level remained organized for six more days as it moved north, then west, and finally southwestward before dissipating over the Celebes Sea after crossing Mindanao.

II. HIGHLIGHT OF EVENTS

- 180600Z - First mentioned on Significant Tropical Weather Advisory as an area of persistent convection with maximum sustained surface winds estimated at 10-15 kt (5-7 m/sec) and a minimum sea-level pressure estimated at 1006 mb.
- 201400Z - Tropical Cyclone Formation Alert based on increased organization of central convection and improved outflow.
- 210600Z - First warning due to improved upper- and lower-level organization, increased deep convection, increasing wind speeds in the synoptic data, and a CI 2.5.
- 211200Z - Upgrade to tropical storm based on continued improvement in organization, increased deep convection, and good symmetrical outflow in all quadrants.
- 211800Z - Upgraded to typhoon following development of a 20 nm (35 km) diameter circular eye, continued rapid intensification, and a CI 4.0.
- 230600Z - Upgrade to a super typhoon based on continued warming of eye temperature and a CI 6.5.
- 240000Z - Downgrade to typhoon intensity based on observed vertical wind shear and restricted outflow in all quadrants except the southwest.
- 270000Z - Upgrade to super typhoon intensity based on warm eye temperature, cold surrounding convective cloud tops and a Dvorak current intensity estimate of 6.5.
- 280000Z - Downgrade to typhoon intensity based on increased vertical wind shear and restricted outflow to the east.
- 300600Z - Downgraded to a tropical storm due to increased vertical wind shear and exposed low-level circulation.
- 020600Z - (December) Downgrade to a tropical depression based on lack of deep central convection and decreased organization.
- 031800Z - Final warning followed further decrease in cloud organization and associated convection.

III. TRACK & MOTION

Owen developed out of a convective cluster 860 nm (1590 km) southwest of Hawaii near Palmyra Island, and was initially mentioned on a CPHC advisory. The system tracked westward across the central Pacific embedded in the tradewinds south of the subtropical ridge. It continued on this track until it reached the western Marshall Islands. Owen then slowed and tracked southwestward on 22 and 23 November as it approached an anticyclone located to the northwest (Figure 3-30-1). By 24 November, the omega block near the date line had dissipated and the mid-latitude westerlies returned to a more zonal flow. Owen then tracked west-southwestward along the

southern side of the subtropical ridge until 28 November. At that time, the typhoon entered an area dominated by broad low-level westerlies flowing into the recurving Typhoon Page (29W) (Figure 3-30-2). Owen's deep convection sheared apart late on 29 November and revealed an exposed low-

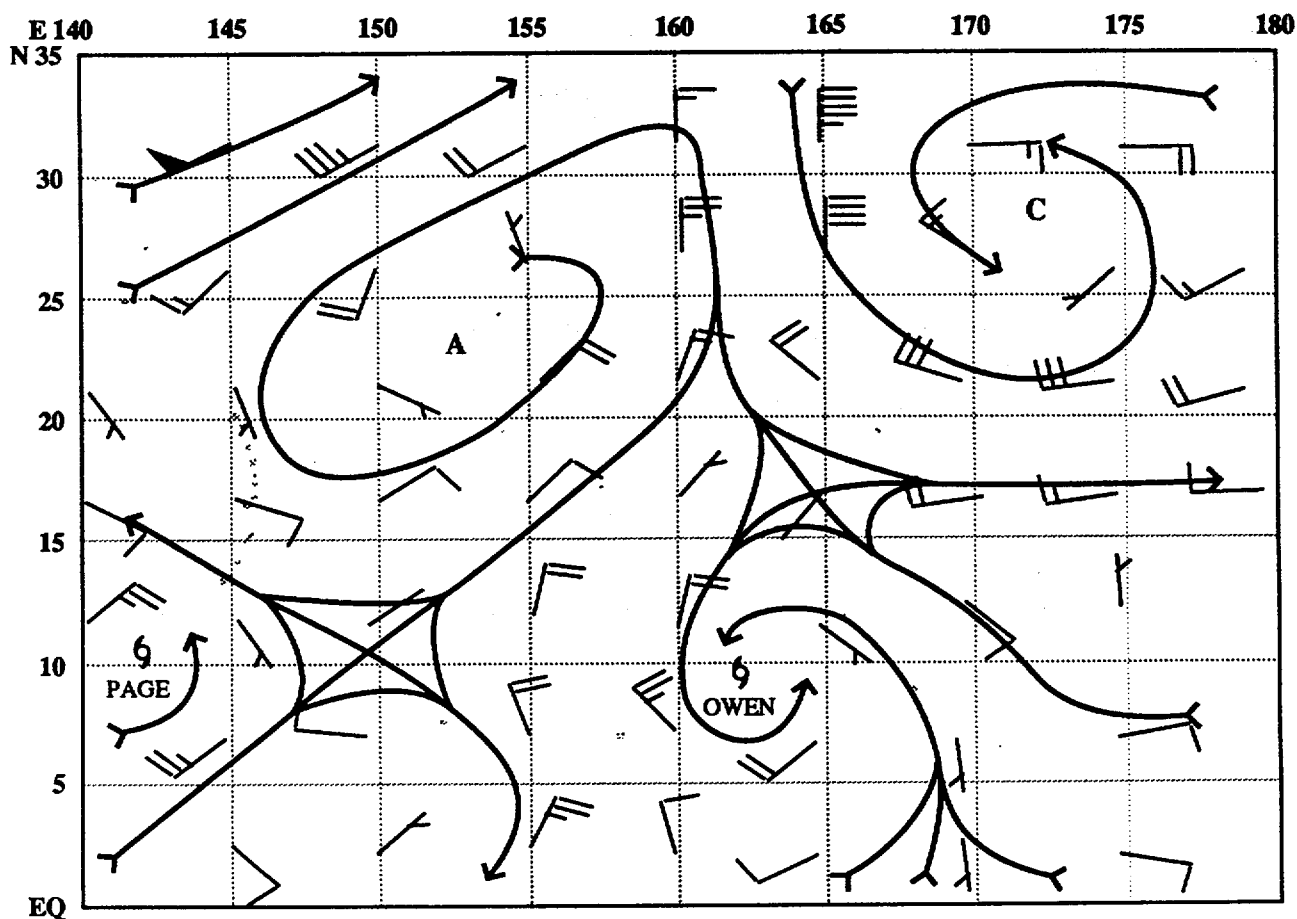
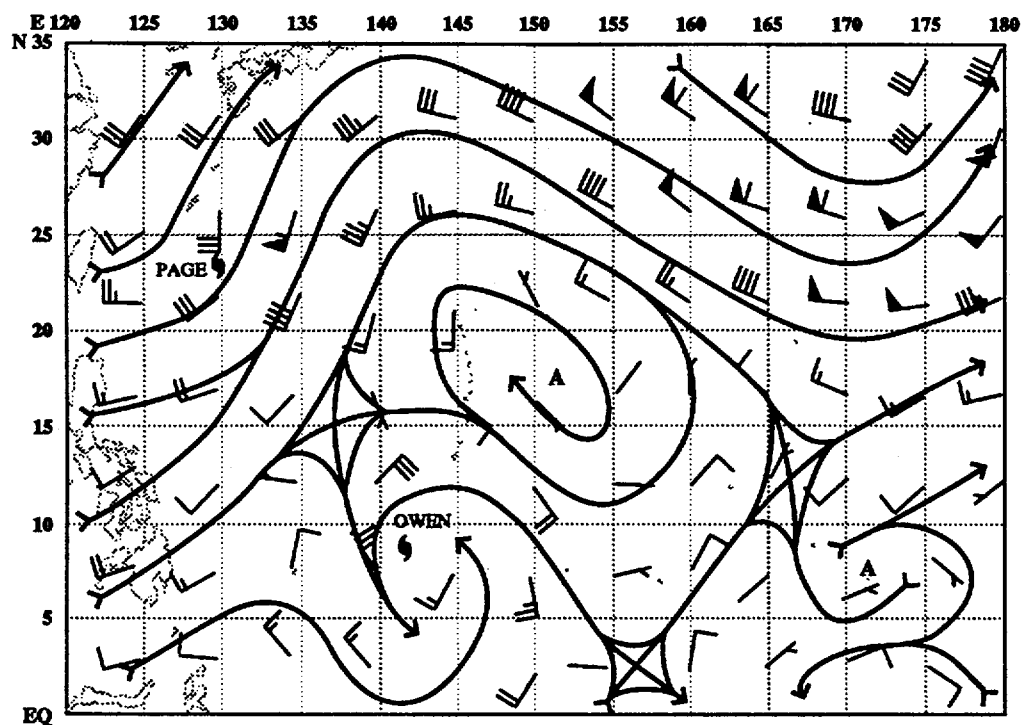
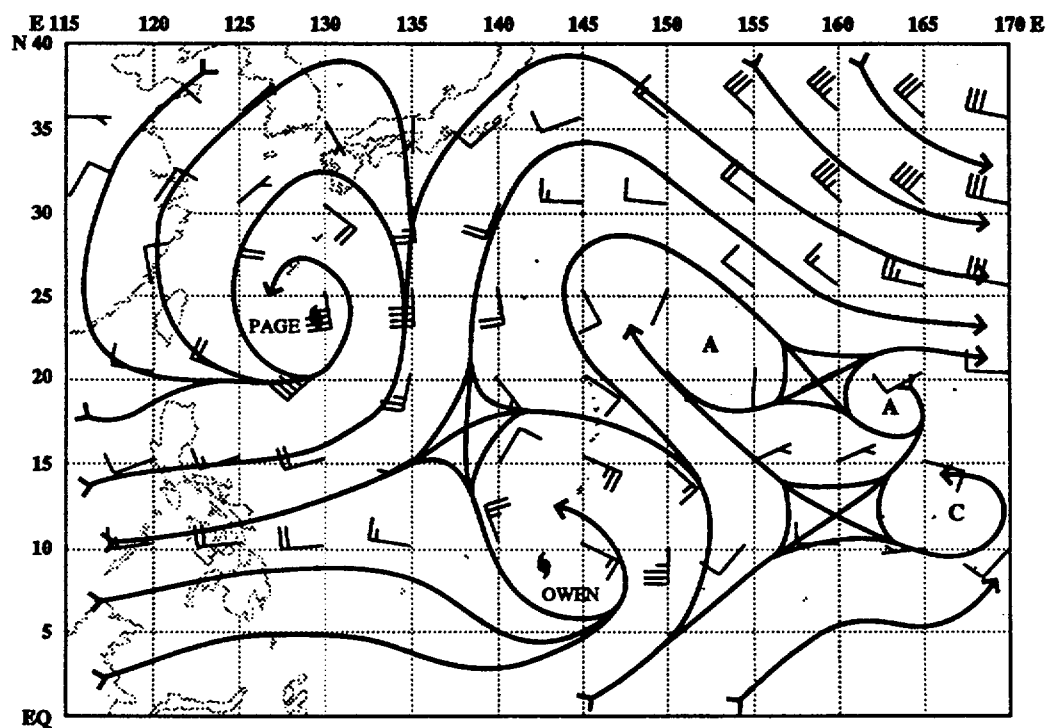


Figure 3-30-1. The 221200Z November NOGAPS 500-mb analysis shows an anticyclone to the northwest of Owen. The cyclonic circulation to the northeast of Owen is part of an omega block.

level circulation to the east of the major convection (Figure 3-30-3). This low-level circulation then tracked around the western periphery of the subtropical high until it encountered a shear line. Then, it turned southwestward, tracked down the shear line, and dissipated over the Celebes Sea on 05 December.



a)



b)

Figure 3-30-2. The 290000Z November NOGAPS analyses: a) for 500 mb depicting Owen near the western periphery of the subtropical high, and b) for 700 mb showing the deep westerly flow associated with Page (29W) which is located to the northwest of Owen.

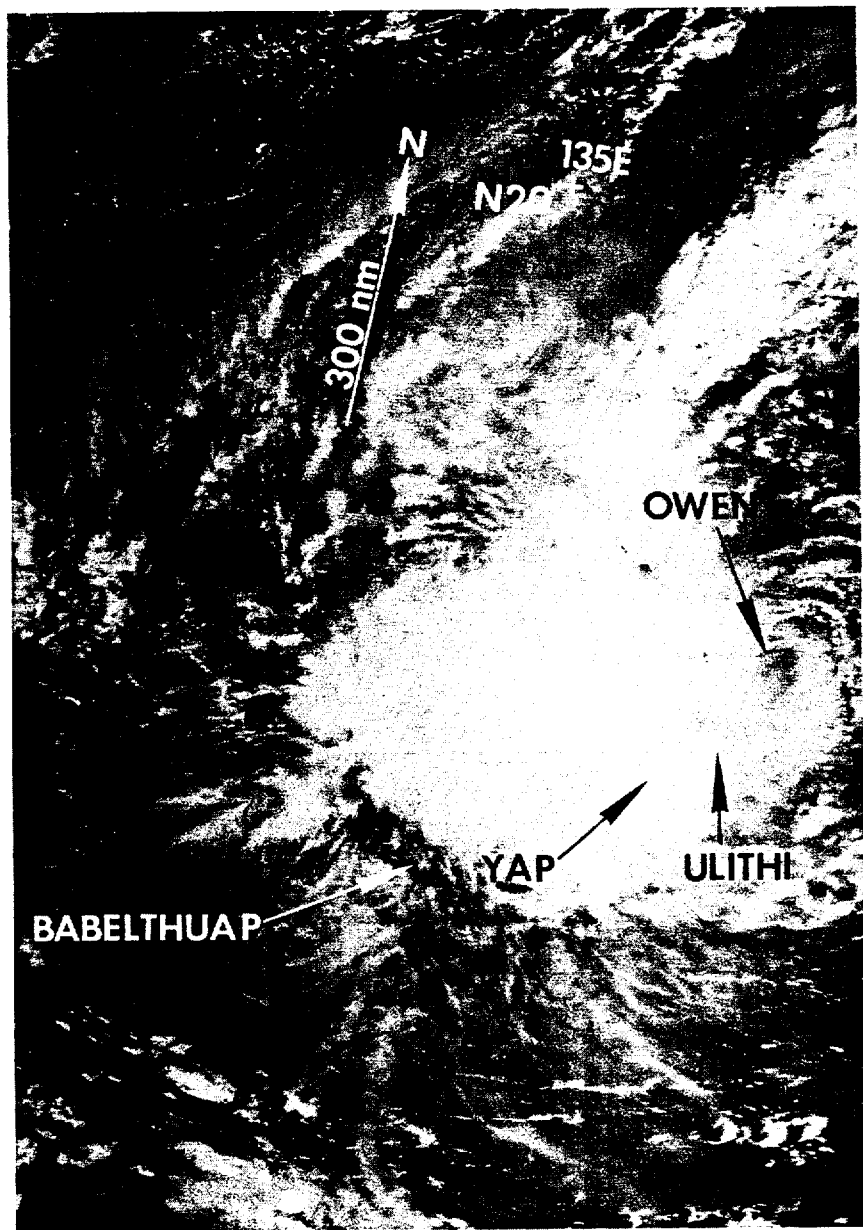


Figure 3-30-3. As Owen shears apart, the low-level circulation center appears to the east of the deep convective mass (300529Z November NOAA visual imagery).

IV. INTENSITY

The convective cloud mass that eventually became Owen formed southwest of Hawaii near Palmyra Island and maintained its continuity as it tracked across the central Pacific and past the date line. A discernible low-level circulation persisted, but the upper-levels did not favor further development. On 20 November, the convection flared and the overall organization started to improve as Owen entered an area of upper-level divergence and lighter winds (Figure 3-30-4). By 21 November, there were signs of an upper-level anticyclone forming over the disturbance and by the time of the second warning, surface pressures started dropping rapidly (Figures 3-30-5). By warning number 3, Owen had developed a 20 nm (35 km) diameter symmetric eye (Figure 3-30-6) and was well into its explosive intensification phase. Although tropical cyclones normally experience explosive intensification after reaching near typhoon intensity (Dunnavan, 1981), Owen commenced explosive intensification as a tropical depression, experiencing a drop in central pressure of 62 mb in 24 hours. Early stage development was supported by surface observations in

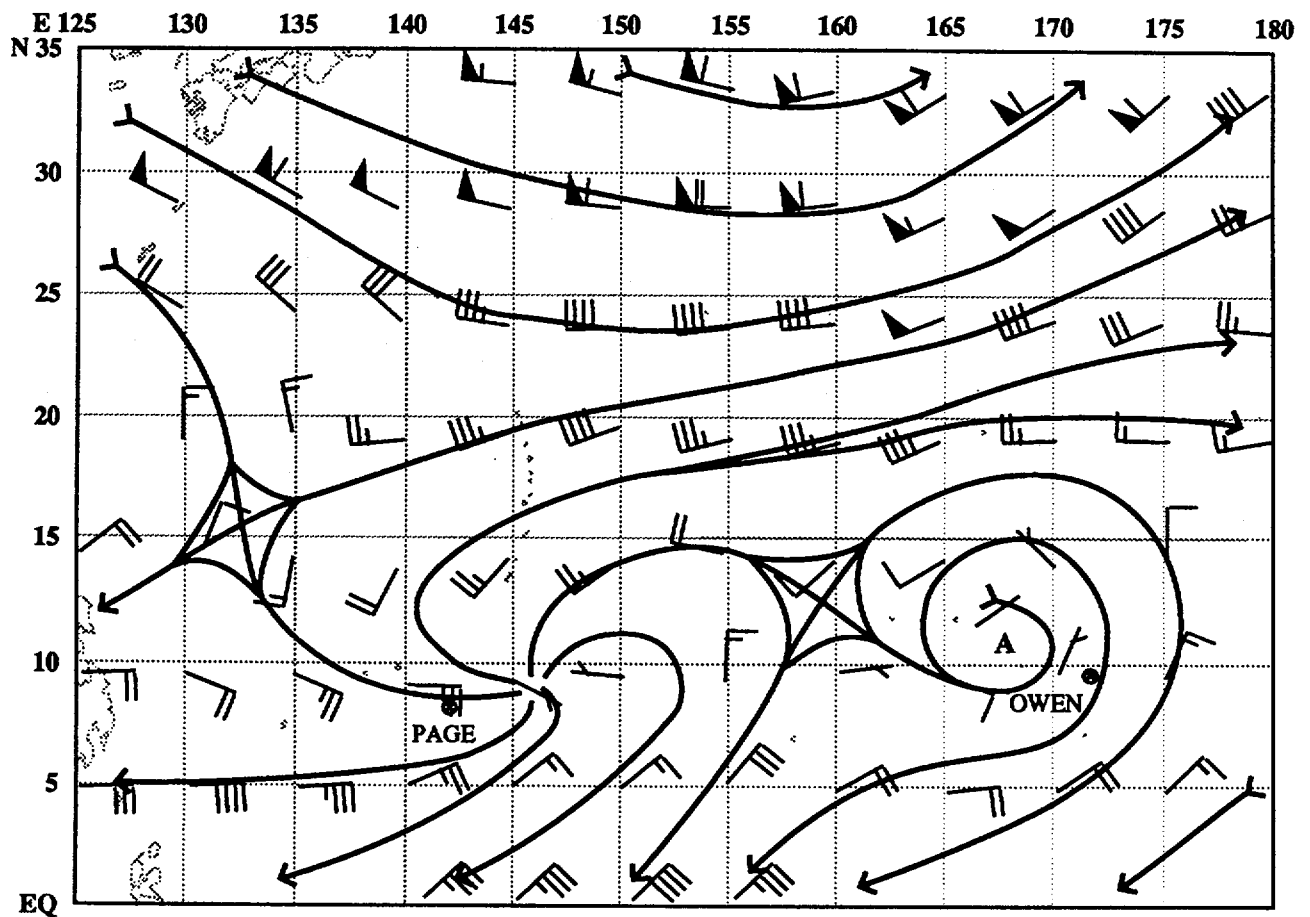


Figure 3-30-4. The 200000Z November 200-mb analysis shows Owen entering an area of lighter winds aloft.

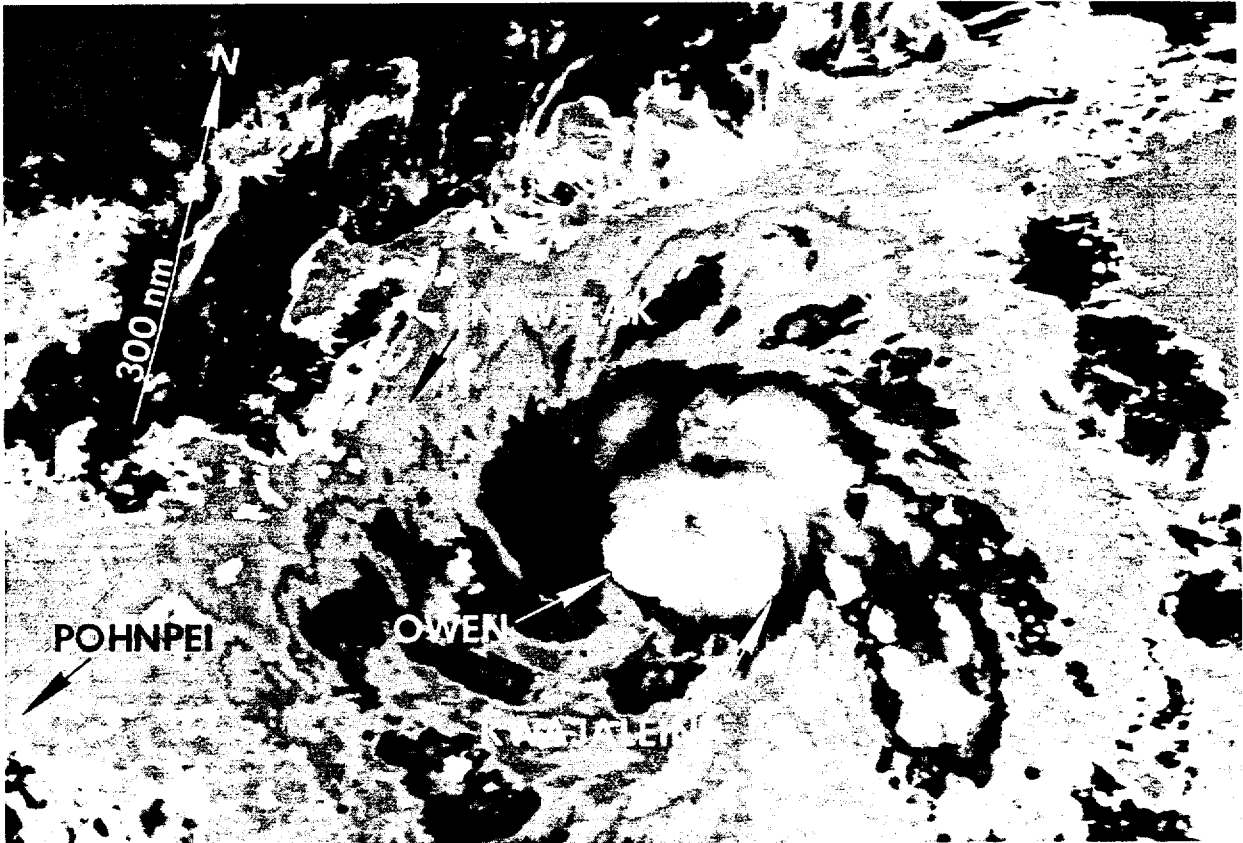


Figure 3-30-5. Tropical Storm Owen starting its explosive intensification phase (210822Z November DMSP enhanced infrared imagery).

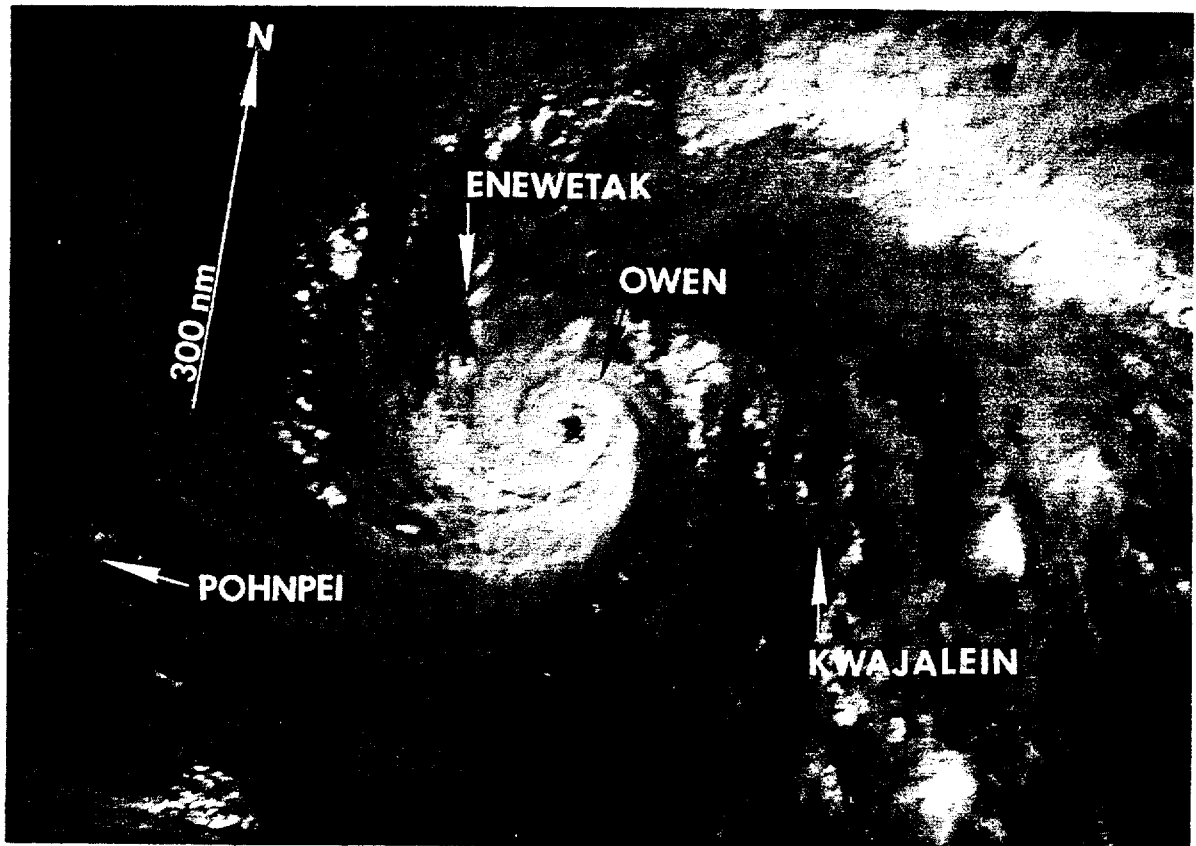


Figure 3-30-6. Fourteen hours after Figure 3-30-5, Owen has a symmetrical 20 nm (35 km) diameter eye and has reached typhoon intensity (212229Z November DMSP visual imagery).

the Marshall Islands and by radar observations from Kwajalein. Owen intensified from 30 kt (15 m/sec) to 105 kt (54 m/sec) in 24 hours, and peaked at 130 kt (67 m/sec) in 48 hours (Figure 3-30-7). Upon reaching super typhoon intensity (Figure 3-30-8), Owen moved into an area of increasing vertical wind shear and its outflow channel to the north was suppressed and eventually cut off by the convergence associated with a passing mid-latitude trough and the eastern side of the anticyclone located between Owen and Typhoon Page (29W). The vertical wind shear eased on 26 November

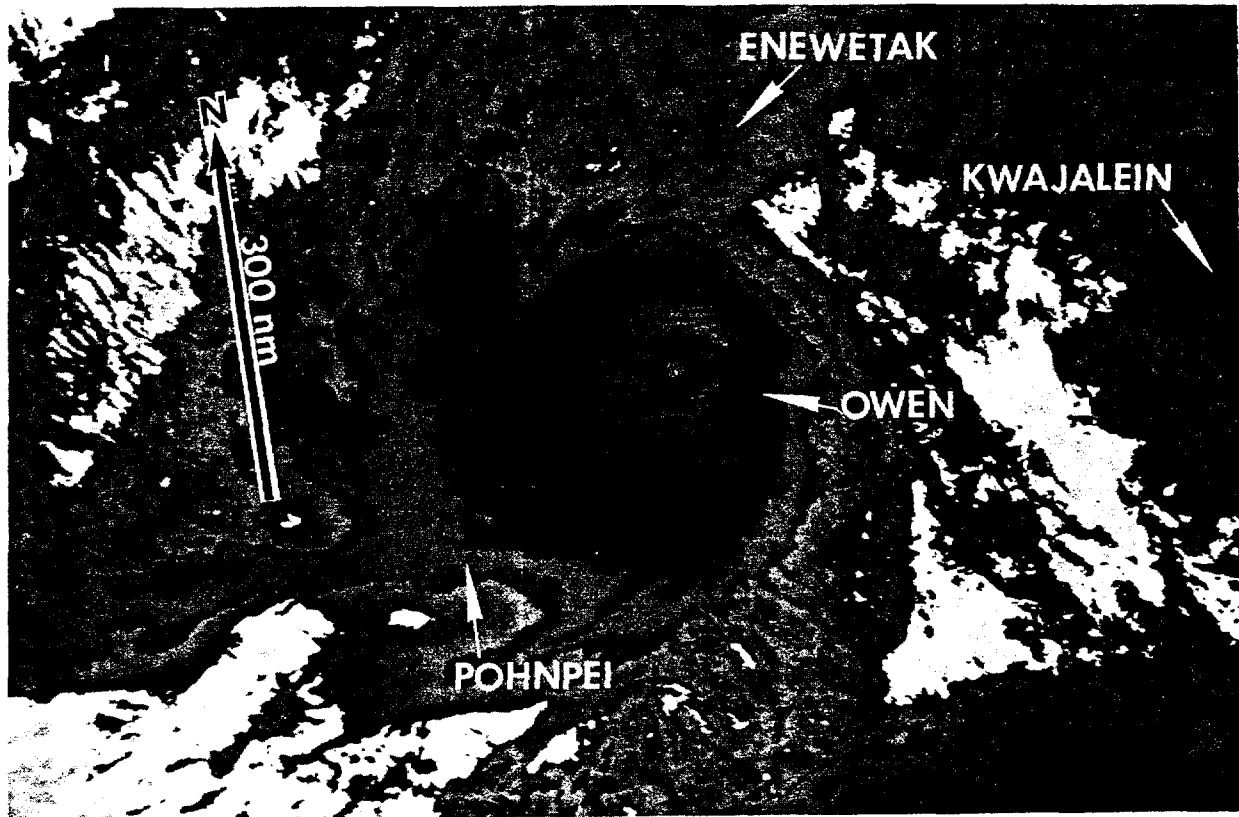


Figure 3-30-7. Super Typhoon Owen near its first peak in intensity (231048Z November DMSP enhanced infrared imagery).

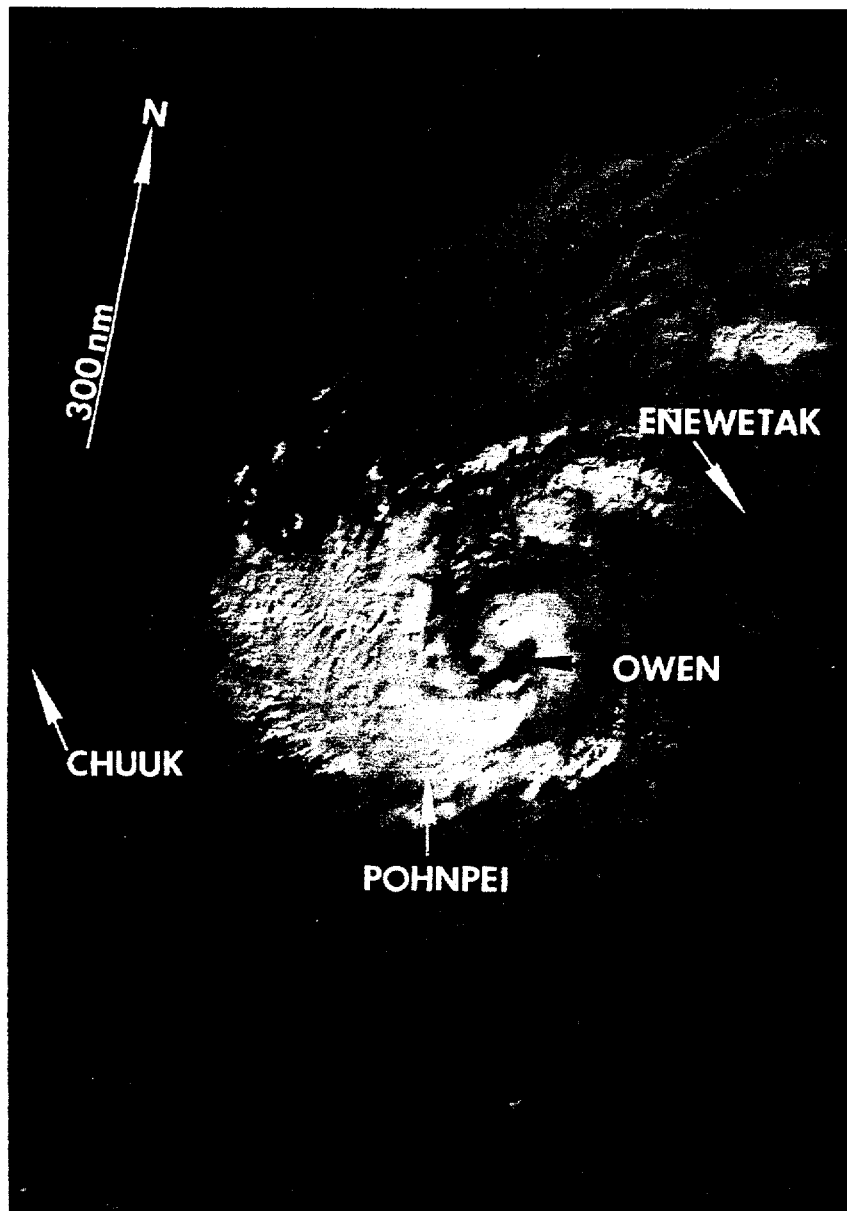


Figure 3-30-8. Shear from the northeast and restricted outflow to the north are evident, as Typhoon Owen weakens (242307Z November DMSP visual imagery).

permitting Owen to reintensify (Figure 3-30-9). The peak intensity of 140 kt (72 m/sec) was reached on 29 November and there was a significant shift in the position of the upper-level anticyclone (Figure 3-30-10). As the anticyclone shifted position, the upper-level shear from the east increased dramatically from approximately 10 kt to 40 kt. This environment persisted until Owen sheared apart late on 29 November with the upper-level convection continuing west-northwestward and the low-level circulation center moving north-northwestward. Owen never moved back into an environment favorable for redevelopment and only maintained scattered convection until it dissipated in the Celebes Sea (Figure 3-30-11).

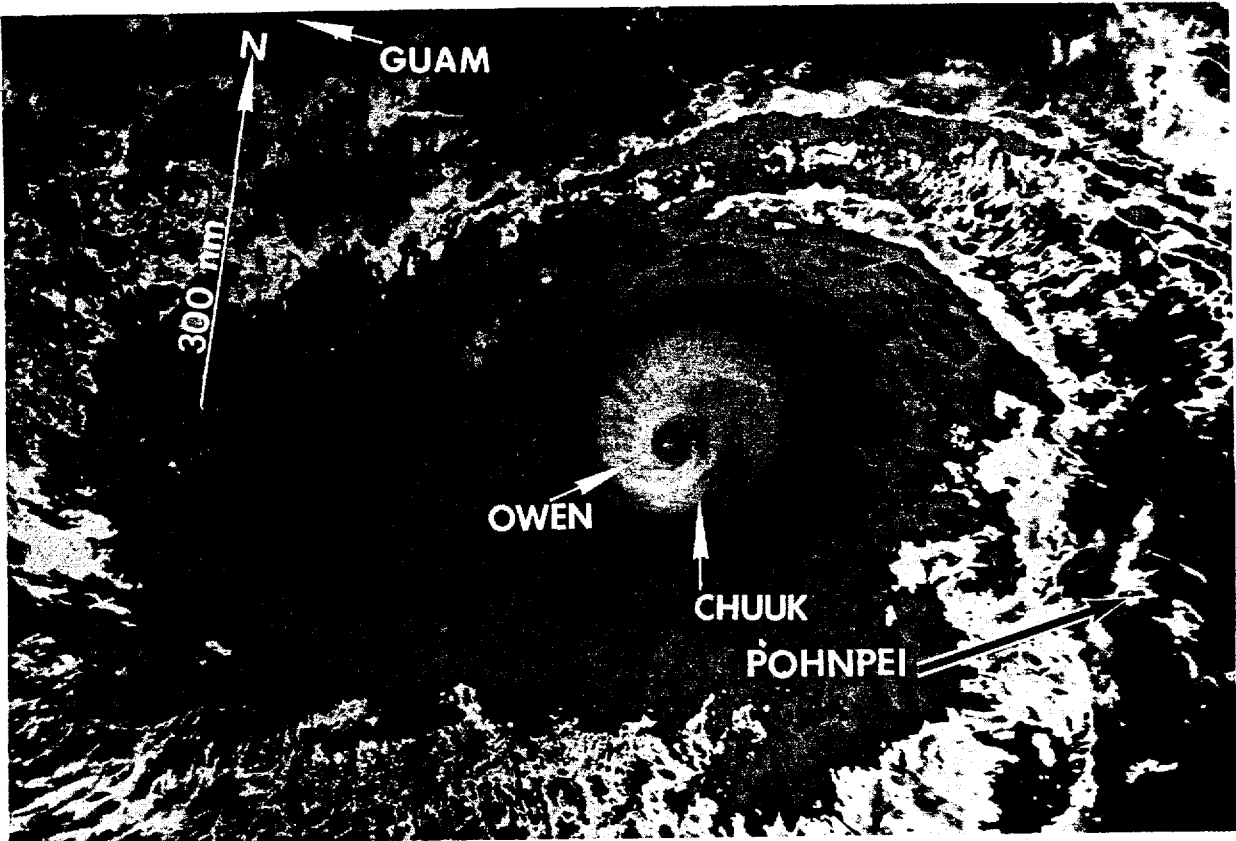
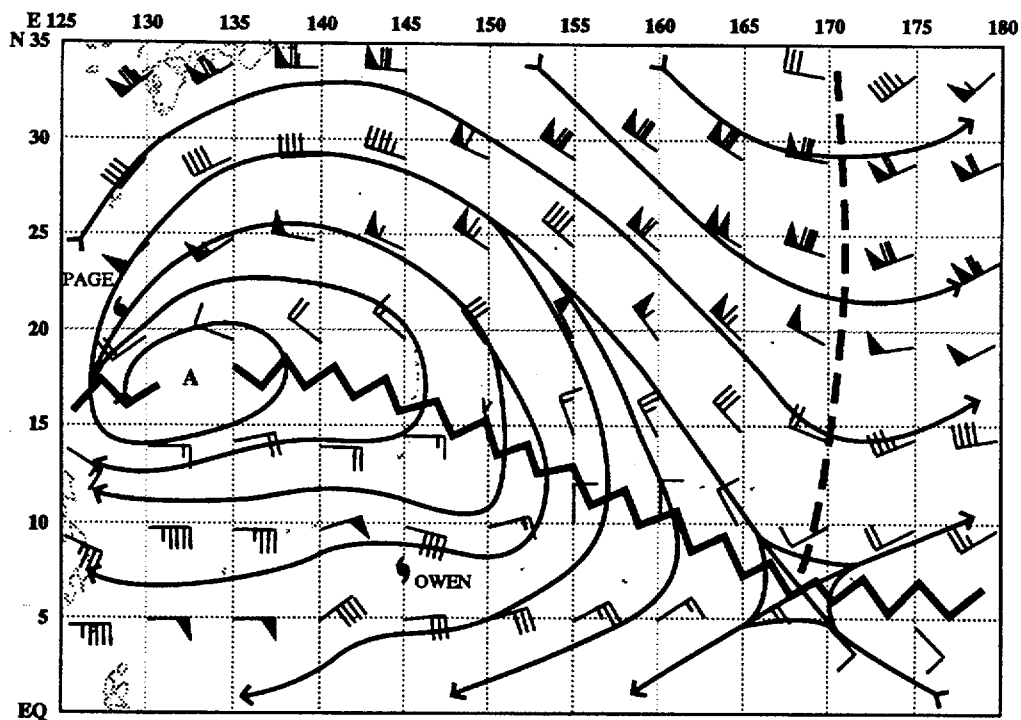
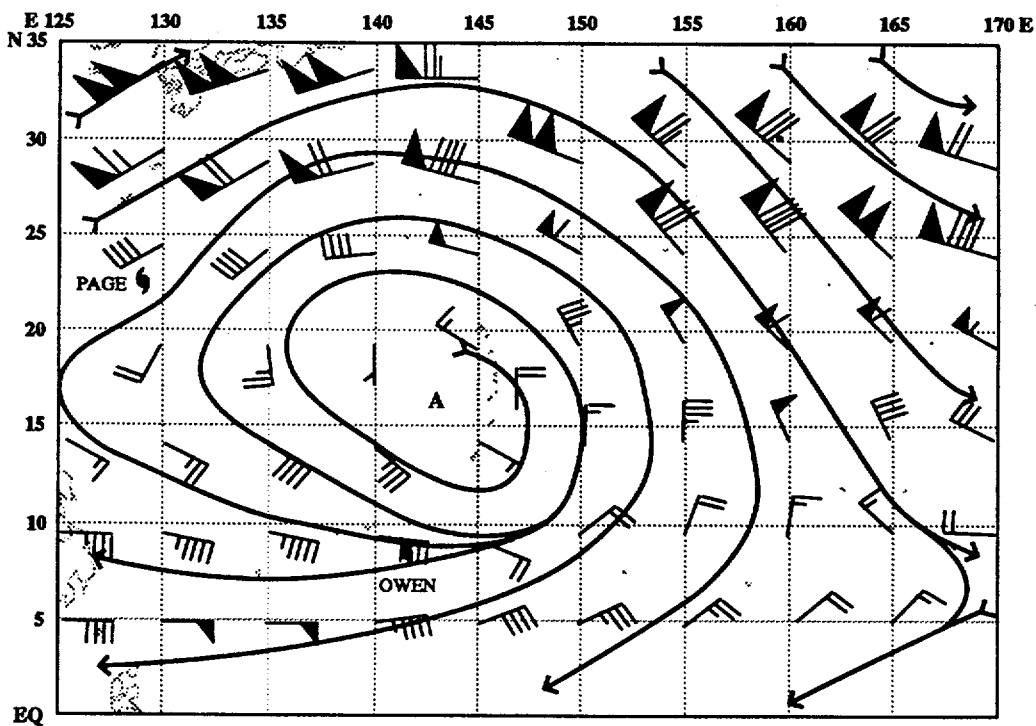


Figure 3-30-9. Owen after reintensifying to super typhoon intensity for a second time (270925Z November NOAA enhanced infrared imagery).



a)

Figure 3-30-10. The November NOGAPS 200-mb analyses: a) for 281200Z November showing Page (29W) and the center of the anticyclone to the northwest of Owen, and b) for 290000Z November showing the relocation of the center of the anticyclone to the north of Owen.



b)

V. FORECASTING PERFORMANCE

JTWC's forecast performance is shown in Figure 3-30-12. Overall errors for this system were well below the long term average because JTWC did well forecasting the speed and speed changes exhibited by Owen. JTWC forecasts were generally right of track until 28 November for a

number of reasons. First, the usually dependable NOGAPS deep layer mean provided guidance that indicated northwestward movement. Second, Owen took an anomalous track to the west-southwest. Finally, NOGAPS after 26 November consistently forecast the anticyclone steering Owen to reposition itself east of Guam sooner than the 29 November timeframe when the shift actually occurred. JTWC forecast recurvature early as a result, and the recurvature forecast had to be adjusted back to the west. Once Owen sheared, forecast guidance was based on the NOGAPS 700 mb and lower levels. JTWC accurately forecast the initial peak intensity and the subsequent weakening. The reintensification was correctly reflected in the forecasts, but the maximum intensity was under forecast. Since Owen was not forecast to shear apart, the final weakening trend was significantly faster than forecast.

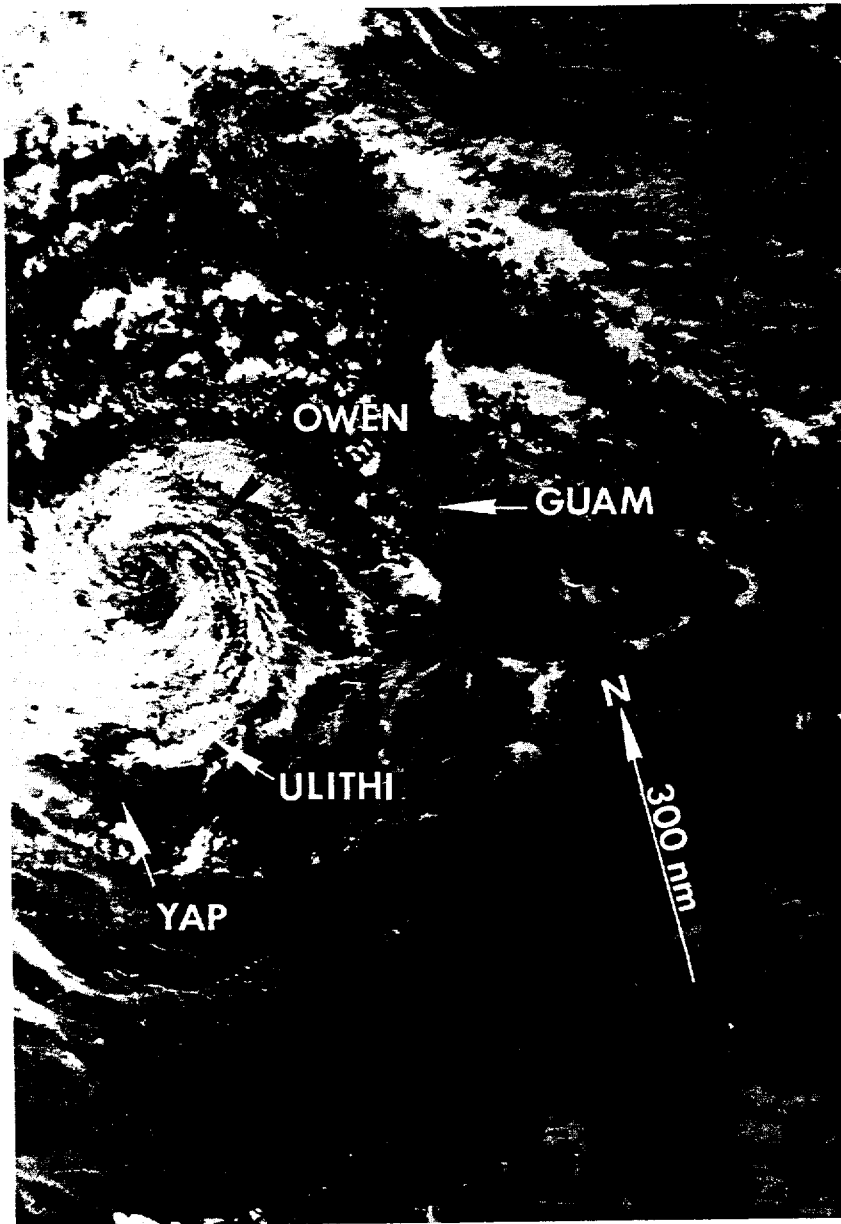


Figure 3-30-11. Owen's low-level circulation is fully exposed (301632Z November NOAA infrared imagery).

VI. IMPACT

POHNPEI

CHUUK STATE

- 2 killed when a live power line fell and struck them.
- declared a U.S. federal disaster area, 1000 people left homeless, major power failures.

HALL ISLANDS

- extensive crop damage, nearly all homes destroyed, all food crops destroyed.

NAMONUITO ATOLL

- extensive crop damage, nearly all homes destroyed, all food crops destroyed.

PULAP ATOLL

- extensive crop damage, 99 percent of homes destroyed.

YAP STATE

- declared a U. S. federal disaster area.

SATAWAL ISLAND

- reported winds in excess of 100 mph, 95 percent food crop destroyed, 90 percent homes damaged, all power lost.

LAMOTREK ATOLL

- reported winds in excess of 100 mph, 85 percent homes destroyed, 95 percent food crop destroyed, all power lost.

ELATO ATOLL

- 99 percent dwellings destroyed, 90 percent food crops destroyed.

IFALIK ATOLL

- dwellings - no report, 95 percent food crops destroyed, 20 percent land eroded.

WOLEAI ATOLL

- 85 percent dwellings, 90 percent food crops destroyed.

FARAULEP ATOLL**

- 20 percent dwellings, 100 percent canoes, 100 percent food crops destroyed, 20 - 30 percent land eroded.

ULITHI ISLAND

- 30 percent dwellings and government buildings, 100 percent food crops destroyed.

** NOTE: AMOS site on Faraulep was lost during passage of Owen. The shore was completely eroded away leaving the tower on its side in 10 feet of water and 20 yards off the beach. Site is now abandoned.

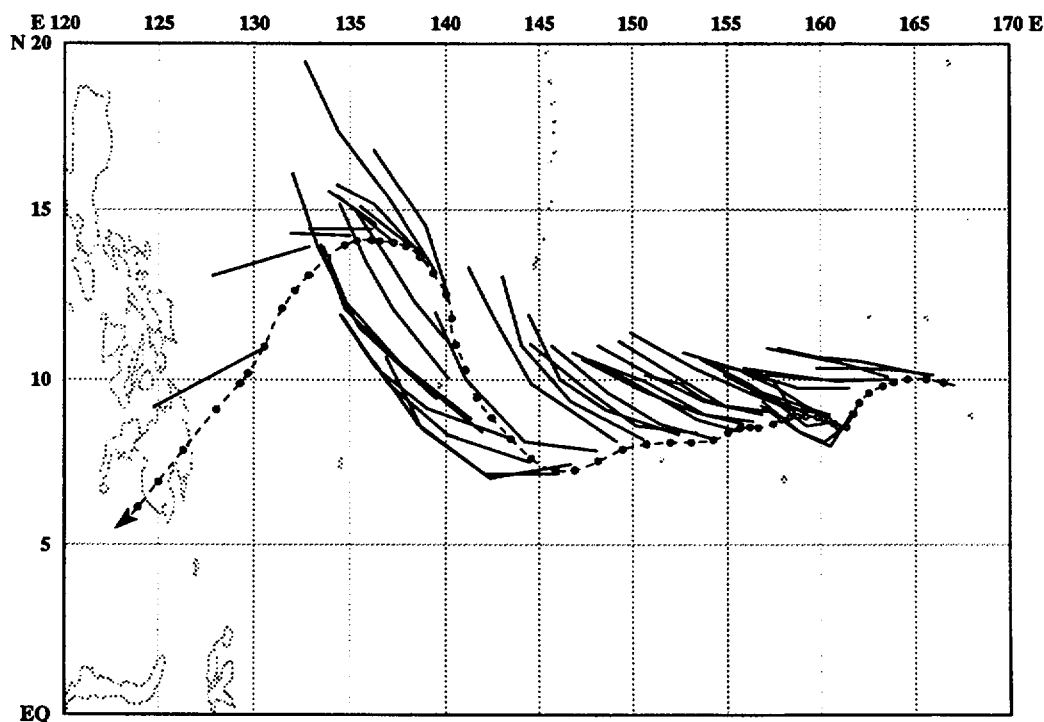
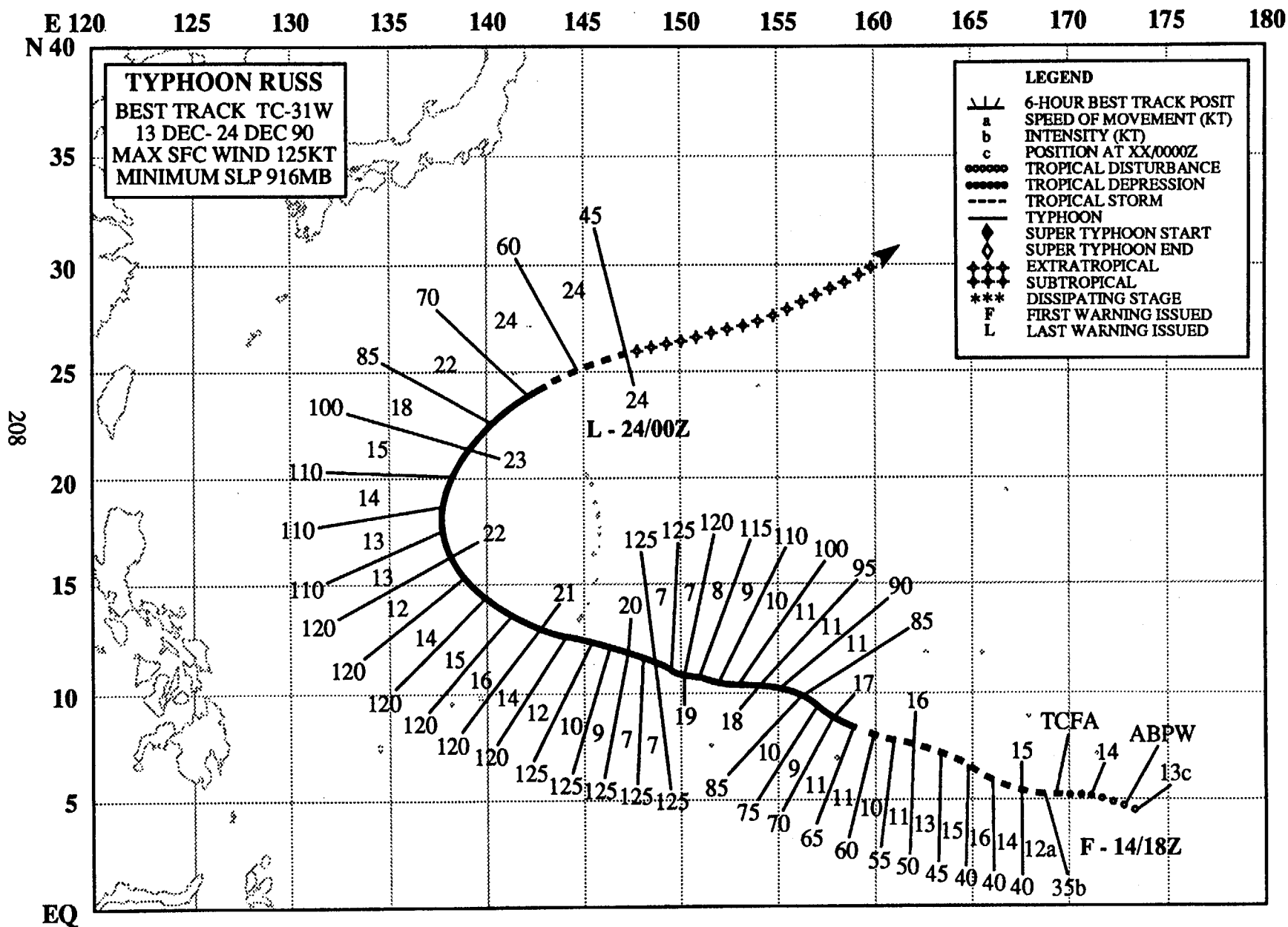


Figure 3-30-12. The JTWC forecast tracks (solid lines) for Owen superimposed on the final best track (dashed line).



TYPHOON RUSS (31W)

I. HIGHLIGHTS

Russ, the last western North Pacific tropical cyclone of 1990, was the most severe to strike Guam in 14 years. Damage was estimated as high as 120 million dollars. Russ formed in the Marshall Islands, tracked west-northwestward and intensified to near super typhoon intensity as it approached Guam. The typhoon passed within 30 nm (55 km) of the southern tip of Guam and brought typhoon force winds which caused extensive damage, especially to the southern portion of the island. After leaving Guam, Russ slowly weakened, recurved and became an extratropical cyclone.

II. CHRONOLOGY OF EVENTS

- 130600Z - First mentioned on a Significant Tropical Weather Advisory as an area of persistent convection associated with a low-level cyclonic circulation and an estimated minimum sea-level pressure of 1004 mb. The potential for significant tropical cyclone development was assessed as poor.
- 140600Z - Second mention on a Significant Tropical Weather Advisory due to persistent convection with an anticyclone developing aloft. Potential for development upgraded to fair.
- 141330Z - Tropical Cyclone Formation Alert issued based on increased curvature in the spiral cloud bands and a 35 kt (13 m/sec) surface wind report from Jaluit Atoll (WMO 91369).
- 141800Z - First Warning issued and Russ upgraded to tropical storm intensity prompted by rapid increase in amount and organization of the central convection.
- 161800Z - Upgraded to a typhoon based on anticipated appearance of an eye and a satellite intensity estimate of 65 kt (35 m/sec).
- 190600Z - Peak intensity - 125 kt (65 m/sec) - followed observation of further drying and warming within the 30 nm (55 km) diameter eye and a satellite intensity estimate of 125 kt.
- 240000Z - Final warning issued with Russ downgraded to tropical storm intensity and transitioning to an extratropical cyclone after the loss of its persistent central dense overcast.

III. TRACK AND MOTION

Russ developed in the near-equatorial trough in the southern Marshall Islands. The tropical cyclone followed a basic recurvature track, passing just south of Guam and recurving through the axis of the subtropical ridge to the northwest of Guam. Although Russ maintained an essentially west-northwestward direction of motion as it approached Guam, significant changes in speed of motion occurred. Beginning on 18 December, Russ began to decelerate in response to the passage of a mid-latitude short wave passing to the north of the subtropical ridge. By 19 December, the typhoon had slowed to 7 kt (13 km/hr) - almost half the 13 kt (24 km/hr) speed expected from climatology. Once the short wave passed to the northeast, the subtropical ridge and the steering flow strengthened, and on 20 December, Russ started to accelerate. Fortunately for Guam, this reduced the time of exposure to Russ' damaging winds. By the time Russ entered the Philippine Sea, another short wave had moved eastward from Asia and caused a break in the subtropical ridge to the northwest of Guam. Russ recurved through this break, accelerated and became an extratropical cyclone on 24 December.

IV. INTENSITY

Russ' initial intensification was surprisingly rapid. As a result, Russ was at minimal tropical storm intensity when the first warning was issued. Although satellite imagery showed poorly

organized convection with multiple circulations (Figure 3-31-1), surface wind reports from Jaluit Atoll (WMO 91369) in the southern Marshall Islands of 35 and 40 kt (17 and 20 m/sec) at 141200Z and 150000Z respectively, revealed that the tropical cyclone was consolidating. After this sudden initial development (Figure 3-31-2), Russ intensified (Figure 3-31-3) at a normal rate until it reached 125 kt (64 m/sec) at 190600Z. The passage of a mid-latitude short wave trough, which weakened the subtropical ridge and caused Russ to slow down, aided intensification by enhancing the typhoon's

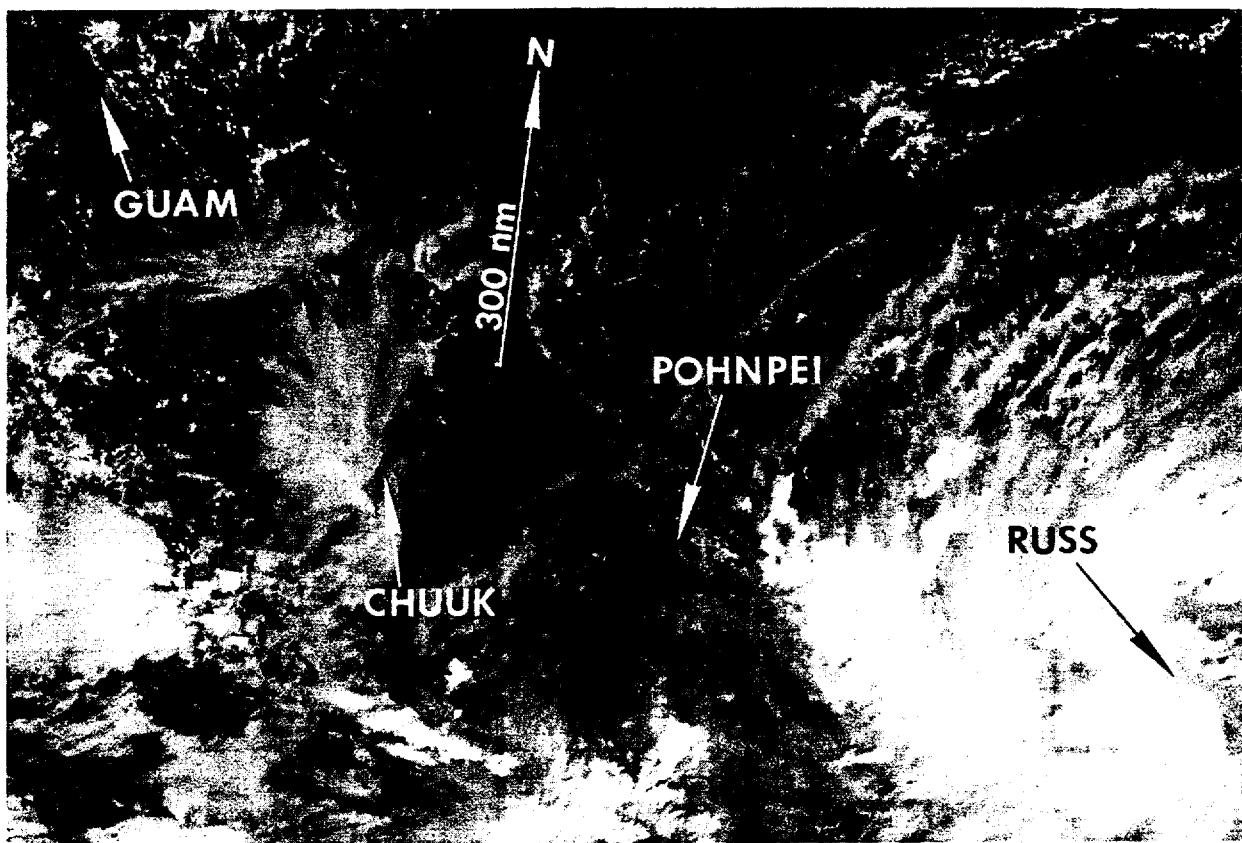


Figure 3-31-1. Russ after reaching tropical storm intensity. Central convection and outflow are well organized (142250Z December DMSP visual imagery).

outflow aloft into the polar westerlies (Figure 3-31-4). The tropical cyclone remained near its peak intensity for three days. During this time, it passed within 30 nm (55 km) of the southern tip of Guam (Figure 3-31-5). The closest point of approach at 201700Z (210300 local time on Guam) was reflected in the lowest pressure (Figure 3-31-6), increased wind (Figure 3-31-7), and increased seas (Figure 3-31-8). Maximum sustained winds experienced on the island, which is only 30 nm (55 km) in length, varied from minimum typhoon intensity in the north to almost double that in the south (Figure 3-31-9).

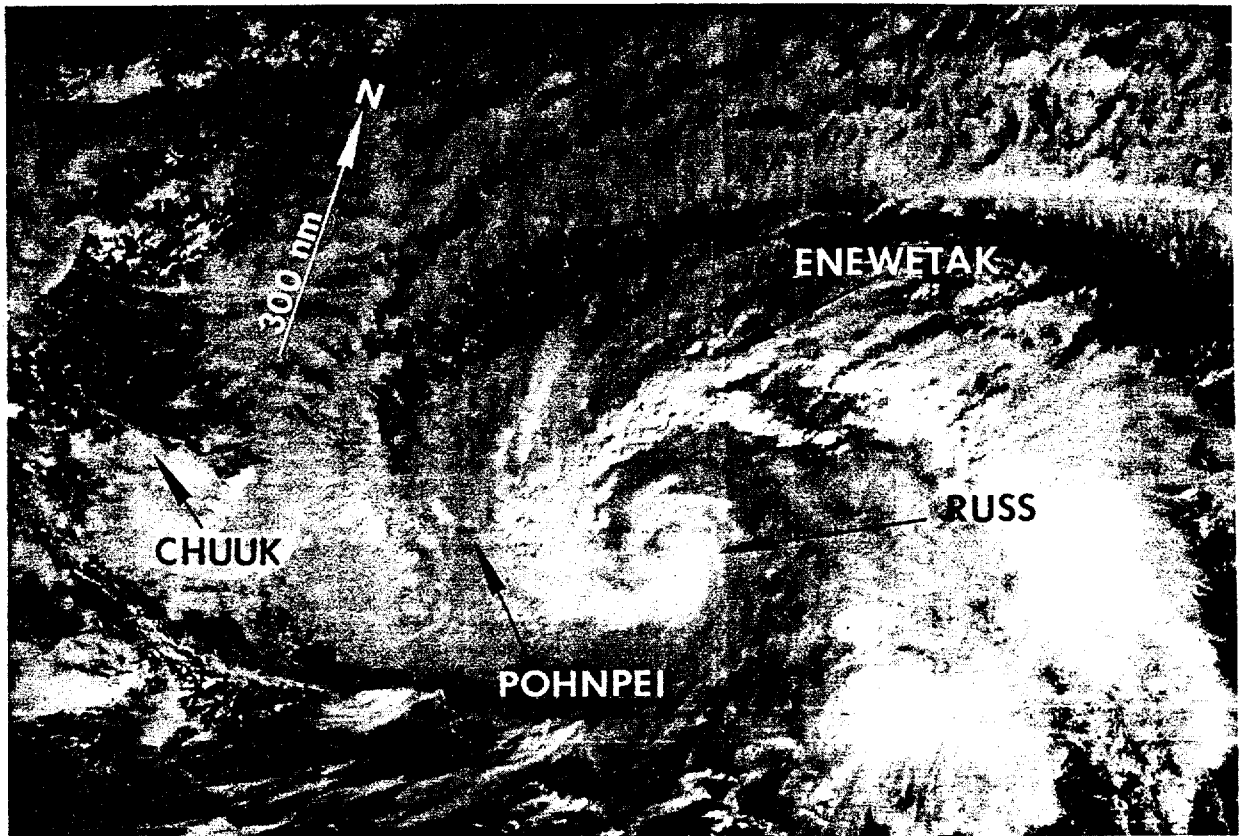


Figure 3-31-2. Spiral cloud band curvature increases as Russ intensifies (152229Z December DMSP visual imagery).

After passing to the west of Guam and into the Philippine Sea, Russ started to slowly weaken as it turned more to the north and interacted with the stronger polar westerly winds aloft (Figure 3-31-10). The typhoon's compact central convection resisted the increased vertical wind shear until 24 December, a day after recurvature. By then, the supporting deep convection was displaced to the north and east of the low-level circulation center and the cyclone (Figure 3-31-11) was extratropical.

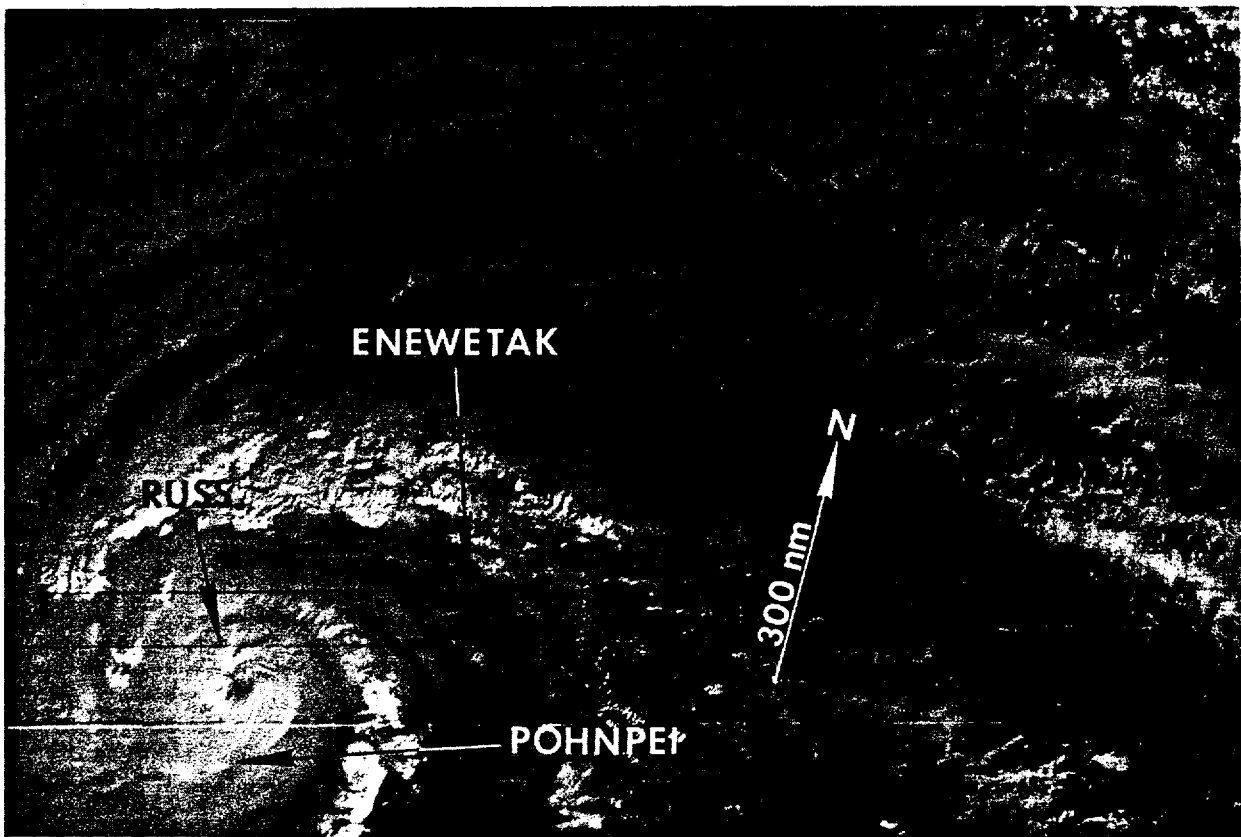


Figure 3-31-3. Russ develops an eye and reaches typhoon intensity (162207Z December DMSP visual imagery).

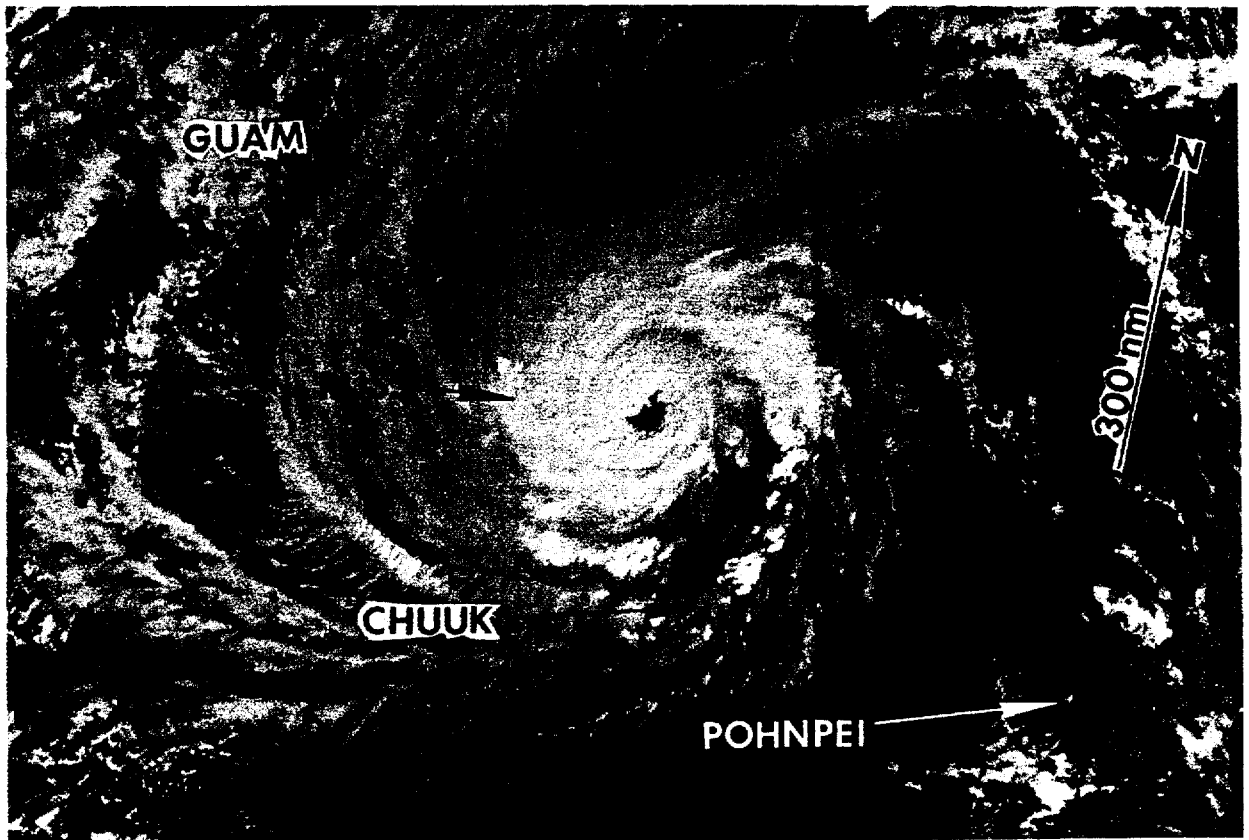


Figure 3-31-4. Russ at near peak intensity after the passage of a mid-latitude short wave to the north enhanced its outflow aloft (182307Z December DMSP visual imagery).

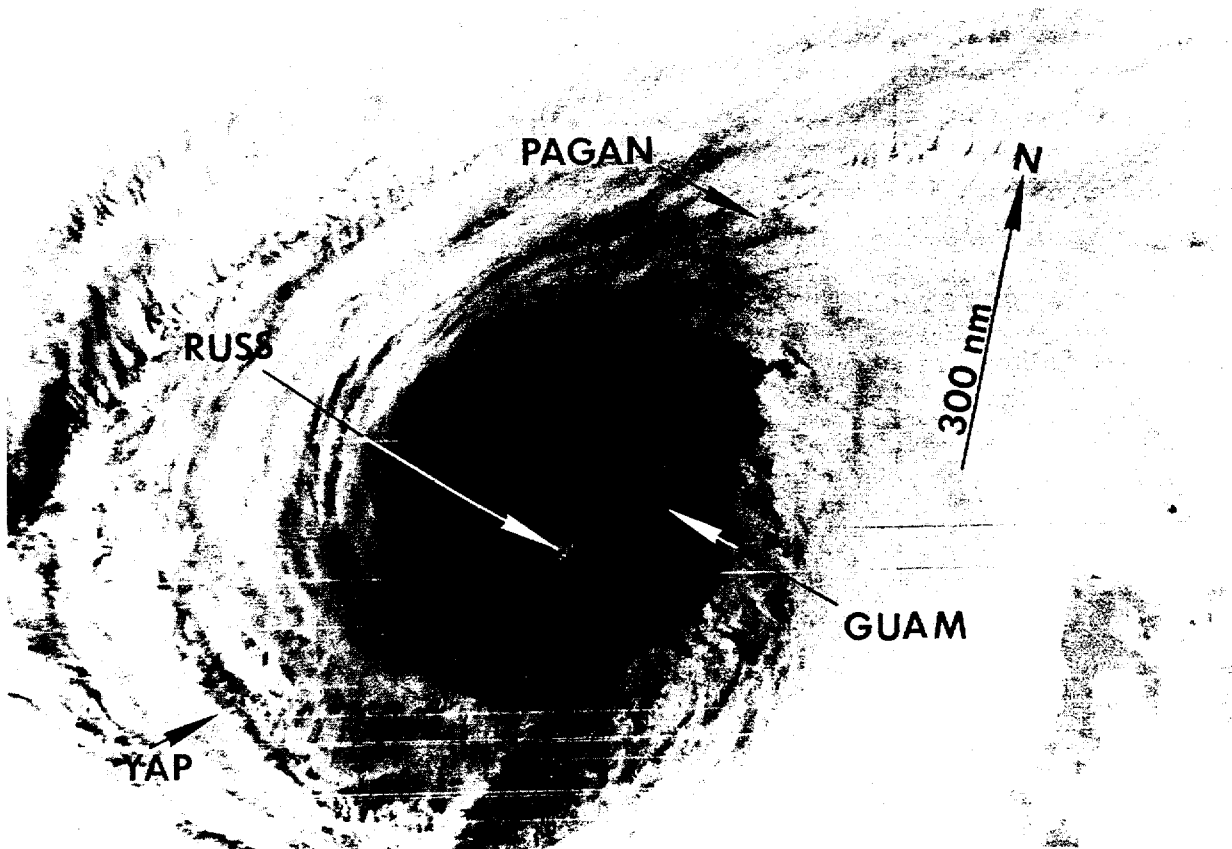


Figure 3-31-5. Russ after its closest point of approach to Guam (202014Z December DMSP infrared imagery).

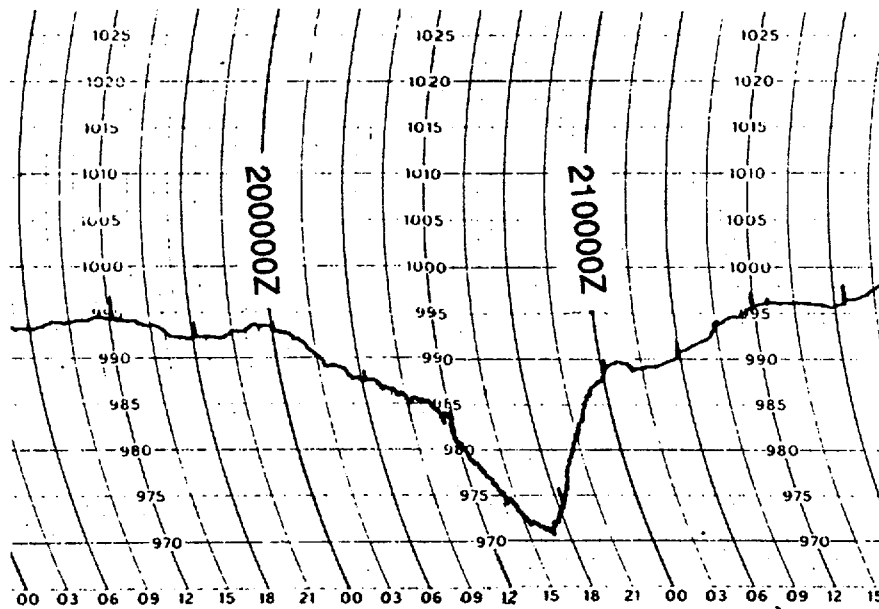


Figure 3-31-6. The microbarograph trace from Naval Air Station (WMO 91212), Agana, Guam shows its lowest pressure of 971 mb, at 210000Z, as Russ is near its closest point of approach to Guam.

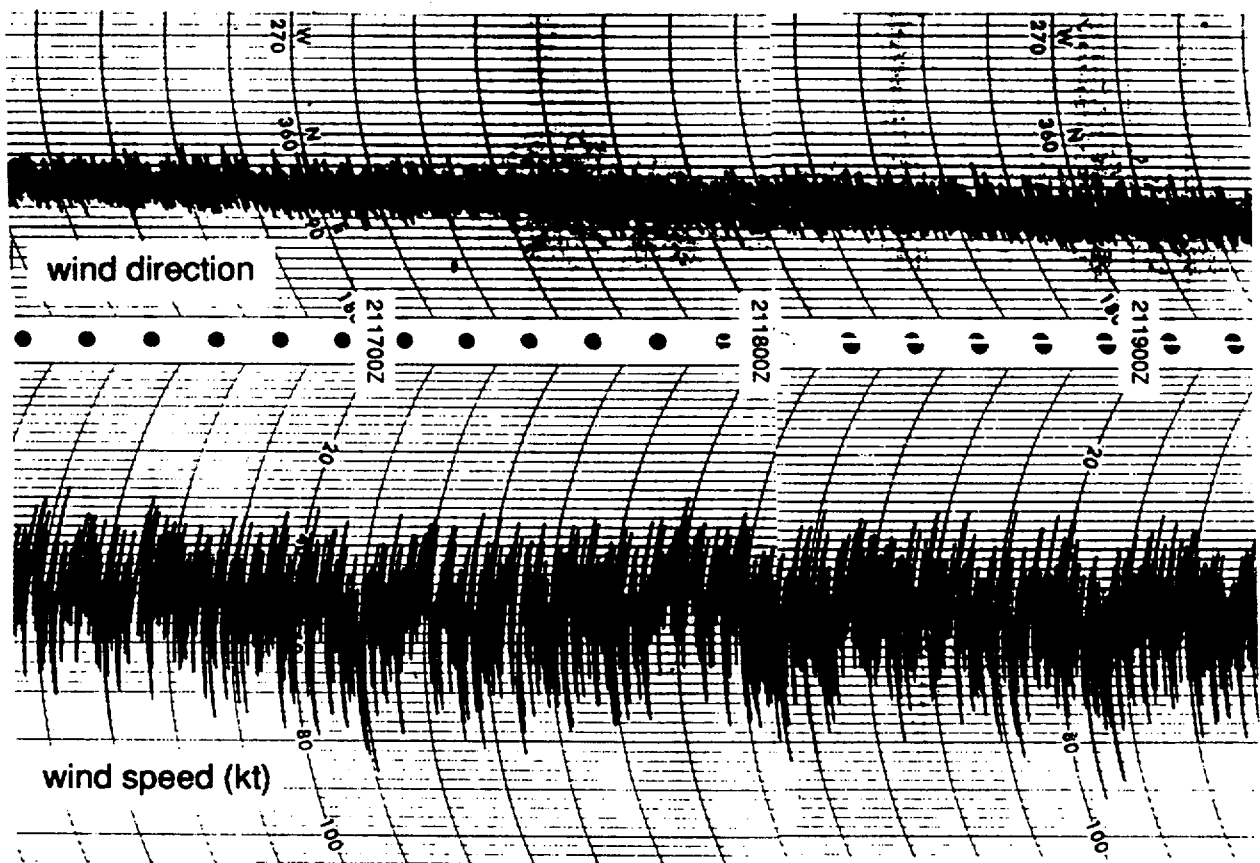


Figure 3-31-7. The wind record from Naval Air Station (WMO 91212), Agana, Guam reflects a steady increase from 211500Z through 211700Z as Russ approaches.

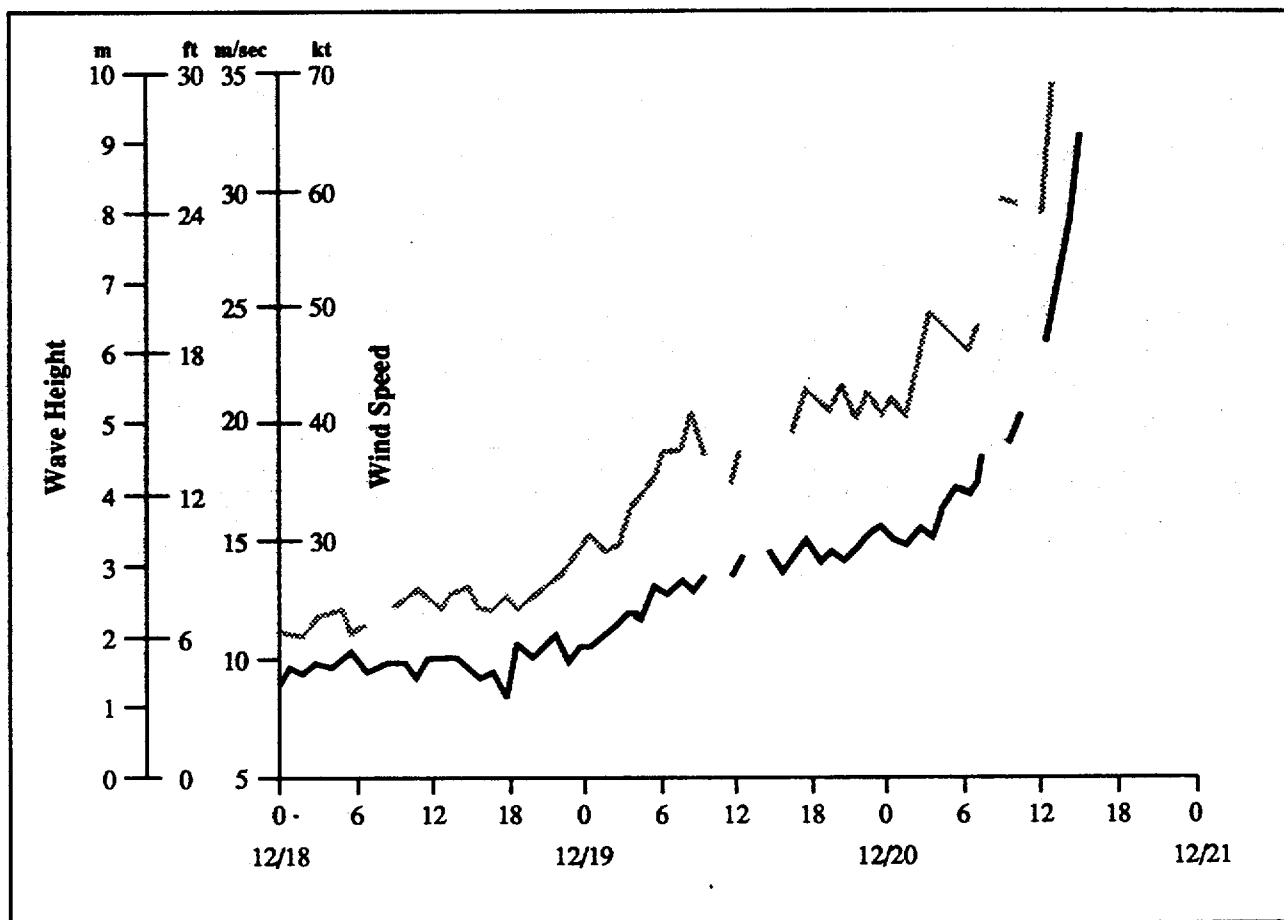


Figure 3-31-8. Time series plot of wave height (gray line) and wind speeds (black line) from a buoy moored 7 nm (13 km) west of the southern tip of Guam shows 30 ft (10 m) seas and 65 kt (33 m/sec) winds. The buoy was in the lee of Guam, but was lost shortly before Russ' CPA. (Data courtesy of the National Data Buoy Center)

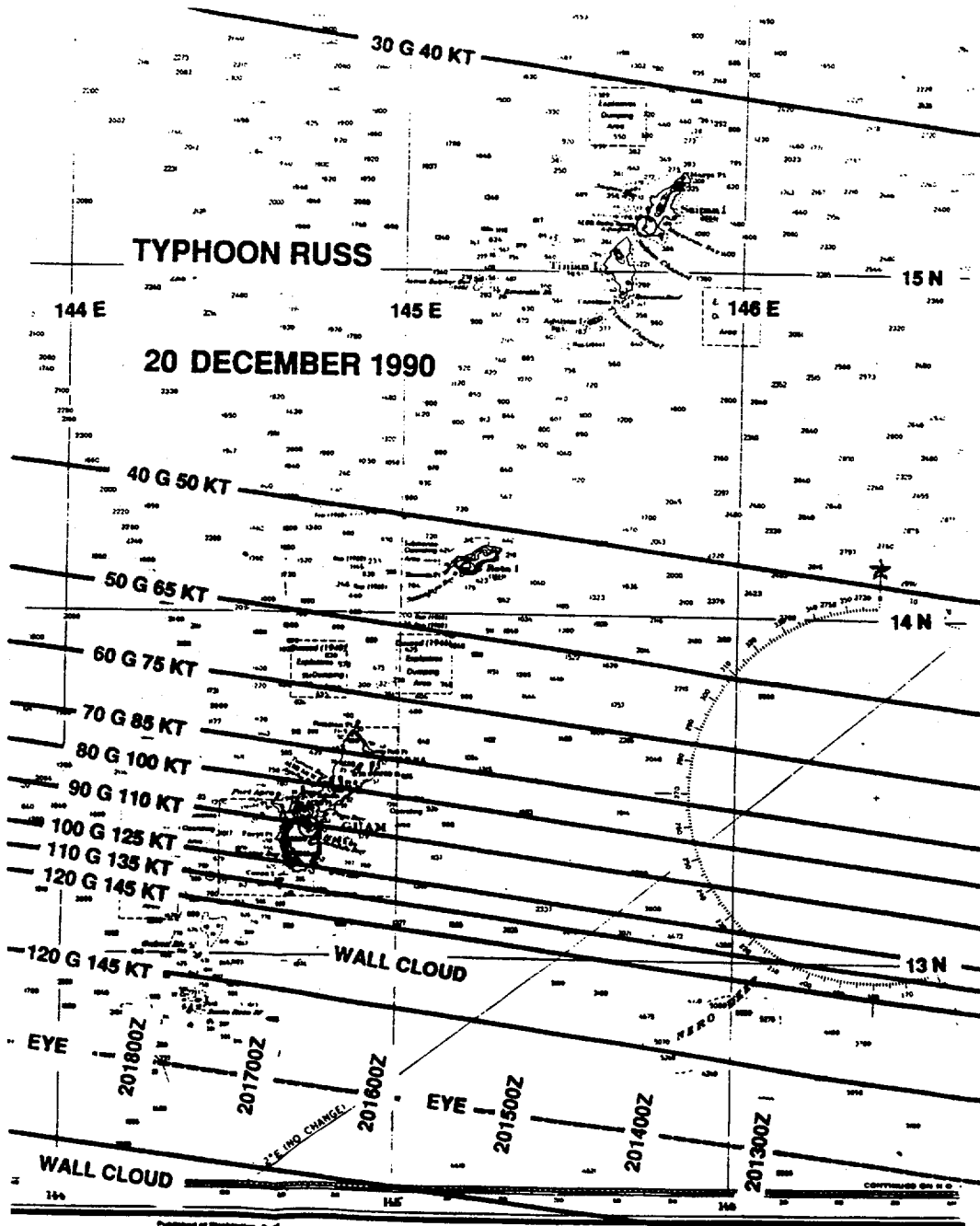


Figure 3-31-9. The post-analysis of the over water winds associated with Russ on 20 and 21 December while its track was nearest Guam. Note the rapid increase of winds near the eye wall.

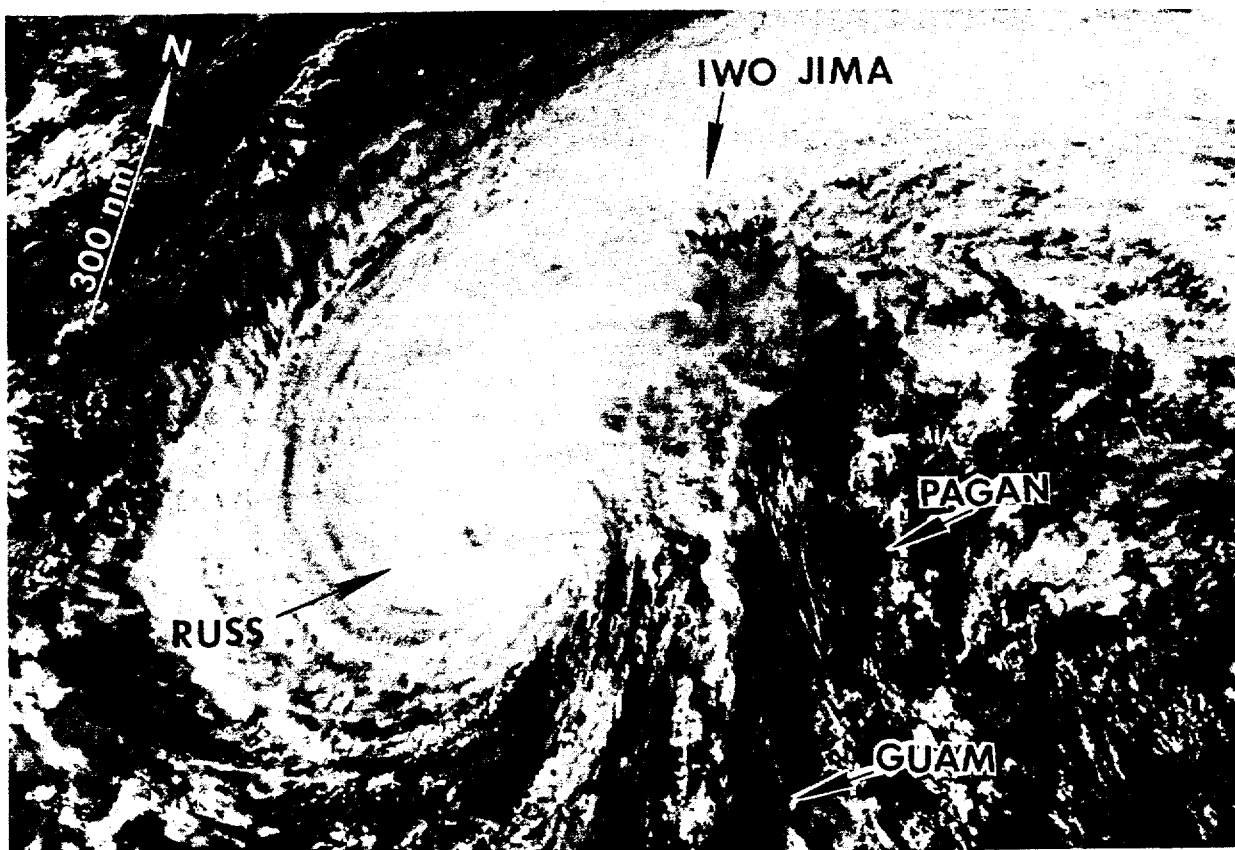


Figure 3-31-10. As Russ starts to move northward, it interacts with the polar westerlies aloft. The eye is still present in a compact central dense overcast, but the typhoon is weakening (220445Z December NOAA visual imagery).

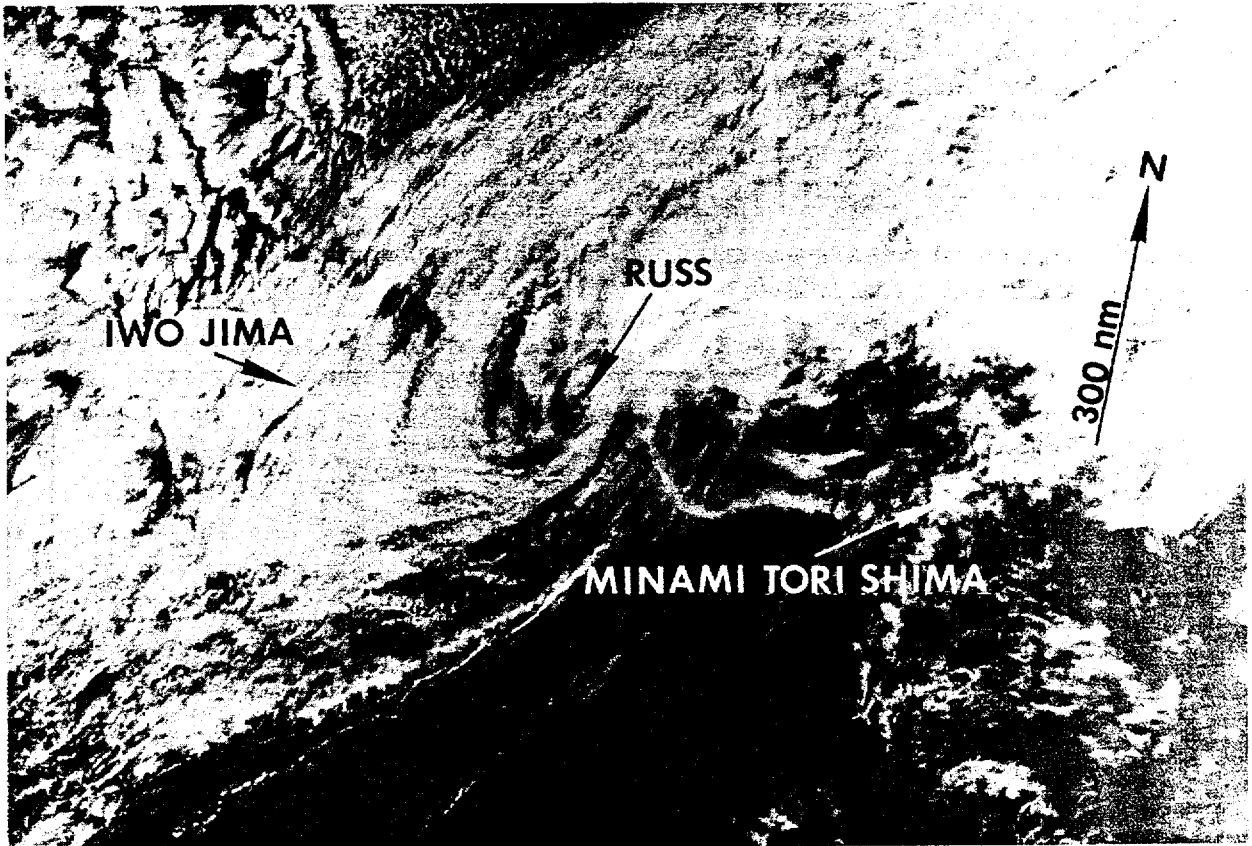


Figure 3-31-11. With the central dense overcast gone, Russ' low-level circulation center is exposed (232302Z December DMSP visual imagery).

V. FORECASTING PERFORMANCE

Overall JTWC forecast performance is shown in Figure 3-31-12. The clustering of the forecasts about the best track indicated that JTWC had a good handle on Russ' direction of motion. The mean cross track (direction) error was roughly one half the magnitude of the along track (speed) errors. This larger mean along track (speed) error was due to problems forecasting slowing and acceleration of Russ east of Guam and its acceleration after recurvature.

Russ influenced JTWC's operations. However, the day before Russ arrived JTWC had anticipated that damage might occur and had transferred all its tropical cyclone data files to the Alternate Joint Typhoon Warning Center (AJTWC) at Pearl Harbor, Hawaii. This transfer paid off because as Russ approached, JTWC began, after 201200Z, to lose most of its data base, including meteorological satellite imagery, analytic and prognostic fields, and objective guidance from Fleet Numerical Oceanography Center at Monterey, California. The increasing winds destroyed the geostationary satellite antenna, the polar orbiting satellite receiver lost power when the back up generator failed and off-island communications were interrupted. In addition, the Andersen AFB weather radar failed at this time, leaving the Federal Aviation Administration's air traffic control radar at Mount Santa Rosa as the only remaining on-island source of fixes. Rather than operate in a degraded mode, JTWC transferred responsibility for warnings to the AJTWC at Pearl Harbor, Hawaii, after the 201800Z warning. A half a day later, JTWC was able to resume normal operations and take the warning responsibility back from AJTWC.

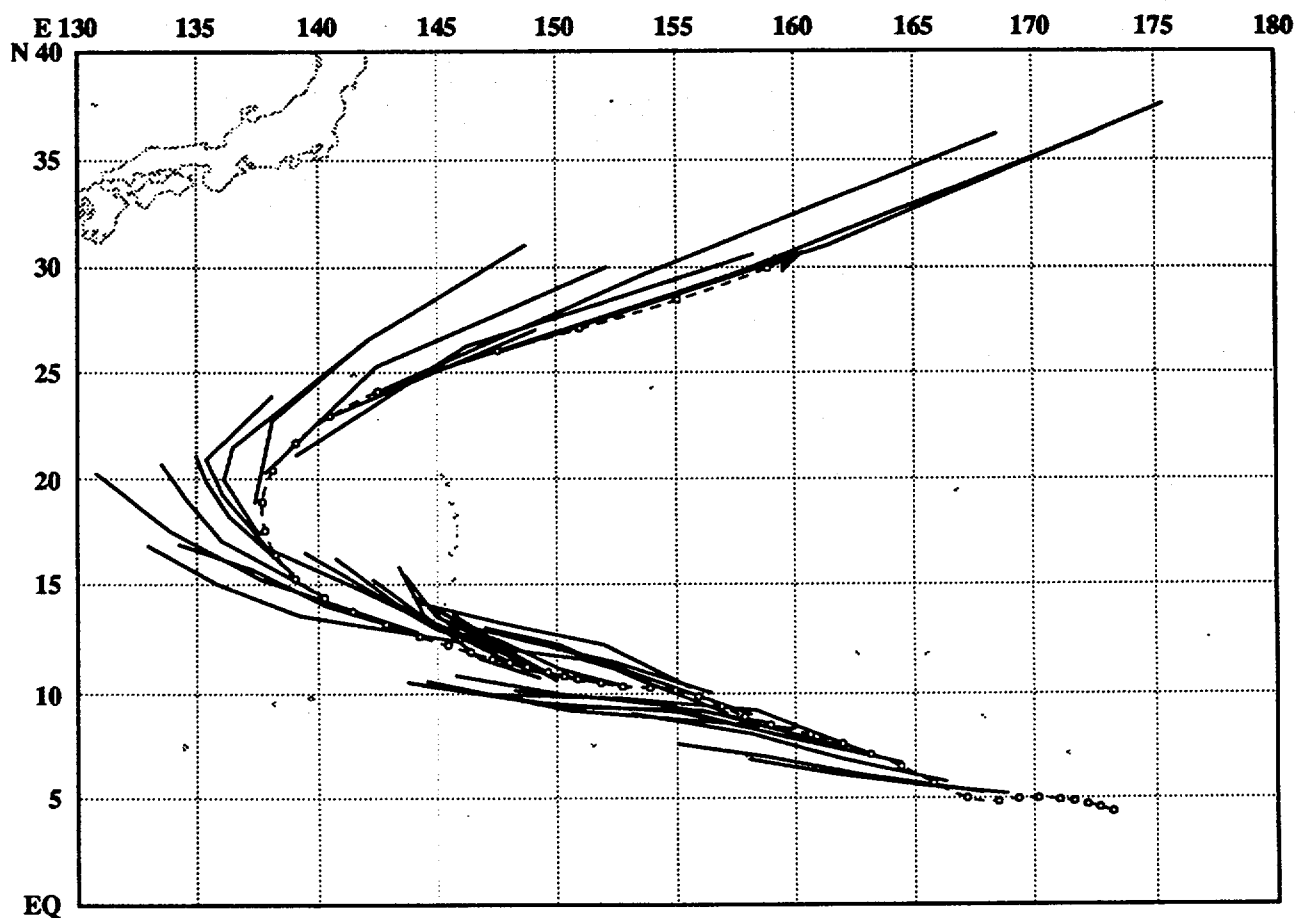


Figure 3-31-12. JTWC forecasts (solid lines) overlaid on the best track (dashed line). The clustering of forecasts shows that the general understanding of the motion toward Guam and of the recurvature track taken by Russ was good.

VI. IMPACT

Russ was the most severe tropical cyclone to hit Guam in 14 years. The island was declared a national disaster area by President Bush on 24 December, and damage estimates were as high as 120 million dollars. Miraculously, no fatalities occurred on Guam and only minor injuries were reported. This was a great credit to the disaster preparedness agencies and communications media which heightened public awareness. At sea, however, one crew member was lost from a Japanese fishing vessel that foundered southeast of Guam, and ten crew members from a South Korean fishing vessel were lost at sea after their 65 ft (20 m) boat apparently broke down south of Guam, directly in the typhoon's path.

The southern end of the Guam experienced the highest sustained winds and the most damage. Russ' winds uprooted many of the island's trees and defoliated much of the island's foliage. Two thousand houses were considered uninhabitable due to unsafe or unhealthy conditions. Of these, 341 houses were destroyed, 460 suffered major damage, and 1210 suffered minor damage. In addition, 10% of the island's total structures sustained some damage. Russ also left most of the island without power and water for several days. On the southern end of Guam, many residences were without power (Figure 3-31-13) and water for more than one week; some experienced outages for several weeks. Most telephones remained in service throughout the typhoon; however, the cable TV network sustained extensive damage. In some place on the southern and southeastern end of the island the combination of storm surge and wave run-up reached levels of 8 to 9 ft (2 to 3 m) above normal and extended inland 240-300 ft (75-90 m). For Guam, Russ was a relatively dry typhoon because the eye wall with its torrential rains passed just to the south, and rain bands were oriented north-south allowing the heavy rain to pass rapidly across the island. Thus, the inhabited part of the island was spared extensive flooding and additional damage.

An estimated 20 million dollars damage was done to civilian housing and 5 million dollars to the infrastructure. Government buildings incurred another estimated 20 million dollars in damage, including an estimated 300,000 dollars at the Oceanview High School to replace the roofing on three classrooms and other school property. The Port Authority of Guam recorded 107,000 dollars in property damage to port service equipment, primarily generators and gantry cranes. Private businesses estimated damage at 31 million dollars. This included 28 million dollars damage to the Cocos Island Resort, located on a small island on the fringing reef at the south end of Guam. The resort will have to be completely rebuilt. In addition, two ships broke their moorings and went aground on the breakwater in Apra Harbor. One of vessels was a three masted dinner cruise ship (Figure 3-31-14); the other was a 220 ft (65 m) commercial fishing vessel. Military losses (Figure 3-31-15) were estimated at over 6.5 million dollars, including 2 million dollars to military housing. It would be months before Guam fully recovered from the fury of Typhoon Russ (Figure 3-31-16).



Figure 3-31-13. Splintered utility pole bares mute testimony to Russ' high winds and termites that never sleep. (Photo courtesy of COMNAVMAR Public Affairs/PH1 Jon Hockersmith)

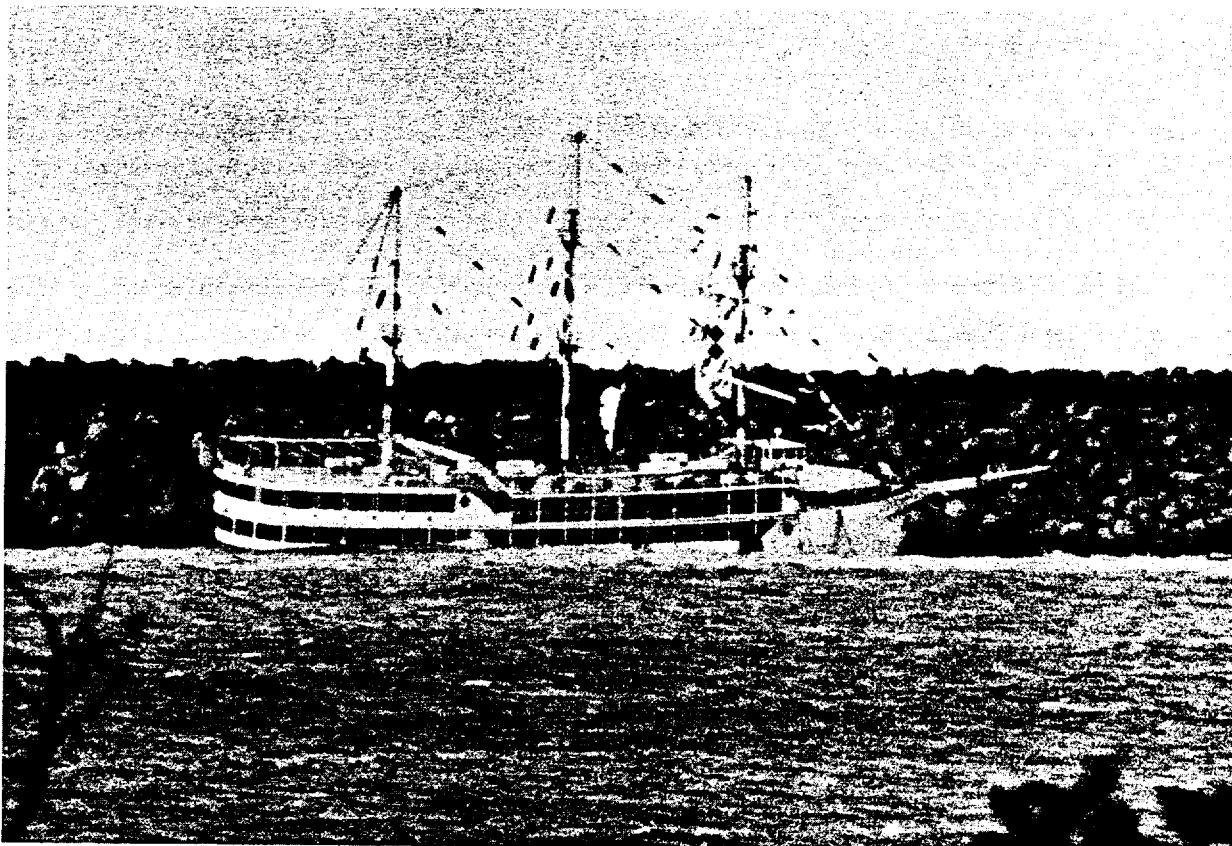


Figure 3-31-14. Dinner cruise ship Courageous aground on the Glass breakwater in Apra Harbor, Guam. (Photo courtesy of COMNAVMAR Public Affairs/PH1 Jon Hockersmith)

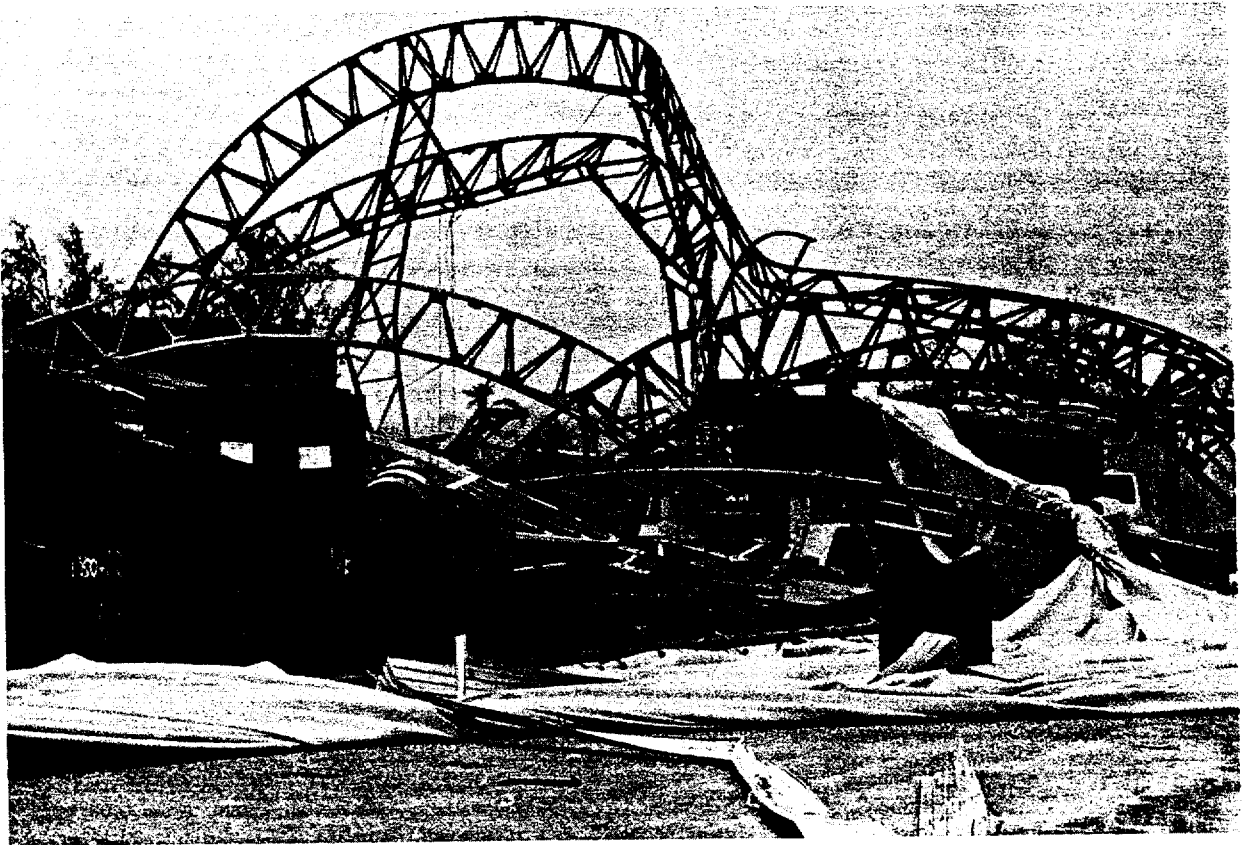


Figure 3-31-15. The steel girders of this temporary warehouse on Naval Station, Guam were twisted by the high winds and collapsed during Russ' passage.



Figure 3-31-16. Concrete power pole on the highway north of Talofofo, Guam. The pole snapped about 5 feet above the ground and fell across the road. It was later pushed out of the road to enable traffic to pass. (Photo courtesy of Det 1, 1 Weather Wing/1Lt Joe Hanser)

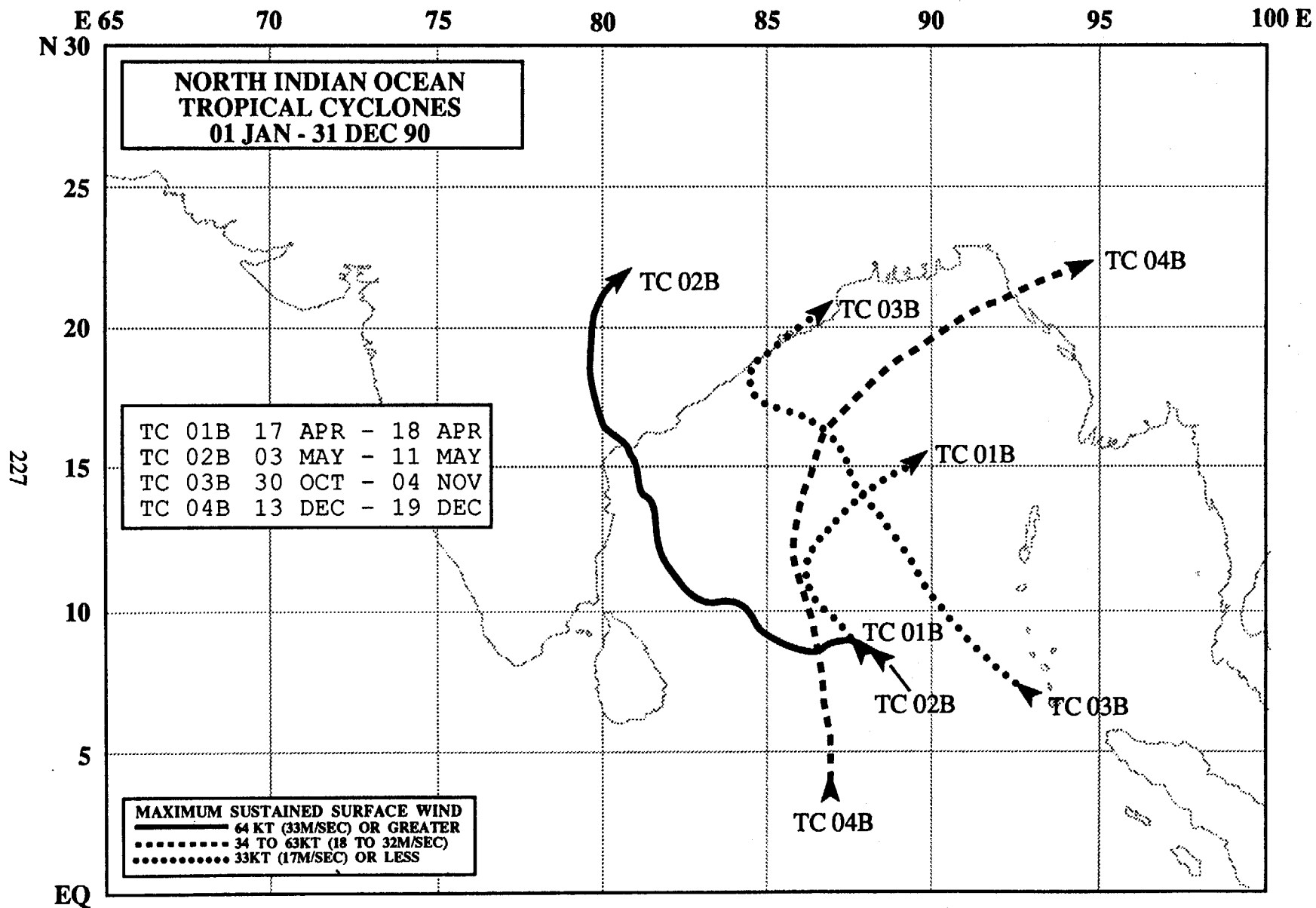
3.3 NORTH INDIAN OCEAN TROPICAL CYCLONES

Spring and fall in the North Indian Ocean are periods of transition between major climatic controls and the most favorable seasons for tropical cyclone activity (Tables 3-5 and 3-6). Two significant tropical cyclones developed in the spring and two in the fall in the Bay of

Bengal, however none occurred in the Arabian Sea. This activity was slightly below the 16-year average of five. Tropical Cyclone 02B was unusually intense - 125 kt (64 m/sec) - and like Tropical Cyclone 32W in 1989 occurred in November.

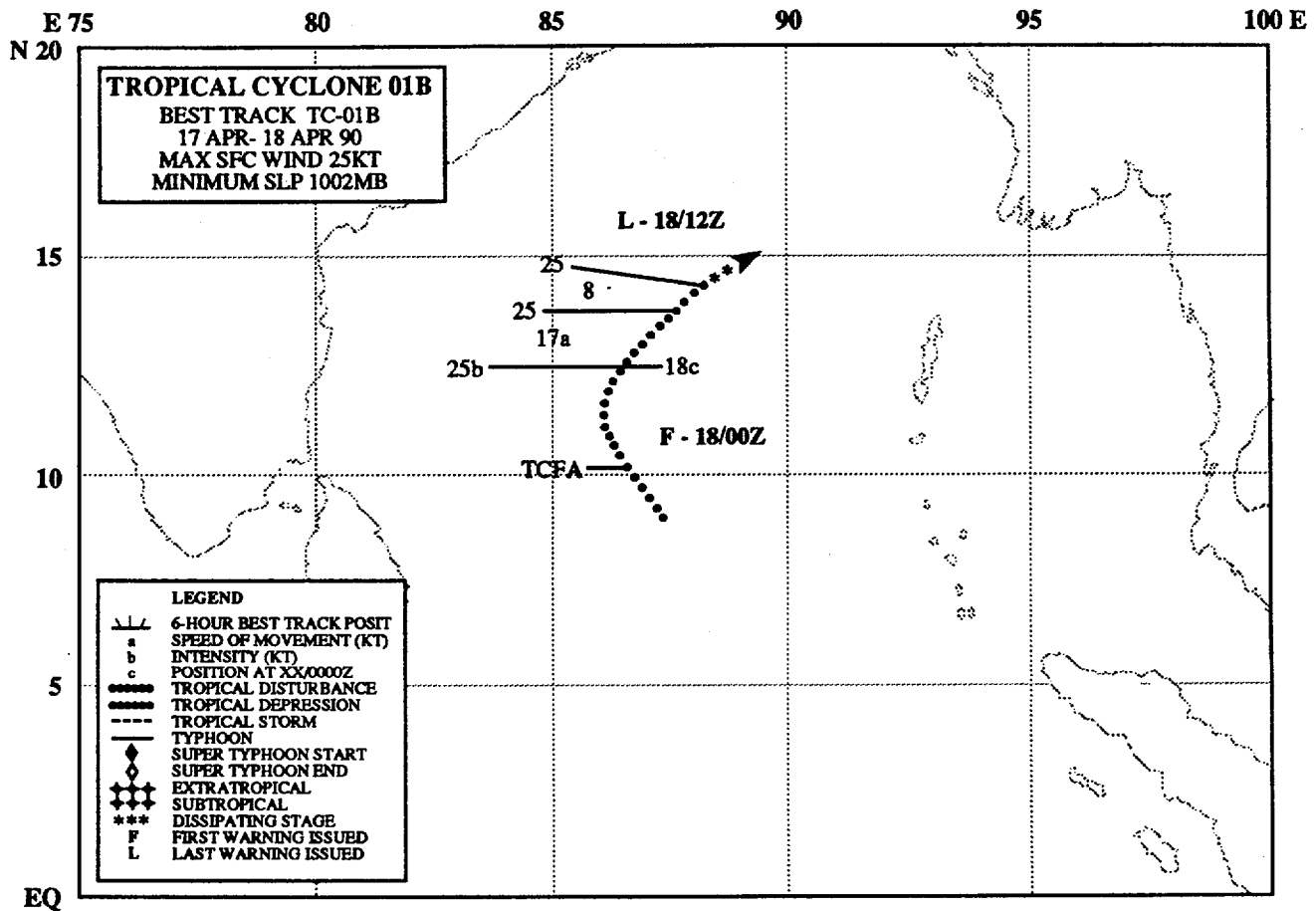
TABLE 3-5. 1990 SIGNIFICANT TROPICAL CYCLONES NORTH INDIAN OCEAN				
TROPICAL CYCLONE	PERIOD OF WARNING	NUMBER OF WARNINGS ISSUED	MAXIMUM SURFACE WINDS-KT (M/SEC)	ESTIMATED MSLP (MB)
TC 01B	18 APR - 18 APR	2	25 (13)	1002
TC 02B	05 MAY - 11 MAY	25	125 (64)	916
TC 03B	02 NOV - 03 NOV	6	30 (15)	1000
TC 04B	15 DEC - 18 DEC	14	45 (23)	991
	TOTAL:	47		

TABLE 3-6. NORTH INDIAN OCEAN TROPICAL CYCLONES DISTRIBUTION													
YEAR	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTAL
1971*	-	-	-	-	-	0	0	0	0	1	1	0	2
1972*	0	0	0	1	0	0	0	0	2	0	1	0	4
1973*	0	0	0	0	0	0	0	0	0	1	2	1	4
1974*	0	0	0	0	0	0	0	0	0	0	1	0	1
1975	1	0	0	0	2	0	0	0	0	1	2	0	6
1976	0	0	0	1	0	1	0	0	1	1	0	1	5
1977	0	0	0	0	1	1	0	0	0	1	2	0	5
1978	0	0	0	0	1	0	0	0	0	1	2	0	4
1979	0	0	0	0	1	1	0	0	2	1	2	0	7
1980	0	0	0	0	0	0	0	0	0	0	1	1	2
1981	0	0	0	0	0	0	0	0	0	1	1	1	3
1982	0	0	0	0	1	1	0	0	0	2	1	0	5
1983	0	0	0	0	0	0	0	1	0	1	1	0	3
1984	0	0	0	0	1	0	0	0	0	1	2	0	4
1985	0	0	0	0	2	0	0	0	0	2	1	1	6
1986	1	0	0	0	0	0	0	0	0	0	2	0	3
1987	0	1	0	0	0	2	0	0	0	1	2	2	8
1988	0	0	0	0	0	1	0	0	0	1	2	1	5
1989	0	0	0	0	1	1	0	0	0	0	1	0	3
1990**	0	0	0	1	1	0	0	0	0	0	1	1	4
(1975-1990)													
AVERAGE:	0.1	0.1	0.0	0.1	0.7	0.5	0.0	0.1	0.2	0.9	1.4	0.5	4.6
TOTAL:	2	1	0	2	11	8	0	1	3	14	23	8	73
* JTWC WARNING RESPONSIBILITY BEGAN ON 4 JUNE 1971 FOR THE BAY OF BENGAL, EAST OF 90° EAST LONGITUDE. AS DIRECTED BY CINCPAC, JTWC ISSUED WARNINGS ONLY FOR THOSE TROPICAL CYCLONES THAT DEVELOPED OR TRACKED THROUGH THAT PART OF THE BAY OF BENGAL. COMMENCING WITH THE 1975 TROPICAL CYCLONE SEASON, JTWC'S AREA OF RESPONSIBILITY WAS EXTENDED WESTWARD TO INCLUDE THE WESTERN PART OF THE BAY OF BENGAL AND THE ENTIRE ARABIAN SEA.													
** JTWC ISSUED EIGHT TROPICAL CYCLONE FORMATION ALERTS. FORMATION ALERTS WERE ISSUED FOR ALL OF THE SIGNIFICANT TROPICAL CYCLONES THAT DEVELOPED IN 1990.													
WARNINGS: NUMBER OF CALENDAR WARNING DAYS: 11 THERE WERE NO CALENDAR WARNING DAYS WITH TWO OR MORE TROPICAL CYCLONES.													



TROPICAL CYCLONE 01B

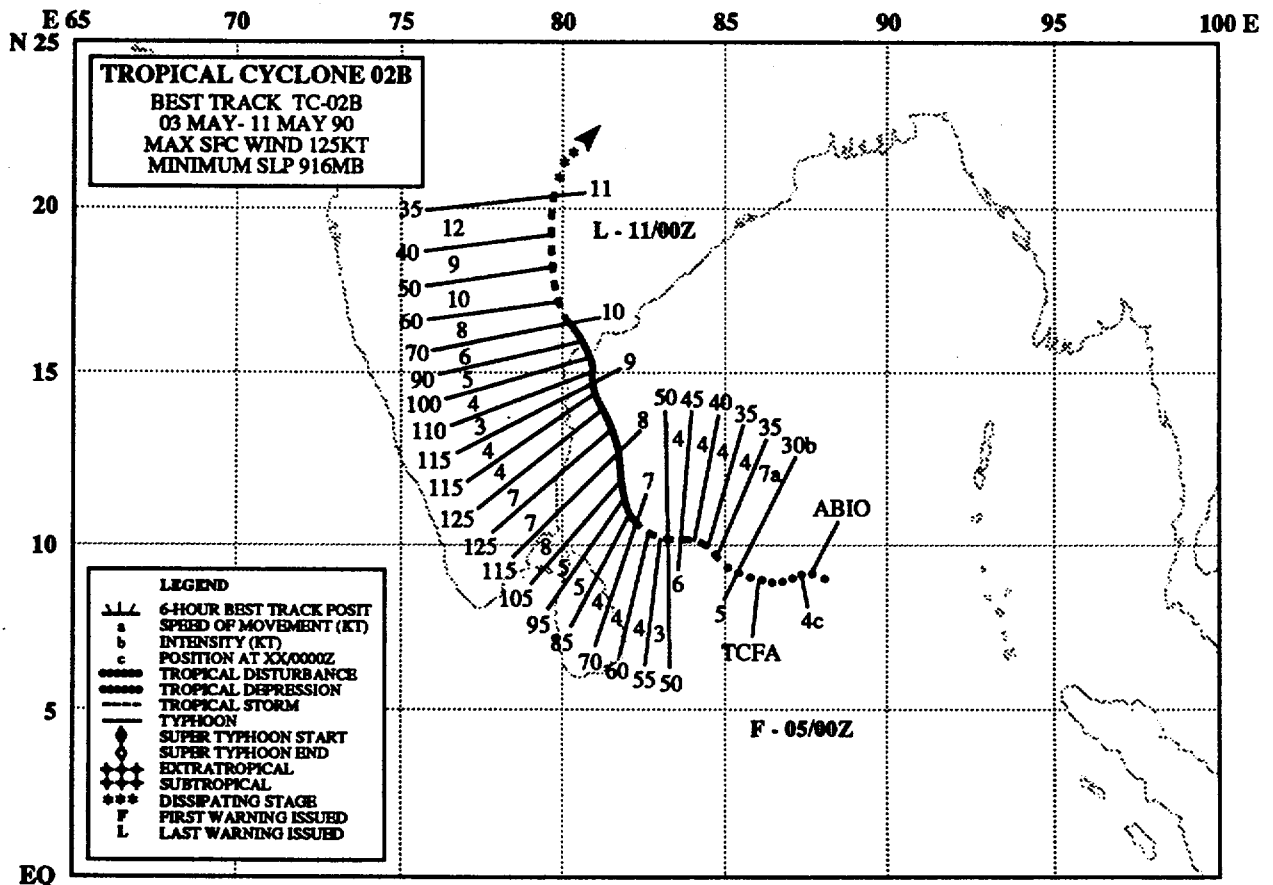
Tropical cyclone 01B was the first of two systems that formed in the Bay of Bengal during the spring transition season. It organized quite rapidly, having existed for less than 12 hours as a region of persistent convective activity, before becoming the subject of a TCFA issued at 171630Z. The brief period of northward motion followed by a short recurvature track to the northeast was related to the cyclone's formation near the axis of the subtropical ridge. The proximity of strong upper-level westerlies to the north, inhibited development beyond tropical depression intensity and brought about the rapid dissipation of 01B over water. Only two tropical depression warnings were issued on the system at 180000Z and 181200Z. Tropical cyclone 01B had no reported impact on military or civilian vessels.



TROPICAL CYCLONE 02B

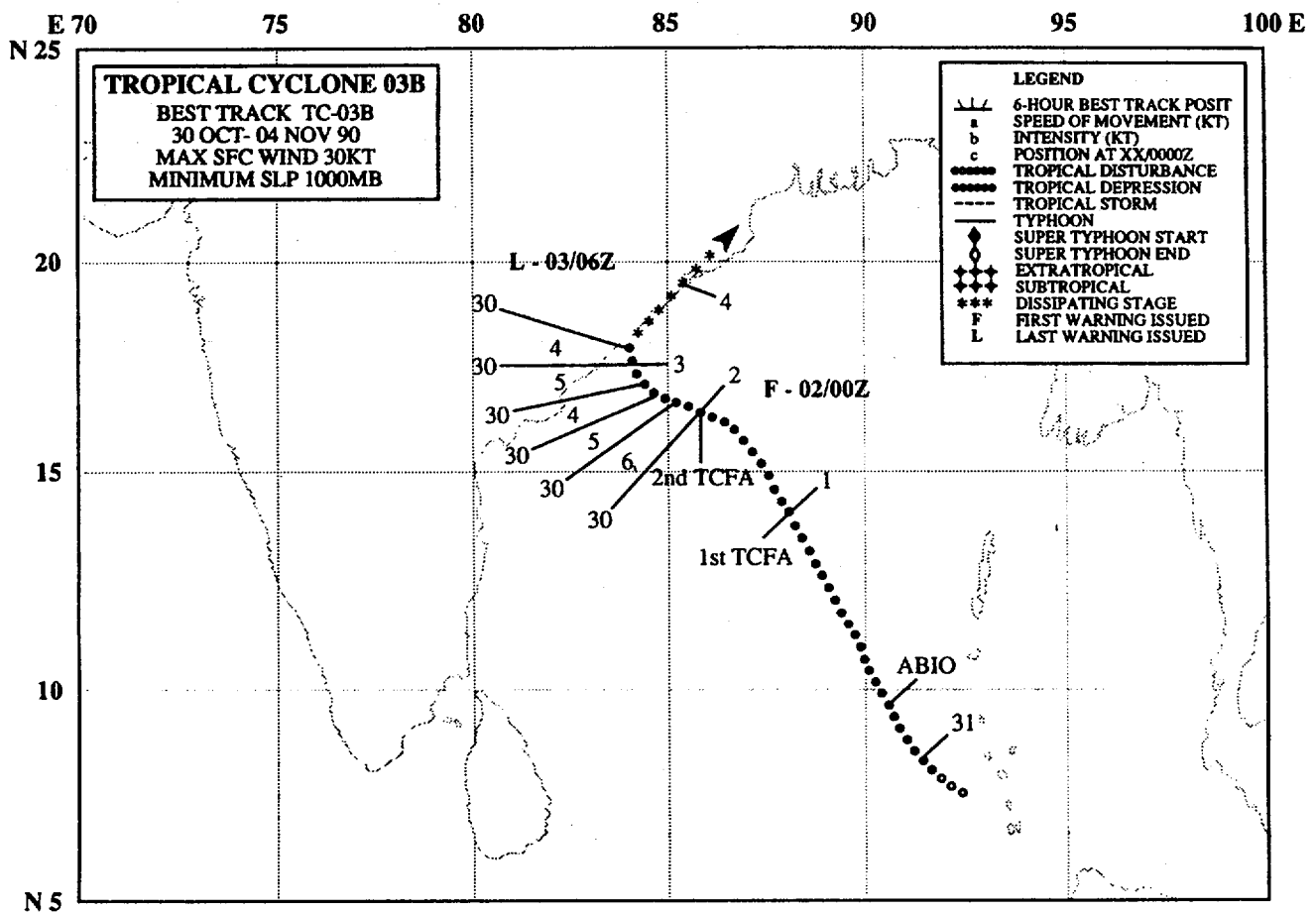
In stark contrast to its predecessor, Tropical Cyclone 01B, Tropical Cyclone 02B achieved near-super typhoon intensity, and had a major impact on the civilian populace of India. Existing as a discrete disturbance for about 36 hours before becoming the subject of a TCFA, the cyclone followed a sinuous west-northwest track under the mid-level subtropical ridge. Although TC02B had good outflow into the upper-level easterlies, landfall in southern India was expected by about 72 hours, prompting early forecasts of intensification to only nominal typhoon intensity followed by weakening due to the approach of landfall. Because of a weakness in the subtropical ridge, a moderate turn toward the northwest was expected, however the actual track change turned out to be much more northward than anticipated. This permitted the cyclone to stay off-shore and to establish strong outflow into the upper-level southwesterlies of a passing 200-mb short-wave trough. As the northward turn began, JTWC modified the intensity forecast to one of rapid deepening beginning with the 070000Z warning. The rapid deepening did in fact occur beginning at 061800Z with a 60 kt (30 m/sec) intensity and peaking at 125 kt (65 m/sec) by 080600Z: an increase of 65 kt (33 m/sec) in 36 hours. At 091200Z, the cyclone, with winds of 100 kt (50 m/sec) made landfall 165 nm (305 km) north of Madras in the vicinity of Machilipatnam in Andra Pradesh State.

In impact of this cyclone on India was substantial. An estimated 150,000 people were evacuated in preparation for landfall. Over 100 villages were destroyed resulting in at least 510 human fatalities. The cyclone also wreaked havoc on the rich agriculture industry of the region killing more than 100,000 farm animals and causing more than \$600 million in damage to crops. Local officials reported that Tropical Cyclone 02B was the worst disaster for southern India since the 1977 cyclone that killed an estimated 10,000 people.



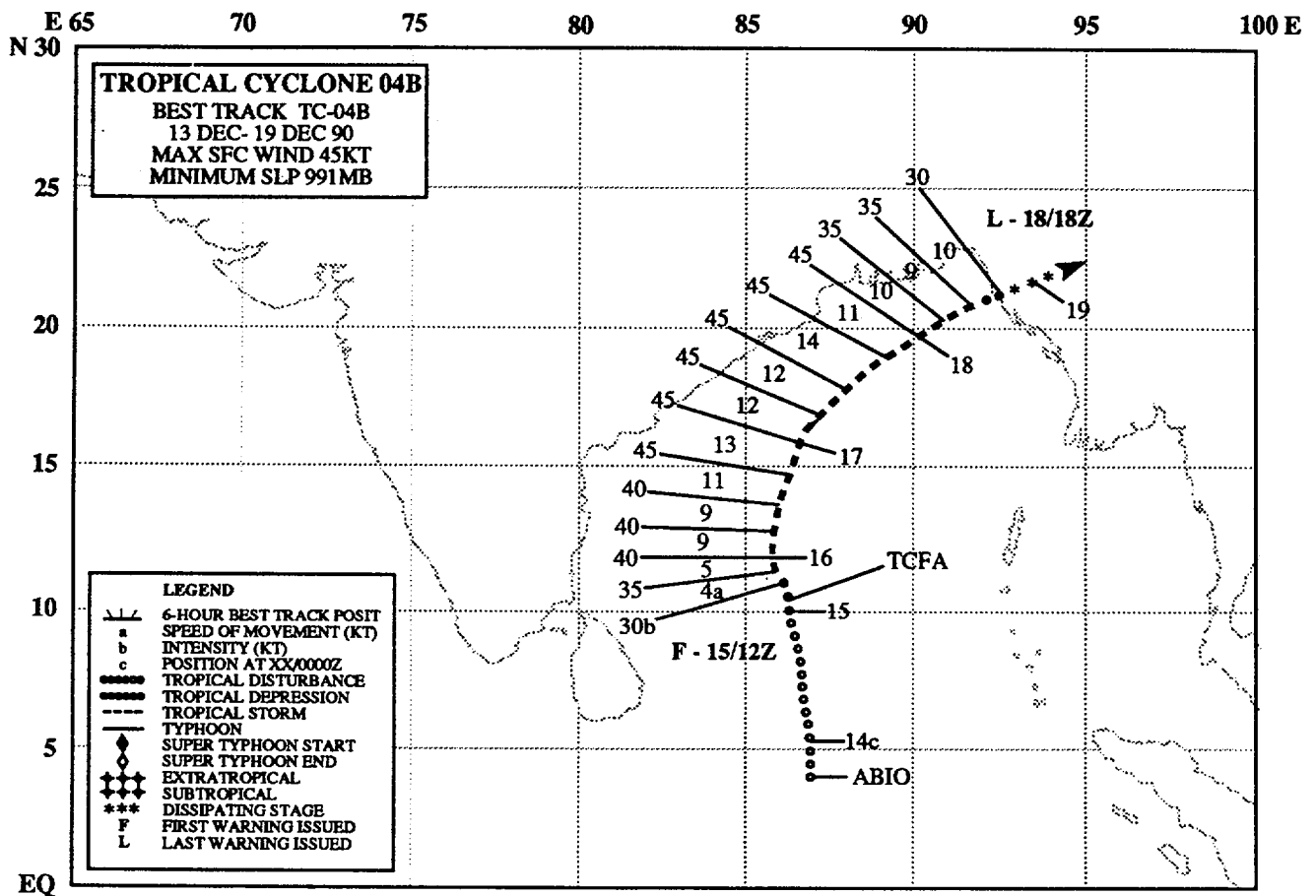
TROPICAL CYCLONE 03B

Tropical Cyclone 03B was the first of two systems that occurred the Bay of Bengal during the fall transition season. The system formed and remained under upper-level east-southeasterly wind shear associated with a 200-mb ridge circulation well to the northeast. Thus, development was strongly inhibited and TC03B never exceeded tropical depression intensity. The cyclone tracked north-northwestward along the western periphery of a broad mid-level subtropical ridge circulation centered over Indochina. It reached the axis of the ridge as it made landfall, then skirted the coast of India on a northeastward track as it dissipated. No reports of impact were received.



TROPICAL CYCLONE 04B

Tropical Cyclone 04B, the final system for the North Indian Ocean for the year, formed just to the south of a col in the mid-level subtropical ridge that typically extends across the Bay of Bengal between semi-permanent ridge circulations over Indochina and the Northeast African/Arabian Sea region. As a result, the cyclone tracked through the break in the subtropical ridge and followed a recurvature track that resulted in landfall in the area between Bangladesh and Burma. The development of TC04B into a significant tropical cyclone coincided with its movement into an area of relatively weak upper-level winds, however further intensification was restricted to a maximum of moderate tropical storm intensity. As the cyclone moved northeastward, it began to encounter increasing upper-level wind shear associated with the mid-latitude westerlies and weakened by the time landfall occurred. No reports of impact were received.



Intentionally left blank.